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Anthropogenic warming has substantially increased the likelihood of July 2017like heat waves over Central-Eastern China

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1 **Capsule summary**

Heat waves in Central-Eastern China like the record-breaking July 2017 event were
rare in natural worlds, but have now become approximately one-in-five-year events
due to anthropogenic forcings.

5 Introduction

6 During July 2017, an unprecedentedly intense heat wave struck Central-Eastern 7 China, resulting in drastically-increased human morbidity/mortality, steeply-reduced 8 agriculture productivity, and serious shortage of electricity and water supply (China 9 Climate Bulletin of 2017). Many meteorological stations registered 15-25 hot days 10 (daily maximum temperature over 35°C), and some even had their record-high July 11 temperatures, such as a new record of 40.9°C amongst historical observations since 12 1873 in Xu-Jia-Hui station in Shanghai (China Climate Bulletin of 2017). The China 13 Meteorological Administration issued 10 high-level warnings against hot weather during 21st-25th July. Such unprecedentedly frequent alarms within only 5 days 14 15 attracted intense scrutiny from policy-makers, media, and the public on the 16 relationship between this heat wave and global warming.

Previous studies usually conducted attribution analyses on seasonal warmth in Central-Eastern China (e.g. the 2013 record-breaking summer, Sun et al. 2014), leaving attribution statements for short-term (synoptic) hot extremes sparsely reported. This study therefore attempts to answer whether and to what extent anthropogenic warming has increased the likelihood of 5-day heat waves as hot or hotter than the 21st–25th July 2017 case over Central-Eastern China.

23 Data and Methods

Homogenized observations of daily maximum temperatures (Tmax) during 1960-2017 from 760 meteorological stations are used (Li et al. 2015; homogenization methods see Szentimrey 1999). Daily observations is interpolated onto the $0.56^{\circ} \times$ 0.83° grid of the model via a 'natural neighbor' scheme (Sibson 1981), following the model's resolution and geography.

29 The upgraded HadGEM3-GA6-N216 model is employed (Christidis et al. 2013; 30 Ciavarella et al. 2018). Model outputs include all-forced simulations conditioned on 31 the observed 2017 sea surface temperature (SST) and sea ice from the HadISST 32 dataset (Rayner et al. 2003), and naturalized simulations with anthropogenic signals 33 removed from observed SSTs and with pre-industrial forcings. Accordingly, 34 occurrence probabilities and resultant attribution conclusions reported in this study are 35 also conditioned on the 2017 SST patterns. The ensemble is generated through 36 physics perturbations of multiple initial conditions with identical external forcings.

37 More specifically, historical simulations (histCLIM) consisting of fifteen 38 members over 1961–2013 are compared with interpolated observations, to evaluate 39 the model's fidelity in simulating climatological statistics (mean and variability) of 40 the strongest 5-day heat waves. Two ensembles of 525-member simulations for the 41 2017 July with (hereafter histALL, as an extension of previous histCLIM runs) and 42 without (hereafter histNAT) anthropogenic forcings are used to estimate the probability of the 21st-25th July heat wave in each scenario. Denoting PALL and PNAT 43 44 as the occurrence probability of events equivalent to or stronger than the targeted case 45 in 525-member histALL and histNAT ensembles, the risk ratio (RR) is expressed as PALL/PNAT. The fraction of attributable risks (FAR) is expressed as 1- PNAT/PALL. 46

47 Reference climatologies over 1961-1990 are formed for both simulations 48 (ensemble mean of 15-member histCLIM) and observations from the hottest 5-day 49 running mean Tmax in July. These pentad climatologies are approximately 2-3°C 50 warmer than July monthly-mean Tmax climatologies in both simulations and 51 observations, and serve to distinguish especially intense 5-day heat waves from more 52 typical 5-day cases (Fig. 1c-d). Respective climatologies are then removed from 53 observations and simulations to create overlapping pentad Tmax anomalies (see Fig. 54 1c, hereafter PTmax). Based on these PTmax anomalies, both the historical distribution of the hottest 5-day heat waves and warm anomalies for the 2017 case 55 56 could be well reproduced by this model (Fig. S1), indicating the suitability of using 57 this model and PTmax anomalies for attributing this 5-day heat wave. Freychet et al. 58 (2018) also reported good performance of this model in simulating characteristics of 59 5-day heat waves in Central-Eastern China, as it is capable of capturing critical 60 mechanisms generating heat waves there. In the reminder of this paper, we used the 61 PTmax anomaly to define the threshold.

62 **Results**

During 21st-25th July, almost the entirety of Central-Eastern China had 63 64 temperatures over 35°C, equivalent to 2–6°C PTmax anomalies (Fig. 1a). Anomalies 65 of these magnitudes produced numerous record- or near-record July PTmax (Fig. 1b). 66 In terms of domain-averaged values, the PTmax in this pentad not only peaked during 67 July 2017, but also set a new record amongst all historical July counterparts (any 5day mean Tmax during July) since 1960 (Fig. 1c-d; note: we consider this pentad 68 instead of 22nd-26th because of its extensive social and economic repercussions). It is 69 70 well-known that heat waves in this area result dynamically from the persistence of anticyclonic circulations which facilitate increased surface solar radiation and adiabatic heating (Freychet et al. 2017; Chen et al. 2015). Specific to this case, an unprecedentedly (all Julys since 1960) strong anomalous anticyclonic cell was centered above Central-Eastern China, dynamically explaining the origin of the "record-breaking" Tmax (Fig. S2) and its exclusive occurrence in this domain (Fig. 1a).

77 The PTmax anomaly from the interpolated observation $(2.52^{\circ}C)$ was used as a 78 threshold to characterize the July 2017-like heat wave. Events of this magnitude are 79 fairly rare (P_{NAT}=2.1%) in natural-forcing simulations (Fig. 2a, green). Without 80 anthropogenic warming, similar heat waves should have been seen one to three times 81 per century (mean return period: 47.7 years, 95% CI: 30.8–75.0 years, Fig. 2b, green). 82 By contrast, the distribution of simulated PTmax anomaly is markedly positive-83 displaced in all-forcing worlds, signifying substantially increased odds (PALL=20.1%) of events this hot. In the current climate, anthropogenic warming has exposed Central-84 85 Eastern China to 2017-like heat waves about twice per decade (mean return period: 4.9 years, 95% CI: 4.3–5.8 years, Fig. 2b, red). 86

87 Quantitatively speaking, the risk of an event as hot or hotter increased by at least 88 10-fold (RR=9.8, 95% CI: 5.9-18.9) due to anthropogenic warming. Translating into 89 FAR, human influence accounted for at least 90% (95% CI: 83.0%-94.7%) for the 90 presence of 2017-like heat waves. To avoid selection bias potentially introduced by 91 using the critical threshold at the very end tail (Stott et al. 2004), we also adopted the 92 second hottest July record (2.09°C in July 2002) as an alternative threshold. 93 Simulated anomalies exceeding this threshold are recorded 5 times more frequently 94 (RR=4.5, 95% CI: 3.4–6.5) in the all-forcing world (PALL=26.8%) than in the natural-95 forcing world (P_{NAT}=5.9%). These results also indicate anthropogenic forcings

96 contributed more to increases in risks of rarer, more extreme heat waves. So, we 97 reiterate that anthropogenic warming played an overarching role (FAR=77.8%, 95%) CI: 70.4%-84.6%) in elevating the risk of heat waves stronger than this second-98 99 hottest threshold (e.g. the July 2017 case).

100

Conclusion and Discussion

101 In Central-Eastern China, heat waves hotter than the July 2017 event should 102 have had a very slim chance to occur in natural-forcing worlds. But now, forced by 103 anthropogenic warming and conditioned on the 2017 SST pattern, a 5-day heat wave 104 like this case has become 10 times more likely, as a one-in-five-year or more common 105 event.

106 Although influences of anthropogenic warming could be detected and were largely attributable, attribution conclusions for a single high-impact case may be 107 108 subject to some uncertainties. Firstly, the estimated RR and FAR may be 109 quantitatively sensitive to the selection of baseline periods (here 1961-1990), as 110 reported by Knutson et al. (2013). Still, sensitivity tests adopting varying baselines for 111 this case indicate that the qualitative statement "increase in the likelihood of a July 112 2017-like heat wave could be largely attributable to anthropogenic warming" robustly 113 holds. Secondly, the estimated RR and FAR only apply to the current climate. As the 114 planet keeps warming, a higher RR of a July 2017-like case would be expected 115 (Perkins and Gibson 2015). Future reductions in aerosols due to increasingly stricter air quality control in this area may also give a greater RR of a July 2017-like case 116 117 (Van Oldenborgh et al. 2018; Wang et al. 2017). This study is based only on factual 118 and counterfactual runs in a single atmosphere-only model, with the intention of 119 exploiting its large ensembles for calculating the statistics of rare events (Otto 2017).

Estimated RRs should still be compared with those derived via other methods/models, such as observation-constrained estimates (van Oldenborgh et al. 2015), alternative atmosphere-only model-based estimates (e.g. weather@home, Massey et al. 2015) and fully-coupled model-based estimates (CMIP5, Sun et al. 2014), to further clarify uncertainties.

125 Comparing temperatures alone in factual and counterfactual simulations, the 126 estimated RR only delivers a general attribution message, leaving physical 127 interpretations about how anthropogenic forcings influenced the likelihood of the heat 128 wave and its preferential occurrence in Central-Eastern China to be addressed. To this 129 end, follow-up efforts will be made to disentangle this general attribution effort into 130 dynamical part (e.g., large-scale circulations) and thermodynamic part (Vautard et al. 131 2016; Schaller et al. 2016). A critical step toward dynamic attribution is to quantify 132 the extent to which anthropogenic warming affected the presence, location, 133 maintenance and amplitude of anticyclonic circulations akin to the 2017 case (Fig. 134 S2). Such a separation could also facilitate to track down and communicate the source 135 of attribution uncertainties from both dynamic and thermodynamic perspectives 136 (Vautard et al. 2016; Wehrli et al. 2018).

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218 Figure Caption List

Fig. 1. (a) Observed pentad-mean (21st -25th July 2017) Tmax anomalies (°C) relative 219 220 to the 1961-1990 climatology for the maximum 5-day mean Tmax. The green contour 221 indicates the 35°C-isoline of mean Tmax during this pentad. Central-Eastern China is shown by the dashed rectangle. (b) Spatial distribution of stations that registered 222 record- and near-record (since 1960) pentad-mean July Tmax during $21^{st} - 25^{th}$ July 223 224 2017. (c) Observed overlapping pentad-mean Tmax anomaly averaged over Central-225 Eastern China during July 2017. Each value is indexed by the first day of the pentad. 226 (d) Observed maximum 5-day mean Tmax anomaly averaged over Central-Eastern China in each July over 1960–2017. The red vertical line labels the 2017 event, and 227 228 the dashed line indicates its anomaly.

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Fig. 2. (a) Distribution of domain-averaged hottest 5-day mean Tmax anomalies 230 231 during July 2017 (histogram), based on 525-member histALL (red) and histNAT 232 (green) ensembles, and their GEV-fitted curves shown by respective colors. (b) 233 Return periods of domain-averaged hottest 5-day mean Tmax anomalies in histALL 234 (red) and histNAT (green) ensembles. The threshold value of 2.52°C is indicated by dashed lines in (a) and (b). In (b), vertical and horizontal bars represent the 5%-95% 235 236 uncertainty interval of temperature anomalies and return periods, derived via the 237 bootstrapping method (N=1000). Grey shadings specify the uncertainty interval of 238 return period of the threshold-exceedance in histNAT and histAll runs.

239 Figures



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