

Selenium uptake onto natural pyrite

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Received: 3 July 2012 / Published online: 25 July 2012
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Abstract ^{79}Se is a major dose-determining redox-sensitive nuclide in safety analysis of radioactive waste disposal sites. In aqueous solutions, selenium forms soluble anionic species ($\text{Se}^{\text{IV}}\text{O}_3^{2-}$ and $\text{Se}^{\text{VI}}\text{O}_4^{2-}$) that hardly sorb on negatively charged surfaces of common host-rock minerals. However, Se is known to have a strong affinity with sulphides and interacts with pyrite, a common minor mineral of argillaceous rocks being considered as host formations for radioactive waste repositories. In this study, we present micro- and bulk X-ray spectroscopy data ($\mu\text{-XRF}$, $\mu\text{-XANES}$, and EXAFS) showing that, under nearly anoxic conditions, dissolved SeO_3^{2-} and SeO_4^{2-} sorb directly onto the pyrite surface and are subsequently reduced to Se^0 with increasing ageing time (up to 8 months). These results suggest that the mobility of $^{79}\text{Se}^{\text{IV}}$ released from radioactive waste could greatly decrease through uptake on the pyrite surface followed by transformation into a sparingly soluble reduced form.

Keywords X-ray spectroscopy · Redox · Selenium · Pyrite · Radioactive waste · Micro-XANES · Micro-XRF · Sorption

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Introduction

Radioactive waste produced worldwide by the operation of nuclear power plants is nowadays a high-priority global environmental problem. The generally accepted option for the safe disposal of radioactive waste is the long-term storage in a system of protective materials (engineered barriers) within a suitable deep geological formation [1]. This multiple barrier system must guarantee the long-term isolation of the waste from the biosphere.

In most scenarios for planned repository sites (see e.g. [2, 3]) ground water will slowly penetrate the repository and aqueous corrosion will ultimately release the long-lived radionuclides from the waste. The mobility of each radionuclide will then largely depend on its chemical properties, specifically on its capability to bind to minerals and materials in the near-field and far-field of the repository.

In safety assessment calculations, long-lived nuclides such as ^{14}C , ^{129}I , ^{36}Cl and ^{79}Se dominate the calculated dose curves in safety assessment calculations of repository sites [2]. These nuclides form anionic species that do not sorb significantly on exposed surfaces of major minerals present in the repository near-field (e.g. clays, carbonates, and sulfates), since their surfaces are negatively charged under the neutral to alkaline pHs prevailing in such environments. They are therefore considered to be mobile compared with fission products and actinides dissolved as cationic species.

However, the case of ^{79}Se (half-life 3.77×10^5 years) is somehow peculiar, because selenium may be stable in geochemical environments either as reduced species ($\text{Se}^{-\text{II}, -\text{I}, 0}$) or as oxidized species ($\text{Se}^{\text{IV}, \text{VI}}$). Whereas reduced Se species are fairly insoluble and immobile, oxidized species form soluble, poorly sorbing $\text{Se}^{\text{IV}}\text{O}_3^{2-}$

(selenite) and $\text{Se}^{\text{VI}}\text{O}_4^{2-}$ (selenate) oxo-anions. In vitrified high-level waste derived from reprocessing of spent nuclear fuel, Se resides in the oxidized Se^{IV} form (see later) and will therefore be released to the aqueous solution as easily soluble selenite anions. For this reason, ^{79}Se is normally treated as soluble anionic nuclide in safety assessments. However, in such calculations the effects of minor minerals are conservatively neglected. For instance, there is increasing evidence that Se may bind, through reductive sorption processes, to minor sulfide minerals such as pyrite (FeS_2) and pyrrhotite (FeS).

In the high-level waste repository planned in Switzerland, pyrite will be present as a minor mineral both in bentonite, foreseen as buffer material to seal the underground tunnels hosting the waste (about 0.3 wt%), and in the surrounding Opalinus clay host formation (about 1 wt%) [2]. Pyrite has a negative surface charge above pH_{pzc} of 1.7 [4], and is therefore not expected to bind anions via electrostatic forces (outer sphere complexes). Nevertheless, significant uptake of selenium on pyrite from selenite/selenate solutions was observed in batch uptake experiments [5]. This quantitative uptake study suggested reduction of dissolved Se^{IV} by sulfides and other Fe(II)-bearing minerals, which was confirmed by subsequent EXAFS/XANES studies [6–8].

In this work, Se^{IV} and Se^{VI} uptake on natural pyrite was studied at 80 °C under nearly anoxic conditions by micro X-Ray fluorescence (μ -XRF), micro X-ray absorption near edge spectroscopy (μ -XANES) and bulk extended X-ray absorption fine structure (EXAFS) spectroscopy, after interaction with Na_2SeO_3 and Na_2SeO_4 solutions during up to 8 months. The aims of this investigation were: (a) to study the distribution of Se on the pyrite surface at the microscopic (micrometer) scale and the kinetics of related redox processes, with special attention to possible differences between the uptake from Se^{IV} and Se^{VI} bearing solutions; (b) to determine the atomic-scale coordination environment of Se sorbed on the pyrite and (c), to understand the mechanisms through which Se is bound to pyrite. By combining the information obtained, consequences in view of the future treatment of ^{79}Se in safety assessment calculations are evaluated.

Materials and methods

Mineral and solutions preparation

The use of natural pyrite was preferred due to the complexity of laboratory procedures for synthesizing pure pyrite (α - FeS_2) free of other Fe_xS_y sulfides [9, 10]. Pyrite samples from the Huanzala (Peru) and Navajun (Spain) mines purchased from Ward's Natural Science, Rochester,

NY, were used. Huanzala is a Zn-Pb strata-bound hydrothermal ore deposit hosted in a Cretaceous limestone formation, with pyrite formation occurring at early stages of mineralization as fine-grained to coarse-grained granular masses. The Navajun ore, hosted in a Jurassic marl formation, is characterized by euhedral cm to dm sized well-shaped cubic pyrite crystals. The X-ray diffraction patterns of the sample powders, recorded on a Phillips Xpert-Pro diffractometer ($\text{CuK}\alpha$ radiation, 2θ from 5° to 70°) confirmed pyrite as the only detectable crystalline phase.

For the μ -XAS/ μ -XRF measurements, Navajun pyrite crystals were cut into $7 \times 7 \times 2 \text{ mm}^3$ chips ($\sim 0.5 \text{ g}$) and polished under normal air. All subsequent steps were carried out in a N_2 -flushed glove box (5 ppm O_2 max). The chips were re-polished with diamond paste to eliminate any surface oxidation artifact and then introduced in glass bottles filled with 50 ml of 0.01 M Na_2SeO_3 or 0.01 M Na_2SeO_4 solution, with or without 0.001 M NaHS. Addition of NaHS was intended to provide completely anoxic conditions. All solutions were prepared in the glove box using analytical grade reagents (Fluka or Sigma-Aldrich) and Milli-Q® water (Millipore, Billerica, USA).

Huanzala pyrite for the bulk Se-K EXAFS experiments was also prepared in the glove box. After grinding in an agate mortar and sieving, the two finer fractions (63–250 μm and $<63 \mu\text{m}$) were washed with concentrated (38 %) HCl during at least 6 h in order to digest iron oxide traces.

Approximately 0.4 g ($<63 \mu\text{m}$ fraction) or 0.8 g (63–250 μm fraction) of the pyrite powder were introduced into glass bottles and filled with 50 or 100 ml, respectively, of Se-solution (0.01 M Na_2SeO_3 or 0.01 M Na_2SeO_4).

All samples were equilibrated at 80 °C in an oven inside the glove box. The pyrite powder suspensions were shaken by hand at least once per week and periodically adjusted to $\text{pH} \sim 8.0$ by addition of HCl or NaOH. Except for one sample where substantial pH drops were observed (Se-23, see later) the pH remained fairly stable and only minimal adjustments were needed. After a reaction time of 2 weeks (sample Se-07) or 2 months (all other samples), 2.5 ml of solution were pipetted and passed through a filter with molecular weight cut-off (MWCO) of 5×10^4 Dalton. The first 0.5 ml aliquot of filtrate served to equilibrate the filter with Se and was therefore discarded. The subsequent 1.8 ml were diluted to 18.0 ml with Milli-Q® water and kept for the wet chemical analyses. Concentrations of Se, Fe and S in the diluted filtrate were measured by inductively coupled plasma—optical emission spectroscopy (ICP-OES).

X-ray spectroscopy

After the desired contact time, the Se-treated pyrite samples (chips or powders) were separated from the solution

bottles for the XRF/XAS measurements. All operations were carried out in the glove box. The polished chips were taken from the Se-solutions with a pair of tweezers and washed with ethanol for a few seconds. Residual ethanol on the surface was eliminated by shortly drying with a tissue. Later tests at the beamline (see section “Selenium sorption”) confirmed that this procedure eliminates any measurable trace of dissolved Se from the pyrite surface (which might cause Se-salt precipitation artifacts due to drying). All samples were mounted on a sample holder equipped with a gas inlet and outlet to allow for XRF/XAS measurements under constant inert gas flushing (N_2 or He) and then transported to the beamline in vacuum desiccators, to prevent oxidation.

The μ -XRF maps and μ -XANES spectra were collected at the micro-XAS beamline, Swiss Light Source (SLS), Switzerland. Measurements were carried out in a special sample holder flushed with inert gas (N_2 or He) at room temperature, in fluorescence mode, using a single channel Si detector and with a beam size focused to diameters of ~ 1 to $5 \mu\text{m}$. Energy calibration was carried out by assigning the value 12,668 eV to the white line of commercial $\text{Na}_2\text{Se}^{\text{IV}}\text{O}_3$ (following [11]). The μ -XRF maps were generated using the Fmap software provided at the Micro-XAS beamline.

Reference samples of $\text{Na}_2\text{Se}^{\text{VI}}\text{O}_4$, $\text{Na}_2\text{Se}^{\text{IV}}\text{O}_3$, $\text{Se}^{\text{IV}}\text{O}_2$, grey Se^0 (Fluka or Sigma-Aldrich) and iron selenide ($\text{FeSe}^{-\text{II}}$) (Alfa Aesar) were prepared for transmission measurements by diluting with appropriate amounts of boron nitride (BN). XRD data showed that the purchased FeSe contains minor amounts of Se^0 ($\sim 6 \text{ wt}\%$). To avoid oxidation, the FeSe reference sample was prepared in the glove box and mounted on the aforementioned sample holder. The other reference samples were not sensitive to oxidation and could therefore be prepared in air. All references and bulk XAS spectra were measured simultaneously in transmission and in fluorescence modes.

In contrast to the polished chips, the pyrite powder samples were measured using a defocused X-ray beam (about $2 \times 2 \text{ mm}$) to obtain bulk-XANES spectra. Two pyrite powder samples treated with Se^{IV} (Se-21/Se22) and one treated with Se^{VI} (Se-23) were prepared after 2 months ageing at the $80 \text{ }^\circ\text{C}$ by extracting a small portion of Se-sorbed pyrite from the solution with a pair of tweezers, absorbing the residual solution with a tissue, introducing the pyrite into a sample holder and finally sealing it with kapton tape in the glove box. The Se^{IV} suspensions were subsequently left at room temperature in the glove box during additional 6 months. After this additional ageing time further samples were prepared in the same way in order to record Se K-edge fluorescence EXAFS spectra. These spectra were measured in a cryostat system operated with N_2 at 90 K at the Dutch Belgian Beamline

(DUBBLE), European Synchrotron Radiation Facility (ESRF), in Grenoble, using a 9-channel Ge solid-state detector. The Si (111) monochromator angle was calibrated by assigning the energy of 11,979 eV to the first inflection point of the K-edge of an Au metal foil. The reference compounds Se^0 -grey and Na_2SeO_3 were measured in transmission mode with ionization chambers filled with 100 % Ar (I_0) and a 100 % Kr (I_L). These reference samples were the same previously measured at the Micro-XAS beamline, thus allowing for careful alignment of XANES spectra obtained at the two beamlines. EXAFS data reduction was performed using the IFEFFIT software package (Athena and Artemis) following standard procedures. The spectra were normalized by fitting a third-degree polynomial. The energy was converted to photoelectron wave vector units k (\AA^{-1}) by assigning the origin E_0 to the first inflection point of the absorption edge. Radial Structure Functions (RSF) were obtained by Fourier transformation of the k^3 -weighted $\chi(k)$ function between 3.7 and 12.7 \AA^{-1} , using a Kaiser-Bessel window and an apodization parameter of 1.0. Fits of the first two shells were performed in real space. Theoretical scattering paths for the fit were calculated with FEFF 8.40 using the structure of trigonal Se^0 -grey and an automated self-consistency option for potential calculations. The amplitude reduction factor (S_0^2) was set to 0.9 based on the method indicated in Koningsberger and Prins [12], using the shake-off probability tables given by Carlson et al. [13].

Because of non-uniform energy calibrations at the various beamlines, systematic E_0 shifts were applied in order to align the Se^0 -grey and Na_2SeO_3 XANES spectra obtained at Micro-XAS and DUBBLE using the same reference samples and to define a set of energy values consistent with those given by Scheinost et al. [8]. These shifts ($+1.0 \text{ eV}$ for DUBBLE data and -3.4 eV for the Micro-XAS data) were applied systematically to all raw XANES data.

Results

Overview of wet chemical analyses

A summary of the collected wet chemistry data is given in Table 1. Selenium ICP-OES analyses of the final solutions (data not shown in Table 1) indicated in all cases that only small fractions of the initially dissolved Se were taken up by the pyrite (1–16 %).

Dissolved Fe concentrations (data not shown in Table 1) were mostly below the detection limit and always $<1 \mu\text{M}$. Except for two samples, the dissolved sulfur concentrations of filtrates in contact with NaHS-free solutions were below the detection limit of $\sim 6 \times 10^{-5} \text{ M}$. For samples Se-21

Table 1 Overview of pyrite interaction experiments with 0.01 M Na₂SeO₃ and 0.01 M Na₂SeO₄ solutions

Pyrite type	Initial Se redox state	Sample code	Ageing time	pH _{min} (80 °C)	pH _{max} (80 °C)	[NaHS] mM	[S] _{final} (μM)
Powder <63 μm (Huanzala)	(IV)	Se-21	2 + 6 months	8.0	8.1	–	3.4 × 10 ²
	(IV)	Se-22	2 + 6 months	7.9	8.0	1.0	1.17 × 10 ³
	(VI)	Se-23	2 months	4.7	8.1	–	7 × 10 ¹
Polished chips (Navajun)	(IV)	Se-07	2 weeks	8.0	8.0	–	<DL
	(IV)	Se-08	2 months	8.0	8.1	–	<DL
	(IV)	Se-09	2 months	8.1	8.4	1.0	7.1 × 10 ²
	(VI)	Se-10	2 months	7.9	8.3	–	<DL
	(VI)	Se-11	2 months	8.1	8.4	1.0	4.4 × 10 ²

All experiments were carried out at 80 °C in a glove box flushed with N₂ (5 ppm O₂ max). The detection limit (DL) of S concentrations is about 60 μM

and Se-23, ~0.1–0.3 mM dissolved sulfur were detected. This probably indicates the presence of sulfate in solution and hence pyrite oxidation, in spite of the precautions taken to minimize reaction with oxygen. This interpretation is supported for the Se-23 sample by the low pH-values (down to pH 4.7) measured during the experiment in spite of repeated adjustments with NaOH. For the three NaHS-bearing solutions, the ICP-OES data indicate ~100 % S recovery in the case of experiment Se-22, whereas significant losses (30 and 60 %, respectively) were measured for samples Se-09 and Se-11. A black precipitate, possibly Se⁰ or S⁰, was observed in the Se-09 solution.

μ-XRF/XAS and bulk EXAFS experiments

Reference spectra

Commercially available Se compounds (Na₂Se^{VI}O₄, Na₂Se^{IV}O₃, Se^{IV}O₂, Se⁰-grey and FeSe^{-II}) were measured as references to discriminate oxidation states based on XANES spectra. In Table 2, the energies obtained for the absorption edges (defined by the maxima of the second derivative) and white lines of the normalized XANES spectra are listed. These values indicate that the oxidized selenium forms (Se^{IV,VI}) are easily distinguished from the reduced species (Se^{0,-II}) due to the large energy shifts (at least 4.7 eV). A distinction is also possible between Se^{IV} and Se^{VI} (relative edge shift: 2.0 eV). In contrast, discrimination between Se⁰ and Se^{-II} based on edge or white line positions is difficult

considering the small and ill-defined energy shifts of -1.3 ± 0.6 eV and -0.7 ± 0.6 eV, respectively. The large uncertainty is a consequence of the low amplitude of the XANES absorption peak for this compound. Moreover, the position of the Se^{-II} edge is known to depend on the neighbouring atoms [6].

Discrimination between Se⁰ and Se^{-II} might be achieved by comparing the normalized heights of the absorption peak maxima (white lines), which is strongly depressed for FeSe^{-II}. However, this method is ambiguous in spots containing a mixture of Se^{IV} and Se⁰, since linear combinations of the Se^{IV} and Se⁰ reference spectra may produce spectra very similar to that of FeSe. Also, a few spots with very high Se concentrations in the pyrite samples showed depressed absorption peaks due to self-absorption effects. The situation is further complicated by the possible formation of ferroselite (FeSe₂). According to recent measurements [14] the Se K-edge energy position of this compound is almost indistinguishable from those of Se⁰ and FeSe, and the normalized height of the absorption peak is intermediate between these two compounds.

Therefore, at the available energy resolution our XANES spectra are not always able to discriminate between Se⁰, Se^{-I} and Se^{-II} species. Only when both edge position and normalized height of our Se-pyrite sample spectra coincide with those of the Se⁰-grey reference can lower oxidation states be excluded in the samples (fortunately, this is a frequent case). For some XANES spectra, typically when the normalized height of the absorption peak is low, the Se^{0,-I,-II} oxidation

Table 2 Assigned edge (E_0) and white line maxima (E_w) for the listed reference compounds

Reference	Oxidation state	E_0 (eV)	Shift from Se ⁰ (eV)	E_w (eV)	Shift from Se ⁰ (eV)
FeSe*	-II	12,654.8 ± 0.6	-1.3 ± 0.6	12,657.0 ± 0.6	-0.7 ± 0.6
Se ⁰ -grey	0	12,656.1	0	12,657.7	0
Na ₂ SeO ₃	+IV	12,660.8	+4.7	12,662.4	+4.7
Na ₂ SeO ₄	+VI	12,662.8	+6.7	12,665.3	+7.6

* Edge and white line positions of FeSe cannot be determined with precision due to the low edge-step amplitude and the broad white line peak

states cannot be distinguished, in which case we will use the generic notation Se^{red} instead of explicit oxidation numbers.

Experiments with Se^{IV}

The μ -XRF maps reveal a highly heterogeneous Se distribution on the pyrite surface after interaction with Na_2SeO_3 solution during 2 weeks or 2 months. In each studied sample in contact with Se^{IV} bearing solutions, highly-concentrated Se clusters of a few μm size were detected, surrounded by areas with lower and uniform Se concentration (Figs. 1, 2).

Figure 3 shows the μ -XANES of the spots labeled in Figs. 1 and 2 and the XANES of the reference compounds. The spectra obtained for the two concentrated spots after 2 weeks interaction (Fig. 2) differed considerably, indicating that spot 5 was dominated by Se^{IV} , whereas spot 6 represented a mixture of Se^{red} and Se^{IV} , with a high proportion of reduced species. Similar mixed oxidation state spots were found in other areas of this sample. In the sample aged during 2 months, the large majority of the μ -XANES spectra indicated the presence of only Se^{red} . Occasionally spots indicating a minor proportion of Se^{IV} were found. The coexistence of Se^{IV} and Se^{red} in the same spots suggests a sequential mechanism by which selenium is first adsorbed as Se^{IV} and subsequently reduced in situ on the pyrite surface.

The progressive reduction of the sorbed selenium was confirmed by the results obtained for Se^{IV} -treated pyrite powder measured after 2 and 8 months ageing time. The bulk XANES spectrum of the Se-21 sample after interaction with Na_2SeO_3 solution during 2 months at 80°C is consistent with a mixture of about 50 mol.% Se^{IV} and 50 mol.% Se^0 (Fig. 4, circles).

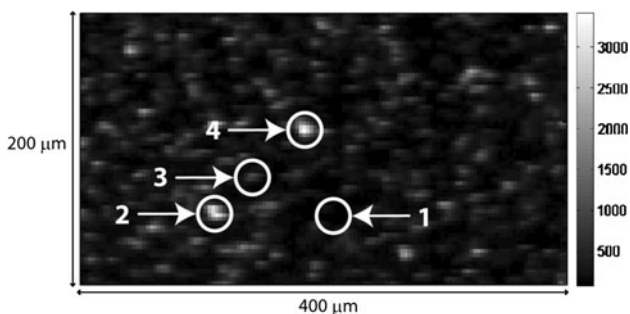


Fig. 1 Micro-XRF map of the Se-K α signal for a polished pyrite chip (sample Se-08) reacted during 2 months with Na_2SeO_3 solution under anoxic conditions, showing the location of μ -XANES collection spots (labeled 1–4). The beam diameter was $\sim 2\ \mu\text{m}$ and the pixel size $3\ \mu\text{m} \times 3\ \mu\text{m}$. The intensity scale gives the detector response in counts per seconds

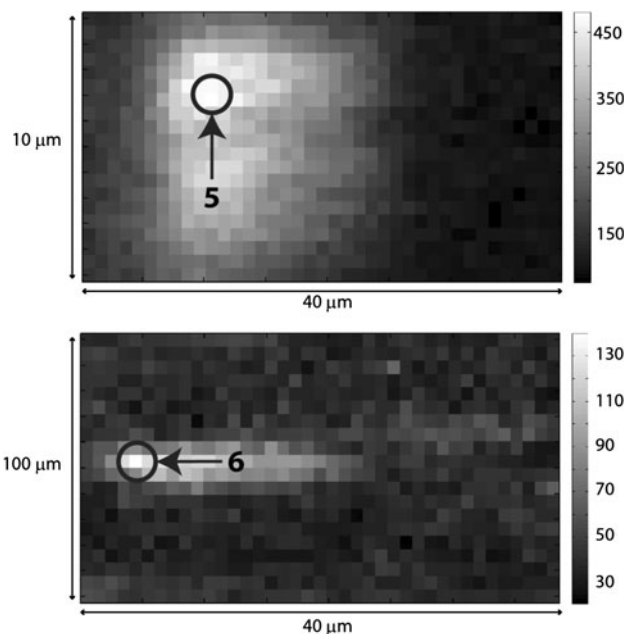


Fig. 2 Micro-XRF elemental maps of the Se-K α signal for the polished pyrite chip reacted during 2 weeks with Na_2SeO_3 solution under anoxic conditions (sample Se-07). Selected analyzed μ -XANES spots are circled. The beam diameter was $\sim 2\ \mu\text{m}$ and the pixel size $3\ \mu\text{m} \times 3\ \mu\text{m}$. The intensity scale gives the detector response in counts per seconds

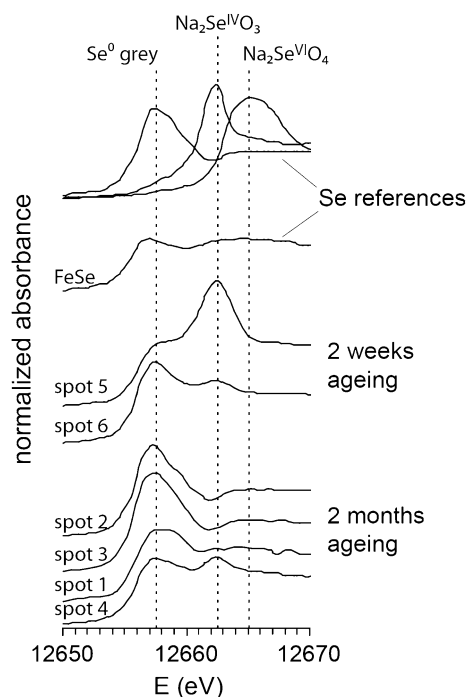


Fig. 3 Se K-edge μ -XANES of selected spots (labeled in Figs. 1, 2) from pyrite chips treated with Na_2SeO_3 solutions, compared to the XANES spectra of reference compounds for $\text{Se}^{-\text{II}}$, Se^0 , Se^{IV} and Se^{VI}

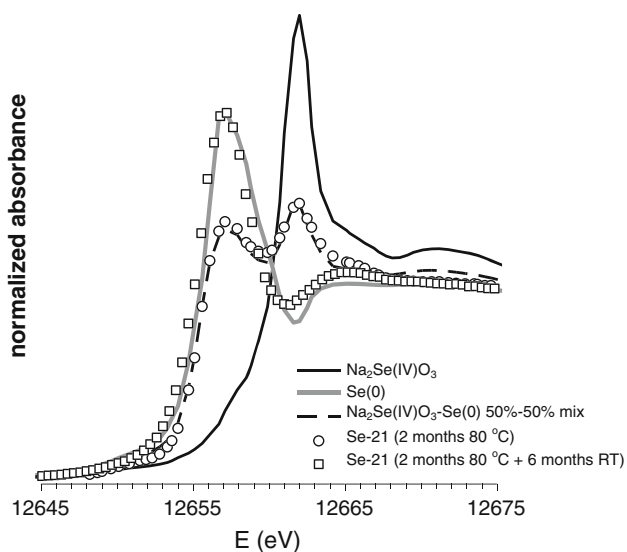


Fig. 4 Bulk XANES of pyrite powder (sample Se-21) treated with Na_2SeO_3 solution during 2 months at 80°C and after additional 6 months ageing at room temperature (RT). The normalized Se-pyrite spectra are superimposed on the spectra of pure Se^0 , of pure Na_2SeO_3 and of an equimolar mixture of the two reference compounds

After additional 6 months ageing time in the glove box at room temperature, complete reduction to Se^0 was observed (Fig. 4, squares), which was confirmed by the EXAFS analysis of the same sample (Se-21 after 8 months total ageing time) summarized in Fig. 5 and Table 3.

For the Se^0 reference, the Radial Structure Function (RSF) showed a major peak at $R + \Delta R \sim 2.3 \text{ \AA}$, and a minor peak at $R + \Delta R \sim 3.2 \text{ \AA}$, which could be fitted assuming the presence of two Se–Se shells (Table 3). The CN of 2.14 was close to the expected value of 2. Further peaks were very weak and could not be fitted satisfactorily.

The EXAFS spectrum and derived fit parameters of the Se-treated pyrite were very similar to those of the Se^0 reference. Also in this case, a satisfactory fit was obtained assuming two Se–Se shells. The interatomic distances (R) determined for the pyrite sample and the Se^0 reference were identical within experimental uncertainties, but the

corresponding coordination numbers (N) differed. In particular, $N_{\text{Se-Se}1}$ for the Se-treated pyrite indicated three neighbouring atoms at $\sim 2.38 \text{ \AA}$ instead of the two expected in most crystalline Se^0 polymorphs. Additionally, the value of $N_{\text{Se-Se}2}$ was lower than expected, which could indicate structural disorder.

In order to test for the possible formation of a solid solution with pyrite or of isolated Se complexes on the pyrite surface, we attempted fits assuming Se–Fe or Se–O shells and using FeSe or Na_2SeO_3 , respectively, as models for FEFF8 calculations. Any attempt to fit the spectrum of the Se-pyrite sample with the mentioned alternative models was unsuccessful. Formation of pure metallic Se on the pyrite surface is therefore conclusively proven by our EXAFS analysis, whereas no evidence was found that Se sorbed on the Se-21 sample did form a solid solution with pyrite or surface complexes, nor a separate Fe–Se or $\text{Se}^{\text{IV,VI}}$ oxide phase. These alternative Se species, if they exist in sample Se-21, must be present in minor amounts.

All attempts to fit our data starting from other Se^0 structural models led to comparable results. The differentiation between various Se^0 polymorphs was therefore not possible. Scheinost et al. [8] studied EXAFS spectra of various Se reference compounds, including an amorphous red Se^0 phase, a crystalline red Se^0 phase and a trigonal grey Se^0 phase. Their fitted N and R values for the first Se–Se shell did not significantly vary among the different compounds. However, their fits of the second Se–Se shell yielded for the red Se^0 compounds longer interatomic distances ($R + \Delta R \sim 3.7 \text{ \AA}$) than for the Se^0 -grey reference ($R + \Delta R \sim 3.4 \text{ \AA}$). Our Se^0 -grey reference material also yielded $R + \Delta R \sim 3.4 \text{ \AA}$ for the second shell, although with a lower coordination number ($N \sim 2.6$) than that found by Scheinost et al. [8] ($N \sim 4$). Based on the high precision of the interatomic distances and considering that coordination numbers determined by EXAFS are frequently inaccurate, we conclude that the Se sorbed on the Se-21 pyrite sample was also trigonal grey Se^0 .

Fig. 5 Se K-edge spectra of Se^0 -grey reference material and of Se sorbed onto pyrite (sample Se-21 after 8 months ageing time): k^3 -weighted normalized spectra (left) and corresponding Fourier Transforms (right)

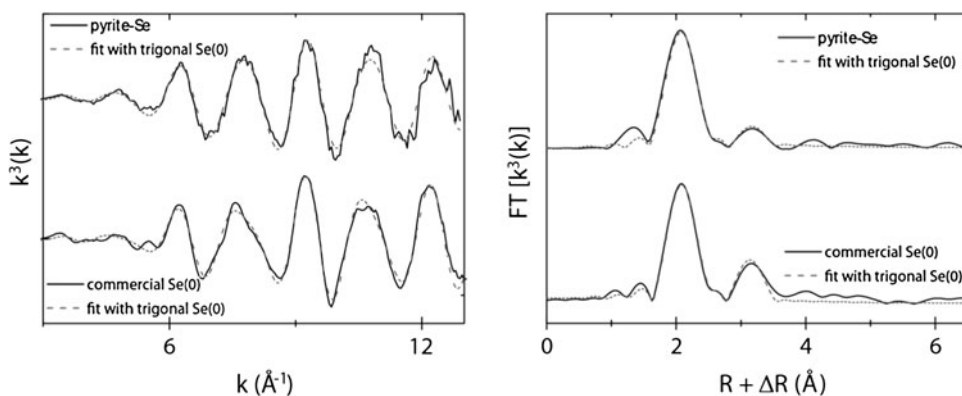


Table 3 EXAFS fit results of Se sorbed onto pyrite after 8 months ageing time (sample Se-21), showing coordination numbers (N), interatomic distances (R), Debye–Waller factors (DW), energy shift (ΔE_0) and the R-factor

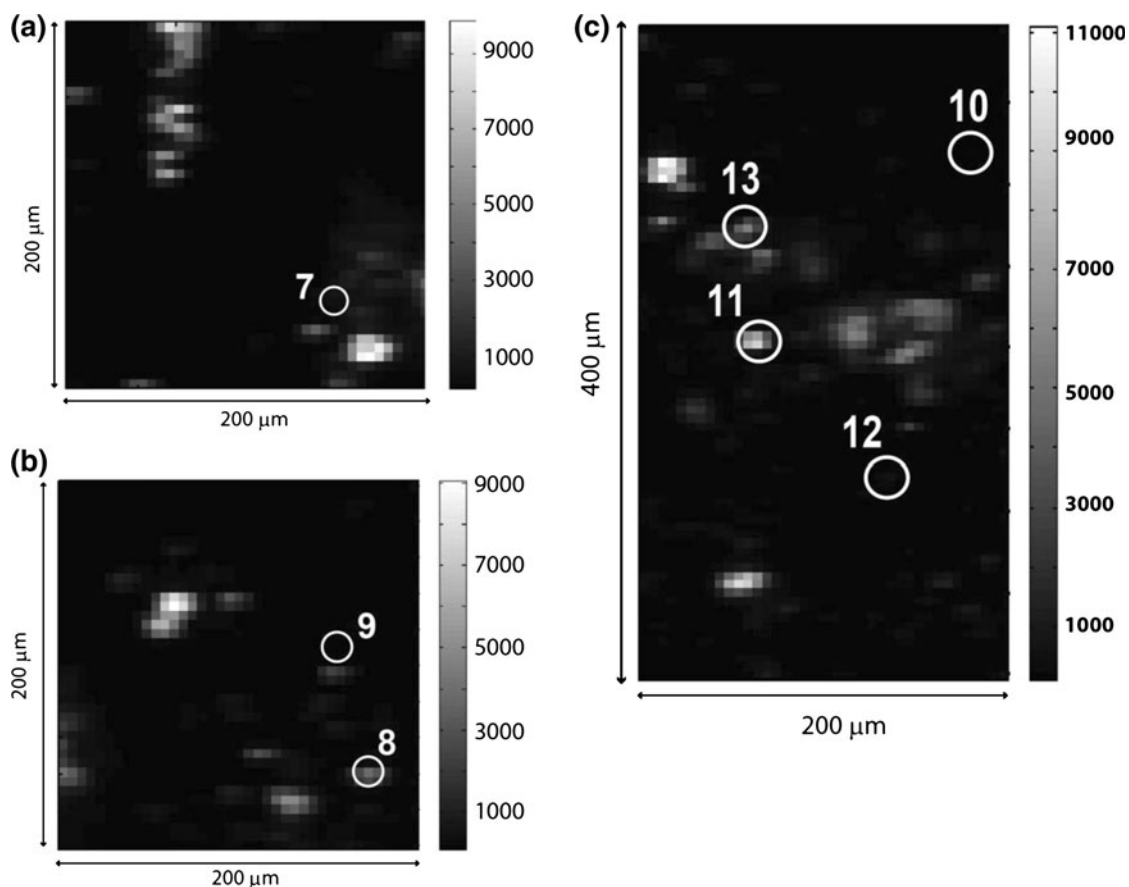
Sample	$N_{\text{Se-Se1}}$	$R_{\text{Se-Se1}} (\text{\AA})$	$DW (\text{\AA}^2)$	$N_{\text{Se-Se2}}$	$R_{\text{Se-Se2}} (\text{\AA})$	$DW (\text{\AA}^2)$	$\Delta E_0 (\text{eV})$	R-factor ^a
Se ⁰ (grey) reference	2.1 ± 0.2	2.39 ± 0.01	0.003	2.6 ± 0.3	3.41 ± 0.01	0.005 ^b	6.1 ± 1.2	0.01
Se sorbed onto pyrite	3.2 ± 0.3	2.38 ± 0.01	0.004	1.3 ± 0.4	3.41 ± 0.02	0.005 ^b	7.2 ± 1.3	0.01

^a R-factor from Artemis code^b Fixed parameter

Experiments with Se^{VI}

The μ -XRF maps recorded on pyrite chips treated with Na₂Se^{VI}O₄ solution during 2 months at 80 °C revealed a highly heterogeneous Se uptake, with distribution patterns very similar to those observed for the Se^{IV}-treated pyrites. In each studied sample in contact with Se^{VI} bearing solutions highly concentrated Se clusters of a few μm size were detected surrounded by areas with lower and uniform Se concentration, yielding μ -XRF maps (Fig. 6) very similar to those obtained after interaction with Na₂Se^{IV}O₃ solution.

Figure 7 shows the Se K-edge XANES of a pyrite powder (Se-23) and the μ -XANES of selected spots from two pyrite chips (Se-10, Se-11) treated with Na₂Se^{VI}O₄ (\pm NaHS) solutions, compared to the XANES spectra of Se reference compounds. In contrast to the measurements carried out with Se^{IV}-treated samples, photoreduction was frequently observed during the XANES data collection, but only for the sample treated with NaHS-bearing solution (Se-11). In such cases, each successive scan showed an increasing component of Se^{red} with simultaneously decreasing contribution of Se^{VI}. In contrast, the spectra from the pyrite sample treated with NaHS-free

**Fig. 6** Micro-XRF map of the Se-K α signal from polished pyrite chips (a, b: sample Se-10; c: sample Se-11) reacted during 2 months with Na₂SeO₄ solution under anoxic conditions, showing the location

of μ -XANES collection spots (labeled 7–13). The beam diameter was $\sim 2 \mu\text{m}$ and the pixel size $5 \mu\text{m} \times 5 \mu\text{m}$. The intensity scale gives the detector response in counts per seconds

Se^{VI}-solution (Se-10) remained stable. In Fig. 7, we included for sample Se-11 only μ -XANES spectra that remained unaffected with increasing collection time. We suspect that the observed photoreduction may be due to the presence of hydrogen sulfide adsorbed on the pyrite surface, owing to the mild washing procedure adopted. Under the X-ray beam, this residual hydrogen sulfide may have acted as a reductant of Se^{VI} sorbed on the pyrite surface.

An inspection of Fig. 7 shows the coexistence of both oxidized and reduced Se species at the same location, similarly to the Se^{IV}-treated pyrites. Nevertheless, the majority of the recorded μ -XANES spectra, including a number of them not included in Fig. 7, is dominated by Se^{red}. Only a few of the collected μ -XANES spectra show the presence of Se^{IV} (spot 7) or Se^{VI} (spot 10 and spot 13).

The bulk XANES spectrum recorded on Huanzala pyrite powder (Se-23) shows a quite different picture, clearly indicating a predominance of both oxidized species Se^{IV} and Se^{VI} with only a minor component of reduced Se species. This spectrum suggests that Se uptake proceeded via sorption of Se^{VI}, followed by subsequent reduction to Se^{IV} and finally to Se^{red}. After 2 months ageing, only a small proportion of the Se was reduced. In summary, our data indicate that at 80 °C Se^{VI} can also be reduced in the presence of pyrite, although the rate of reduction may have been slower than for the Se^{IV}-treated pyrites.

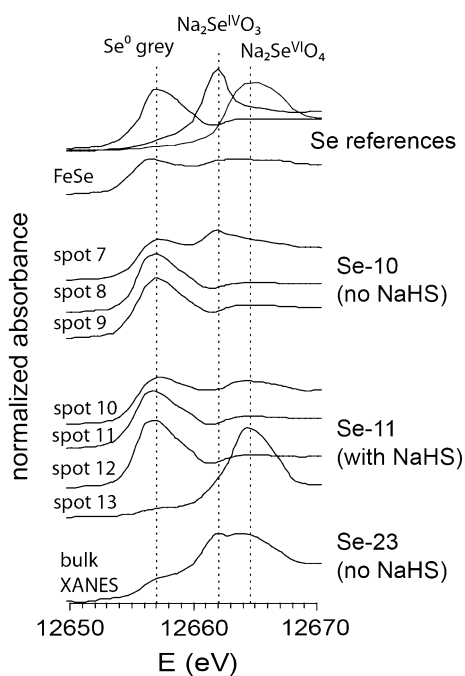


Fig. 7 Se K-edge XANES of a pyrite powder (Se-23) and μ -XANES of selected spots from pyrite chips (Se-10, Se-11) treated with Na₂Se^{VI}O₄ solutions, compared to the XANES spectra of reference compounds for Se^{-II}, Se⁰, Se^{IV} and Se^{VI}

Discussion

Thermodynamic stability

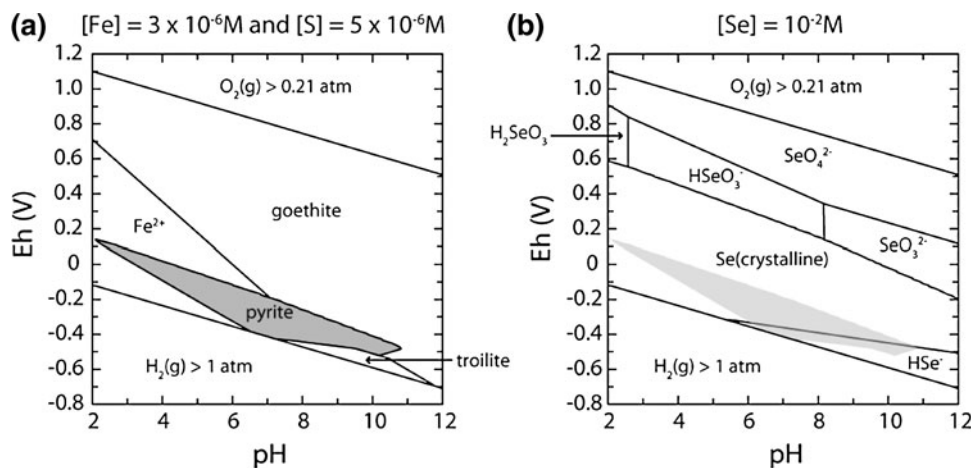
Pourbaix diagrams for the S–Fe–O–H system and for the Se–O–H systems were computed using the software PhreePlot (beta version, <http://www.phreeplot.org/>), which is based on the geochemical speciation code PHREEQC (Fig. 8). Total S and Fe concentrations were fixed at values approaching those calculated or measured in our systems ([S]_{tot} = 5 μ M and [Se]_{tot} = 0.01 M). Underlying thermodynamic data were taken from the Nagra-PSI 01/01 database [15] except for selenium, for which the data selected in Olin et al. [16] were used. The diagrams were calculated at 25 °C (1 bar) because many of the required data cannot be consistently extrapolated to the temperatures of the batch experiments (T = 80 °C).

The diagrams show that, provided no Fe–Se solid phases form (e.g. FeSe, FeSe₂ or a solid solution with pyrite), the expected coexisting stable phases under reducing conditions are pyrite and Se⁰. Although this result is consistent with the observations, oxidation of pyrite in aqueous media is frequent even in laboratory tests performed at trace levels of O₂ [17]. For instance, a moderate drop in pH was observed in pyrite-iodide experiments carried out under similar conditions [18]. Our experiments with pyrite powders (Se-21, Se-23) also suggest oxidation of pyrite in the glove-box atmosphere. On the other hand, due to the high pyrite:solution ratio (4–16 g/l), only a very limited fraction of the exposed mineral surface could have been affected by the oxidation process. The large excess of pyrite ensures that the Eh–pH conditions at the mineral surface remain within the grey-shaded field in Fig. 8b. Thus, in spite of thermodynamic disequilibrium between the (O₂-bearing) N₂ gas in the glove box and solid phase, the Eh–pH conditions of the solution in contact with the mineral surface are buffered by pyrite. Dissolved oxygen coming in contact with pyrite will lead to oxidation, but since the amount of oxygen available is small compared to that of pyrite, only a minor fraction of the mineral surface will be affected. Under such conditions, and provided that no solid solution between Fe selenide and Fe sulfide forms, the predicted thermodynamic stable Se species in our system is metallic selenium, which is consistent with our observation.

Selenium sorption

Although the observed surface precipitation of sparingly soluble reduced Se is in agreement with thermodynamic predictions, the explanation of the mechanisms leading to surface adsorption of the anionic dissolved Se^{IV,VI} species, which is a necessary intermediate step required for the formation of Se^{red} on the pyrite surface, is not trivial, since

Fig. 8 Eh–pH diagrams at 25 °C for **a** the S–Fe–O–H system with total Fe = 10^{-6} M and total S = 5×10^{-6} M, **b** Se–O–H with total Se = 10^{-2} M. On **b** the stability field of pyrite from **a** is superimposed as *shaded area*



the pyrite surface is negatively charged at pH ~ 8 . Our μ -XANES results showed however without any doubt that in some cases Se^{IV} and/or Se^{VI} coexist, at least transiently, with Se^{red} at the same location. We can exclude artifacts due to incomplete washing of the solution, since μ -XRF maps of pyrite immersed shortly (< 1 min) in 0.01 M Na_2SeO_3 solution and then briefly washed with ethanol in the same way as the samples, showed only background Se $K\alpha$ signals. Thus, the presence of $\text{Se}^{\text{IV,VI}}$ on the pyrite surface indicates that sorption of these species indeed occurred. Similar results were obtained in batch experiments at lower concentrations (10^{-6} M total Se), by Bruggeman et al. [5] who studied $^{75}\text{Se}^{\text{IV}}$ radiotracer uptake by a natural pyrite powder in a synthetic clay water. After 7 days of equilibration, R_d values varying from ~ 0.1 to ~ 10 m^3/kg were determined. After 2 months ageing, bulk EXAFS on powder samples showed that the Se sorbed on the pyrite from a Se^{IV} solution was completely reduced to the metallic state, probably in the form of trigonal (grey) Se^0 .

In analogy with the results obtained on iodide uptake on pyrite [18], one could assume that the cluster-like distribution of Se in the μ -XRF maps reflects adsorption of the dissolved SeO_3^{2-} and SeO_4^{2-} on the oxidized parts of the pyrite surface. Such regions are probably similar to ferric oxy-hydroxides ($\text{pH}_{\text{PZC}} > 8$, see [19]) and thus positively charged at pH ~ 8 , offering suitable surface sites for outer-sphere complexation of oxy-anions. According to this model, Se uptake would take place only on minor patches of the exposed pyrite that have undergone surface oxidation, due to the trace amounts of O_2 present in the solution prepared under quasi-anoxic conditions. The negatively charged part of the surface would remain free of sorbed Se. However, this interpretation, although appealing, can hardly explain the considerable amounts of Se (up to 16 % of the initially dissolved Se) sorbed on the pyrite surface. Indeed, a simple outer or inner-sphere adsorption

process can be ruled out for the high-intensity spots, since the very high Se- $K\alpha$ counting rates and the anticorrelation with the Fe- $K\alpha$ signal (maps not shown) indicate formation of surface precipitates of reduced Se (probably Se^0 -grey, as indicated by the bulk EXAFS data) rather than of a monolayer of adsorbed Se. Moreover, the diffuse low intensity regions between such spots (black to dark grey in Figs. 1, 2, 6) are not Se-free, as shown by detector counting rates of up to $\sim 1,000$ cps. Thus, in spite of the very heterogeneous distribution, Se appears to have been taken up over the entire pyrite surface. The nature of Se in the diffuse low-intensity region is not clear. It cannot be excluded, that $\text{Fe}(\text{S,Se})_2$ solid solution was formed there.

In spite of the aforementioned unresolved issues, our data indicate without any doubt that also selenate (not only selenite) was sorbed and then progressively reduced. This process apparently proceeded via an intermediate stage as Se^{IV} , as indicated by the bulk XANES of the Se-23 sample, where at least three redox states (Se^{red} , Se^{IV} and Se^{VI}) were found to coexist. Reduction of selenate has been observed in experiments carried out in contact with FeS under quasi-anoxic conditions [20], green rust [21] and recently a pyrite-greyite nanocomposite [14]. The evidence emerging from the present and the cited studies is that both selenite and selenate are readily reduced even at low temperature in the presence of Fe^{II} -sulfides and other Fe^{II} -bearing minerals.

In the view of the available knowledge on the geochemical cycle of Se, these results are not surprising. As pointed out by Stanton [22] the redox behavior of selenium in geochemical systems is quite different from that of sulfur, in spite of the reciprocal affinity of these two elements. Whereas dissolved sulfate is ubiquitous in near-surface environments (e.g. ocean water), only trace amounts of dissolved selenium are usually found in natural waters [23], indicating that Se^{VI} and Se^{IV} are more easily reduced to insoluble solids than the corresponding sulfur

species. Indeed, most of the natural Se in the earth crust is found as minor component in sulfides and only a few selenite/selenate minerals exist, which are stable only under exceptional conditions [24]. This fact is also reflected in the higher standard reduction potentials of Se redox couples compared to the corresponding S couples [25].

The behavior of Se in the geochemical cycle has important consequences on the fate of the long-lived radionuclide ^{79}Se when released to aqueous solutions from corroding radioactive waste. In most safety analyses of radioactive waste disposal sites ^{79}Se is considered to be transported in aqueous media as soluble selenite/selenate, leading to a high contribution of this nuclide to the calculated radiological dose [2]. This assumption is in line with the fact that ^{79}Se in radioactive waste is mainly present in oxidized form. For instance, Se is present as 100 % Se^{IV} in the simulated vitrified waste denoted MW [26] (Fig. 9) and will therefore be released as selenite upon aqueous corrosion. However, in the light of the present results, the global-scale geochemical evidence and considering the highly reducing conditions that will prevail in the repository environment (e.g. buffered by steel/copper canisters), there is little doubt that ^{79}Se will be readily reduced to insoluble Se^{red} forms before it can reach the biosphere. This is also indicated by kinetic sorption tests of Se^{IV} on clay rocks, which clearly show higher R_d -values under anaerobic conditions compared to similar tests conducted in air [27]. Therefore, the assumption made in safety analyses that ^{79}Se will migrate as soluble oxidized nuclide seems to be exceedingly conservative.

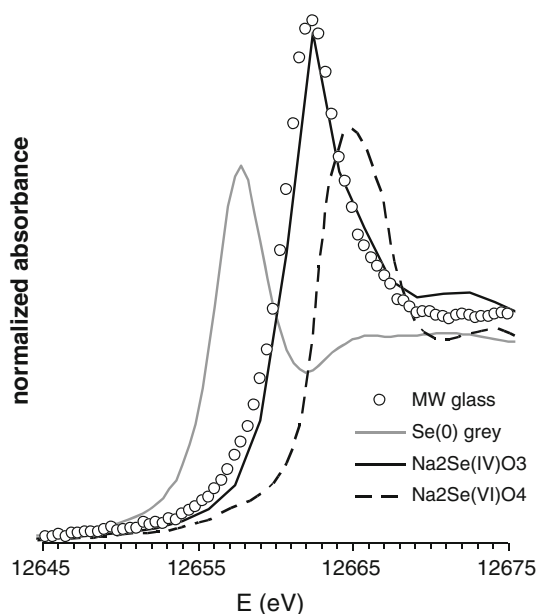


Fig. 9 XANES spectrum of a vitrified radioactive waste simulate (MW), compared to the spectra of reference compounds for Se^0 , Se^{IV} and Se^{VI}

Conclusion

Batch experiments at $\text{pH} \sim 8$ and 80°C showed that dissolved selenite and selenate are sorbed and subsequently reduced to insoluble forms (at least partly as Se^0) on the surface of natural pyrite in the time scale of weeks to months. The Se sorbed on the pyrite is distributed heterogeneously on the mineral surface and forms micrometer-sized clustered surface precipitates, surrounded by homogeneous regions with lower, but still elevated Se-concentration. In most cases, μ -XANES data indicate reduction to Se^0 , although further reduction to $\text{Se}^{-\text{I,II}}$ cannot always be excluded. Evidence was found for locations where reduced Se coexists with Se^{IV} and/or Se^{VI} . In a few cases, contributions from all three species were found, suggesting the stepwise reduction $\text{Se}^{\text{VI}} \geq \text{Se}^{\text{IV}} \geq \text{reduced Se}$. In two cases, bulk EXAFS analyses of pyrite powder reacted with $\text{Na}_2\text{Se}^{\text{IV}}\text{O}_3$ solution allowed us to identify ordered Se^0 grey phase as the end-product of the reduction process (after 8 months ageing time).

Thermodynamic models, though possibly incomplete because of the lack of Fe–Se phases, were in good agreement with spectroscopic results. According to the calculated Pourbaix diagrams $\text{Se}^{\text{IV,VI}}$ should indeed be reduced to crystalline, sparingly soluble Se^0 at equilibrium with pyrite. Owing to reductive sorption on pyrite, the migration of long-lived ^{79}Se released as oxidized oxo-anion (selenite/selenate) from vitrified waste or spent fuel may be significantly retarded under repository conditions.

Acknowledgments The beamline scientists at the Micro-XAS beamline at the Swiss Light Source (Dr. D. Grolimund, Dr. C. Borca) and at the DUBBLE beamline at ESRF (Dr. S. Nikitenko) are warmly thanked for the experimental assistance during the synchrotron-based investigations. Many thanks are due to B. Meyer from the Micro-XAS team for developing a sample holder allowing anoxic measurements. We are also indebted to Dr. Rainer Dähn and Dr. Marika Vespa for assistance during the measurement campaigns and suggestions during manuscript preparation. The authors wish also to thank Dr. Urs Eggenberger, Verena Jakob, Christine Lemp, and Dr. Christoph Wanner from the University of Bern, for sample preparation, BET measurements and for XRD assistance. This work was partially funded by the Helmholtz Virtual Institute of “Advanced Solid-Aqueous Radiogeochemistry”.

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