



Title	Large-Eddy-Simulation Study of the Effects of Building- Height Variability on Turbulent Flows over an Actual Urban Area
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¹ Large-Eddy-Simulation Study on the Effects of

² Building-Height Variability on Turbulent Flows over

3 an Actual Urban Area

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Abstract A large-eddy simulation (LES) was conducted to investigate the ef-7 fects of building-height variability on turbulent flows over an actual urban area, 8 the city of Kyoto, which was reproduced using a 2-m resolution digital surface 9 dataset. Comparison of the morphological characteristics of Kyoto with those of 10 European, North American, and other Japanese cities indicates a similarity to 11 European cities but with more variable building heights. The performance of the 12 LES model is validated and found to be consistent with turbulence observations 13 obtained from a meteorological tower and Doppler lidar. We conducted the follow-14 ing two numerical experiments: a control experiment using Kyoto buildings, and 15 a sensitivity experiment in which all the building heights are set to the average 16 height over the computational region h_{all} . The difference of Reynolds stress at 17 Toshiya yoshida \cdot Tetsuya Takemi \cdot Mitsuaki Horiguchi

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height $z = 2.5 h_{all}$ between the control and sensitivity experiments is found to 18 increase with the increase in the plan-area index (λ_p) for $\lambda_p > 0.32$. Thus, values 19 of λ_p of around 0.3 can be regarded as a threshold for distinguishing the effects of 20 building-height variability. The quadrant analysis reveals that sweeps contribute to 21 the increase in the Reynolds stress in the control experiment at height $z = 2.5 h_{all}$. 22 The exuberance in the control experiment at height $z = 0.5 h_{all}$ is found to de-23 crease with an increase in the building-height variability. Although the extreme 24 momentum flux at height $z = 2.5h_{all}$ in the control experiment appears around 25 buildings, it contributes little to the total Reynolds stress and is not associated 26 with coherent motion. 27

Keywords Actual urban building · Large-eddy simulation · Atmospheric
 turbulence · Roughness parameter · Reynolds stress · Quadrant analysis

30 1 Introduction

Atmospheric processes over urban areas are affected not only by meteorological 31 disturbances, such as thunderstorms, fronts, and cyclones, but also by the rough-32 ness and thermal effects of buildings and man-made structures. The geometrical 33 features of buildings and structures determine the roughness effects of an urban 34 area, while human activities and the material characteristics of buildings play a 35 role in defining the thermal effects of such areas. The complex geometrical nature 36 of urban surfaces results in highly complex turbulent flows. To properly under-37 stand the physical processes of momentum and heat transfer in urban areas and 38 develop parametrizations for urban environments in numerical weather prediction 39

⁴⁰ models, it is important to reveal relationships between the effects of actual urban
⁴¹ buildings and turbulent flows.

The characteristics of turbulent flows over urban surfaces have been examined 42 in numerous previous studies. Oke (1988) categorized the airflow over roughness 43 obstacles as a function of obstacle density as isolated flows, wake-interference flows, 44 and skimming flows. Macdonald et al. (1998) derived a theoretical relation for the 45 aerodynamic roughness length z_0 and displacement height d for flows over rough-46 ness blocks. While these studies examined turbulent flows over roughness blocks 47 with constant height and regular distribution, the recent focus has shifted to the 48 effects on turbulent flows of roughness blocks with variable height and inhomo-49 geneous arrangement. Wind-tunnel experiments conducted by Cheng and Castro 50 (2002) demonstrated that the roughness sublayers over block arrays with random 51 height are thicker than those over uniform-height arrays. Xie et al. (2008) con-52 ducted a large-eddy simulation (LES) of turbulent flows over block arrays with 53 random height, and found that the tall blocks significantly contribute to the total 54 drag of such arrays. Nakayama et al. (2011) performed LES investigations over 55 building arrays with different height variability and found that the vertical pro-56 files of the mean velocity and Reynolds stress depend significantly on the building-57 height variability. Zaki et al. (2011) performed wind-tunnel experiments with block 58 arrays of buildings with variable height distributed randomly, and showed that the 59 drag coefficient C_d increases with the building density and the standard deviation 60 of the building height for high building densities. Numerical simulations of plume 61 dispersion over urban surfaces have revealed that the turbulence is significantly 62 affected by the source location and wind direction because of the strong depen-63 dence on the building height and distribution. (Xie and Castro 2009; Xie 2011; 64

⁶⁵ Nakayama et al. 2016). The parametrizations of z_0 and d have been improved by ⁶⁶ taking into account roughness parameters associated with actual urban buildings, ⁶⁷ such as the maximum, standard deviation, and skewness of the building height ⁶⁸ (Nakayama et al. 2011; Kanda et al. 2013; Zhu et al. 2017). Giometto et al. (2016) ⁶⁹ suggested that the dispersive flux derived from spatial variations of temporal mean ⁷⁰ flows around buildings should be considered to improve conventional urban-canopy ⁷¹ parametrizations.

To fully understand the effects of roughness obstacles on turbulent flows, it 72 is helpful to investigate the relationships between turbulent organized structures 73 and obstacles, because organized structures are associated with downwards mo-74 mentum transfer in the form of ejection and sweep events based on a quadrant 75 analysis for the turbulent momentum flux. The results of wind-tunnel experiments 76 on flows over rough surfaces conducted by Raupach (1981) indicate that sweeps 77 are dominant for the total momentum flux near surfaces, and that the contribu-78 tion of ejection to the momentum flux increases with height. Studies in which 79 turbulence was observed over actual urban areas have revealed the characteristics 80 of momentum transfer and coherent motion. Oikawa and Meng (1995) observed 81 turbulent structures associated with ejections and sweeps over an urban area, and 82 found that turbulent structures correlate with heat transfer within and above the 83 urban canopy. Christen et al. (2007) analyzed field experimental data obtained 84 from sonic-anemometer measurements within and above a street canyon in Basel, 85 Switzerland, and found that sweeps are mostly dominant up to a height of approxi-86 mately twice the average building height in a street canyon. Numerical simulations 87 of flows over building arrays have revealed the spatial characteristics of turbulent 88 organized structures. Kanda et al. (2004) carried out LES investigations of tur-89

bulent flows over uniform-height block arrays to investigate turbulent organized 90 structures over such arrays. They found low-speed streaks and streamwise vortices 91 similar to those in flows over flat-wall boundary layers. Kanda (2006) indicated 92 that streak structures are a common feature over various types of block arrays. 93 Using direct numerical simulations, Coceal et al. (2007a,b) revealed that hairpin 94 vortices associated with ejections and sweeps are generated over uniform block 95 arrays, and that the low-speed streaks identified above such arrays are composed 96 of large numbers of hairpin vortices aligned in the streamwise direction. Park et al. 97 (2015) used LES results to analyze turbulent-flow structures over an actual urban 98 area in Seoul, Korea, and showed that turbulent structures behind high-rise build-99 ings are characterized by streamwise vortices with strong ejections. They focused 100 on small regions containing high-rise buildings, and demonstrated the significant 101 influence of high-rise buildings on wake flows. The majority of studies presented 102 thus far have focused on the characteristics of turbulent flows over idealized or 103 specific buildings, while only a few have examined the urban-scale effects on the 104 characteristics of turbulent momentum transfer produced by the complex geomet-105 rical features of actual urban surfaces. 106

The geometrical characteristics of actual urban surfaces can be reproduced from digital surface datasets. Ratti et al. (2002) calculated the roughness parameters of North American and European cities, and found that parameters differ significantly by city. Bou-Zeid et al. (2009) indicated that turbulent flows are dependent on the building representation over the actual urban surface. To understand the characteristics of turbulent flows over urban areas, it is therefore important to use the geometry of actual buildings in simulations and experiments. We investigate here the effects of building-height variability in an actual urban area on turbulent flows at an urban scale, focusing on the airflow within and above an urban-canopy layer, where turbulent flows are strongly influenced by individual buildings.

We simulate the turbulent flow over the urban area of Kyoto, which is charac-118 terized by the presence of both business districts with high buildings and densely 119 built residential districts. Furthermore, a meteorological observation tower owned 120 by Kyoto University and located in the southern part of the city can be used for 121 the validation of simulations. In Sect. 2, the building morphological characteris-122 tics of Kyoto are evaluated using roughness parameters. The details of our LES 123 model are described in Sect. 3. The study area of the LES investigation is defined 124 to include the meteorological tower site at which turbulence was measured by a 125 sonic anemometer and Doppler lidar, so that LES results may be compared with 126 the observations (see Sect. 4). Along with a control simulation, we conduct a sen-127 sitivity test assuming a constant building height to reveal the effects of building 128 height-height variability, with the differences between the control and sensitivity 129 experiments examined in Sect. 5. Finally, Sect. 6 concludes the paper. 130

¹³¹ 2 Building Morphological Characteristics of Kyoto

Our study area covered both business districts and suburban areas in Kyoto. Figure 1 shows the area of interest in Kyoto, which extends 11 km in a north-south direction and by 2 km in an east-west direction. A digital surface model (Kokusai Kogyo Co., Ltd.) was used to reproduce the actual urban buildings within a numerical model. The original 2-m-resolution data are smoothed and converted to a 4-m resolution, which is used as the horizontal grid spacing of the numerical
experiments as described in Sect. 3.2.

Figure 2a shows the height of the actual buildings in the analysis area. The north-south and west-east directions are referred to as the x and y directions, respectively. The region with x = 0 - 4 km corresponds to the city centre of Kyoto. The heights of almost all buildings in the region are up to 50 m, and there are no high-rise building clusters of the type seen in the centre of Tokyo. The region for x = 7 - 11 km is primarily occupied by suburban areas and rivers.

The difference between the building heights over these two regions is clearly indicated in Fig. 2b, which shows the frequency distributions of building heights over the entire analysis area and in the x = 0 - 4 km and x = 7 - 11 km regions. In calculating the frequency distributions, all buildings are defined as having heights of at least 1 m to distinguish between the buildings and the ground. It is seen that most of the buildings taller than 25 m are located in the former region.

To quantitatively indicate the morphological characteristics of buildings in Ky-151 oto, we use roughness parameters such as the average building height H_{ave} , the 152 standard deviation of the building height σ_H , the plan-area index λ_p (the ratio of 153 the plan area occupied by buildings to the total surface area), and the frontal-area 154 index λ_f (the ratio of the frontal area of buildings to the total surface area). These 155 parameters are calculated for each 1 km by 1 km area following the analysis of 156 Kanda et al. (2013). Figure 3a shows λ_p calculated in the areas of 1 km by 1 km 157 for the buildings shown in Fig. 2a, with the values of the roughness parameters in 158 the 1 km by 1 km areas summarized in Fig. 3b. The average values of H_{ave} , σ_H , 159 λ_p , and λ_f over the x = 0 - 4 km region are 10.8, 7, 0.41, and 0.25, respectively, 160 while the corresponding averages over x = 7 - 11 km are 9.8, 5.3, 0.2, and 0.16, 161

respectively. Thus, the x = 0 - 4 km region is more densely built than the x = 7 - 11 km. Using building data from Tokyo and Nagoya, Japan, Kanda et al. (2013) derived the following empirical relationships between λ_p and λ_f , and between H_{ave} and σ_H ,

$$\lambda_f = 1.42\lambda_p^2 + 0.4\lambda_p,\tag{1}$$

$$\sigma_H = 1.05 H_{ave} - 3.7. \tag{2}$$

Figure 3c and d indicates the respective relationships between λ_p and λ_f , and 166 between H_{ave} and σ_H , based on the data given in Fig. 3b. Also shown in the 167 panels are the empirical relationships of Kanda et al. (2013) and the data for North 168 American and European cities found in Ratti et al. (2002). For $\lambda_p > 0.3$, the λ_f 169 values for Kyoto tend to be smaller than in the empirical profile. This feature 170 of Kyoto appears to be similar to those seen in European cities, and indicates 171 that the fraction of high buildings in Kyoto is limited relative to those in major 172 metropolitan cities in Japan and North America. The relationship between H_{ave} 173 and σ_H for Kyoto is in good agreement with those of Tokyo and Nagoya, but 174 differs from those of European cities. Finally, the magnitudes of H_{ave} and σ_H in 175 Kyoto are smaller than those of Los Angeles by a factor of 5 - 10. 176

According to these results, Kyoto can be morphologically characterized as having densely distributed buildings with widely varying heights. The Kyoto dataset was used for the numerical simulations described in the next section.

¹⁸⁰ 3 Numerical Model and Experimental Design

¹⁸¹ 3.1 Numerical Model

Our LES model is effectively the same as the one used in Nakayama et al. (2011), 182 except that it neglects the molecular viscosity term, and employs a bottom bound-183 ary condition based on Monin-Obukhov similarity theory, as described later. In 184 Nakayama et al. (2011), the performance of the LES model reproducing turbu-185 lent statistics was validated using data obtained from wind-tunnel experiments; 186 as a close agreement was found, the model developed by Nakayama et al. (2011) 187 has subsequently been applied to simulate turbulent flows over actual urban cities. 188 Nakayama et al. (2012) conducted LES investigations of turbulent flows over Tokyo 189 by coupling their model with a mesoscale meteorological model, and found that 190 observed gust factors are accurately reproduced by the model. The model was 191 also used to successfully reproduce the wind speeds and directions at the ground 192 level in the Fukushima Daiichi Nuclear Power Plant during the Great East Japan 193 Earthquake and its aftermath in March 2011 (Nakayama et al., 2015). Nakayama 194 et al. (2016) further applied their LES model for the simulation of turbulent flows 195 and plume dispersion over Oklahoma City, and showed that the observed charac-196 teristics of turbulence and dispersion are reproduced despite the fact that small 197 differences in wind direction caused by the building distribution significantly in-198 fluenced the plume dispersion. Thus, our LES model has been widely tested and 199 is applicable for the analysis of the turbulent flow over Kyoto. 200

The LES model solves the filtered continuity and Navier–Stokes equations in Cartesian coordinates with the subgrid-scale stress parametrized by the standard Smagorinsky model (Smagorinsky, 1963). The governing equations are

$$\frac{\partial \tilde{u}_i}{\partial x_i} = 0,\tag{3}$$

$$\frac{\partial \tilde{u}_i}{\partial t} + \tilde{u}_j \frac{\partial \tilde{u}_i}{\partial x_j} = -\frac{1}{\rho} \frac{\partial \tilde{p}^*}{\partial x_i} - \frac{\partial \tau_{ij}}{\partial x_j} + f_i, \qquad (4)$$

$$\tau_{ij} - \frac{1}{3} \delta_{ij} \tau_{kk} = -2(C_s \Delta)^2 (2\tilde{S}_{ij} \tilde{S}_{ij})^{1/2} \tilde{S}_{ij}, \tag{5}$$

$$\tilde{S}_{ij} = \frac{1}{2} \left(\frac{\partial \tilde{u}_i}{\partial x_j} + \frac{\partial \tilde{u}_j}{\partial x_i} \right),\tag{6}$$

where t denotes time, \tilde{u}_i is the filtered air velocity in the direction $i, \tilde{p}^* =$ 204 $\tilde{p} + \frac{1}{3}\rho\tau_{kk}$ is the modified pressure, \tilde{p} is the filtered pressure, ρ is the density of air, 205 τ_{ij} is the subgrid-scale stress, δ_{ij} is the Kronecker delta, \tilde{S}_{ij} is the filtered stress 206 tensor, and f_i is the external force exerted by roughness obstacles. The parameter 207 x_i represents the coordinate system, with components i = 1, 2, and 3 referring 208 to the streamwise (x), spanwise (y) and vertical (z) directions, respectively. In 209 addition, $\Delta = (\Delta_x \Delta_y \Delta_z)^{1/3}$ is the filter width, where Δ_x , Δ_y , and Δ_z are the 210 streamwise, spanwise, and vertical grid spacings, respectively. The Smagorinsky 211 coefficient C_s is set to 0.14. Note that the viscous term is neglected because our 212 target is the simulation of turbulent flow with a high Reynolds number. 213

The external force f_i is used to simulate the effects of buildings on the flow, for which we employ the feedback forcing by Goldstein et al. (1993) who give

$$f_i = \alpha \int_0^t u_i(t')dt' + \beta u_i(t), \quad \alpha < 0, \ \beta < 0, \tag{7}$$

where α and β are negative constants. The stability limit is given by $\Delta t < \frac{-\beta - \sqrt{(\beta^2 - 2\alpha k)}}{\alpha}$, where k is a constant of order one. Following Nakayama et al. (2011), these constants are set as $\alpha = -10$, $\beta = -1$, and k = 1.

The governing equations are discretized on a staggered-grid system. The velocity and pressure fields are solved using a coupling method based on the marker-andcell method (Chorin, 1967). The successive over-relaxation method is used to solve the Poisson equation for pressure, and the Adams–Bashforth scheme is adopted for the time integration. A second-order accurate, central-differencing scheme is employed for spatial discretization. The code is parallelized using a Message Passing Interface library to reduce the computational time.

²²⁶ 3.2 Experimental Design

The governing equations are numerically solved in two computational domains: 227 the driver region, which features regularly arrayed obstacles, and the main region, 228 which contains the actual buildings of Kyoto. To ensure the flow field of the LES 229 is turbulent, a turbulent flow is generated in the driver region and imposed as 230 the inflow at the boundary of the main region. The concept involved in setting 231 the driver and the main regions is demonstrated in Fig. 4. The size of the driver 232 region is 6 km (streamwise) \times 2.4 km (spanwise) \times 1.015 km (vertical), with a 233 grid spacing of 4 m in the horizontal directions, and a grid spacing stretched with 234 increasing altitude from 1 m to 16 m in the vertical direction. The total number 235 of grid points is $1500 \times 600 \times 105$. In the driver region, there is one rectangular 236 block aligned in the spanwise direction, and an array of roughness blocks staggered 237 with $\lambda_p = 0.04$. The individual rectangular and roughness block sizes are 50 m \times 238 2400 m \times 50 m and 16 m \times 16 m \times 10 m, respectively. The purpose of setting 239 the size of the rectangular block is to enhance perturbations near the inlet of the 240 driver region. The λ_p value chosen for the block array is set to be a little larger 241 than that in Nakayama et al. (2014) to reduce the generation of turbulence. The 242

height of the blocks is chosen according to the mean building height in the mainregion.

A uniform flow with a velocity magnitude of 5 m s⁻¹ is imposed at the inflow boundary of the driver region. The Sommerfeld radiation condition is imposed at the outflow boundary, while a periodic condition is set at the lateral boundaries. At the top boundary, free-slip and zero-speed conditions are imposed for the horizontal and vertical velocity components, respectively. At the ground, a boundary condition based on Monin–Obukhov similarity theory is employed. The stress at the first vertical grid $\tau_{i3}(x, y, t)$ is calculated as (Stoll and Porté-Agel, 2006)

$$\tau_{i3}(x, y, t) = -\left[\frac{\tilde{u}_r(x, y, z_s, t)\kappa}{\ln(z_s/z_0)}\right]^2 \frac{\tilde{u}_i(x, y, z_s, t)}{\tilde{u}_r(x, y, z_s, t)},\tag{8}$$

where $\tilde{u}_r(x, y, z_s, t) = [\tilde{u}_1(x, y, z_s, t)^2 + \tilde{u}_2(x, y, z_s, t)^2]^{1/2}$ is the instantaneous resolved velocity magnitude, z_s is the altitude at the first vertical grid, z_0 is the roughness length, and κ is the von Kármán constant. Here, $z_0 = 0.1$ m (Bou-Zeid et al., 2009) and $\kappa = 0.4$.

The ratio of the boundary-layer height δ of the generated outflow to the rough-256 ness block height in the driver region is 27.9. Note that, here, δ is defined as the 257 height at which the mean streamwise velocity component at the outflow indicates a 258 peak value. In Nakayama et al. (2011), the ratio of δ to the roughness block height 259 in the driver region is 13. In addition, we confirmed that the vertical profiles of the 260 standard deviation of each velocity component and Reynolds stress are in reason-261 able agreement with those obtained from wind-tunnel experiments, although the 262 LES results underestimate the spanwise and vertical components and Reynolds-263 stress values relative to the wind-tunnel results (see Online Resource 1, Figure 1). 264

These results suggest that well-developed, deep turbulent flows are generated in the driver region.

In the main region, the domain size and the total number of grid points are 12 267 km \times 2.4 km \times 1.015 km and 3000 \times 600 \times 105, respectively. The main region 268 includes the actual buildings and structures in Kyoto, as shown in Fig. 2a. For 269 computational purposes, we set a buffer area spanning 500 m and 200 m in the 270 streamwise and spanwise directions, respectively, surrounding the actual building 271 area in the main region (not shown in Fig. 2a). The streamwise width of the area 272 was determined based on Nakayama et al. (2012), who carried out an LES inves-273 tigation of the airflow over Tokyo. Whereas Nakayama et al. (2012) did not set a 274 buffer area in the spanwise direction, we decided that a spanwise buffer is necessary 275 to avoid building discontinuities arising from the periodic boundary conditions. In 276 this buffer area, the same roughness blocks used in the driver region are applied to 277 maintain a turbulent flow over roughness surfaces. Note that the coordinates x = 0278 km and y = 0 km are set to the northern and western boundaries, respectively, of 279 the actual building area in the main region. Correspondingly, the inflow boundary 280 condition provided by the driver region is set at x = -500 m in the main region. 281 Outside of the inflow boundary, the boundary conditions of the main region are 282 the same as those in the driver region, and all grid spacings are identical to those 283 in the driver region. 284

Hereafter, the simulation using the actual buildings in Kyoto is referred to as the control experiment (CTL). To reveal the effects of building-height variability, we conducted an additional experiment referred to as the uniform experiment (UNI) in which all building heights are set to the average of the actual building heights in the main region ($h_{all} = 10.3$ m). The integration time for each of the

two experiments is 7,200 s, with the results obtained from the last 1,800 s used 290 for the analysis of turbulent statistics. In Sect. 4.3, we confirm that the flows were 291 in equilibrium states during this analysis period, as shown Fig. 5. In addition, 292 as seen in Fig. 1 of Online Resource 1, the second-order moments of the inflow 293 profiles are relatively small compared with those of the wind-tunnel experiments, 294 which possibly influences the results presented here. However, as the same inflow 295 condition was applied in both the CTL and UNI experiments, we can assume that 296 any differences in the respective experimental results are unaffected by this issue. 297

²⁹⁸ 4 Comparison with Observations

²⁹⁹ 4.1 Observational Setting

The observations were performed at the Ujigawa Open Laboratory of the Disaster 300 Prevention Research Institute, Kyoto University, during the period from 12 Jan-301 uary to 12 February 2016. The laboratory is located in the southern part of Kyoto, 302 and is surrounded by low-rise buildings and structures. The location of the obser-303 vation site is shown in Fig. 1, which includes a meteorological observation tower 304 of height 55 m. This tower is a unique facility first deployed in 1978 (Nakajima 305 et al., 1979), and is currently one of the few meteorological towers operating in 306 Japan. 307

A sonic anemometer (DA-600, Kaijo Co.) installed on the tower at a 25-m height measures the three velocity components as well as the air temperature at a 10-Hz sampling rate. The surrounding area up to 500 m north of the tower has only low building heights (< 25 m), enabling the assumption that observations taken by the sonic anemometer are not influenced by the strong wakes of tall buildings. We also installed a Doppler lidar (WINDCUBE WLS-7, Leosphere) at the ground near the tower, from which we obtained three-component velocity measurements at heights ranging from 40 m to 200 m with a 20-m interval at a sampling rate of 1 Hz.

317 4.2 Data Selection

The observation site was included in the main region assessed in the numerical 318 experiment for the purpose of directly comparing the LES results in the CTL 319 experiment with the observations. As the sonic anemometer installed on the tower 320 faces northwards, we analyzed data for dominant northerly wind directions to 321 minimize the interference from the tower. To extract suitable periods from the 322 observational data, we imposed two criteria for sorting values obtained from the 323 sonic anemometer. First, a northerly flow condition was adopted by classifying 10-324 min averaged wind directions into 16 classes and extracting periods when northerly 325 wind directions $(348.25^{\circ} - 360^{\circ}, 0^{\circ} - 11.25^{\circ})$ were sustained for at least 30 min. 326 Note that the time period for the analysis of the LES data was also 30 min. Second, 327 a neutrally stratified condition was chosen based on the Monin–Obukhov stability 328 parameter 329

$$\frac{z}{L} = -\frac{(g/\overline{T})\overline{w'T'}}{u_*^3/\kappa z},\tag{9}$$

³³⁰ so that the assumption of turbulent flows under a neutrally stratified condition in ³³¹ the LES model is valid. Here, L is the Obukhov length (m), g is the acceleration ³³² due to gravity (m s⁻²), T is the air temperature (K), $\overline{w'T'}$ is the sensible heat flux ³³³ (K m s⁻¹), and u_* is the friction velocity (m s⁻¹). An overbar and prime denote a temporal average and fluctuation, respectively. A period for $|z/L| \le 0.05$ (Roth, 2000) is regarded as fulfilling the neutrally stratified condition.

By imposing the above conditions on the observational data, we obtained the following four 30-min periods: 0720 - 0750 LT (local time = UTC + 9 h) 22 January; 1650 - 1720 LT 30 January; 0740 - 0810 LT 2 February; and 1830 -1900 LT 10 February, which are referred to as the D1, D2, D3, and D4 periods, respectively. The wind directions for each period calculated from the averaged horizontal velocity components are 4.9° , 358.8° , 353.8° , and 351.5° for the D1 to D4 periods, respectively.

To compare the LES results with the observations, it is necessary to use airflows 343 observed at the Ujigawa Open Laboratory coming from the northern boundary of the analysis region of Kyoto passing through the analysis region, and not from the 345 western or eastern boundaries. Because of the periodic conditions at the western 346 and eastern boundaries, the flow through these lateral boundaries is unlikely to 347 be accurately simulated by the LES model. This condition requires that wind 348 directions be within a range of between approximately 355° and 5° based on the 349 streamwise length and half the spanwise length of the analysis region (i.e., arctan(1 350 km/11 km)). Overall, the wind directions in the periods D1 – D4 are almost 351 within the range of this condition, although those in the periods D3 and D4 are 352 slightly shifted westwards from the condition. We confirmed that the area within 353 at least 1 km westwards from the analysis region is dominated by land-use and 354 building types similar to those in the analysis region. Thus, we concluded that the 355 anemometer data taken during the four periods described above are appropriate 356 for comparison with the LES results. However, the wind directions measure by the 357 Doppler lidar deviate from those recorded by the sonic anemometer. The directions 358

of the Doppler lidar in the D1 and D3 periods become more westerly with height, 359 reaching 330° at a height of 200 m, while those in the D2 and D4 periods are 360 relatively constant with altitude and within a range between approximately 350° 361 and 0° . We discuss the possible influences of the variation of wind direction in 362 Sect. 4.3. As explained above, none of the observed wind directions were oriented 363 in a truly northerly fashion. Correspondingly, we rotated the streamwise directions 364 to the mean of the wind directions measured by the sonic anemometer and the 365 Doppler lidar. 366

367 4.3 Results

Figure 5a and b shows the time series of streamwise and spanwise velocity com-368 ponents produced by the LES model and measured by the sonic anemometer at a 369 25-m height, respectively. To avoid interference from the tower on the wind-speed 370 profiles, the LES results are shown for a grid point 16 m north of the tower. It 371 is seen that the LES turbulent fluctuations in both the streamwise and spanwise 372 directions are quite comparable to those from the anemometer. Note that average 373 spanwise velocity components are nearly zero, as indicated in Fig. 5b. The stream-374 wise velocity component is stronger in the D2 period than in the other periods. 375 Comparison of the respective weather charts for the four time periods reveals the 376 stronger wind speeds in the D2 period to be caused by a large low-pressure system 377 passing through the northwest Pacific Ocean off the coast of the Japanese Islands. 378 Figure 5c shows a comparison of the LES and observed vertical profiles of the 379 mean streamwise velocity component. Both datasets are averaged over time, and 380 the time-averaged LES data are averaged horizontally over a 16 m by 16 m area to 381

the north of the tower to increase the representativeness of the simulated flows for 382 the observation site. Note that, given the logarithmic scales used on both axes, the 383 slopes of the mean streamwise velocity component in Fig. 5c suggest a power-law 384 profile. According to Counihan (1975), the slopes of suburban and urban areas 385 range between 0.21 and 0.28, making a power-law exponent of 1/4 suitable for ref-386 erence, where it is seen that the slopes of the observations and the LES results are 387 very similar to this value. We also examined the respective vertical profiles of the 388 mean streamwise velocity component normalized by the mean streamwise velocity 389 component at the 25-m height (see Online Resource 1, Figure 2) and found that 390 the LES and observed mean streamwise velocity components are quantitatively 391 consistent. We conclude that this result is also good evidence for the reasonable 392 performance of our LES model. In contrast, the slopes above approximately the 303 150-m height in the D1, D2, and D3 periods appear to deviate from the reference 394 slope. In the case of the D1 and D3 periods, we assume this occurs because of the 395 change in wind direction from northerly to westerly, as described in the previous 396 subsection. Another possible explanation for the deviation at the higher levels is 397 that the stability conditions may not have been neutral at these heights during 398 the observed periods. Because there were no observational data available to clas-399 sify the stability condition above height of the sonic anemometer at 25 m, it is 400 impossible to quantitatively reveal the stability above that height. 401

The vertical profiles of Reynolds stress in both the observations and the LES results are shown in Fig. 5d. Note that the Reynolds stress is normalized by the mean streamwise velocity components at each height. The Reynolds stress of the LES data is averaged horizontally over the same 16 m by 16 m area used for the mean streamwise velocity component. It is seen that the vertical profile of the LES data is within the range of differences found in the observation periods, which is a feature similar to that of the profiles normalized by the mean streamwise velocity component at the 25-m height (see Online Resource 1, Figure 2). However, it is necessary to be careful in comparing the LES results with the Doppler lidar data because the latter might include some errors in representing perturbations of the wind speed as discussed below.

We now compare the results for the turbulence intensity, which is the ratio 413 of the standard deviation of each velocity component σ_i to the mean streamwise 414 velocity component. As previously mentioned, the turbulence intensity was also 415 averaged horizontally in the 16 m by 16 m area. Figure 6 compares the vertical 416 profiles of turbulence intensity in the LES results and observations with the em-417 pirical form of the ESDU (1985), which is a database providing the turbulence 418 characteristics of a neutrally stratified atmospheric boundary layer based on var-419 ious field measurements from around the world. In Fig. 6, all sonic-anemometer 420 components fall within the rough-surface category given by the ESDU, which in-421 dicates suburban areas with z_0 between 0.1 and 0.5. Each component simulated 422 by the LES model appears to capture the vertical distribution of that obtained by 423 the ESDU within or around its upper and lower limits, at least below about the 424 height of 150 m, while being slightly smaller than those of the sonic-anemometer 425 observations. In fact, the values obtained from the sonic anemometer lie near the 426 upper limit of the ESDU profile, suggesting that the LES results within the ESDU 427 range are generally more favourable. 428

In contrast, there appears to be large discrepancies between the Doppler lidar observations and the LES results in terms of the u and v velocity components. The turbulence intensities for these components measured by the Doppler lidar

are even larger than the upper limits of the ESDU, suggesting that the measure-432 ments may include an overestimating bias for the turbulence intensities. It is in 433 fact commonly understood that the Doppler lidar measurements overestimate the 434 turbulence intensities for the streamwise component. This characteristic was noted 435 in Cañadillas et al. (2011), who showed that the results produced by the Doppler 436 lidar observations are larger than those of sonic anemometers at various wind 437 speeds and altitudes, and the deviations become larger with the decrease in wind 438 speed. A close look at Figs. 5c and Fig. 6a indicates that the difference between 439 Doppler lidar and the ESDU in terms of the streamwise turbulence intensity be-440 low 100 m decreases as the streamwise velocity conponent increases, in apparent 441 confirmation of the finding of Cañadillas et al. (2011). For the Doppler lidar data 442 above 100 m, changes in the wind direction and uncertainty in the stability, as 443 revealed in the mean streamwise velocity component, may contribute to this over-444 estimation. An overestimating tendency in the lidar data can also be found for the spanwise component, which has a mean value of nearly zero. 446

The vertical component produced by the LES results appears to be consistent with both the lidar data and the ESDU profile, but the lidar tends to underestimate the vertical turbulence intensity, particularly in weaker wind-speed conditions.

Figure 7 shows the power spectra of the time series of each velocity component obtained from the LES results and the sonic anemometer at a height of 25 m. The spectra were calculated from the time series shown in Fig. 5, and the frequency f and velocity spectra E(f) are normalized in dimensionless form. The figure includes the empirical reference from Kaimal et al. (1972) derived from observations over a rural region. A close agreement is seen between the sonic anemometer and the reference results for all three components. The spectra from the sonic anemometer clearly represent an inertial subrange with a ?2/3 slope. Comparison of the LES spectra with the observations and empirical reference reveals that the spectra of the u and v components of the LES data are similar to those of the sonic anemometer data except in the highest frequency range. The lower frequency portion of the inertial subrange appears to be well reproduced for these components in the LES results.

However, the LES model is able to reproduce the vertical velocity components in only the lowest frequency portion of the inertial subrange. It is possible that the grid spacing used in our modelling is insufficient for resolving the smallest eddies and their corresponding vertical motion. Further increases in the vertical resolution may be required to represent the small-scale vertical motion likely to be induced at the edges of buildings. However, we note that the spectral peak of the w component in the LES results agrees well with that of the sonic anemometer.

From the above comparisons, we conclude that the use of our LES model leads 470 to a reasonable reproduction of the turbulent boundary-layer flow over actual 471 buildings under a neutral stability condition, at least up to a height of about 150 472 m. We emphasize that, in general, the results produced by our LES model agree 473 favourably with the observations within the range of differences among the chosen 474 periods (D1 – D4), even though our inflow condition employed an idealized turbu-475 lent flow generated in the driver region without realistic meteorological conditions. 476 These results are sufficient here because our analysis of building-height variabil-477 ity focuses on altitudes below approximately 25 m (i.e., at height $z = 2.5 h_{all}$), 478 where the LES results show an especially close agreement with the observations, 479 as shown in Figs. 6 and 7. 480

⁴⁸¹ 5 Sensitivity to Building-Height Variability

⁴⁸² 5.1 General characteristics of turbulent flows

We now focus on the overall characteristics of turbulent flows in the CTL and UNI 483 experiments, starting with the differences between the respective experiments. 484 Figure 8a and b shows the vertical profiles of the space- and time-averaged 485 streamwise velocity component $\langle \overline{u} \rangle_{all}$ and Reynolds stress $-\langle \overline{u'w'} \rangle_{all}$ over 486 the entire main region for the CTL and UNI experiments, respectively. Here, the 487 angled brackets denote a spatial average, while the subscript *all* refers to the overall 488 main region. Note that the values are normalized by the mean streamwise velocity 489 component U_{∞} at the height of the boundary-layer (δ). The mean streamwise ve-490 locity components above height $z = h_{all}$ (i.e., above the canopy layer) are lower in 491 the CTL experiment than in the UNI experiment. In contrast, the velocities below 492 height $z = h_{all}$ for the CTL experiment are higher than in the UNI experiment. 493 The Reynolds stress above height $z = h_{all}$ in the CTL experiment is larger than 494 that in UNI experiment. Furthermore, the level of peak Reynolds stress is higher 495 in the CTL experiment than in the UNI experiment. 496

These differences between the CTL and UNI results can be attributed to the 497 effects of building-height variability. Using the LES results of flows over idealized 498 arrays of roughness blocks, Nakayama et al. (2011) showed that the mean velocity 499 above the building height decreases with increasing building-height variability, and 500 that the magnitude and height of the peak of the Reynolds stress both increase 501 with building-height variability. Our results in terms of the streamwise velocity 502 component and Reynolds stress are consistent with the results of Nakayama et al. 503 (2011).504

Xie et al. (2008) carried out an LES investigation over block arrays with ran-505 dom and uniform heights, and found that both types of arrays produced similar 506 turbulent kinetic energies below the average building height. The Reynolds stresses 507 produced in the CTL and UNI experiments below height $z = h_{all}$ are consistent 508 with their results. From Fig. 8b, it is seen that the Reynolds stress in the UNI 509 experiment sharply increases around height $z = h_{all}$, which is likely caused by 510 the presence of the uniform tops of buildings in the UNI experiment, resulting in 511 sharp wind shear and the generation of turbulence. 512

In Coceal et al. (2006), the velocity components u_i were decomposed as

$$u_i = \langle \overline{u}_i \rangle + \overline{u}_i'' + u_i', \tag{10}$$

where $\langle \bar{u}_i \rangle$ are the time- and space-averaged velocities, \bar{u}_i'' is the spatial vari-514 ation of the time-averaged velocity, and u'_i is the turbulent fluctuation. Coceal 515 et al. (2006) showed that dispersive flux, which is defined as $<\overline{u}''\overline{w}''>$, signifi-516 cantly contributes to the total momentum flux in the canopy layer in which the 517 time-averaged velocities are spatially inhomogeneous. The vertical profiles of the 518 dispersive flux normalized by U_{∞} in the CTL and UNI experiments are shown 519 in Fig. 8c. Although the dispersive fluxes for both experiments have peaks just 520 below height $z = h_{all}$, the magnitude of the peak in the UNI experiment is larger 521 than that in the CTL experiment. The UNI profile decreases sharply with height 522 above the height of the peak. Above height $z = h_{all}$, the dispersive flux in the 523 CTL experiment is larger than that in the UNI experiment up to about height 524 $z = 3.5 h_{all}$. Xie et al. (2008) performed an LES investigation to compare the 525 dispersive flux in random and uniform block arrays. Their results suggest that 526 both types of dispersive flux have peaks near the average building height, that 527

the peaks obtained from uniform block arrays are stronger than those for random 528 block arrays, and that the dispersive flux of uniform block arrays decreases much 529 more abruptly with increasing height above the height of the peak than that of 530 random block arrays. These characteristics are qualitatively consistent with our 531 results. The dispersive fluxes in both the CTL and UNI experiments appear not to 532 decrease linearly with height because the time-averaged velocities are not spatially 533 homogeneous at heights above the canopy layer. Based on the results shown in Fig. 534 8a and b, we focus on the height of $z = 0.5 h_{all}$ at which the difference between 535 the CTL and UNI experiments is small, and the height of $z = 2.5 h_{all}$ where clear 536 differences are seen between the respective experiments. 537

Figure 9 shows the fields of time-averaged streamwise velocity component nor-538 malized by U_{∞} for the CTL and UNI experiments over an upstream region (x =539 1-5 km) in which the business districts are located. The difference between 540 the respective experimental results for the region appears to be small at height 541 $z = 0.5h_{all}$ except in areas along a major street around y = 1.3 km. This is likely 542 caused by a stronger convergence of the streamwise velocity components on the 543 street in the UNI experiment owing to enhancements arising from the presence of 544 uniform-height buildings (i.e., in the UNI experiment, all lower building heights are 545 raised to $z = h_{all}$). The velocity-deficit regions are reproduced at height $z = 2.5 h_{all}$ 546 behind buildings in the CTL experiment, which contrasts to the smooth field of 547 time-averaged streamwise velocity components at height $z = 2.5h_{all}$ in the UNI 548 experiment. 549

Figure 10 shows the fields of Reynolds stress normalized by U_{∞} for the CTL and UNI experiments over the upstream region. While the features are quite similar at height $z = 0.5h_{all}$, the field at height $z = 2.5h_{all}$ in the CTL results has larger values behind the buildings than in the UNI results, which indicates the important role of sparsely and randomly distributed buildings at and above height $z = 2.5h_{all}$ in generating turbulence in the CTL experiment.

556 5.2 Analysis of Roughness Parameter

To quantitatively reveal the effects of building-height variability, we examined the relationships between the turbulent statistics and roughness parameters. The planarea index λ_p is used for this analysis because the CTL and UNI experiments have the same values for this parameter. Turbulent statistics were derived in each 1 km by 1 km area in a manner similar to that used to find the roughness parameters in Sect. 2.

563 5.2.1 Reynolds stress

Figure 11 shows how the Reynolds stress normalized by U_{∞} in the CTL and 564 UNI experiments changes as a function of λ_p at the heights of $z = 0.5 h_{all}$ and 565 $z=2.5h_{all}.$ The brackets with subscript $1~{\rm km}^2$ indicate spatial averaging over a 1566 km by 1 km area. The Reynolds stress at height $z = 0.5h_{all}$ is very similar for the 567 two experiments, which is consistent with the features shown in Fig. 10a and b. 568 By contrast, the values at height $z = 2.5 h_{all}$ in the CTL experiment increase with 569 λ_p , while those in the UNI experiment are nearly independent of λ_p . In addition, 570 the differences between the CTL and UNI results at height $z = 2.5 h_{all}$ are more 571 apparent when $\lambda_p > 0.32$. 572

As shown in Figs. 9 and 10, the difference between the CTL and UNI experiments in terms of building distributions at height $z = 2.5h_{all}$ has a significant

effect on the turbulent flow results. To interpret this difference, we calculated the 575 respective plan-area indices λ_p at this altitude; i.e., for each experiment, if the 576 building height in a grid cell is below $z = 2.5 h_{all}$, the grid cell is regarded as 577 having no buildings. Figure 12a and b shows λ_p at heights $z = 0.5 h_{all}$ (denoted 578 by $\lambda_{p, 0.5h_{all}}$ and $z = 2.5h_{all} (\lambda_{p, 2.5h_{all}})$, respectively, plotted against λ_p at the 579 surface for both the CTL and UNI experiments. Note that, in the UNI experiment, 580 the value of $\lambda_{p, 0.5h_{all}}$ is the same as that of λ_p at the surface, and that $\lambda_{p, 2.5h_{all}}$ 581 is zero for this experiment. The difference between the CTL and UNI experiments 582 in terms of $\lambda_{p, 0.5h_{all}}$ is not very large, confirming the similarity of the respec-583 tive Reynolds stresses at height $z = 0.5h_{all}$ in Fig. 10. In the CTL experiment, 584 $\lambda_{p, 2.5h_{all}}$ rapidly increases if λ_p exceeds 0.32, which appears to be consistent with 585 the Reynolds-stress feature in the CTL experiment at height $z = 2.5 h_{all}$ as seen 586 in Fig. 10. Based on these results, we suggest that the Reynolds stress from the 587 CTL experiment at height $z = 2.5 h_{all}$ becomes stronger at $\lambda_p > 0.32$ because 588 some building clusters are still present at height $z = 2.5 h_{all}$ in this experiment. 589

The frontal-area index λ_f is another important parameter for describing the ge-590 ometrical characteristics of urban areas. Here we examine the frontal area of build-591 ings above the height h_{all} . Figure 12c shows λ_f above height $z = h_{all} (\lambda_f, h_{all})$ 592 plotted against λ_p for the CTL experiment (the figure does not include the corre-593 sponding values for the UNI experiment owing to the absence of buildings at that 594 altitude). It is seen that $\lambda_{f, h_{all}}$ increases with λ_p , and sharply increases when 595 $\lambda_p > 0.32$. These features agree well with the characteristics determined above 596 for $\lambda_{p, 2.5h_{all}}$ and the Reynolds stress. According to these results, the effects of 597 building-height variability on the Reynolds stress increase with λ_p when λ_p is 598

greater than 0.32, and are closely linked to the higher values of $\lambda_{p, 2.5h_{all}}$ and $\lambda_{f, h_{all}}$ at such values of λ_{p} .

Interestingly, Zaki et al. (2011) found that the drag coefficient C_d in wind-601 602 tunnel experiments, which is relevant to the Reynolds stress, increases with λ_p when $\lambda_p > 0.32$ in flows over block arrays with random heights. A similar feature 603 can also be found for the Reynolds stress and C_d in the LES investigation by 604 Nakayama et al. (2011). According to Zaki et al. (2011), this is because taller 605 buildings, which contribute largely to the total drag in a block array (Xie et al., 606 2008), tend to be sparsely distributed and, therefore, despite the increase in λ_p , 607 the flow pattern does not enter a skimming flow regime (Oke, 1988). Based on 608 these previous studies and our results, $\lambda_p \approx 0.3$ can be regarded as a threshold 609 at which the effects of building-height variability on the turbulent flow become 610 apparent in various cities. 611

⁶¹² 5.2.2 Momentum transfer according to a quadrant analysis

As described in Sect. 1, turbulent coherent structures over urban surfaces are re-613 lated to the physical process of turbulent momentum transfer. A quadrant analysis 614 is a useful method for identifying the characteristics of the momentum transfer as-615 sociated with coherent structures, and has been used in numerous studies of wall 616 turbulence (Wallace, 2016). This method divides the Reynolds stress into four 617 components based on the signs of u' and w': outwards interaction (quadrant 1, 618 u' > 0, w' > 0; ejection (quadrant 2, u' < 0, w' > 0); inwards interaction 619 (quadrant 3, u' < 0, w' < 0) and sweep (quadrant 4, u' > 0, w' < 0). Raupach 620 (1981) introduced a conditional averaging using the threshold H to investigate the 621

 $_{622}$ contribution to the Reynolds stress from the *i*th quadrant as

$$< u'w'>_{i, H} = \lim_{T \to \infty} \frac{1}{T} \int_0^T u'(t)w'(t)I_{i, H}[u'(t), w'(t)]dt,$$
 (11)

623 where the trigger indicator $I_{i, H}$ is defined as

$$I_{i, H}(u', w') = \begin{cases} 1, \text{ if } (u', w') \text{ is in quadrant } i \text{ and if } |u'w'| \ge H |\overline{u'w'}|, \\ 0, \text{ otherwise.} \end{cases}$$
(12)

The fraction of stress exceeding the threshold, which indicates the relative quantity of the *i*th quadrant, is

$$S_{i, H} = \langle u'w' \rangle_{i, H} / \overline{u'w'}.$$
 (13)

626 It is noted that the relationship

$$S_{1,0} + S_{2,0} + S_{3,0} + S_{4,0} = 1$$
(14)

holds only for H = 0. When the Reynolds stress is negative (as is normally seen in the boundary layer), $S_{2, 0}$ and $S_{4, 0}$ are positive, while $S_{1, 0}$ and $S_{3, 0}$ are negative. Ejections and sweeps contribute to the downwards momentum flux, and are considered to be associated with organized turbulent motions as indicated in Sect. 1. Thus, the magnitude of ejections and sweeps is a good indicator for determining the characteristics of turbulent flows.

To further reveal the relative roles of ejections and sweeps in vertical momentum transfer, we introduce the two parameters

$$\Delta S_0 = S_{4,0} - S_{2,0},\tag{15}$$

$$Ex = (S_{1,0} + S_{3,0}) / (S_{2,0} + S_{4,0}),$$
(16)

where ΔS_0 is the difference between sweeps and ejections, and Ex, which is called the exuberance (Shaw et al., 1983), is the ratio of unorganized (S_1 and S_3) motions to organized (S_2 and S_4) motions. The exuberance indicates the efficiency of the vertical momentum flux. Christen et al. (2007) used these parameters to
investigate vertical momentum exchange in an urban district and elucidated the
roles of coherent structures in momentum transport.

Figure 13 shows the vertical profiles of ΔS_0 and Ex for the CTL and UNI 641 experiments, which are averaged temporally and spatially in a manner similar to 642 the profiles in Fig. 8, where ΔS_0 in the CTL experiment is generally larger than in 643 the UNI experiment, except for heights around h_{all} . This feature of ΔS_0 contrasts 644 to the vertical profile of the Reynolds stress shown in Fig. 8b, which indicates that 645 the Reynolds stress is nearly identical in the CTL and UNI experiments below 646 height $z = 0.5h_{all}$. This suggests that, despite the similarities in the Reynolds 647 stress seen in the two experiments, the building-height variability in the CTL 648 experiment changes the ratio of ejections to sweeps within the building canopy 649 layer. In the upper layer from heights $z = 2.5h_{all}$ to $z = 10h_{all}$, both ΔS_0 and the 650 Reynolds stress are larger in the CTL experiment than in the UNI experiment. 651 We consider that the increased Reynolds stress in this upper layer in the CTL 652 experiment is caused by a sweep-dominated vertical flux. 653

Figure 13b shows the value of Ex below $z = 2.5h_{all}$ in the CTL experiment 654 to be smaller than that in the UNI experiment. Below $z = 0.5h_{all}$, the decrease 655 in Ex appears to be more pronounced in the CTL experiment than in the UNI 656 experiment even though the respective Reynolds stresses are similar, as shown in 657 Fig. 8b. This indicates that the efficiency of the vertical momentum flux in the 658 canopy layer is reduced by building-height variability. In contrast, the values of 659 Ex in the CTL and UNI experiments are very similar at altitudes above height 660 $z = 2.5 h_{all}$, indicating that the efficiency of the momentum flux above these 661 altitudes is similar for both experiments. 662

Based on the differences between the vertical profiles shown in Fig. 13, we 663 focus on the heights $z = 0.5h_{all}$ and $z = 2.5h_{all}$ to reveal the relationship between 664 building-height variability and turbulent-flow characteristics. Figure 14a and b 665 shows variations in ΔS_0 against λ_p in the CTL and UNI experiments at these two 666 altitudes. At height $z = 0.5h_{all}$, sweeps are dominant among the contributions 667 to the Reynolds stress for both experiments, which is consistent with previous 668 results showing a stronger contribution of sweeps to the total momentum flux 669 than ejections near and below the tops of block arrays (Raupach 1981; Coceal 670 et al. 2007a). From Fig. 13a, it is seen that the contribution of sweeps in the 671 CTL experiment is larger than in the UNI experiment. By contrast, the value of 672 ΔS_0 at height $z = 0.5 h_{all}$ appears to be independent of λ_p in both experiments. 673 However, at height $z = 2.5 h_{all}$, the value of ΔS_0 in the CTL experiment increases 674 with λ_p when $\lambda_p > 0.32$, while in the UNI experiment, it is independent of λ_p . 675 The increase in ΔS_0 in the CTL experiment is consistent with the Reynolds-stress 676 results shown Fig. 11b, thus suggesting that sweeps contribute to the increase in 677 Reynolds stress for $\lambda_p > 0.32$. Similar results were noted by Kanda (2006). 678

Figure 14c and d shows Ex plotted against λ_p in the CTL and UNI experiments at heights z = 0.5 and $2.5h_{all}$, respectively, where the difference in the value of Exat height $z = 0.5h_{all}$ increases with λ_p , suggesting the dominance of unorganized structures as λ_p increases. As shown in Fig. 13b, at height $z = 2.5h_{all}$ the values of Ex in both experiments are practically independent of λ_p .

⁶⁸⁴ By setting H in Eq. 12 to a value larger than zero, we evaluate the extent to ⁶⁸⁵ which extreme instantaneous momentum fluxes contribute to the total Reynolds ⁶⁸⁶ stress in a certain period. We define the percentage contribution to the Reynolds ⁶⁸⁷ stress of a value of u'w' larger than the Reynolds stress by a factor of H using

$$E_H \equiv \sum_{i=1}^4 S_{i, H} = \sum_{i=1}^4 \langle u'w' \rangle_{i, H} / \overline{u'w'}.$$
 (17)

⁶⁶⁸ Unlike in Raupach (1981) in which each component of the momentum flux was ⁶⁶⁹ evaluated, all of the components in Eq. 17 are added to assess the total of the ⁶⁹⁰ extreme momentum fluxes. We set H = 20 here to extract extreme values of u'w', ⁶⁹¹ with qualitatively similar results also found with H = 15 and H = 10. Thus, H =⁶⁹² 20 is assumed to be a representative value.

Figure 14e and f shows the variations of E_{20} with λ_p at heights z = 0.5 and 693 $= 2.5 h_{all}$, respectively, for both experiments. The results at height $z = 0.5 h_{all}$ 7. 694 reveal small differences between the CTL and UNI experiments and are, in general, 695 larger than those at height $z = 2.5 h_{all}$. This indicates that the flow is highly 696 turbulent at height $z = 0.5h_{all}$, and that the extreme values of the momentum 697 flux contribute more significantly to the total momentum flux at this altitude. 698 However, as the magnitude of u'w' itself is low at height $z = 0.5h_{all}$, the effects of 699 the fluctuation itself may not be very strong. It is seen that, at height $z = 2.5 h_{all}$, 700 the value of E_{20} in the CTL experiment increases with λ_p , but is independent of 701 λ_p in the UNI experiment. Moreover, the shape of the relationship between E_{20} 702 at height $z = 2.5 h_{all}$ and λ_p in the CTL experiment appears to be quite similar to 703 that between $\lambda_{p, 2.5h_{all}}$ and λ_p shown in Fig. 12b. This suggests that increasing 704 the number of buildings at height $z = 2.5h_{all}$ generates highly turbulent flows at 705 higher values of λ_p . 706

The increase in the contribution from extreme values of u'w' to the Reynolds stress at height $z = 2.5h_{all}$ in the CTL experiment occurs because the buildingheight variability in this experiment leads to a higher momentum flux at this

altitude as clearly indicated in Fig. 15a, which shows the horizontal cross-section 710 of E_{20} over a 1 km by 1 km area within one of the business districts. It is seen 711 that high values of E_{20} appear in areas around randomly and sparsely distributed 712 buildings. In contrast, areas with higher E_{20} values also correspond to areas with a 713 weak Reynolds stress and small value of Ex (see Fig. 15b and c), which indicates 714 the small contribution of the extreme momentum flux around buildings to the 715 total momentum flux, and is not related to organized turbulent motion. From 716 the features demonstrated in Figs. 14 and 15, it is seen that the turbulent flow 717 characteristics and contributions of extreme momentum fluxes are significantly 718 influenced by the presence of buildings with significant height variability. 719

We have shown the qualitative consistency of the Reynolds stress and quadrant analysis results, if averaged both in time and space, with that over block arrays with variable height. In contrast, the inhomogeneous profiles of the turbulent-flow characteristics (Fig. 15) suggest that the local characteristics of the turbulent flow over urban surfaces are significantly influenced by the inhomogeneity of actual urban buildings, and would not be expected to be similar to that over idealized block arrays.

727 6 Summary and Conclusions

An LES investigation of the turbulent flow over the city of Kyoto has been conducted to investigate the effects of building-height variability on the turbulence in the lower part of the urban boundary layer. A digital surface model data has reproduced the actual buildings of Kyoto in the LES model.

We used roughness parameters such as H_{ave} , σ_H , λ_p , and λ_f to evaluate the 732 morphological characteristics of buildings, and compared these parameters with 733 those derived for Tokyo, Nagoya as well as for North American and European 734 cities. For $\lambda_p > 0.3$, the value of λ_f for Kyoto is small compared with the em-735 pirical values for Tokyo and Nagoya, but similar to those obtained for European 736 cities. The relationship between H_{ave} and σ_H in Kyoto agrees closely with the em-737 pirical profile. From these comparisons, the building morphological characteristics 738 of Kyoto indicate a dense distribution, and buildings with a variety of heights. 739

We compared the LES results with observations of atmospheric turbulence obtained using a sonic anemometer and a Doppler lidar at the Ujigawa Open Laboratory, which is an area included in the main region of the LES model. For this comparison, certain periods were extracted from the total set of observations to meet the weather conditions assumed in the LES model. The model is to reproduce the observed characteristics of turbulence up to a height of about 150 m.

We carried out two experiments: one modelling the actual buildings of Kyoto 746 (CTL), and one (UNI) in which all building heights were set to the average building 747 height in the main region of the city h_{all} . We find small differences between the 748 CTL and UNI experiments in terms of the mean streamwise velocity component 749 and the Reynolds stress at height $z = 0.5h_{all}$, but large differences at height 750 $z = 2.5 h_{all}$. The spatial fields of time-averaged streamwise velocity components 751 and Reynolds stresses produced in the CTL experiment indicate regions of reduced 752 velocity and strong Reynolds stress behind sparsely and randomly distributed 753 buildings at height $z = 2.5 h_{all}$; this contrasts with the UNI results, in which 754 these fields at height $z = 2.5 h_{all}$ are smooth. We investigated the relationships 755 between turbulent statistics and λ_p evaluated over 1 km by 1 km areas to reveal 756

the differences between the CTL and UNI experiments. The Reynolds stress in the 757 CTL experiment at height $z = 2.5 h_{all}$ is larger than that in the UNI experiment 758 when $\lambda_p > 0.32$, while the Reynolds stress at height $z = 0.5h_{all}$ is similar for 759 both experiments. We suggest that the increase in the Reynolds stress at height 760 $z = 2.5 h_{all}$ is caused by the presence of some building clusters at height $z = 2.5 h_{all}$ 761 in the CTL experiment, and that a value of λ_p of about 0.3 is the threshold above 762 which the effects of building-height variability become obvious over various urban 763 surfaces. 764

A quadrant analysis was used to investigate the characteristics of turbulent 765 coherent flows. Sweeps in the CTL experiment at height $z = 2.5 h_{all}$ are found 766 to increase with λ_p for $\lambda_p > 0.32$, which is similar to that seen in the Reynolds 767 stress for $\lambda_p > 0.32$, suggesting the increase in Reynolds stress is caused by the 768 presence of sweeps. The vertical momentum flux in the CTL experiment is less 769 efficient than that in the UNI experiment at height $z = 0.5h_{all}$, which indicates 770 that the building-height variability in the CTL experiment reduces the efficiency 771 of the flux in the canopy layer. 772

The contributions of the extreme instantaneous momentum flux to the total 773 Reynolds stress were also investigated. The amount of extreme momentum flux 774 in the CTL experiment at height $z = 2.5h_{all}$ depends strongly on the presence 775 of buildings at this altitude. Examination of horizontal cross-sections reveals that 776 areas with extreme momentum fluxes are distributed around buildings. However, 777 the efficiency of the Reynolds stress and momentum flux are small in areas with 778 an extreme momentum flux, implying its negligible contribution around build-779 ings to the net Reynolds stress, as well as the lack of association with coherent 780 turbulent motions. The relationships between turbulent coherent structures and 781

⁷⁸² building-height variability were investigated through the use of space- and time-⁷⁸³ averaged profiles. However, future research on turbulent coherent structures over ⁷⁸⁴ urban surfaces should focus on instantaneous and local structures, such as vortex ⁷⁸⁵ structures behind high, isolated buildings (Park et al., 2015), and flow patterns ⁷⁸⁶ in block arrays associated with coherent structures above blocks (Inagaki et al., ⁷⁸⁷ 2012).

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798 References

- 799 Bou-Zeid E, Overney J, Rogers BD, Parlange MB (2009) The effects of build-
- ing representation and clustering in large-eddy simulations of flows in urban
- canopies. Boundary-Layer Meteorol 132(3):415–436
- ⁸⁰² Cañadillas B, Westerhellweg A, Neumann T (2011) Testing the performance of
- a ground-based wind lidar system: One year intercomparison at the offshore
- ⁸⁰⁴ platform fino1. Dewi Mag 38:58–64

- ⁸⁰⁵ Cheng H, Castro IP (2002) Near wall flow over urban-like roughness. Boundary-
- 806 Layer Meteorol 104(2):229–259
- ⁸⁰⁷ Chorin AJ (1967) A numerical method for solving incompressible viscous flow
 ⁸⁰⁸ problems. J Comput Phys 2(1):12–26
- ⁸⁰⁹ Christen A, van Gorsel E, Vogt R (2007) Coherent structures in urban roughness
 ⁸¹⁰ sublayer turbulence. Int J Climatol 27(14):1955–1968
- 811 Coceal O, Thomas T, Castro I, Belcher S (2006) Mean flow and turbulence
- statistics over groups of urban-like cubical obstacles. Boundary-Layer Meteorol 121(3):491-519
- ⁸¹⁴ Coceal O, Dobre A, Thomas T, Belcher S (2007a) Structure of turbulent flow over
- regular arrays of cubical roughness. J Fluid Mech 589:375–409
- 816 Coceal O, Dobre A, Thomas TG (2007b) Unsteady dynamics and organized struc-
- tures from dns over an idealized building canopy. Int J Climatol 27(14):1943– 1953
- ⁸¹⁹ Counihan J (1975) Adiabatic atmospheric boundary layers: a review and analysis
 of data from the period 1880–1972. Atmos Environ 9(10):871–905
- ESDU (1985) Characteristics of atmospheric turbulence near the ground. part ii: single point data for strong winds (neutral atmosphere). ESDU International,
- 823 London
- ⁸²⁴ Giometto M, Christen A, Meneveau C, Fang J, Krafczyk M, Parlange M (2016)
- ⁸²⁵ Spatial characteristics of roughness sublayer mean flow and turbulence over a
- realistic urban surface. Boundary-Layer Meteorol 160(3):425–452
- ⁸²⁷ Goldstein D, Handler R, Sirovich L (1993) Modeling a no-slip flow boundary with
- an external force field. J Comput Phys 105(2):354–366

- Inagaki A, Castillo MCL, Yamashita Y, Kanda M, Takimoto H (2012) Large-eddy
 simulation of coherent flow structures within a cubical canopy. Boundary-Layer
 Meteorol 142(2):207–222
- Kaimal J, Wyngaard J, Izumi Y, Coté O (1972) Spectral characteristics of surfacelayer turbulence. Q J R Meteorol Soc 98(417):563–589
- Kanda M (2006) Large-eddy simulations on the effects of surface geometry of
 building arrays on turbulent organized structures. Boundary-Layer Meteorol
 118(1):151–168
- Kanda M, Moriwaki R, Kasamatsu F (2004) Large-eddy simulation of turbulent
 organized structures within and above explicitly resolved cube arrays. Boundary Layer Meteorol 112(2):343–368
- 840 Kanda M, Inagaki A, Miyamoto T, Gryschka M, Raasch S (2013) A new aero-
- dynamic parametrization for real urban surfaces. Boundary-Layer Meteorol 148(2):357–377
- Macdonald R, Griffiths R, Hall D (1998) An improved method for the estimation of surface roughness of obstacle arrays. Atmos Environ 32(11):1857–1864
- Nakajima C, Mitsuta Y, Tanaka M (1979) Ujigawa meteorological tower for
 boundary layer monitoring. Annuals of Disaster Prevention Research Institute
 22B(2):127-141
- Nakayama H, Takemi T, Nagai H (2011) Les analysis of the aerodynamic surface
 properties for turbulent flows over building arrays with various geometries. J
- 850 Appl Meteorol Clim 50(8):1692–1712
- Nakayama H, Takemi T, Nagai H (2012) Large-eddy simulation of urban boundary-
- layer flows by generating turbulent inflows from mesoscale meteorological sim-
- ⁸⁵³ ulations. Atmos Sci Lett 13(3):180–186

854	Nakayama H, Leitl B, Harms F, Nagai H (2014) Development of local-scale high-
855	resolution atmospheric dispersion model using large-eddy simulation. part 4:
856	turbulent flows and plume dispersion in an actual urban area. J Nucl Sci Technol
857	51(5):626-638
858	Nakayama H, Takemi T, Nagai H (2015) Large-eddy simulation of turbulent winds
859	during the fukushima daiichi nuclear power plant accident by coupling with a
860	meso-scale meteorological simulation model. Adv Sci Res $12(1){:}127{-}133$
861	Nakayama H, Takemi T, Nagai H (2016) Development of lo cal-scale h igh-
862	resolution atmospheric di spersion m odel using l arge-e ddy s imulation. part 5:
863	detailed simulation of turbulent flows and plume dispersion in an actual urban
864	area under real meteorological conditions. J Nucl Sci Technol 53(6):887–908
865	Oikawa S, Meng Y (1995) Turbulence characteristics and organized motion in a
866	suburban roughness sublayer. Boundary-Layer Meteorol $74(3){:}289{-}312$
867	Oke TR (1988) Street design and urban canopy layer climate. Energy Buil
868	11(1):103–113
869	Park SB, Baik JJ, Han BS (2015) Large-eddy simulation of turbulent flow in a
870	densely built-up urban area. Environ Fluid Mech $15(2){:}235{-}250$
871	Ratti C, Di Sabatino S, Britter R, Brown M, Caton F, Burian S (2002) Analysis
872	of 3-d urban databases with respect to pollution dispersion for a number of
873	european and american cities. Water Air Soil Pollut Focus 2(5-6):459–469
874	Raupach M (1981) Conditional statistics of reynolds stress in rough-wall and
875	smooth-wall turbulent boundary layers. J Fluid Mech 108:363–382
876	Roth M (2000) Review of atmospheric turbulence over cities. Q J R Meteorol Soc
877	126(564):941-990

- Shaw RH, Tavangar J, Ward DP (1983) Structure of the reynolds stress in a canopy
 layer. J Clim Appl Meteorol 22(11):1922–1931
- ⁸⁸⁰ Smagorinsky J (1963) General circulation experiments with the primitive equa-
- tions: I. the basic experiment. Mon Weather Rev 91(3):99–164
- 882 Stoll R, Porté-Agel F (2006) Effect of roughness on surface boundary conditions
- for large-eddy simulation. Boundary-Layer Meteorol 118(1):169–187
- Wallace JM (2016) Quadrant analysis in turbulence research: history and evolution. Annu Rev Fluid Mech 48:131–158
- ⁸⁸⁶ Xie ZT (2011) Modelling street-scale flow and dispersion in realistic windsto-
- wards coupling with mesoscale meteorological models. Boundary-Layer Meteorol
 141(1):53-75
- Xie ZT, Castro IP (2009) Large-eddy simulation for flow and dispersion in urban
 streets. Atmos Environ 43(13):2174–2185
- Xie ZT, Coceal O, Castro IP (2008) Large-eddy simulation of flows over random
 urban-like obstacles. Boundary-Layer Meteorol 129(1):1–23
- ⁸⁹³ Zaki SA, Hagishima A, Tanimoto J, Ikegaya N (2011) Aerodynamic parameters
- of urban building arrays with random geometries. Boundary-Layer Meteorol
 138(1):99–120
- ⁸⁹⁶ Zhu X, Iungo GV, Leonardi S, Anderson W (2017) Parametric study of urban-like
- topographic statistical moments relevant to a priori modelling of bulk aerody-
- ⁸⁹⁸ namic parameters. Boundary-Layer Meteorol 162:231–253



Fig. 1 The study area in which the LES model and observations were carried out is indicated by the red box. The observational site of the Disaster Prevention Research Institute, Kyoto University, is indicated by the white circle. The white arrow indicates the streamwise wind direction. The satellite picture is taken from Google Earth.



Fig. 2 (a) Distribution of building and structure heights in the analysis region of Kyoto. (b) Frequency distribution of building heights in the analysis region. The black bar indicates the frequency distribution of buildings in the overall region, while the grey and hatched bars indicate the frequency distributions of buildings in the regions with x = 0 - 4 km and with x = 7 - 11 km, respectively.



Fig. 3 (a) λ_p calculated for 1 km by 1 km areas over the analysis region. (b) Roughness parameters calculated for 1 km by 1 km areas over the analysis region. In each box, the first row is H_{ave} , the second is σ_H , the third is λ_p , and the fourth is λ_f . Scatter plots (c) between λ_p and λ_f and (d) between σ_H and H_{ave} . The black lines in (c) and (d) indicate the empirical relationships derived from Tokyo and Nagoya, respectively, by Kanda et al. (2013). The values of Salt Lake City and Los Angeles in North America and London, Toulouse, and Berlin in Europe, as indicated by the lower legend, are obtained from Ratti et al. (2002).



Fig. 4 Schematic of turbulent flows formed in the driver region and imposed on the main region as the inflow condition.



Fig. 5 Time series of (a) streamwise and (b) spanwise velocity components at 25-m height observed by the sonic anemometer during periods D1 - D4 and simulated by the LES model. Vertical profiles of (c) mean streamwise velocity component and (d) Reynolds stress normalized by the mean streamwise velocity component in the observations and the LES results. Note that the profiles in (c) are plotted on logarithmic axes. A line with a slope 1/4 is also plotted for reference. Here, 'ane' and 'dop' refer to the observations by the anemometer and Doppler lidar, respectively.



Fig. 6 Vertical profiles of turbulence intensity from the observations, the LES model, and the empirical profiles provided by the ESDU (1985) for the (a) u, (b) v, and (c) w components. The dashed and dashed-dotted lines indicate the upper and the lower limits, respectively, of the rough-surface class based on the ESDU (1985).

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Fig. 7 Power spectra obtained from the sonic anemometer and the LES model at 25-m height plotted on logarithmic axes: (a) u, (b) v, and (c) w components. The dashed line indicates the empirical profile over a rural surface proposed by Kaimal et al. (1972).



Fig. 8 Vertical profiles of (a) time-averaged streamwise velocity component, (b) Reynolds stress, and (c) dispersive flux averaged spatially over the main region. These values are normalized by U_{∞} . Red and blue lines denote the result of the CTL and UNI experiments, respectively. The vertical axis is normalized by $z = h_{all}$. Note that a logarithmic scale is used for the vertical axis.



Fig. 9 Horizontal cross sections of the time-averaged streamwise velocity component normalized by U_{∞} in (a) the CTL experiment at height $z = 0.5h_{all}$, (b) the UNI experiment at height $z = 0.5h_{all}$, (c) the CTL experiment at height $z = 2.5h_{all}$, and (d) the UNI experiment at height $z = 2.5h_{all}$. An upstream part of the main region is shown. The legend indicating the wind speed is present to the right of each panel. The grey shading indicates buildings.



Fig. 10 As Fig. 9, except with the corresponding Reynolds-stress results.



Fig. 11 Variations of Reynolds stress normalized by U_{∞} with λ_p at heights (a) $z = 0.5h_{all}$ and (b) $z = 2.5h_{all}$.



Fig. 12 Variations of λ_p calculated (a) at height $z = 0.5h_{all}$ ($\lambda_{p, 0.5h_{all}}$), (b) at height $z = 2.5h_{all}$ ($\lambda_{p, 2.5h_{all}}$), and (c) λ_f calculated at height $z = h_{all}$ ($\lambda_{f, h_{all}}$) with λ_p at the surface. Note that the scale of the vertical axis differs by panel.



Fig. 13 Vertical profiles of (a) ΔS_0 and (b) Ex averaged spatially over the main region. Red and blue lines denote the results of the CTL and UNI experiments, respectively. The vertical axis is normalized by h_{all} . Note that a logarithmic scale is used for the vertical axis.



Fig. 14 Variations of ΔS_0 with λ_p at (a) heights $z = 0.5h_{all}$ and (b) $z = 2.5h_{all}$, Ex at (c) heights $z = 0.5h_{all}$ and (d) $z = 2.5h_{all}$, and E_{20} at (e) heights $z = 0.5h_{all}$ and (f) $z = 2.5h_{all}$.



Fig. 15 Horizontal cross section of (a) E_{20} , (b) Reynolds stress normalized by U_{∞} , and (c) Ex at height $z = 2.5h_{all}$ over a 1 km by 1 km area within the business district in the CTL experiment. The grey shading indicates buildings.