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Paleolimnological proxies reveal continued eutrophication issues in the St. Lawrence River Area of Concern

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1 **Title Page**

2 Paleolimnological proxies reveal continued eutrophication issues in the St. Lawrence River Area
3 of Concern

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Abstract

Recent surface-water surveys suggest that high nutrient concentrations and nuisance algae remain issues in the St. Lawrence River Area of Concern (AOC) at Cornwall, Ontario, specifically in the tributaries and nearshore zones of Lake St. Francis (LSF). In particular, it is unclear whether management actions designed to reduce nutrient inputs, first implemented in the 1990s as part of the Remedial Action Plan for the AOC, have reduced algal production or influenced assemblage composition. To address this issue, a paleolimnological approach was used to provide a historical context for the present-day nutrient concentrations and to quantify the extent of change in water quality in LSF since the early 1990s. A sediment core was collected near the north shore of LSF and was examined for changes in the concentrations and compositions of fossil diatoms and pigments, as well as stable isotope ($\delta^{15}\text{N}$ and $\delta^{13}\text{C}$) values. Analyses of diatom and pigment concentrations indicated that overall algal abundance has risen in the last few decades, including trends of increasing occurrences of potentially toxic cyanobacteria, despite ongoing remediation efforts. Temporal patterns of stable isotope signatures in the core suggest a steady increase in nutrient influx since the mid-20th century, with the post-1990 increase in algal production likely attributable to recent inputs associated with land-use changes in local contributing watersheds. These patterns suggest that the AOC delisting goals for the LSF tributaries will not be reached without a drastic change in land management practices.

Keywords

Eutrophication, paleolimnology, St. Lawrence River, pigments, diatoms, stable isotopes

44

Introduction

45 Within the Laurentian Great Lakes Basin, 43 Areas of Concern (AOCs) have been
46 identified by the International Joint Commission as regions that have experienced environmental
47 degradation as a result of biological, chemical, or physical changes in the aquatic ecosystem
48 (Dreier et al., 1997; International Joint Commission, 2003a). The St. Lawrence River near
49 Cornwall, ON, and Massena, NY, is the easternmost AOC where environmental issues arose
50 from intensive industrial and agricultural activities, habitat loss and degradation, as well as from
51 hydrodynamic changes from anthropogenic modifications to the waterway such as the
52 construction of the St. Lawrence Seaway (Anderson et al., 1992). Two Remedial Action Plans
53 (RAPs) were developed for the St. Lawrence River AOC at Cornwall and Massena, serving to
54 identify and remediate beneficial use impairments (BUIs; International Joint Commission, 2012)
55 in the Canadian and U.S. portions of the AOC, respectively. Within the Canadian section of the
56 AOC, many of the identified environmental stressors have been mitigated through regulations
57 and local action, including reductions in the concentrations of harmful bacteria, improved
58 management of fish populations, and restrictions on industrial discharges to the waterway
59 (Environment Canada and Ontario Ministry of the Environment, 2010). However, three BUIs
60 remain impaired in this AOC, including eutrophication and the presence of undesirable algae
61 (e.g., toxic cyanobacterial blooms), a problematic issue in the nearshore zones and tributaries of
62 the fluvial lake known as Lake St. Francis (LSF; Environment Canada and Ontario Ministry of
63 the Environment, 2010).

64 Increased nutrient loadings from the LSF watersheds, faulty septic systems in nearshore
65 communities, changes to the hydraulics of the system from seaway construction, and climate
66 change have all been suggested as contributing sources of the nuisance eutrophication and algal

67 blooms in the AOC (Anderson et al., 1992; The St. Lawrence River (Cornwall) RAP Team,
68 1995). Although only 5% of the water in LSF originates in its tributaries (Anderson et al., 1992),
69 the large proportion of agricultural land in the contributing watersheds could disproportionately
70 affect nutrient loadings to LSF and impair water quality in nearshore areas. Remediation goals
71 for eutrophication in the AOC originally included mean summer tributary and nearshore total
72 phosphorus (TP) concentrations $\leq 30 \mu\text{g/L}$ and no eutrophication-related fish kills (Dreier et al.,
73 1997). The targets for TP concentration in the tributaries were updated in 2009 to reflect
74 proportional goals based on the amount of agricultural activity in each watershed, ranging from
75 35-60 $\mu\text{g/L}$ (AECOM Canada Ltd., 2009; J. Ridal, pers. comm.). The TP target for the main
76 body of LSF, beyond the 2-m isopleth, remains at 20 $\mu\text{g/L}$ and is not currently considered
77 impaired.

78 Since the early 1990s, efforts to reduce eutrophication in LSF have primarily targeted
79 nutrients emanating from local farms and those from the city of Cornwall. Actions have included
80 tributary restoration programs (including tree planting along tributary banks and fencing to
81 restrict cattle access to streams), upgrades to septic systems, reductions in agricultural runoff
82 through the *Nutrient Management Act* (2002), upgrades to the city of Cornwall wastewater
83 treatment plant, and reductions in the number of combined sewers in the city of Cornwall
84 (Environment Canada and Ontario Ministry of the Environment, 2010). Unfortunately,
85 monitoring of the water-quality and ecological responses to these actions has been limited,
86 hampering the ability to assess potential eutrophication declines in LSF. Monitoring data for both
87 TP and algal abundance and community structure are sparse prior to the last decade (Pilon and
88 Chrétien, 1991; Reavie et al., 1998; Richman et al., 1997), and it remains unclear if and how

89 algal communities in the tributaries and nearshore zones of LSF have responded to remedial
90 actions.

91 Provided that the sediment has remained relatively undisturbed, paleolimnological
92 approaches can be applied to LSF to examine how algal assemblages have responded to the
93 implementation of the RAP, and how those communities have changed over time. Previous
94 paleolimnological characterisations of the eastern end of LSF, collected in the early 1990s,
95 suggested that diatom communities responded to the known period of eutrophication in the Great
96 Lakes in the mid-20th century, and additionally responded to well-documented macrophyte
97 growth in the region (Reavie et al., 1998). However, less is known about historical changes in
98 other groups of primary producers, including potentially toxin-producing cyanobacteria such as
99 *Anabaena* and *Microcystis* (Carmichael, 2001), occurrences of which have been reported in this
100 region in recent years (Bramburger, 2014; Waller et al., 2016). Paleolimnological techniques
101 have been successfully applied to other AOCs (e.g., Alexson et al., 2017; Dixit et al., 1998),
102 providing valuable information to stakeholders regarding historical environmental changes to the
103 impacted systems.

104 The objective of the current study is to assess the degree to which the abundance and
105 composition of algal communities in the nearshore areas of LSF have changed since the
106 implementation of the RAP in the early-1990s. Although some surface-water sampling has been
107 conducted in recent years, the response of algal assemblages to actions implemented as part of
108 the RAP has not been examined, despite ongoing concerns regarding high nutrient
109 concentrations and algal blooms in the AOC, including occurrences of toxin-producing
110 cyanobacteria (Bramburger, 2014; Environment Canada and Ontario Ministry of the
111 Environment, 2010; Savard et al., 2013, 2015). To address this issue, we quantified sedimentary

112 concentrations of photosynthetic pigments known to reliably indicate historical changes in
113 abundances of primary producers (Hall et al., 1999; Leavitt and Findlay, 1994), fossil diatom
114 assemblages to infer past environmental conditions along the impacted northern shore of LSF
115 (Battarbee et al., 2002; Reavie and Edlund, 2010), and carbon (C) stable isotopes to evaluate
116 temporal changes in production and C sources (Hodell and Schelske, 1998; Savage et al., 2010).
117 In addition, stable isotopes of nitrogen (N) were used to infer historical changes in nutrient
118 sources arising from changes in aquatic production (N₂ fixation), agriculture within the
119 watershed, or regional urban development (Bunting et al., 2016; Leavitt et al., 2006). These
120 proxies can be used to provide a comprehensive overview of changes to algal abundance,
121 production, and composition, suitable to evaluate water quality status. This information is
122 valuable to the St. Lawrence River AOC, as beneficial uses must be restored to all 14 BUIs prior
123 to delisting (International Joint Commission, 2012), including reductions in symptoms of
124 eutrophication and the presence of undesirable algae.

125 **Methods**

126 *Study Area*

127 The St. Lawrence River at Cornwall, Ontario, Canada, marks the eastern end of the
128 international section of the waterway and is located just downstream of the Moses-Saunders
129 Power Dam. East of the city of Cornwall, the river widens into Lake St. Francis for 50 km before
130 narrowing again as it passes around Grande-Île, near Salaberry-de-Valleyfield, Quebec (Fig. 1a).
131 Lake St. Francis covers approximately 233 km², with a mean depth of 6 m (maximum 26 m),
132 short hydraulic residence time (3 days) and a total volume of 2.8 km³ (Anderson et al., 1992;
133 Fortin et al., 1994). Water level is controlled in this portion of the St. Lawrence River by the
134 Moses-Saunders Power Dam upstream and the Coteau works and Beauharnois hydroelectric

135 generating station downstream (Anderson et al., 1992). Water levels in the St. Lawrence River
136 are regulated by the International Joint Commission to stabilise Lake Ontario and to ensure
137 adequate capacity for navigation, hydroelectric power generation, and flood control (Yee et al.,
138 1990). In LSF, Hydro Quebec manages the downstream discharge through the Beauharnois dam
139 such that water level variation is typically <20 cm (Morin and Leclerc, 1998). Approximately
140 95% of the flow in LSF comes from Lake Ontario, with the remainder originating from
141 tributaries on the north and south shores (Anderson et al., 1992). Little mixing occurs across the
142 main shipping channel, which divides the north and south portions of LSF, each of which is
143 differently influenced by local inflow tributaries (International Joint Commission, 2003b). As a
144 result, the main channel and the flows north and south thereof can be considered to be three
145 distinct water bodies (Dreier et al., 1997). On the northern shore, nine Ontario watersheds drain
146 into LSF, the largest of which, the Raisin River watershed, covers over 500 km² (Fig. 1b, c).
147 Across the northern watersheds, the dominant agricultural products are corn and soybeans,
148 accounting for 15% and 14% of land use, respectively, with other dominant land cover including
149 forest (43%), pasture and forages (15%), and urban and developed areas (8%; 2015 annual crop
150 inventory data from Agriculture and Agri-Food Canada,
151 <http://open.canada.ca/data/en/dataset/3688e7d9-7520-42bd-a3eb-8854b685fef3>, accessed 25
152 July, 2017).

153 In the deep, fast-flowing channels of the river, sedimentation does not reliably occur
154 (Carignan and Lorrain, 2000), making the collection of a sediment core representative of past
155 conditions unlikely from deeper sites. Several areas in LSF also have been disturbed previously
156 by dredging activities when the shipping channel was created as part of the construction of the
157 St. Lawrence Seaway in the 1950s (Morin and Leclerc, 1998); such areas were avoided for the

158 current study to ensure a continuous, undisturbed sedimentary record. In the AOC, five
159 sedimentation basins have been described (Lorrain et al., 1993), two of which are on the northern
160 side of the main channel of the St. Lawrence River and are likely to be influenced by flows from
161 the northern tributaries. Sediment cores with reliable, continuous dating profiles have previously
162 been collected from both of these basins (Carignan and Lorrain, 2000). The more westerly of
163 these two basins, located just east of Lancaster, Ontario, is in a portion of the river that has seen
164 extensive water-quality monitoring take place since 2010 (Bramburger, 2014; Savard et al.,
165 2013, 2015). Both the availability of recent monitoring data and the known sedimentation
166 characteristics of the basin influenced the selection of this site for sample collection.

167 *Sediment Collection*

168 A sediment core was collected on May 5, 2016 from the St. Lawrence River near
169 Lancaster, Ontario, Canada (74°27'47"W, 45°08'07"N; Fig. 1b) using a modified gravity corer
170 (Glew, 1989) with an internal diameter of 7.62 cm. The collection site was located
171 approximately 900 m from shore, 2.5 km downstream from the outlet of the Raisin River (Fig.
172 1b). The core was collected from a depth of 5 m to minimise sediment mixing (Carignan and
173 Lorrain, 2000; Lepage et al., 2000). Additionally, this location was selected for sample
174 collection to best achieve proximity to tributary inputs while remaining deep enough to
175 experience permanent sediment deposition without resuspension.

176 The collected core was sectioned in the field into 0.5-cm increments which were bagged,
177 transported in the dark to Queen's University, and stored in the dark at ~4°C until analysis.
178 Subsamples were taken for determination of sediment ages using gamma spectroscopy, pigment
179 concentrations using high performance liquid chromatography (HPLC), organic matter content

180 via loss-on-ignition (LOI), stable isotope analyses using mass spectrometry, and diatom
181 assemblages using light microscopy.

182 *Chronology*

183 Gamma spectroscopy was used to measure the activities of total ^{210}Pb , ^{214}Pb and ^{214}Bi
184 (proxies of supported ^{210}Pb), and ^{137}Cs following the methods of Schelske et al. (1994) in 25
185 intervals throughout the core. Sediments were freeze-dried, then approximately 1 g dry mass was
186 sealed into counting tubes using 2-Ton[®] Epoxy. Samples were left for 2 weeks for in situ decay
187 of ^{226}Ra to stabilise. The constant rate of supply (CRS) calculation of Appleby and Oldfield
188 (1978) was used to estimate sediment ages using unsupported ^{210}Pb activity in conjunction with
189 ^{137}Cs activity, an independent indicator of the year 1963 (Appleby, 2002).

190 *Pigments*

191 Frozen subsamples of whole sediments were taken for determination of photosynthetic
192 pigment concentrations from 37 intervals throughout the core. Pigments were extracted and
193 quantified using high performance liquid chromatography (HPLC) following the protocol
194 outlined in Leavitt and Hodgson (2001). Briefly, frozen samples were freeze-dried, and
195 approximately 0.05 g of dried sediment was extracted using a mixture of acetone:methanol:water
196 (80:15:5, by volume) to extract chlorophylls, carotenoids, and their derivatives. Extracts were
197 evaporated to dryness under a stream of N_2 , then redissolved in injection solvent containing
198 Sudan II dye as an internal standard. Pigment concentrations are reported as nmoles pigment/g
199 organic matter. HPLC analyses were restricted to common taxonomically-diagnostic pigments
200 including fucoxanthin (siliceous algae), diatoxanthin (mainly diatoms), alloxanthin
201 (cryptophytes), phaeophytin *b* (chlorophytes), echinenone (total cyanobacteria), canthaxanthin
202 (Nostocales cyanobacteria), and β -carotene (all phytoplankton). In addition, lutein (chlorophytes)

203 and zeaxanthin (cyanobacteria) were not separable on our HPLC system and were used as an
204 index of bloom-forming taxa (Leavitt et al., 2006; Leavitt and Hodgson, 2001).

205 *Organic Matter and Stable Isotopes*

206 Percent organic matter was determined through standard LOI procedures (Dean, 1974) in
207 25 intervals throughout the sediment core. Briefly, a known mass (~0.08 g) of freeze-dried
208 sediment was tared and combusted at 550°C for 4 hours to determine organic content, then
209 ignited at 950°C for 2 hours to determine carbonate content. Stable isotopes of N ($\delta^{15}\text{N}$) and C
210 ($\delta^{13}\text{C}$), as well as elemental N and C contents, were determined using isotope ratio mass
211 spectrometric (IRMS) analysis of 0.01-0.015 g of freeze-dried sediment following Savage et al.
212 (2010). IRMS was performed using a Thermoquest (Finnigan-MAT) Delta Plus^{XL} mass
213 spectrometer coupled with a Carlo Erba NC2500 elemental analyser (Savage et al., 2010).
214 Isotope values are presented as per mille (‰) differences of samples to standard references for
215 each element (Savage et al., 2010). Sediment elemental composition is reported as the mass ratio
216 of C:N, as determined through the elemental analyser.

217 *Diatoms*

218 Diatom slurries were prepared for enumeration after removal of organic matter using an
219 acid digestion procedure. Briefly, approximately 0.2-0.3 g of whole wet sediment was
220 subsampled at 1-cm intervals into 20-mL glass scintillation vials. Known masses of sediment
221 were mixed with a 50:50 (molar) solution of sulphuric and nitric acids overnight, then digested in
222 a hot water bath at 70°C for 8 hours. Diatoms were allowed to settle for 24 hours, after which the
223 supernatant above the settled diatoms was aspirated. Scintillation vials were then refilled with
224 double-deionised water, and samples were agitated to resuspend the diatoms. Samples were
225 rinsed until the pH was the same as deionised water, as verified with litmus paper (typically eight

226 rinses). Samples were then spiked with a solution of microspheres (mean diameter = 7.9 μm) of
227 known concentration (34,000 microspheres/mL). Samples were plated on coverslips in a series of
228 four dilutions and allowed to evaporate, after which they were fixed permanently to slides using
229 Naphrax[®], a medium with a high refractive index (>1.7).

230 Diatom valves were identified and enumerated using a Leica (DMRB model) microscope
231 fitted with a 100x fluotar objective (numerical aperture of objective = 1.3) and using differential
232 interference contrast optics at 1000x magnification. Diatoms were identified to species wherever
233 possible, or to the lowest possible taxonomic classification. Valves were counted until a
234 minimum of 400 valves were enumerated, or, if the concentration of valves was exceptionally
235 low, until five transects were completed. Primary taxonomic keys used for diatom identification
236 were Krammer and Lange-Bertalot (1991a, 1991b, 1988, 1986) and Reavie and Smol (1998).

237 The main chronological zones of diatom species assemblages were estimated using a
238 constrained incremental sum of squares analysis (CONISS; Grimm, 1987), performed in the R
239 computing environment (R Core Team, 2015) and the *rioja* (Juggins, 2015) and *vegan* (Oksanen
240 et al., 2015) packages. Diatom abundances were Hellinger-transformed (Rao, 1995) prior to
241 CONISS analysis using Euclidean distance, to minimise distortions that can occur when zero
242 values are present (Legendre and Legendre, 2012). A broken stick model (Bennett, 1996) was
243 used to determine the number of significant zones in the stratigraphic sequence.

244 **Results**

245 *Chronology*

246 The total ²¹⁰Pb activity decreased from the top of the sediment core and followed an
247 exponential decay ($r^2 = 0.83$; Fig. 2). Both ²¹⁴Pb and ²¹⁴Bi activities remained relatively constant
248 throughout the core, and are consistent with previously collected sediment cores from LSF which

249 have reported supported ^{210}Pb activities of approximately 20 Bq/kg (Carignan and Lorrain,
250 2000). ^{137}Cs activity reached a distinct peak at a depth of 18.25 cm.

251 Application of the CRS model to determine sediment ages and sedimentation rates from
252 the unsupported ^{210}Pb activities suggested that dates were reliable until approximately 30-cm
253 burial depth (ca. 1940; Fig. 2). Although local error estimates are large, sedimentation rates
254 appear to have increased substantially between depths of approximately 25 and 20 cm (1955-
255 1959) in the core. The depth at which the year 1963 occurred was agreed-upon by the ^{210}Pb
256 model and the analysis of ^{137}Cs activity (~18 cm). Given this dating profile, approximate depths
257 in the core for the designation of the AOC (1987) and the release of the Stage 1 (1992) and Stage
258 2 (1997) RAP reports are 9.75 cm, 8.75 cm, and 7.25 cm, respectively.

259 *Pigments*

260 Analysis of concentrations of all pigment biomarkers suggested a progressive increase in
261 lake production during the 20th century (Fig. 3). In general, total phytoplankton abundance (as β -
262 carotene) was stable from the base of the core to ca. 1960, after which time inferred abundance
263 increased approximately twofold to an irregular plateau after ca. 1980. As ratios of labile to
264 stable pigments (chlorophyll *a*:phaeophytin *a*) did not change in the ca. 1960-1980 interval, we
265 infer that elevated concentrations of pigments reflect actual increases in mean water column
266 standing stock, rather than alterations in the preservation environment. Although timing of the
267 concentration change varies slightly among algal groups, concentrations of chemically-stable
268 biomarkers from total (echinenone) and colonial cyanobacteria (canthaxanthin), chlorophytes
269 (phaeophytin *b*), diatoms (diatoxanthin) and cryptophytes (alloxanthin) all exhibited similar
270 patterns, with elevated abundance of these groups after the mid-1970s. In contrast, levels of
271 labile pigments from siliceous (fucoxanthin) and total algae (chlorophyll *a*) declined

272 exponentially with burial depth, suggesting rapid degradation following deposition, especially in
273 the most recent 5 years. Finally, concentrations of many chemically-stable pigments increased
274 sharply after ca. 2005, suggesting a recent increase in either algal abundance or changes in
275 sedimentary preservation.

276 *Organic Matter and Stable Isotopes*

277 The organic matter content remained relatively constant throughout the core, varying
278 between 8% and 13% of dry mass (Fig. 4). In contrast, $\delta^{15}\text{N}$ values were stable ($\sim 5\text{‰}$) from the
279 bottom of the core until the early 1960s ($\sim 20\text{-cm}$ depth) and then increased steadily towards 7‰
280 at the top of the core, in a pattern similar to the changes in total algal abundance (Fig. 3). The
281 $\delta^{13}\text{C}$ values increased rapidly from approximately -23‰ at $\sim 30\text{-cm}$ depth (early 20th century) to -
282 16‰ at 25 cm (ca. early 1950s), before declining to consistent values of approximately -20‰ in
283 sediments deposited in the upper 20 cm of the core (after 1960; Fig. 4). The C:N mass ratio
284 followed a similar trajectory to that of $\delta^{13}\text{C}$, starting at approximately 15 at the base of the core,
285 increasing to ~ 23 at a depth of 25 cm (ca. 1960), then gradually decreasing until the top of the
286 core where it reached a value of 10 (Fig. 4).

287 *Diatoms*

288 Concentrations of diatoms increased exponentially from $\sim 1 \times 10^5$ valves/g dry mass prior
289 to ca. 1960 to greater than 4×10^6 valves/g dry mass in surface deposits (Fig. 5). After ca. 1970,
290 diatom concentrations increased in a pattern moderately correlated ($r^2 = 0.57$, $p < 0.001$) with
291 fossil concentrations of the labile biomarker of siliceous algae (fucoxanthin). Constrained cluster
292 analysis using CONISS distinguished two main zones of diatom species assemblages, which
293 appear to correspond to intervals before and after construction of the Moses-Saunders Power
294 Dam (1954-1959; 23.25-cm depth). The older assemblage was characterised by higher

295 abundances of *Fragilaria construens* (Ehrenberg) Grunow, *Sellaphora submuralis* (Hustedt)
296 Wetzel, Ector, Van de Vijver, Compère, & Mann, *Achnanthes clevei* Grunow, and *Cocconeis*
297 *neothumensis* Krammer. Above 23.25 cm, several taxa became much more prevalent, including
298 *Achnantheidium minutissimum* (Kützing) Czarnecki, *F. capucina* Desmazières, *C. placentula*
299 (Ehrenberg) Grunow, *Navicula cryptotenella* Lange-Bertalot, and *Nitzschia fonticola* Grunow.
300 Other taxa, such as *Amphora pediculus* (Kützing) Grunow ex Schmidt, *Staurosirella pinnata*
301 (Ehrenberg) Williams & Round, *Pseudostaurosira brevistriata* (Grunow) Williams & Round,
302 and *Planothidium lanceolatum* (Brébisson ex Kützing) Lange-Bertalot, were present in relatively
303 high abundances throughout the core. Diatoms identified were predominantly benthic taxa.

304

Discussion

305 Analyses of fossil pigments, diatoms, and stable isotopes revealed a progressive increase
306 in the abundance of primary producers in this portion of LSF during the late-20th century,
307 continuing into the 21st century (Figs. 3-5). In general, algal abundance and community
308 composition were relatively stable prior to the 1954-1959 construction of the Moses-Saunders
309 Power Dam and the St. Lawrence Seaway, with low and constant concentrations of biomarker
310 pigments from diatoms (diatoxanthin) and chlorophytes (phaeophytin *b*, lutein-zeaxanthin), and
311 lower abundances of total (echinenone) and colonial (canthaxanthin) cyanobacteria. Diatom
312 assemblage composition changed after dam construction (ca. 1960), though assemblages before
313 and after this period were both characterised by benthic taxa, none of which indicated a change
314 to LSF trophic status. Total algal abundance appears to have increased after ca. 1970, with an
315 approximate two-fold increase in fossil concentrations of most pigment biomarkers (Fig. 3), but
316 no marked change in the preservation environment, as recorded by the degradation index of
317 labile chlorophyll *a* to stable phaeophytin *a* (Leavitt and Hodgson, 2001). This increase in

318 abundance occurs in parallel with elevated nutrient supply inferred from the $\delta^{15}\text{N}$ signal (Fig. 4).
319 Microfossil and labile pigments from diatoms were particularly abundant after ca. 2005, as were
320 concentrations of chemically stable carotenoids from total cyanobacteria (echinenone, lutein-
321 zeaxanthin) and cryptophytes (alloxanthin) but not those from chlorophytes (phaeophytin *b*) or
322 total algae (β -carotene). Overall, this pattern shows that water quality along the north shore of
323 LSF did not improve as a result of local and regional remedial actions implemented in the early
324 1990s and suggests that substantial additional measures to curb nutrient influxes from regional
325 and headwater sources are required if the AOC delisting goals relating to eutrophication and
326 undesirable algae are to be achieved.

327 *Baseline Conditions (pre-1950s)*

328 Prior to the construction of the Moses-Saunders Power Dam and the St. Lawrence
329 Seaway in the mid-1950s, conditions were stable, with relatively constant concentrations of most
330 photosynthetic pigments, low and steady concentrations of diatoms, and diatom assemblages
331 characterised by predominantly benthic taxa. Organic material in the aquatic environment was
332 supplied by both autochthonous and allochthonous sources, as indicated by the moderate and
333 stable molar ratio of C:N (Meyers and Ishiwatari, 1993; Fig. 4). An exact chronology is difficult
334 to assign to this portion of the sediment core, as errors associated with sediment dating are large
335 (Fig. 2); however, we are confident that the bottom four intervals represent a period of time prior
336 to the 1950s. Although we cannot consider these records to represent pristine conditions, as
337 industrial activity had been occurring upstream of our site in Cornwall, Ontario, since the late-
338 19th century (Stein, 1995), we will refer to them as baseline conditions which represent a period
339 prior to the major anthropogenic changes that occurred in our system during the second half of
340 the 20th century.

341 *St. Lawrence Seaway and Power Dam (mid-1950s)*

342 Between 1954 and 1959, two major construction projects occurred in this portion of the
343 St. Lawrence River: the construction of the Moses-Saunders Power Dam and the dredging of the
344 St. Lawrence Seaway. These concurrent events appear to be represented in our paleolimnological
345 record through marked changes in $\delta^{13}\text{C}$ values and C:N ratios, as well as diatom species
346 composition. For example, although the $\delta^{15}\text{N}$ values remained relatively stable through the
347 1950s, the C:N ratio increased quickly at this time, indicating a substantial increase in the
348 terrestrial fraction of organic matter entering the system (Meyers and Ishiwatari, 1993). At the
349 same time, a sharp increase in the $\delta^{13}\text{C}$ signal (Fig. 4) to values characteristic of regional
350 terrestrial plants suggests an increase in organic matter subsidies from adjacent farms (Meyers
351 and Ishiwatari, 1993). Elevated influxes of terrestrial organic matter most likely arose from the
352 construction of the Moses-Saunders Power Dam, which flooded more than 75 km² of land, much
353 of it agricultural, on July 1, 1958 (Macfarlane, 2014). High sedimentary $\delta^{13}\text{C}$ values in the 1950s
354 may be additionally driven by an elevated proportion of C4 plants such as corn in the watershed.
355 At present, corn is a predominant crop within the local catchment area (2015 annual crop
356 inventory data from Agriculture and Agri-Food Canada,
357 <http://open.canada.ca/data/en/dataset/3688e7d9-7520-42bd-a3eb-8854b685fef3>, accessed 25
358 July, 2017), although we recognise that it is difficult to distinguish among potential plant sources
359 of C from an analysis of bulk sediment isotopic values.

360 Analysis of the fossil diatom assemblages using stratigraphically constrained hierarchical
361 cluster analysis revealed only a single transition in species assemblages, which occurred in the
362 late-1950s, coinciding with the construction of the Moses-Saunders Power Dam and the St.
363 Lawrence Seaway (Fig. 5). Previous research at the eastern end of LSF has suggested that an

364 increase in epiphytic diatom taxa and inferred higher macrophyte coverage occurred in the early-
365 to mid-20th century, possibly attributable to a decrease in the variability of the water level
366 resulting from Seaway construction and the construction of water control structures at the eastern
367 end of LSF (Reavie et al., 1998). Seasonal water level variability in LSF was known to exceed
368 0.5 m in the first half of the 20th century, but this variability was reduced to less than 0.2 m after
369 the construction of the Moses-Saunders Power Dam (Morin and Leclerc, 1998). Although some
370 epiphytic diatom taxa (e.g., *Cocconeis placentula*) were observed in the current study to be more
371 abundant after the construction of the Moses-Saunders Power Dam, possibly attributable to
372 higher macrophyte coverage due to stabilisation of water levels, many of the diatom taxa that
373 were abundant following dam construction (e.g., *Fragilaria capucina*, *Amphora pediculus*) can
374 be commonly found on other substrata (e.g., rocks; Reavie and Smol, 1997). The diatom
375 assemblages before and after the late-1950s share many characteristics, such as being
376 predominantly benthic taxa with no strong trophic status affiliations, with some epiphytic taxa
377 present. It seems likely that the two major construction projects in the St. Lawrence River in the
378 1950s caused a substantial disturbance to the aquatic environment (as indicated by the abrupt
379 terrestrial loading suggested by the C:N ratio), which allowed a slightly different diatom
380 assemblage to settle and thrive once the disturbance was over.

381 *20th Century Eutrophication (1960s-1970s)*

382 In the 1960s and 1970s, the lower Laurentian Great Lakes were characterised by
383 intensive eutrophication and related algal blooms and hypoxia (Beeton, 1965; Mortimer, 1987;
384 Schelske, 1991), events which are represented in our core by increases in photosynthetic
385 pigments and diatoms. The pigment data suggest that production increased first in the early- to
386 mid-1970s, indicated particularly by biomarkers derived from bloom-forming cyanobacteria

387 (canthaxanthin), chlorophytes (phaeophytin *b*), diatoms (diatoxanthin), and total production (β -
388 carotene). This trend is supported by an increase in diatom production between the early-1960s
389 and 1970s, during which a 4-fold increase in fossil frustule concentration occurred, and a
390 previously reported mid-20th century increase in eutrophic diatom taxa at the eastern end of LSF
391 (Reavie et al., 1998). At present, we cannot easily distinguish between elevated production in
392 LSF due to inputs of nutrient- and phytoplankton-rich waters from the upstream Great Lakes,
393 and elevated production due to eutrophication of the LSF basin from local nutrient influx, as
394 historical surface water-quality monitoring data are limited. However, nutrient monitoring of the
395 Raisin River, a major tributary near the sampling location of the current study, indicates that,
396 after 1976, high ($>30 \mu\text{g/L}$) and variable TP concentrations have occurred (data from Ontario
397 Ministry of the Environment and Climate Change, [https://www.ontario.ca/data/provincial-](https://www.ontario.ca/data/provincial-stream-water-quality-monitoring-network)
398 [stream-water-quality-monitoring-network](https://www.ontario.ca/data/provincial-stream-water-quality-monitoring-network), site 12007300302, accessed 19 December, 2016).

399 N influx to LSF appears to have increased markedly during the 1960s and 1970s, as
400 indicated by persistent increases in sedimentary $\delta^{15}\text{N}$ values (Fig. 4). As noted elsewhere
401 (Leavitt et al., 2006; Savage et al., 2010; Bunting et al., 2016), the addition of anthropogenic
402 reactive N to terrestrial and aquatic systems often results in the enrichment of adjoining water
403 bodies due to microbial processing of added N and loss of depleted N to the atmosphere.
404 Similarly, the values of $\delta^{15}\text{N}$ in aquatic food webs (Anderson and Cabana, 2005) and stream
405 nitrate (Harrington et al., 1998) have been positively correlated with agricultural land use in the
406 surrounding catchment. Consistent with this mechanism, the sharp increase in fertilisation of
407 Ontario farmlands with N between the 1960s and 1980s (Smith, 2015) is expected to have
408 favoured the elevation of both flux and isotopic values of N in runoff into the St. Lawrence
409 River.

410 *Response to Modern Management (post-1990s)*

411 Since the designation of the St. Lawrence River at Cornwall, Ontario, as an AOC,
412 remediation efforts have successfully targeted many of the BUIs. For example, water quality has
413 improved through upgrades to the Cornwall wastewater treatment plant, remediation of
414 decommissioned industrial sites (e.g., chemical manufacturing facilities), and legislation
415 restricting concentrations of harmful substances in wastewater effluent from industrial facilities
416 (Environment Canada et al., 2007; Environment Canada and Ontario Ministry of the
417 Environment, 2010). Similarly, progress has been made on improving fish and wildlife
418 biodiversity and condition through wetland construction, habitat protection programs, and
419 changes to fishing regulations (Environment Canada et al., 2007). As well, many remedial
420 actions in the AOC have targeted the issue of eutrophication and undesirable algae. For example,
421 a tributary restoration program initiated in 1994 has led to the planting of over 300,000 trees in
422 riparian areas, increased buffer zones and cattle exclusion fencing along waterways, upgraded
423 manure storage facilities and rural septic systems, increased well protection projects, and more
424 (Environment Canada et al., 2007). However, despite these efforts, our data suggest that the
425 eutrophication and undesirable algae BUI remains impaired, with continuously elevated algal
426 abundance (as pigment and diatom concentrations) since the 1990s and no evidence of recovery
427 to lower abundances.

428 Pronounced increases in sedimentary pigment content during the past 10 years were
429 observed for echinenone, a chemically stable biomarker for total cyanobacteria (Leavitt and
430 Hodgson, 2001), and are consistent with increased reports of potentially toxic cyanobacteria in
431 recent years (Bramburger, 2014; Savard et al., 2013, 2015). Elevated cyanobacterial abundance
432 could be attributable to several factors, including high nutrient concentrations (Downing et al.,

433 2001); although lotic TP concentrations from the Raisin River have not increased in recent years,
434 values have remained persistently high ($> 30 \mu\text{g/L}$) since the early-1990s (data from Ontario
435 Ministry of the Environment and Climate Change, [https://www.ontario.ca/data/provincial-](https://www.ontario.ca/data/provincial-stream-water-quality-monitoring-network)
436 [stream-water-quality-monitoring-network](https://www.ontario.ca/data/provincial-stream-water-quality-monitoring-network), site 12007300302, accessed 19 December, 2016), and
437 similarly high TP concentrations have been reported in the nearshore areas of tributary mouths
438 (Savard et al., 2015). Given these consistently high TP values and the evidence of persistent
439 increases in sedimentary nutrient influx in LSF ($\delta^{15}\text{N}$ in Fig. 4), it appears likely that recent
440 cyanobacterial growth has been influenced by nutrient inputs. Upstream of our study location,
441 surface-water chlorophyll *a* concentrations collected at Kingston and Brockville have dropped
442 substantially since the 1980s, from approximately 2-5 $\mu\text{g/L}$ to $< 1 \mu\text{g/L}$ (data from Ministry of
443 the Environment and Climate Change, [https://www.ontario.ca/data/lake-water-quality-drinking-](https://www.ontario.ca/data/lake-water-quality-drinking-water-intakes)
444 [water-intakes](https://www.ontario.ca/data/lake-water-quality-drinking-water-intakes), stations 020170010 and 020180011, accessed 3 September 2017), suggesting that
445 production has not increased in the main river channel flowing into LSF and supporting local
446 nutrient inputs as an influence to cyanobacterial growth. In general, modelling of nutrient fluxes
447 in St. Lawrence River catchments shows that net anthropogenic inputs of nitrogen and
448 phosphorus have increased throughout the 20th century, with pronounced effects of agricultural
449 fertilisers during the past 50 years (Goyette et al., 2016), which may be particularly relevant
450 given the high proportion of agricultural lands in the contributing watersheds to our study site
451 (Fig. 1c). Unfortunately, without refined hydrologic models of water flow in the LSF nearshore
452 region, it is difficult to identify which catchments may be fertilising waters in the AOC. In
453 particular, combining flow modelling with nutrient modelling (Goyette et al., 2016) might allow
454 for the determination of priority areas for nutrient monitoring in the LSF nearshore.

455 It is possible that other factors have also affected the recent cyanobacterial growth in our
456 study area, though these influences are difficult to assess without in-depth analyses. Recently,
457 persistent eutrophication issues have been described in areas where monitoring data suggest that
458 nutrient concentrations have declined (Alexson et al., 2017), which may be explained in lakes by
459 hypoxia-induced internal phosphorus loading resulting from stronger thermal stratification
460 related to climate change (North et al., 2014). Although it is unlikely that internal phosphorus
461 loading is providing an additional source of nutrients at our shallow, well-oxygenated, fluvial
462 site, hydrologic changes relating to catchment land use and climate change may have altered
463 nutrient delivery to LSF. For instance, pronounced land-use changes in the Raisin River
464 watershed have occurred between 1990 and 2010, with urban areas increasing by 12% and the
465 extent of treed and forested areas decreasing by 20% (data from Agriculture and Agri-Food
466 Canada, <http://open.canada.ca/data/en/dataset/02202cec-b4a1-4a1d-9997-edcbaca92d41>,
467 <http://open.canada.ca/data/en/dataset/9e1efe92-e5a3-4f70-b313-68fb1283eadf>, accessed 3
468 September, 2017). Although land use devoted to crops in this watershed has only increased by
469 0.5%, deforestation and urbanisation could have increased runoff (Hundecha and Bárdossy,
470 2004), facilitating the transport of nutrients from the watershed to LSF. Climate change may
471 have also favoured the proliferation of cyanobacteria, either directly (i.e., increased water
472 temperatures; Paerl and Huisman, 2008) or indirectly (e.g., increased frequency of droughts and
473 floods; Paerl et al., 2011). Temperature and precipitation trends in the Great Lakes-St. Lawrence
474 River Basin have climbed throughout the 20th century (Magnuson et al., 1997), and temperature
475 increases and hydrologic changes are expected to continue to occur in St. Lawrence River
476 tributaries throughout the 21st century, particularly in the winter and spring months (Boyer et al.,
477 2010). Though it is impossible to assess the influences of land-use change and climate change on

478 cyanobacterial abundance within the confines of the current study, it is unlikely that these factors
479 have not affected cyanobacterial growth and should be more thoroughly investigated.

480 **Conclusion**

481 Though numerous actions have targeted reducing nutrient inputs to LSF in the past 20
482 years, we found that algal abundance has not decreased in response to remediation efforts, and
483 that, in fact, populations of cyanobacteria appear to have expanded during the past decade. The
484 causal mechanism for this increase is not immediately clear, but is likely related to continuously
485 high nutrient concentrations in major LSF tributaries, possibly combined with major land-use
486 changes and climate change. The potential for toxin-producing cyanobacterial blooms is
487 particularly troubling for both local and downstream residents, and the cyanobacterial
488 communities of LSF and its tributaries should be monitored closely for the presence of
489 potentially toxin-producing species. As the AOC committee works toward delisting, it is
490 important to recognise that, despite successes in other areas, the sediment record demonstrates
491 continuing impacts to water quality in LSF over the past two decades, indicating that the
492 eutrophication and undesirable algae BUI remains in need of remediation.

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Figure Captions

Figure 1. A) The Laurentian Great Lakes and St. Lawrence River with inset indicating towns of interest and the Area of Concern (the hatched area). B) Sampling location (indicated by the “X”), Lake St. Francis bathymetry, and nearby tributaries. C) Bathymetry of Lake St. Francis within the Area of Concern, land use of the nine major watersheds in Ontario that contribute to the St. Lawrence River Area of Concern, and locations of nearby dams and Cornwall wastewater treatment plant (WWTP); the extent of panel B is indicated by the rectangle.

Figure 2. A) Activities and errors of the four radioisotopes, by depth in the sediment core, measured through gamma spectroscopy. B) Inferred year, sedimentation rate, and associated errors as calculated through the constant rate of supply (CRS) model.

Figure 3. Concentrations of photosynthetic pigments (per gram organic matter) throughout the sediment core. Secondary y-axis indicates year inferred from the constant rate of supply (CRS) dating model. Top panel: more stable pigments, defined as a category 1 (Leavitt and Hodgson, 2001). Bottom panel: more labile pigments, defined as a category 2, 3, or 4 (Leavitt and Hodgson, 2001) and ratio of chlorophyll *a* to phaeophytin *a*, an indicator of the extent of pigment degradation.

Figure 4. Percent organic matter, per mille ratios of stable isotopes, and mass ratio of carbon to nitrogen by depth and year inferred through the constant rate of supply (CRS) dating model.

734 *Figure 5.* A) Diatom valve concentrations by depth and year inferred through the constant rate of
735 supply (CRS) dating model. B, C) Relative abundances of the dominant (B) and subdominant (C)
736 diatom species observed in the core. The dotted line represents the significant assemblage change
737 identified by the broken stick model and constrained incremental sum of squares.

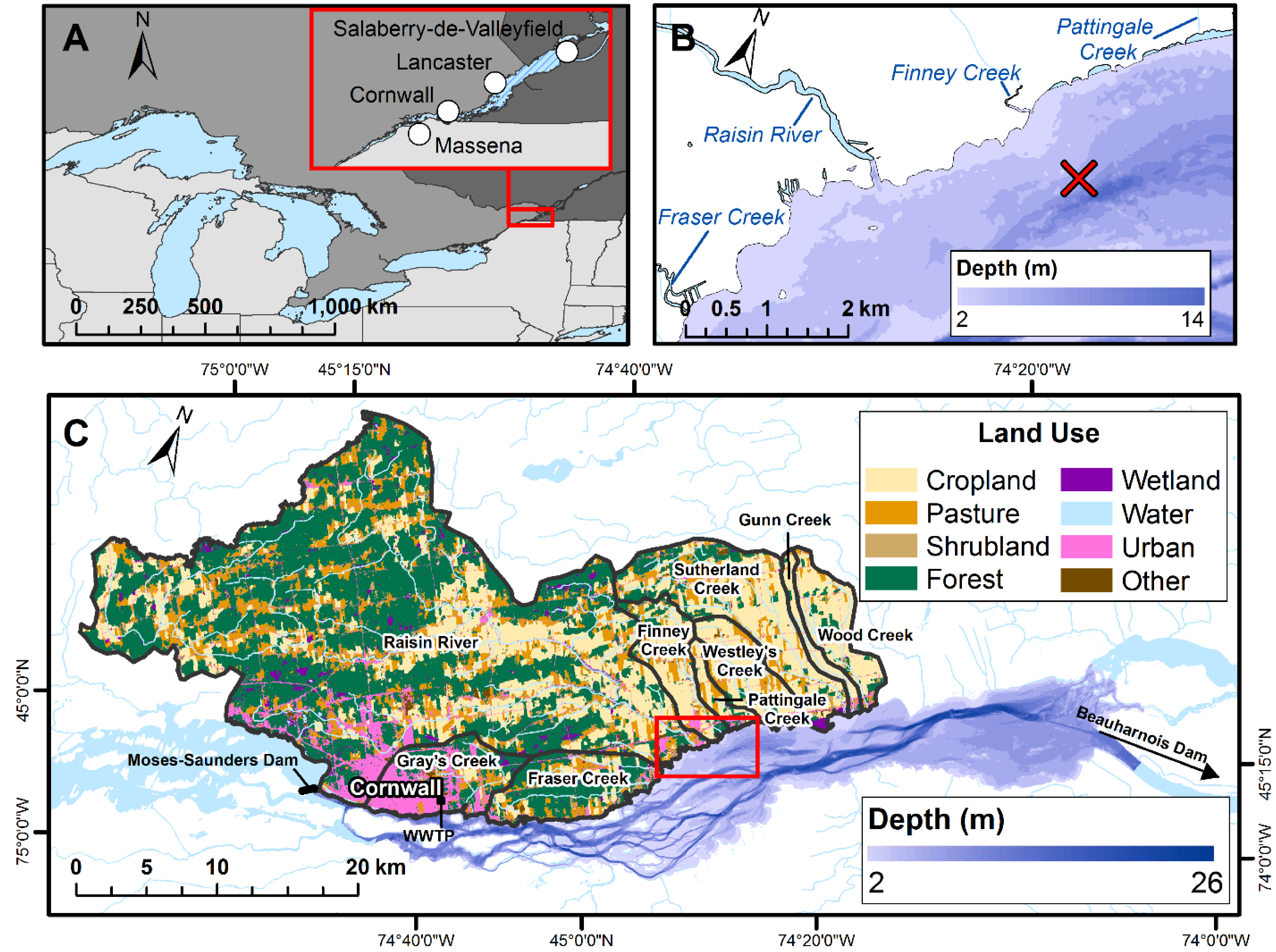


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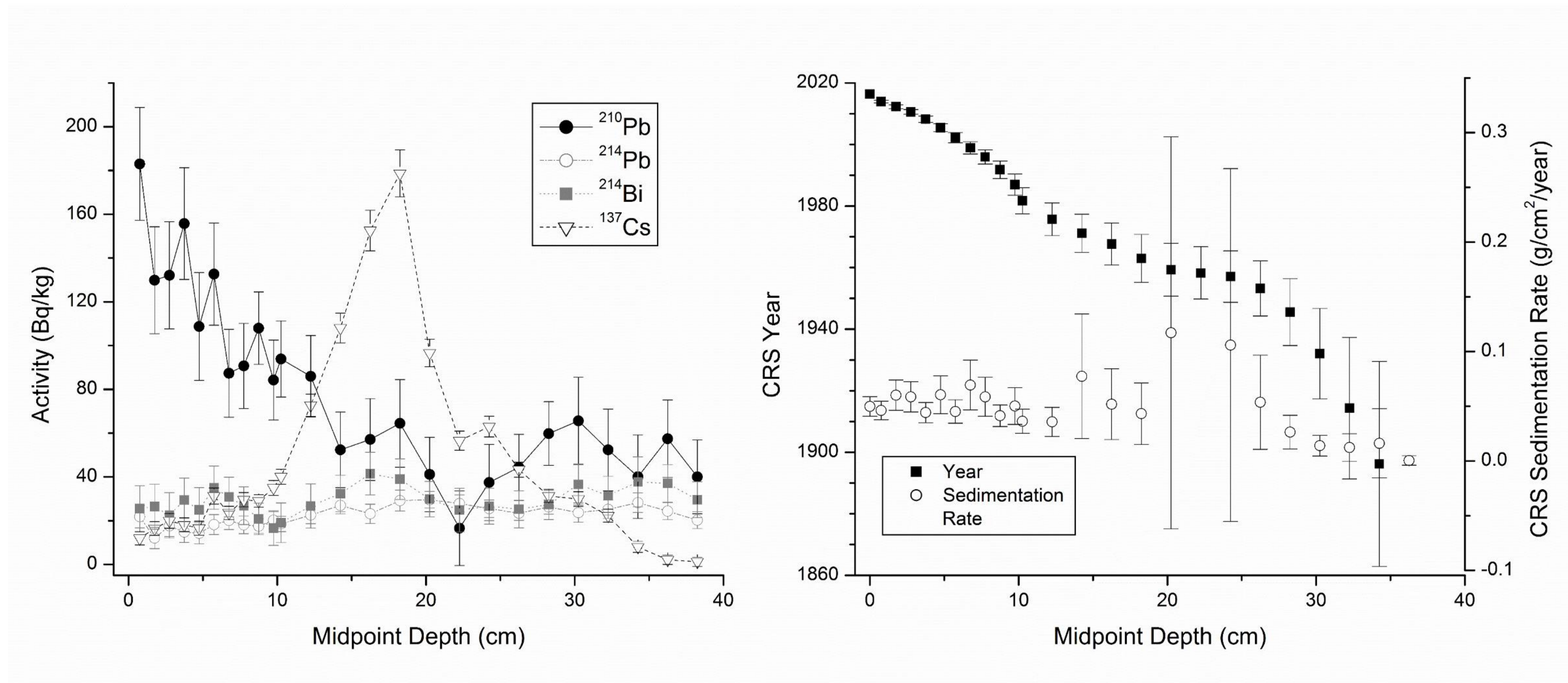


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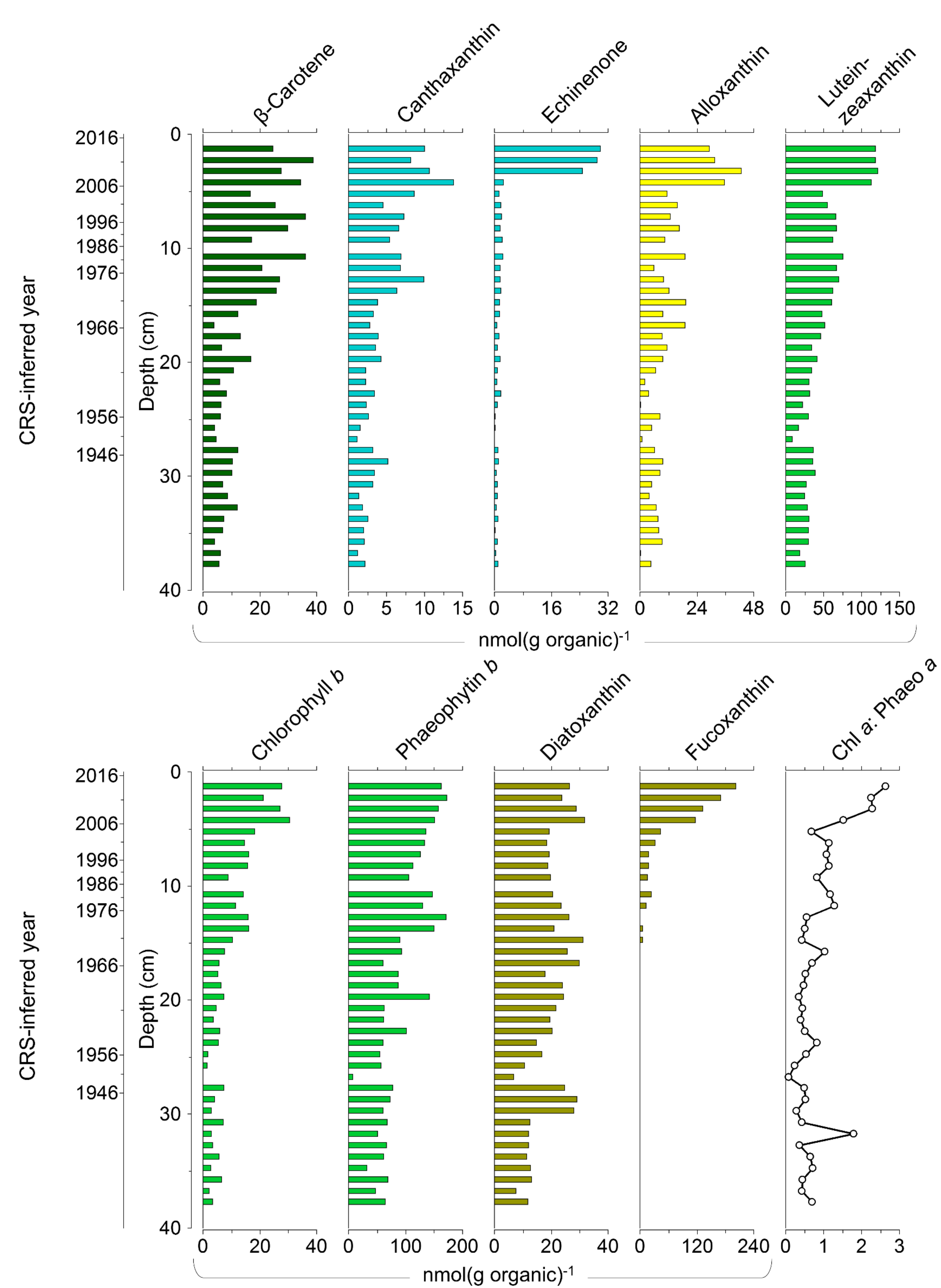


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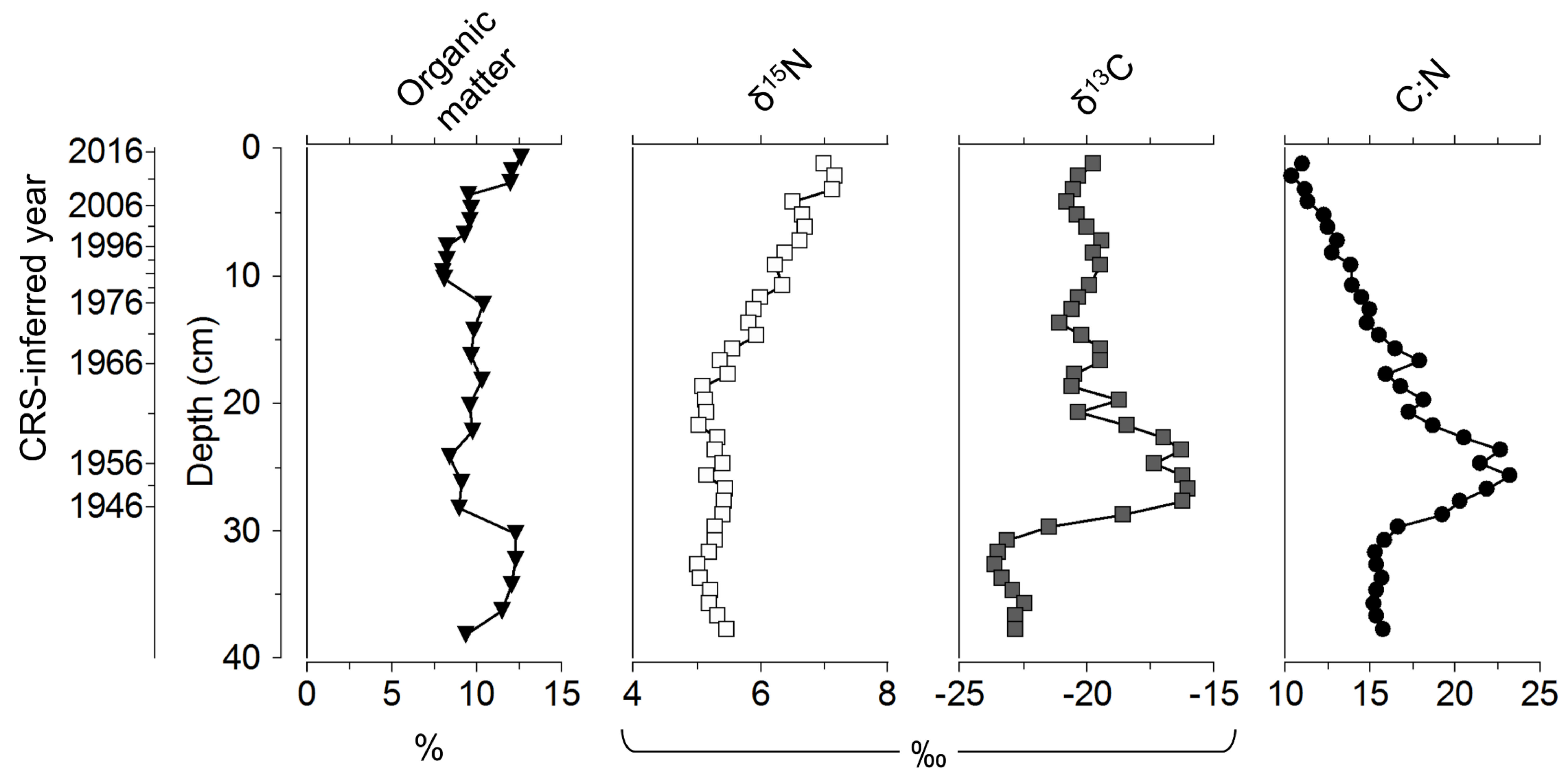


Figure 4. Percent organic matter and ratios of stable isotopes by depth and year inferred through the constant rate of supply (CRS) dating model. The box indicates inferred years during which the Moses-Saunders Power Dam was constructed (1954-1959).

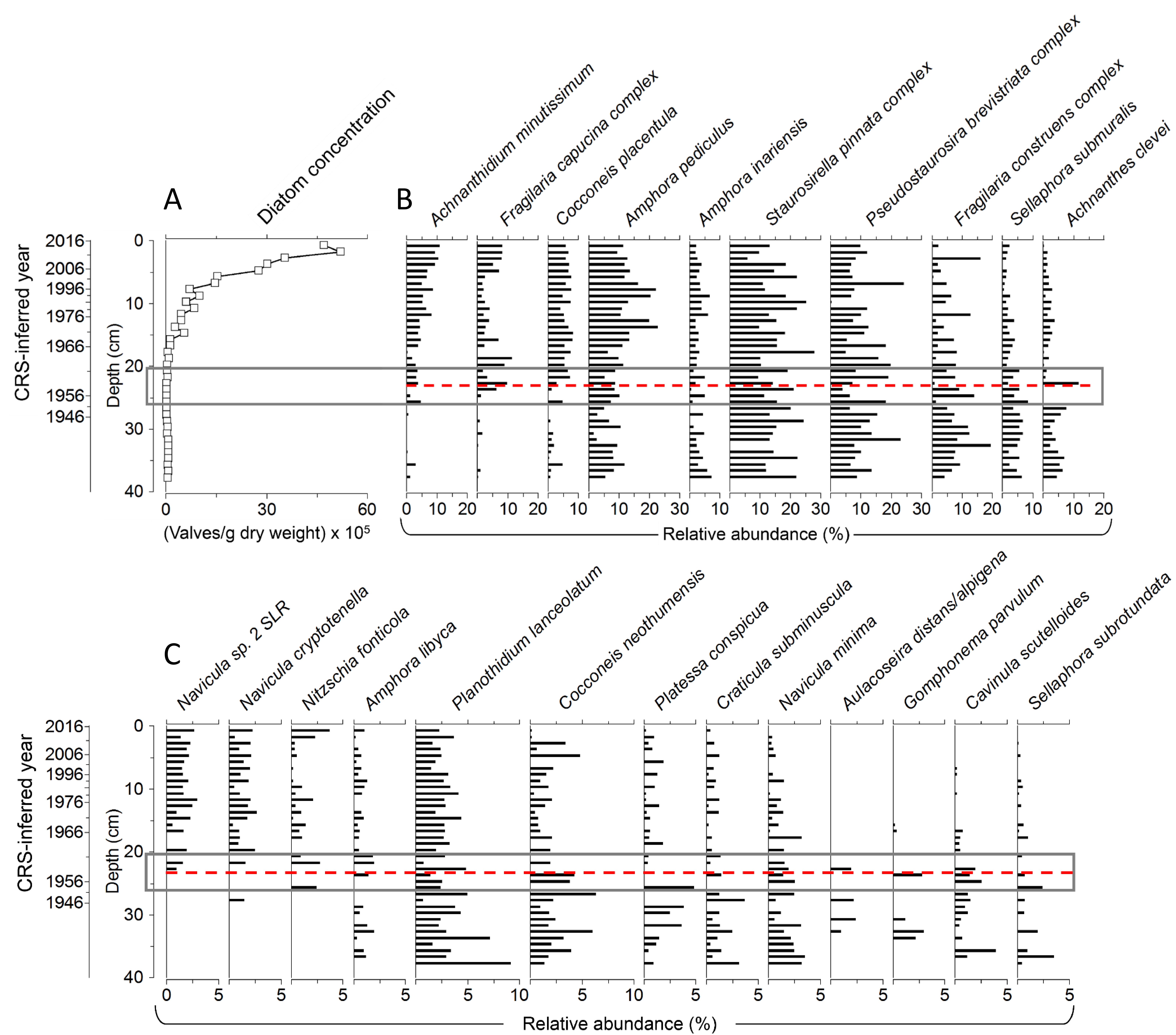


Figure 5. A) Diatom concentrations by depth and year inferred through the constant rate of supply (CRS) dating model. B, C) Relative abundances of the dominant (B) and subdominant (C) diatom species observed in the core. The dotted line represents the significant assemblage change identified by the broken stick model and constrained incremental sum of squares. The box indicates inferred years during which the Moses-Saunders Power Dam was constructed (1954-1959).