

*Citation for published version:*

Liu, Z, Koh, KL, Mezentsev, A, Enno, SE, Sugier, J & Fullekrug, M 2018, 'Lightning Sferics: Analysis of the Instantaneous Phase and Frequency Inferred From Complex Waveforms', *Radio Science*, vol. 53, no. 4, pp. 448-457. <https://doi.org/10.1002/2017RS006451>

*DOI:*

[10.1002/2017RS006451](https://doi.org/10.1002/2017RS006451)

*Publication date:*

2018

*Document Version*

Peer reviewed version

[Link to publication](#)

This is the peer-reviewed version of the following article: 2017RS006451 Liu, Z, Koh, KL, Mezentsev, A, Enno, SE, Sugier, J & Fullekrug, M 2018, 'Lightning Sferics: Analysis of the Instantaneous Phase and Frequency Inferred From Complex Waveforms' *Radio Science*, which has been published in final form at: <https://doi.org/10.1002/> This article may be used for non-commercial purposes in accordance with Wiley Terms and Conditions for Self-Archiving.

## University of Bath

**General rights**

Copyright and moral rights for the publications made accessible in the public portal are retained by the authors and/or other copyright owners and it is a condition of accessing publications that users recognise and abide by the legal requirements associated with these rights.

**Take down policy**

If you believe that this document breaches copyright please contact us providing details, and we will remove access to the work immediately and investigate your claim.

1                   **Lightning Sferics: Analysis of the Instantaneous Phase and**  
2                   **Frequency Inferred from Complex Waveforms**

3  
4   **Zhongjian Liu<sup>1</sup>, Kuang Liang Koh<sup>1</sup>, Andrew Mezentsev<sup>2</sup>, Sven-Erik Enno<sup>3</sup>, Jacqueline**  
5   **Sugier<sup>3</sup>, and Martin Fullekrug<sup>1</sup>**

6  
7  
8  
9  
10  
11  
12  
13   <sup>1</sup> Centre for Space, Atmospheric and Oceanic Science, University of Bath, Bath, United  
14   Kingdom.

15   <sup>2</sup> Birkeland Centre for Space Science, University of Bergen, Bergen, Norway.

16   <sup>3</sup>Remote Sensing (Lightning, Clouds, Aerosols), and Aircraft Based Observation R&D, Met  
17   Office, Exeter, United Kingdom.

18  
19   Corresponding author: Zhongjian Liu ([z.liu@bath.ac.uk](mailto:z.liu@bath.ac.uk))

20  
21  
22   **Key Points:**

- 23       • A complex waveform bank and a spectral waveform bank has been produced from  
24       lightning sferics.
- 25       • Long-range lightning location systems achieve sub-sampling time accuracy using the  
26       instantaneous phase of complex lightning sferic waveform.
- 27       • The instantaneous frequencies at maximum amplitudes are distance dependent.
- 28

**29 Abstract**

30 Analysis of VLF lightning waveforms, or radio sferics, can contribute to research into lower  
31 ionosphere perturbations and the corresponding atmospheric chemistry. Lightning waveforms  
32 can also be characterized on the basis of their propagation distance from receivers in order to  
33 study radio wave propagation. A bank of average waveforms, i.e. the waveform bank, <1,000  
34 km with a spatial resolution of 10 km has been produced, based on the lightning waveforms  
35 recorded in Europe on 8<sup>th</sup> August 2014. These average lightning waveforms at different dis-  
36 tances exhibit a sequence of consecutive maxima resulting from ionospheric reflections,  
37 named sky waves. The spectral waveform bank shows a sequence of consecutive modal max-  
38 ima at different frequencies depending on distance. The Hilbert transform is applied to pro-  
39 duce complex lightning waveforms, which provide additional information to the original real  
40 waveforms alone, i.e., the instantaneous phase and frequency. The time differences calculated  
41 from the instantaneous phases of complex lightning waveforms gives the minimum arrival  
42 time difference error when compared to other analyzed signal processing methods. The de-  
43 rivative of the instantaneous phase, i.e. the instantaneous frequency, represents the amplitude-  
44 weighted average of frequency components at maximum amplitude according to theory and  
45 numerical simulation. In real experiments, the instantaneous frequency can be understood as  
46 the median value of the real frequency distribution calculated at maximum amplitude. It is  
47 found that the instantaneous frequencies at maximum amplitudes are distance dependent.  
48 This finding might enable the development of a novel method to determine lightning distanc-  
49 es in the future.

**50 1 Introduction**

51 Lightning is the strongest natural electromagnetic radiation source in the earth's atmosphere  
52 and it emits electromagnetic waves in the frequency range from ~1 Hz to >300 MHz [e.g.,  
53 *Rakov and Uman, 2003; Rison et al., 2016*]. Ground-based lightning location systems are es-  
54 sentially based on the received electromagnetic waves and subsequent signal processing [e.g.,  
55 *Dowden et al., 2002; Holler et al., 2009; Bitzer et al., 2013; Rakov, 2013; Stock et al., 2014;*  
56 *Lyu et al., 2014; Nag et al., 2015; Wang et al., 2016; Sun et al., 2016*]. For example, the azi-  
57 muth of an incident radio wave from a lightning discharge is determined by calculating the  
58 voltage ratio received by two magnetic loop antennas, named magnetic direction finding  
59 (MDF) [e.g., *Horner, 1954, 1957; Krider et al., 1976; Fullekrug et al., 2000*]. In the most  
60 commonly used lightning location method, arrival time differences (ATD) are extracted by  
61 different signal processing techniques from the lightning waveforms recorded at different ra-  
62 dio receiver stations [e.g. *Lee, 1986, 1990; Cummins et al., 1998, 2009*]. The different signal  
63 processing techniques result in slightly different arrival times with different corresponding  
64 lightning locations [*Liu et al., 2016*]. Using a different time extraction point, e.g. the wave-  
65 form peak or rising edge, and different data pre-processing, e.g. wide or narrow bandwidth or  
66 complex envelope, can all introduce slightly varying time differences. As a result, a better  
67 understanding of the sferic, i.e. the received broadband lightning waveform, is a prerequisite  
68 for improving lightning location accuracy.

69 Much research into ionosphere perturbations and atmospheric chemistry utilize broadband  
70 VLF waveforms, in particular lightning waveforms [e.g. *Cheng et al., 2007; Shao et al.,*  
71 *2013*]. The sferic is a broadband electromagnetic impulse generated by a lightning discharge,  
72 which propagates through the earth-ionosphere waveguide. The perturbations in the D-region

73 caused by the variation of electron densities will change the received sferic in the time and  
74 frequency domain. A collection of average lightning waveforms at different distances, named  
75 waveform bank, was introduced by Said et al. [2010]. This method offers an opportunity to  
76 characterize sferics and to estimate arrival times. This idea has been adopted and it is extend-  
77 ed in this work to produce a new type of waveform bank for further analysis.

78 The time-dependent sferic signal is treated as an analytic signal, or complex trace, to process  
79 the received signal to determine the instantaneous phase and frequency at each sample in the  
80 time series [e.g. *Taner et al., 1979*]. It is more common to use the Fast Fourier Transform  
81 (FFT), or a Short-Time Fourier Transform (STFT), to calculate phases from the spectral coef-  
82 ficients at the harmonic frequencies  $f_k = k/T$ , where  $k$  is the harmonic number and  $T$  is the  
83 length of the time interval used for the FFT or STFT. As a result, the phases inferred from the  
84 FFT or STFT represent average values over the time interval  $T = N\Delta t$ , where  $N$  is the number  
85 of samples in that time interval and  $\Delta t$  is the sampling time interval. Given that the phase of a  
86 lightning sferic is constantly changing (*Fullekrug et al., 2016, Fig 2, right*), an averaged phase  
87 can include unwanted information that is introduced, for example, by the phase change dur-  
88 ing the time interval  $T$ , or by unwanted interference during that time interval. It is therefore  
89 interesting to investigate how useful the instantaneous phase of a lightning sferic is which is  
90 determined for each individual sample and which should have the smallest possible bias in-  
91 troduced by unwanted information. Therefore, this method will be applied here to a new  
92 lightning waveform bank (<1000 km) produced from data collected by a long-range lightning  
93 location system described below. The timing accuracy of the instantaneous phase is com-  
94 pared with other methods by using the speed of light as a reference. The derivative of the in-  
95 stantaneous phase is the instantaneous frequency, which is used to study the relationship with  
96 the real frequency of the sferic determined by its spectrum. The observed instantaneous fre-  
97 quency changes in the complex waveform bank are discussed and related to the radio wave  
98 propagation with distance.

## 99 **2 The Complex Waveform Bank**

### 100 **2.1 The waveform bank**

101 Previous research associated with sferic waveform characteristics, including polarity estima-  
102 tion, cycle errors and peak current, has shown that received lightning waveforms that origi-  
103 nate from a certain storm cluster exhibit similar features [*Said et al., 2010, Figure 1*]. Thus, a  
104 small deviation of lightning waveform shape indicates a propagation over similar distances.  
105 A representative waveform is calculated by averaging each of the lightning waveforms in one  
106 distance bin to reduce the noise in each average waveform and to acquire a more pure light-  
107 ning waveform that includes subtle propagation effects. Comparing the representative wave-  
108 forms at different distances is an effective method to study the propagation effects on light-  
109 ning sferics.

110 Lightning locations are reported by the French lightning detection network *Meteorage*, which  
111 covers south-western Europe and the western Mediterranean Sea. This network distinguishes  
112 cloud to ground lightning and intra-cloud lightning based on a linear discriminant analysis  
113 (LDA) of different waveform parameters, and the classification accuracy reaches 95% for  
114 negative cloud to ground lightning in France [*Kohlmann et al., 2017*]. The electromagnetic  
115 waveforms of the lightning discharges were recorded from ~21:00 UT 8<sup>th</sup> to ~03:00 UT on 9<sup>th</sup>  
116 August 2014 with four wideband digital low-frequency radio receivers located in Bath

117 (BTH), Orleans (ORL), Lannemezan (LMZ), and Rustrel (RST) [Liu *et al.*, 2016]. The radio  
118 receiver records the electric field from a capacitive probe with a sampling frequency of 1  
119 MHz, an effective bandwidth from ~4 Hz to ~400 kHz and a timing accuracy of ~12 ns pro-  
120 vided by a GPS disciplined frequency standard [e.g. Fullekrug, 2010,; Fullekrug *et al.* 2006,  
121 2014, 2015; Mezentsev *et al.*, 2013; Soula *et al.*, 2014].

122 During the 7 hour long recording, more than 150,000 cloud to ground (CG) lightning stroke  
123 waveforms from a mesoscale convective system over central France were recorded by the  
124 four sensors with peak currents ranging from -4 kA to -40 kA, which were located 0-1000 km  
125 away from each radio receiver [Liu *et al.*, 2016]. The lightning waveforms ranging from -1  
126 ms to +4 ms around the occurrence of the lightning discharges were extracted from the digital  
127 recordings based on the lightning locations and occurrence times reported by Meteorage. The  
128 time  $t=0$  of the lightning waveforms is referenced to the propagation time at the speed of light  
129 calculated from the great circle distance between the source and the receiver. This referencing  
130 procedure enables the calculation of one average waveform for each distance bin with a width  
131 of 10 km. Each distance bin consists of at least one hundred lightning waveforms. The result-  
132 ing 100 average waveforms form the lightning waveform bank form the basis for the subse-  
133 quent data analysis (Figure 1, left). The difference between an average lightning source cur-  
134 rent and the interfering signals from the local noise environment, man-made noise and radio  
135 transmissions are thereby minimized. The ionospheric conditions for the propagation of these  
136 lightning waveforms are similar, because all of the lightning sferic waveforms propagated  
137 during nighttime from the source to the receiver.

138 The waveform bank exhibits an initial pulse from the ground wave and a sequence of subse-  
139 quent maxima from the ionospheric reflections or multi-hop sky waves. The first arrival sky  
140 wave can be observed after ~100 km propagation and additional sky wave hops can be ob-  
141 served for longer distance propagation. The energy of the lightning signals propagated along  
142 the ground path attenuate with distance due to ground conductivity, while the energy of light-  
143 ning signals from the ionospheric reflections are attenuated by longer propagation distances  
144 and the ionospheric D-layer conductivity. The waveform bank shows distance-dependent ar-  
145 rival times of the ground wave and ionospheric reflections, which can be explained by ray  
146 theory [e.g. Schonland *et al.*, 1940; Carvalho *et al.*, 2017; Qin *et al.*, 2017]. The waveform  
147 bank shows the ground wave arrival at 0 ms and the sky wave arrivals at increasing time de-  
148 lays for shorter distances, which is in agreement with theoretical calculations that use a flat-  
149 earth model or a spherical earth model [Schonland *et al.*, 1940].

150 A spectral waveform bank can be calculated by averaging the complex spectra of all wave-  
151 forms at the same distance bin or from the spectra of the average lightning waveforms at each  
152 distance bin. Both averaging procedures produce exactly the same result (Figure 1, right).  
153 The spectral waveform bank exhibits a sequence of consecutive maxima in the frequency  
154 range up to 100 kHz. These consecutive relative maxima result from the constructive super-  
155 position of numerous wave propagation modes, named modal maxima in the following text.  
156 The modal maxima are separated from each other by distinct spectral minima that are charac-  
157 teristic for the distance between the lightning discharges and the radio receivers. These fea-  
158 tures of the spectral waveform bank result mainly from the lightning sferics and propagation  
159 effects because averaging the waveforms eliminates most of the interfering noise such that  
160 the best possible average lightning sferic waveform is obtained. The spectral waveform bank

161 is useful for understanding the propagation effects with distance and it can be used for theo-  
162 retical modeling of radio wave propagation.

## 163 **2.2 Complex Waveforms**

164 The time dependent radio signal can be treated as an analytic signal or complex trace. This  
165 allows the extraction of the envelope and instantaneous phase for each sample [*e.g. Taner et*  
166 *al., 1979; Liu et al., 2016*]. The complex trace can be obtained from the real valued record-  
167 ings using the Hilbert transform. In practice, the complex waveform can be calculated from  
168 the real signal by doubling the positive frequency and by eliminating the negative frequency.  
169 This complex waveform is subsequently down converted by multiplying with a frequency  
170 shift operator  $e^{-j\Delta\omega t}$  that centers the spectrum at zero frequency. The shift frequency  $\Delta\omega$  is  
171 normally set as the harmonic frequency which contains most of the energy of the target sig-  
172 nal. For example, the shift frequency of a radio transmission is normally the center frequency  
173 of its modulation. The shift frequency of lightning can be set as 10 kHz as the return stroke  
174 deposits most of its energy around this frequency [*Fullekrug et al., 2013*]. The final complex  
175 waveform can be determined after applying a low-pass filter to the down converted signal in  
176 order to increase the signal to noise ratio

$$F(t) = \left( f(t) + jH(f(t)) \right) e^{-j\Delta\omega t} = A(t)e^{j\varphi(t)} \quad (1)$$

177 where,  $A(t)$  is the time dependent amplitude envelope and  $\varphi(t)$  is the time dependent instantaneous  
178 phase.

179 In this way, the complex waveform bank can be calculated from a real waveform bank by use  
180 of the Hilbert transform. The complex waveform from 2-18 kHz of an average lightning sig-  
181 nal at 300 km is shown in figure 2. This three-dimensional trace of a lightning waveform il-  
182 lustrates the amplitude envelope and instantaneous phase variation over time. In the begin-  
183 ning, the phase of the complex trace is chaotic and the amplitude is low when there is no  
184 lightning signal. The arrival of ground wave and sky waves lead to large signals, which are  
185 well above the noise. These large signals result from a strong lightning discharge, so that the  
186 amplitude is increased and the phase is moved towards a specific value. In this example, the  
187 complex waveform rotates anticlockwise, which means that the instantaneous phase is de-  
188 creasing during the pulse arrival. The signal returns to chaotic behavior after the lightning  
189 event.

## 190 **3 Instantaneous Phase**

191 Different information can be extracted from the complex waveform than from the real-valued  
192 signal, including the envelope of the complex trace and the instantaneous phase. The time  
193 differences calculated from the instantaneous phase between two lightning waveform peaks  
194 can be extracted by use of the transfer function calculated from the ratio between the two  
195 complex values at the peak of the waveform. Normally, the time difference between two  
196 lightning waveforms is constrained by the sampling frequency. In order to achieve the sub-  
197 sample time difference  $\delta t = \delta\varphi/\omega$ , the two waveforms are time shifted first in order to su-  
198 perpose the amplitude peaks. The instantaneous phase difference  $\delta\varphi$  can then be extracted  
199 from the transfer function at the peak samples of the two waveforms  $\delta\varphi = \varphi_{\max 1} - \varphi_{\max 2}$ .  
200 This time difference inferred from the instantaneous phase achieves sub-sampling accuracy  
201 after taking into account cycle ambiguities, if any.

202 The propagation time delay relative to the speed of light using different signal processing  
 203 methods can be compared by using the lightning waveform bank. The average lightning  
 204 waveforms from 310 km to 600 km are compared with the average waveform at 300 km by  
 205 using four different time extraction methods (Figure 3). The time offset  $\delta t$  is the measured  
 206 time difference  $\Delta t$  minus the nominal time difference calculated for a propagation velocity at  
 207 speed of light  $\Delta t = \Delta d/c$ , where  $\Delta d$  is the distance difference with respect to the lightning  
 208 spheric waveform at 300 km distance, i.e.,  $\Delta d = d - 300$  km with  $d$  ranging from 310 km to 600  
 209 km. All these methods determine the time offset that is measured with reference to the propa-  
 210 gation time between the source and the receiver at the speed of light. The first three methods  
 211 extract the occurrence times by waveform cross correlation, from the peak of the filtered data  
 212 (5-15 kHz), and from the peak envelope of the complex trace. The average values of the ab-  
 213 solute time offsets over all distance differences from 10-300 km with respect to the speed of  
 214 light by using these three methods are  $\sim 2.33$   $\mu\text{s}$ ,  $\sim 2.63$   $\mu\text{s}$  and  $\sim 4.23$   $\mu\text{s}$ , respectively, with  
 215 corresponding ranges of  $[-3, 6]$   $\mu\text{s}$ ,  $[-4, 5]$   $\mu\text{s}$  and  $[-8, 11]$   $\mu\text{s}$ . The average value of the abso-  
 216 lute time offset by measuring time differences with the instantaneous phase is only  $\sim 2.08$   $\mu\text{s}$   
 217 with a corresponding range  $[-2.78, 5.62]$   $\mu\text{s}$  (Figure 3). In the case of a lightning location sys-  
 218 tem using a propagation velocity at the speed of light, the time accuracy is therefore slightly  
 219 improved by calculating the instantaneous phase from the complex waveform. The time off-  
 220 sets at different distances are neither constant nor distance dependent, which is suggested to  
 221 be investigated in future analyses. One possible reason is the inappropriate presupposed wave  
 222 propagation velocity in comparison, i.e. the speed of light, because it is found that the light-  
 223 ning electromagnetic wave may propagate at a varying phase propagation velocity [Liu et al.,  
 224 2016].

#### 225 **4 Instantaneous Frequency**

226 The different rotation direction of complex waveforms may indicate a different elevation an-  
 227 gles of the incident lightning spheric [Fullekrug et al., 2016, Figure 2, right]. The rotation direc-  
 228 tion is the polarity of the derivative of the instantaneous phase in a complex waveform. This  
 229 time dependent derivative  $f_i = d\varphi/dt$ , is called the instantaneous frequency [Taner et al.,  
 230 1979]. The instantaneous frequency indicates the phase change for each sample with refer-  
 231 ence to the center frequency of the signal. The rotation direction is anticlockwise when the  
 232 instantaneous frequency is smaller than the center frequency and it is clockwise when the in-  
 233 stantaneous frequency is larger than the center frequency. The benefit of using instantaneous  
 234 frequency for range estimation is the high time resolution and also the frequency resolution.  
 235 For example, the frequency resolution of the result from the STFT of a 1ms long time interval  
 236 is only 1 kHz, without zero padding, while the instantaneous phase has a temporal resolution  
 237 with the sampling time interval  $\Delta t = 1$   $\mu\text{s}$ . The calculated result has no constraints on the fre-  
 238 quency resolution within the passband of the low-pass filter applied to the down converted  
 239 signal. A convenient way of computing the instantaneous frequency  $f_i$  from the complex trace  
 240  $C$  is to compute the derivative of the arctangent function

$$f_i = \frac{1}{2\pi} \frac{d}{dt} \arctan\left(\frac{\text{Im}(C)}{\text{Re}(C)}\right), \quad (2)$$

241 which results in

$$f_i = \frac{\operatorname{Re}(C) \frac{d(\operatorname{Im}(C))}{dt} - \operatorname{Im}(C) \frac{d(\operatorname{Re}(C))}{dt}}{2\pi(\operatorname{Re}(C)^2 + \operatorname{Im}(C)^2)}. \quad (3)$$

242 A signal that consists of two frequency components,  $f_1$  and  $f_2$ , is analysed in order to ex-  
 243 plore the relationship between the instantaneous frequency and the two frequency compo-  
 244 nents. Assuming  $C_1$  and  $C_2$  are two single sinusoid complex signals, the instantaneous fre-  
 245 quency of the superposed signal is

$$f_i = \frac{1}{2\pi} \frac{d}{dt} \arctan \left( \frac{\operatorname{Im}(C_1) + \operatorname{Im}(C_2)}{\operatorname{Re}(C_1) + \operatorname{Re}(C_2)} \right), \quad (4)$$

246 which results in

$$f_i = \frac{1}{2\pi} \frac{d}{dt} \arctan \left( \frac{A_1 \sin \alpha + A_2 \sin \beta}{A_1 \cos \alpha + A_2 \cos \beta} \right), \quad (5)$$

247 where  $A_1$  and  $A_2$  are the amplitudes of the two sinusoid signals and  $\alpha$  and  $\beta$  are the phases of  
 248 the sinusoids  $C_1$  and  $C_2$ . The derivatives of  $\alpha$  and  $\beta$  are  $f_1$  and  $f_2$ . The explicit calculation  
 249 of the derivative in equation 5 yields

$$f_i = \frac{(A_1 \cos \alpha + A_2 \cos \beta)(A_1 f_1 \cos \alpha + A_2 f_2 \cos \beta) + (A_1 f_1 \sin \alpha + A_2 f_2 \sin \beta)(A_1 \sin \alpha + A_2 \sin \beta)}{(A_1 \cos \alpha + A_2 \cos \beta)^2 + (A_1 \sin \alpha + A_2 \sin \beta)^2}. \quad (6)$$

250

251 In order to simplify this result, we assume that  $\alpha = \beta$  when the amplitude of the superposed  
 252 signal is maximal. In this case, it follows that

$$f_i = \frac{A_1 f_1 + A_2 f_2}{A_1 + A_2}. \quad (7)$$

253 Similarly, we assume that  $\alpha = \pi + \beta$  when the amplitude of the superposed signal is mini-  
 254 mal. In this case, it follows that

$$f_i = \frac{A_1 f_1 - A_2 f_2}{A_1 - A_2}. \quad (8)$$

255 In the case where the instantaneous frequency of a signal consists of  $N$  frequency components

$$f_i = \frac{1}{2\pi} \frac{d}{dt} \arctan \left( \frac{\sum_j^N A_j \sin \alpha_j}{\sum_j^N A_j \cos \alpha_j} \right), \quad (9)$$

256 which results in

$$f_i = \frac{(\sum_j^N A_j f_j \cos \alpha_j)(\sum_j^N A_j \cos \alpha_j) + (\sum_j^N A_j \sin \alpha_j)(\sum_j^N A_j f_j \sin \alpha_j)}{2\pi((\sum_j^N A_j \cos \alpha_j)^2 + (\sum_j^N A_j \sin \alpha_j)^2)}, \quad (10)$$

257 Where  $A_j$  are the amplitudes of the sinusoid signals and  $\alpha_j$  are the phases of these sinusoids.  
 258 The derivatives of  $\alpha_j$  are the instantaneous frequencies  $f_j$  of the sinusoids. If the phases of all  
 259 sinusoids are the same, i.e.,  $\alpha_j = \alpha \forall j$ , then



$$f_i = \frac{\sum_j^N A_j f_j}{\sum_j^N A_j}. \quad (11)$$

260 These results show that the instantaneous frequency is equal to the amplitude-weighted fre-  
 261 quency in the frequency domain when the amplitude of the complex signal is maximal. The  
 262 instantaneous frequency may be singular when the denominator with the sum of amplitudes is  
 263 zero.

264 The instantaneous frequency of two frequency components is confirmed by simulating a su-  
 265 perposed signal  $y = \sin(2\pi * 100t) + 2\sin(2\pi * 120t)$  (Figure 4). The signal is down con-  
 266 verted by 110 Hz to compute the complex waveform. The instantaneous frequency is calcu-  
 267 lated directly from the derivative of the arctangent function (equation 3). The instantaneous  
 268 frequency of a single sinusoid signal is a real constant frequency. The instantaneous frequen-  
 269 cy of the superposed signal varies depending on the amplitude of the complex waveform. The  
 270 instantaneous frequency is equal to the amplitude-weighted average value calculated by equa-  
 271 tion 7 when the amplitude is maximal. The instantaneous frequency is not within the range of  
 272 the two frequency components and equal to the result calculated by equation 8 when the am-  
 273 plitude is minimal. This simulation result confirms that the instantaneous frequency repre-  
 274 sents the amplitude-weighted average of the true frequencies in each sample when the ampli-  
 275 tude is maximal. The instantaneous frequency provides no meaningful information about ei-  
 276 ther of the two frequency components when the amplitude is minimal. Therefore, for a real  
 277 wideband signal, such as lightning, the instantaneous frequency is only a reliable indicator of  
 278 the true frequency spectrum when the amplitude is maximal.

## 279 **5 Instantaneous Frequencies of Lightning Waveforms**

280 The instantaneous frequency of the average lightning waveform at a distance of 300 km is  
 281 calculated for different frequency bandwidths (Figure 5). At the beginning, the instantaneous  
 282 frequency is chaotic due to the noise and large sensitivity of the instantaneous phase, and it is  
 283 only relatively stable during the lightning pulses' arrival, which is confirmed by the simula-  
 284 tion result in the previous section. In order to emphasize the results during maximum ampli-  
 285 tude, the amplitude weighted average instantaneous frequency has been calculated by averag-  
 286 ing  $n=20$  samples of the dot product between the amplitudes and their instantaneous frequen-  
 287 cies divided by the sum of all amplitudes. This amplitude weighted average of the instantane-  
 288 ous frequency with empirically selected  $n$  essentially is a moving average to avoid the fre-  
 289 quency variation due to the sudden phase jump. While this moving average has tiny effect  
 290 during lightning pulses, because the instantaneous frequency is stable at that period. This am-  
 291 plitude weighted average instantaneous frequency is likely to be of benefit for a detailed  
 292 monitoring of the frequency distribution during the lightning period.

293 The instantaneous frequency of the average lightning waveform propagated over 300 km cal-  
 294 culated from 2-18 kHz (Figure 2) shows that the instantaneous frequency is smaller than the  
 295 center frequency of 10 kHz when the amplitude is maximum (Figure 5, upper left). This ex-  
 296 plains why the rotation direction of the complex waveform during lightning period is anti-  
 297 clockwise. The result means that the median value of the amplitude-weighted instantaneous  
 298 frequency from 2-18 kHz during the lightning period is below 10 kHz according to the theo-  
 299 retical analysis explained in the previous section. However, there are several modal maxima

300 in the spectra within this frequency range (Figure 1, right). Each peak in the spectrum may  
301 vary differently during the lightning period. As a result, it is better to concentrate on one  
302 modal maximum in the spectrum, so that the small variation of each peak during the lightning  
303 period can be observed individually. The instantaneous frequency calculated from a narrow-  
304 band frequency range around one peak in the spectrum is much less variable (Figure 5, lower  
305 left). The instantaneous frequency around the lightning pulse is almost constant, indicating  
306 that the frequency distribution during the lightning period is stable. It is noted that the stabil-  
307 ity of instantaneous frequency is more important than the value of instantaneous frequency,  
308 because the averaging value of the frequency distribution may be more close to the center  
309 frequency when there is no strong signal input.

310 As a result, the instantaneous frequency at the maximum amplitude is selected to represent  
311 the median frequency during the lightning period. The instantaneous frequencies at the max-  
312 imum amplitudes of the lightning waveforms from similar distances are compared to deter-  
313 mine the differences between individual events (Figure 5, right). As discussed above, a nar-  
314 row frequency bandwidth of 4 kHz is chosen with a varying center frequency, in order to  
315 constrain the target frequency range around the same modal maximum in the spectrum. These  
316 varying center frequencies are selected from the maxima of the second modal peak, which  
317 always fall between 10-20 kHz in the spectrum, for example at 300 km, 350 km and 400 km  
318 distance. The instantaneous frequencies at the maximum amplitudes are calculated for all the  
319 waveforms recorded by one station, and the distribution of the instantaneous frequencies at  
320 the same distance bin is a clearly peaked distribution (Figure 5, right). This indicates that the  
321 instantaneous frequencies at one distance bin are range limited. The center frequencies asso-  
322 ciated with the main peaks of the distributions are clearly distance dependent as a result of the  
323 radio wave propagation. In other words, the instantaneous frequency inferred from the aver-  
324 age lightning waveform can be used to represent the source receiver distance.

325 This idea can be tested by extending the analysis with the average lightning waveforms from  
326 300-600 km at each distance bin separated by 10 km in the frequency range 10-20 kHz (Fig-  
327 ure 6). The calculated instantaneous frequencies are obviously distance dependent and follow  
328 the second modal maximum in the spectra well. This excellent result strongly suggests that  
329 the instantaneous frequency has a promising potential application to determine the lightning  
330 distance from a single radio receiver.

## 331 **6 Discussion and conclusion**

332 The lightning waveform bank produced for distances up to 1000 km with a spatial resolution  
333 of 10 km is well suited for applications to study long range lightning location systems and  
334 electromagnetic wave propagation. For most lightning location systems, the baseline is  
335 smaller than 1000 km so that we can simulate and examine a new location algorithm or new  
336 site deployments. For wave propagation and ionospheric research, this waveform bank is val-  
337 uable as a reference for modeling [*e.g. Pasko and Fullekrug, 2011*]. In particular, spectra  
338 have been calculated across a lightning waveform bank, that reveal a sequence of consecutive  
339 modal maxima depending on distance and frequency. This waveform bank is generated from  
340 lightning recordings of a thunderstorm in Europe, which may not be identical to waveform  
341 banks calculated for other geographical areas. The geometry of the experiment limits the  
342 waveform bank to propagation distances with less than 1000 km. However, the general meth-  
343 od of producing the waveform bank and the spectral waveform bank is applicable to other

344 locations and longer distance than 1000 km in the future. In addition, this method is also ap-  
345 plicable to studies of other types of lightning, such as intra-cloud (IC) lightning discharges.

346 To the best of our knowledge, the complex waveform bank analysis of lightning is used for  
347 the first time, and it provides an opportunity to extract the instantaneous phase and instanta-  
348 neous frequency. The distance discrimination with the instantaneous frequency is just one  
349 potential application of this complex waveform bank. The instantaneous frequency may also  
350 be discriminated by different arrival azimuths, elevation angles or different times of day, giv-  
351 en more data. For example, it has been observed that a different incident elevation angle indi-  
352 cates a different rotation direction in the complex waveform of the lightning, i.e., a different  
353 instantaneous frequency [Fullekrug *et al.*, 2016, Figure 2, right]. By using the instantaneous  
354 frequency for distance determination, the lightning signal can be first approximated within  
355 <50 km, because the instantaneous frequency can vary, e.g. between 400 km and 450 km dis-  
356 tance (Figure 6). This uncertainty is very likely due to the lack of data at these distances,  
357 which could be improved by collecting more data with longer recordings. On the other hand,  
358 the instantaneous frequency is calculated from lightning waveforms that include interference  
359 from the local radio noise environment of each station. It is shown that the distribution of ap-  
360 parent frequencies from lightning at similar distances is a clearly peaked distribution if they  
361 are recorded at the same station (Figure 5, right). Between different stations, the distributions  
362 may differ slightly, most probably because of varying local radio environments and/or differ-  
363 ent propagation paths. As a result, the distance determination by using instantaneous frequen-  
364 cy may be more accurate if the instantaneous frequencies are derived from each station sepa-  
365 rately. It is noted that the instantaneous frequency in maximum amplitude is equal to the am-  
366 plitude weighted average frequency, such that waveform spectra inferred from STFT's with  
367 suitable parameters might result in similar distance dependencies.

368 In summary, this study has offered several results: (1) The average lightning spheric wave-  
369 forms from different distances exhibit a sequence of consecutive maxima resulting from the  
370 ionospheric reflections, which can be used for radio propagation studies, lightning modeling,  
371 lightning detection simulation, etc. (2) In the spectral waveform bank, the sequence of con-  
372 secutive modal maxima is separated by distinct minima at different frequencies and distances.  
373 (3) Long-range lightning location can achieve sub-sampling time accuracy by using the in-  
374 stantaneous phase of the complex lightning spheric waveform. (4) The instantaneous frequency  
375 calculated from average lightning waveforms has been shown to be distance dependent, and it  
376 therefore has the potential to be used for lightning distance determination.

377 **Acknowledgments.** The work of Z.L. is sponsored by the University of Bath, UK MetOffice  
378 (EA-EE1077) and the China Scholarship Council (CSC) (File No.201408060073). The work  
379 of K.K. is sponsored by the Engineering and Physical Sciences Research Council (EPSRC)  
380 under DTA contract EB-EE1151. The work of M.F. and A.M. is sponsored by the Natural  
381 Environment Research Council (NERC) under grants NE/L012669/1 and NE/H024921/1.  
382 This work was partly inspired by the SAINT project of the European Commission (H2020-  
383 MSCA-ITN-2016, 722337). The data used for this publication is available from  
384 <https://doi.org/10.15125/BATH-00404>. Z.L. wrote the paper and performed the data analysis,  
385 K.K. advised on the use of the Hilbert transform, A.M. assisted with the installation of the  
386 receiver network, S.E. and J.S. helped with the interpretation of the results, and M.F. super-  
387 vised the work of Z.L. and advised on the concepts for the data analyses. The authors wish to  
388 thank Serge Soula, Jean-Louis Pincon, Stephane Gaffet and their teams for hosting the radio

389 receivers in Lannemezan, Orleans and Rustrel. Z.L. wants to thank Dirk Klugmann and Ivan  
390 Astin for encouragement toward this project and acknowledges helpful discussions and sup-  
391 port from Andrew Moss as well as the constructive suggestions of the anonymous reviewers.

392

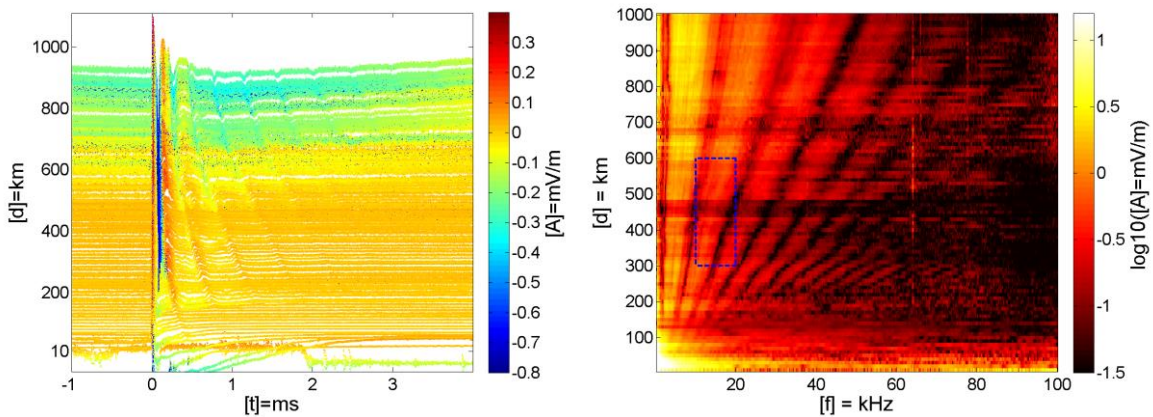
393 **References**

- 394 Bitzer P. M., H. J. Christian, M. Stewart, J. Burchfield, S. Podgorny, D. Corredor, J. Hall, E.  
395 Kuznetsov, and V. Franklin (2013), Characterization and applications of VLF/LF  
396 source locations from lightning using the Huntsville Alabama Marx Meter Array, *J.*  
397 *Geophys. Res. Atmos.*, 118,3120–3138, doi:10.1002/jgrd.50271.
- 398 Carvalho, F. L., M. A. Uman, D. M. Jordan, J. D. Hill, S. A. Cummer, D. A. Kotovsky, and  
399 R. C. Moore (2017), Triggered lightning sky waves, return stroke modeling, and iono-  
400 sphere effective height, *J. Geophys. Res. Atmos.*, 122, 3507–3527,
- 401 Cheng, Z., S. A. Cummer, H.-T. Su, and R.-R. Hsu (2007), Broadband very low frequency  
402 measurement of D region ionospheric perturbations caused by lightning electromag-  
403 netic pulses, *J. Geophys. Res.*, 112, A06318, doi:10.1029/2006JA011840.
- 404 Cummins, K. L., M. J. Murphy, E. A. Bardo, W. L. Hiscox, R. B. Pyle, and A. E. Pifer  
405 (1998), A Combined TOA/MDF Technology Upgrade of the U.S. National Lightning  
406 Detection Network, *J. Geophys. Res.*, 103(D8), 9035–9044, doi:10.1029/98JD00153.
- 407 Cummins, K., and M. Murphy (2009), An overview of lightning locating systems: History,  
408 techniques, and data uses, with an in-depth look at the US NLDN, *IEEE Trans. Elec-*  
409 *tromag. Compat.*, 51(3), 499–518, doi:10.1109/TEM.2009.2023450.
- 410 Dowden, R. L., Brundell, J. B., Rodger, C. J. (2002), VLF lightning location by time of group  
411 arrival (TOGA) at multiple sites. *Journal of Atmospheric and Solar-Terrestrial Phys-*  
412 *ics* 64(7): 817-830. doi:10.1016/S1364-6826(02)00085-8.
- 413 Füllekrug, M. and S. Constable (2000), "Global triangulation of intense lightning discharges."  
414 *Geophysical Research Letters* 27(3): 333-336, doi: 10.1029/1999GL003684.
- 415 Füllekrug, M., Mareev, E. and Rycroft, M. (2006), Sprites, elves and intense lightning dis-  
416 charges, Springer Science & Business Media, doi: 10.1007/1-4020-4629-4.
- 417 Füllekrug, M. (2010), Wideband digital low-frequency radio receiver, *Meas. Sci. Technol.*,  
418 21 (015901), 1–9, doi:10.1088/0957-0233/21/ 1/015901.
- 419 Füllekrug, M., et al. (2013), Electron acceleration above thunderclouds. *Environmental Re-*  
420 *search Letters*, 8 (3), 035027, doi: 10.1088/1748-9326/8/3/035027.
- 421 Füllekrug, M., A. Mezentsev, R. Watson, S. Gaffet, I. Astin, and A. Evans (2014), Array  
422 analysis of electromagnetic radiation from radio transmitters for submarine communi-  
423 cation, *Geophys. Res. Lett.*, 41, 9143–9149, doi:10.1002/2014GL062126.
- 424 Füllekrug, M., A. Mezentsev, R. Watson, S. Gaffet, I. Astin, N. Smith, and A. Evans (2015),  
425 Map of low-frequency electromagnetic noise in the sky. *Geophys. Res. Lett.*, 42,  
426 4648–4653. doi: 10.1002/2015GL064142.
- 427 Füllekrug, M., Z. Liu, K. Koh, A. Mezentsev, S. Pedebay, S. Soula, S.-E. Enno, J. Sugier,  
428 and M. J. Rycroft (2016), Mapping lightning in the sky with a mini array, *Geophys.*  
429 *Res. Lett.*, 43, 10,448–10,454, doi:10.1002/2016GL070737.
- 430 Höller, H., Betz, H.-D., Schmidt, K., Calheiros, R. V., May, P., Houngrinou, E., and Sci-  
431 alom, G. (2009), Lightning characteristics observed by a VLF/LF lightning detection  
432 network (LINET) in Brazil, Australia, Africa and Germany, *Atmos. Chem. Phys.*, 9,  
433 7795-7824, doi: 10.5194/acp-9-7795-2009.
- 434 Horner, F. (1954), The accuracy of the location of sources of atmospheric by radio direction-  
435 finding, *Proc IEEE* 101: 383-390, doi: 10.1049/pi-3.1954.0091.

- 436 Horner, F. (1957), Very-low-frequency propagation and direction-finding, Proc IEEE 104:  
437 73-80, doi: 10.1049/pi-b-1.1957.0118.
- 438 Kohlmann, H., W. Schulz, and S. Pedebay (2017), Evaluation of EUCLID IC/CG classifica-  
439 tion performance based on ground-truth data. 2017 International Symposium on  
440 Lightning Protection (XIV SIPDA), Natal, Brazil, 2-6 Oct. 2017, doi:  
441 10.1109/SIPDA.2017.8116896
- 442 Krider, E. P., R. Carl Noggle, and Martin A. Uman (1976), A Gated, Wideband Magnetic Di-  
443 rection Finder for Lightning Return Strokes. J. Appl. Meteor., 15, 301–306. doi:  
444 10.1175/1520-0450(1976)015<0301:AGWMDF>2.0.CO;2.
- 445 Lee, A.C. (1986), An experimental study of the remote location of lightning flashes using a  
446 VLF arrival time difference technique, Q. J. Roy. Meteor. Soc., 112, 203–229,  
447 doi:10.1002/qj.49711247112.
- 448 Lee, A.C. (1990), Bias Elimination and Scatter in Lightning Location by the VLF Arrival  
449 Time Difference Technique. J. Atmos. Oceanic Technol., 7, 719–733, doi:  
450 10.1175/1520-0426(1990)007<0719:BEASIL>2.0.CO;2
- 451 Liu, Z., K. L. Koh, A. Mezentsev, S.-E. Enno, J. Sugier, and M. Füllekrug (2016), Variable  
452 phase propagation velocity for long-range lightning location system, Radio Sci., 51,  
453 1806–1815, doi:10.1002/2016RS006058.
- 454 Lyu, F., S. A. Cummer, R. Solanki, J. Weinert, L. McTague, A. Katko, J. Barrett, L. Zi-  
455 goneanu, Y. Xie, and W. Wang (2014), A low-frequency near-field interferometric-  
456 TOA 3-D Lightning Mapping Array, Geophys. Res. Lett., 41, 7777–7784,  
457 doi:10.1002/2014GL061963.
- 458 Mezentsev, A., and M. Füllekrug (2013), Mapping the radio sky with an interferometric net-  
459 work of low-frequency radio receivers, J. Geophys. Res., 118, 8390–8398,  
460 doi:10.1002/jgrd.50671.
- 461 Nag, A, Murphy, MJ, Schulz, W, and Cummins, KL (2015), Lightning locating systems: In-  
462 sights on characteristics and validation techniques. Earth and Space Science, 2, 65–  
463 93. doi: 10.1002/2014EA000051.
- 464 Pasko, V.P. and Füllekrug, M. (2011), Waveforms of Nighttime Atmospherics as a Measure  
465 of the Lower Ionospheric Electron Density Profiles over UK and France on August  
466 31, 2008', paper presented to 30th URSI General Assembly (No. GHE2-7), Istanbul,  
467 Turkey, 13-20 August, viewed 24 July 2017,  
468 <<http://www.ursi.org/proceedings/procGA11/ursi/GHE2-7.pdf>>
- 469 Qin, Z., M. Chen, B. Zhu, and Y. Du (2017), An improved ray theory and transfer matrix  
470 method-based model for lightning electromagnetic pulses propagating in Earth-  
471 ionosphere waveguide and its applications, J. Geophys. Res. Atmos., 122, 712–727,  
472 doi:10.1002/2016JD025599.
- 473 Rakov, V. A. and M. A. Uman (2003), Lightning: physics and effects, Cambridge University  
474 Press. doi:10.1256/wea.168/03.
- 475 Rakov, V. (2013), "Electromagnetic methods of lightning detection." Surveys in Geophysics  
476 34(6): 731-753. doi:10.1007/s10712-013-9251-1.
- 477 Rison, W., Krehbiel, P. R., Stock, M. G., Edens, H. E., Shao, X. M., Thomas, R. J., Stanley,  
478 M. A., Zhang, Y. (2016), Observations of narrow bipolar events reveal how lightning  
479 is initiated in thunderstorms." Nat Commun 7. doi:10.1038/ncomms10721.

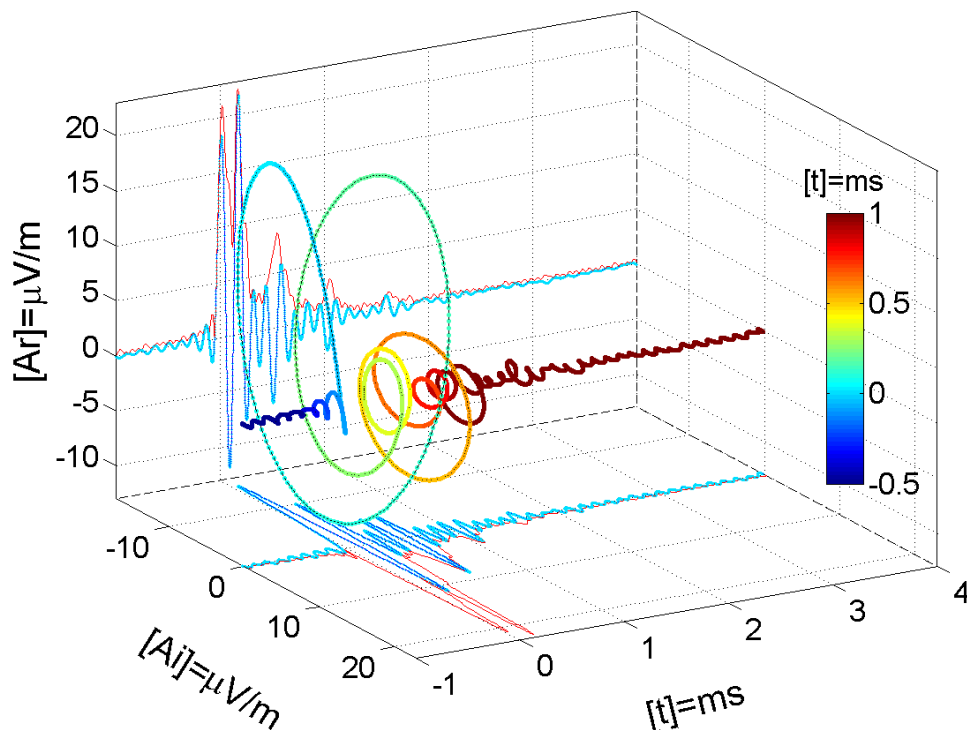
- 480 Said, R. K., U. S. Inan, and K. L. Cummins (2010), Long-range lightning geolocation using a  
481 VLF radio atmospheric waveform bank, *J. Geophys. Res.*, 115, D23108,  
482 doi:10.1029/2010JD013863.
- 483 Schonland, B. F. J., J. S. Elder, D. B. Hodges, W. E. Phillips, J. W. van Wyk (1940), The  
484 wave form of atmospherics at night. *Proc. R. Soc. Lond. A* 1940 176 180-202. doi:  
485 10.1098/rspa.1940.0085.
- 486 Shao, X. M., Lay, E. H., & Jacobson, A. R. (2013). Reduction of electron density in the  
487 night-time lower ionosphere in response to a thunderstorm. *Nature Geoscience*, 6(1),  
488 29-33. doi:10.1038/ngeo1668
- 489 Soula, S., Iacovella, F., van der Velde, O., Montanya, J., Füllekrug, M., Farges, T., Bór, J.,  
490 Georgis, J.-F., NaitAmor, S. and Martin, J.-M., (2014), Multi-instrumental analysis of  
491 large sprite events and their producing storm in southern France. *Atmospheric Re-*  
492 *search* 135: 415-431. doi:10.1016/j.atmosres.2012.10.004.
- 493 Stock, M. G., M. Akita, P. R. Krehbiel, W. Rison, H. E. Edens, Z. Kawasaki, and M. A. Stan-  
494 ley (2014), Continuous broadband digital interferometry of lightning using a general-  
495 ized cross-correlation algorithm, *J. Geophys. Res. Atmos.*, 119, 3134-3165,  
496 doi:10.1002/2013JD020217.
- 497 Sun, Z., X. Qie, M. Liu, R. Jiang, Z. Wang, and H. Zhang (2016), Characteristics of a nega-  
498 tive lightning with multiple-ground terminations observed by a VHF lightning loca-  
499 tion system, *J. Geophys. Res. Atmos.*, 121, 413–426, doi:10.1002/2015JD023702.
- 500 Taner, M. T., F. Koehler, R. E. Sheriff (1979), Complex seismic trace analysis. *Geophysics*  
501 44(6): 1041-1063. doi: 10.1190/1.1440994.
- 502 Wang, Y., X. Qie, D. Wang, M. Liu, D. Su, Z. Wang, D. Liu, Z. Wu, Z. Sun, Y. Tian (2016),  
503 Beijing Lightning Network (BLNET) and the observation on preliminary breakdown  
504 processes, *Atmos. Res.* 171: 121-132. doi:10.1016/j.atmosres.2015.12.012  
505

506



507

508 **Figure 1.** The average night-time waveforms and the spectra of the average waveforms of negative  
 509 lightning discharges at distances from 10-1000 km. (left) Each average waveform at a given distance  
 510 is calculated from more than one hundred events. The time axis is referenced to  $t=0$  corresponding to  
 511 a propagation at the speed of light from the source to the receiver. A sequence of consecutive maxima  
 512 resulting from the ionospheric reflections (or multi-hop sky waves) appear from  $\sim 100$  km distance  
 513 onwards and the time differences between ground wave and sky waves are smaller for larger distances.  
 514 (right) The sequence of consecutive modal maxima (yellow and red) is separated by distinct minima  
 515 (black) which are characteristic for the distance between the radio receiver and the lightning discharge.  
 516 The area of the blue square is used for the analysis of the instantaneous frequencies (compare to Fig-  
 517 6).

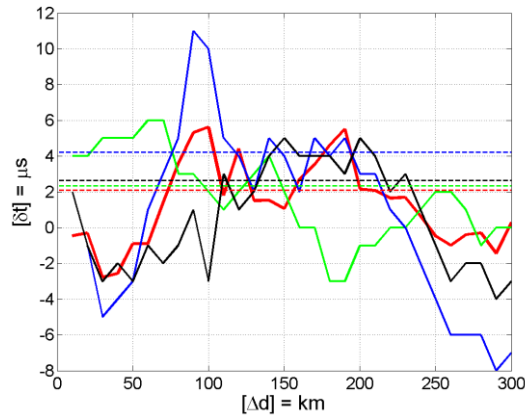


518

519 **Figure 2.** Isometric diagram of an average complex waveform of lightning at a distance of 300 km.  
 520 The 3D trace of the complex lightning signal from 2 kHz to 18 kHz exhibits large amplitudes when  
 521 the ground wave and sky waves are arriving. The rotation direction of the complex trace is anticlock-  
 522 wise, which means the instantaneous frequency is smaller than centre frequency. The blue lines show

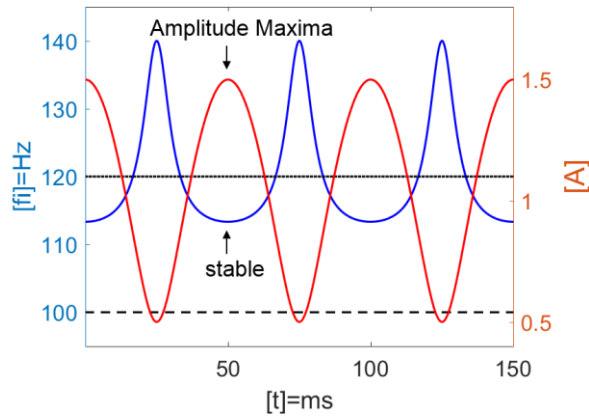


523 the real part ( $A_r$ ) and imaginary part ( $A_i$ ) of the complex lightning waveform. The thin red line is the  
 524 amplitude envelope of the complex waveform.



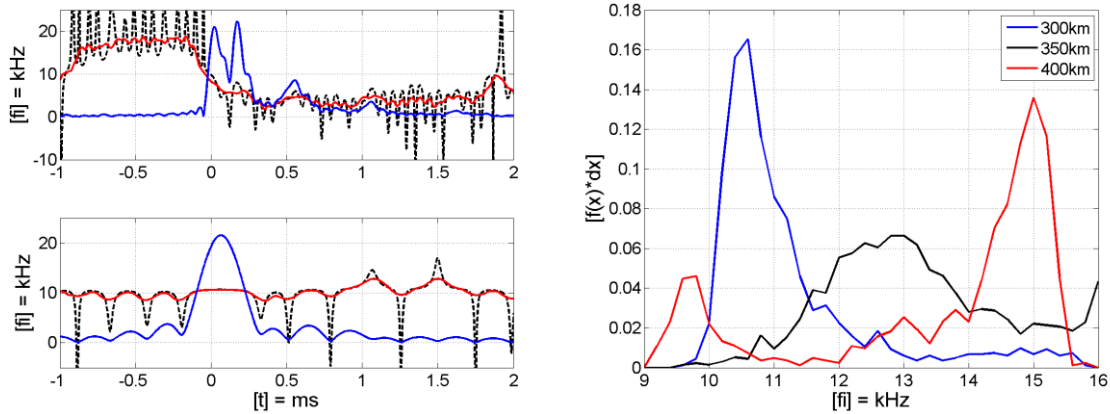
525

526 **Figure 3.** The time offsets between different propagation distance differences with respect to a speed  
 527 of light propagation. The average lightning waveforms from 310 km to 600 km are compared with the  
 528 average lightning waveform at 300 km distance. The time offset  $\delta t$  is the measured time difference  $\Delta t$   
 529 minus the nominal time difference inferred from a wave propagation velocity at speed of light  $\Delta d/c$ ,  
 530 and the distance differences  $\Delta d$  is measured relative to the lightning waveform at 300 km distance, i.e.,  
 531  $\Delta d = d - 300$  km with  $d$  ranging from 310 km to 600 km. By comparing different signal processing  
 532 methods, the average value of the absolute time offset by using the instantaneous phase of the complex  
 533 lightning waveform (red) is less than just using the amplitude envelope (blue) or using the ampli-  
 534 tude of the real signal (black) or cross correlating lightning waveforms (green). The corresponding  
 535 average values of the absolute time offset are shown by the dashed lines.



536

537 **Figure 4.** The instantaneous frequency of a simulated superposed signal that consists of two frequen-  
 538 cy components at 100 Hz and 120 Hz (black solid and dashed lines). The calculated instantaneous  
 539 frequency of the averaged signal (blue line) is more stable and equal to the amplitude-weighted aver-  
 540 age value calculated by equation 7 when the amplitude is maximal (red line).

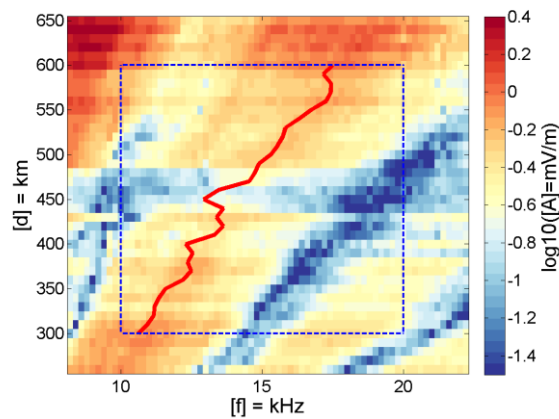


541

542

543 **Figure 5.** Analyses of instantaneous frequencies inferred from lightning waveforms. (left) The instan-  
 544 tantaneous frequency (black dashed line) calculated from a lightning waveform with two frequency  
 545 bandwidths, 2-18 kHz (top) and 8.8-12.8 kHz (bottom), confirm the theoretical result that the instan-  
 546 tantaneous frequency is more stable when the amplitude of the signal is maximal (blue line). The ampli-  
 547 tude weighted average instantaneous frequency (red line) smoothens the original instantaneous fre-  
 548 quency. The instantaneous frequency calculated from a 4 kHz bandwidth around one modal maximum  
 549 in the spectrum is less variable than using a bandwidth from 2-18 kHz that contains many modal max-  
 550 ima. (right) The distributions of the instantaneous frequencies at maximum amplitudes inferred from  
 551 all the lightning waveforms recorded by one station at several distance bins are clearly peaked distri-  
 552 butions.

553



554

555 **Figure 6.** The instantaneous frequencies (red line) at maximum amplitudes of the average lightning  
 556 waveforms at distances ranging from 300-600 km (blue square) follow the modal maximum in the  
 557 spectra well.

558