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## Sensitivity of an ecosystem model to hydrology and temperature

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**Abstract** We tested the sensitivity of a dynamic ecosystem model (LPJ-GUESS) to the representation of soil moisture and soil temperature and to uncertainties in the prediction of precipitation and air temperature. We linked the ecosystem model with an advanced hydrological model (JULES) and used its soil moisture and soil temperature as input into the ecosystem model. We analysed these sensitivities along a latitudinal gradient in northern Russia. Differences in soil temperature and soil moisture had only little influence on the vegetation carbon fluxes, whereas the soil carbon fluxes were very sensitive to the JULES soil estimations. The sensitivity changed with latitude, showing stronger influence in the more northern grid cell. The sensitivity of modelled responses of both soil carbon fluxes and vegetation carbon fluxes to uncertainties in soil temperature were high, as both soil and vegetation carbon fluxes were strongly impacted. In contrast, uncertainties in the estimation of the amount of precipitation had little influence on the soil or vegetation carbon fluxes. The high sensitivity of soil respiration to soil temperature and moisture suggests that we should strive for a better understanding and representation of soil processes in ecosystem models to improve the reliability of predictions of future ecosystem changes.

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## 1 Introduction

The Arctic and Sub-Arctic regions have experienced substantial regional warming in recent decades, with an average temperature increase of about 2 to 3°C since the 1950s (ACIA 2005). Climate projections for the next 100 years suggest a continuation of this warming trend, with the largest changes expected in the winter months (IPCC 2001). The Arctic Climate Impact Assessment (ACIA 2005) predicted an increase in temperature of 1.2°C for the 2011–2030 period, 2.5°C for 2041–2060 and 3.7°C for 2071–2090 and a mean increase in precipitation by 4.3, 7.9 and 12.3% for the respective periods.

Warming as such might be favourable for most Arctic plant species, but the responses to warming are critically controlled by indirect effects such as the availability of moisture and nutrients and competition from invading plants. Warming experiments show a general increase in productivity for most arctic ecosystems (Rustad et al. 2001; Dormann and Woodin 2002; van Wijk et al. 2003; Walker et al. 1998), whereas increasing summer precipitation produced few responses in the Arctic environment, except for mosses (Dormann and Woodin 2002; Sonesson et al. 2002). Still, soil moisture has an important influence on patterns of tundra plant communities and tundra ecosystem processes (Walker 2000).

Although arctic ecosystems have a low primary productivity, they tend to accumulate organic matter because decomposition and mineralization processes are limited more by environmental factors than is the fixation of carbon in photosynthesis (Jonasson et al. 2001). These factors include soil temperature (Lloyd and Taylor 1994; Rustad et al. 2001), soil moisture (McGuire et al. 2000b; Clein et al. 2000), changes in NPP (net primary production), i.e. substrate availability (Raich and Potter 1995), population dynamics of plants (Raich and Schlesinger 1992) and vegetation substrate quality (Raich and Schlesinger 1992). As a result of this mismatch between rates of carbon capture in photosynthesis and decomposition rate, high latitude areas contain about 40% of the world's soil carbon (McGuire et al. 1995) and therefore any change in rates of carbon fluxes will have a direct impact on the global climate system (ACIA 2005). In areas of permafrost, soils in the active layer can be waterlogged; in some areas, when the permafrost thaws, drainage is impeded and the soil remains wet. In other areas associated with good drainage, disappearance of permafrost will lead to aridification, which results in a loss of tundra ponds and soil drying (Callaghan et al. 2005).

The dynamic ecosystem model LPJ-GUESS (Smith et al. 2001) concentrates on processes at the plant level and has proven capable of predicting vegetation distribution, net primary production and net ecosystem exchange in many different ecosystems (Smith et al. 2001; Hickler 2004; Morales 2006). It has been also successfully used for predictions of forest productivity and biomass in Sweden (Koca 2006) and the Barents Region (Wolf et al. 2008).

LPJ-GUESS uses a very simple representation of soil processes and hydrology. To make LPJ-GUESS more applicable to the Arctic, we link it with the complex hydrological model JULES, which simulates the freezing and melting of soil water, and takes account of the dependence of soil thermal characteristics (Cox et al. 1999). Both models were driven with the same regional climate model (REMO, Jacob 2001). We compare the responses of soil and vegetation carbon fluxes, when JULES soil temperature, JULES soil moisture or both were implemented in the ecosystem model with the control run using the simple representation of LPJ-GUESS only. We ran the simulation in a sequential way in order to be able to isolate the impact of the two soil properties on the carbon fluxes. We focused on a transect in northern Russia, because the ecosystem model has been used in this region (Wolf et al. 2008) and strong responses of vegetation were indicated for a future warming

climate there. We analysed the changes along a latitudinal gradient, to cover the response of different ecosystems ranging from boreal needle leaved forest to tundra.

We also tested whether the use of improved estimations of soil properties (temperature and moisture) influence the output of the ecosystem model LPJ-GUESS. Secondly, we investigated whether, and to what extent, uncertainties in precipitation and air temperature predictions influence the estimations of ecosystem carbon fluxes.

## 2 Materials and methods

### 2.1 Ecosystem model

We used the LPJ-GUESS ecosystem modelling framework (Smith et al. 2001), which combines the mechanistic representation of plant physiological and biogeochemical processes with detailed representation of vegetation dynamic processes (Sitch et al. 2003; Smith et al. 2001). The growth of cohorts of PFTs (plant functional types) is simulated in a number of replicated patches. In the standard version, PFTs are either trees or herbaceous vegetation, but in this study we included shrubs (deciduous and needle leaved shrubs) as they are characteristic of the Arctic and Sub-Arctic (for a detailed description see Wolf et al. 2008).

Photosynthesis, water uptake and plant respiration are modelled at daily time steps. At a yearly time step, the resulting annual NPP is allocated to reproduction and growth as determined by a set of prescribed allometric relationships. The yearly leaf and root turnover as well as plant mortality is transferred to the litter pool. Litter and soil carbon dynamics follow first-order kinetics and are sensitive to soil temperature and soil water (modified Arrhenius equation, Lloyd and Taylor 1994). A full description of LPJ-GUESS is found in Smith et al. (2001). Further details of the physiological, biophysical and biogeochemical components of the model are given by Sitch et al. (2003).

The treatment of soil thermal and moisture regime in the dynamic vegetation model LPJ-GUESS is based on the simple “bucket” model. It was recently improved as described in Gerten et al. (2004). The “capacity” approach is employed assuming that all water above the field capacity is removed immediately from the soil. Two layers of 0.5 and 1 m depth are specified. Water movement between the layers is described by the empirically defined percolation rate. Soil temperature is calculated based on the Fourier analytical solution of the heat diffusion equation. Thermal conductivity is a function of water content, but no effect of ice below the active layer or snow on the top of the soil is taken into consideration. Evapotranspiration is calculated based on the Priestley–Taylor approach with an underlying assumption that atmospheric specific humidity does not vary.

The model has been shown in earlier studies to predict correctly the dominant PFT composition in different sites (Badeck et al. 2001; Smith et al. 2001; Hickler et al. 2004; Morales et al. 2005) and for Northern Europe (Koca 2006; Wolf et al. 2008). The closely related LPJ-DGVM has been subject to extensive validation of variation in ecosystem carbon balance (Lucht et al. 2002; Sitch et al. 2003).

### 2.2 Climate and CO<sub>2</sub> data

The model is driven by monthly averages of temperature, precipitation, percentage of sunshine hours, annual atmospheric CO<sub>2</sub>-concentration and soil-type derived from the FAO global soil data set (FAO 1991).

We used a 1,000-year spin-up period to allow the vegetation, soil and litter pools to reach equilibrium with the long-term climate. For this spin-up period, we used the global CRU-data set (Mitchell and Jones 2005, [http://ipcc-ddc.cru.uea.ac.uk/obs/cru\\_climatologies.html](http://ipcc-ddc.cru.uea.ac.uk/obs/cru_climatologies.html)). The time series from 1901 to 1930 was used repeatedly to provide the climate input for the spin-up. For the historical period 1901–1960 we also used data from the CRU-data set. For the period 1961–2000, we used the data from the regional climate model REMO (Jacob 2001). The driving fields from the regional climate model REMO for the vegetation, which was developed at the Max-Planck-Institute for Meteorology (Jacob 2001), were taken from an already existing climate simulation using REMO5.3 at  $0.5^\circ$  ( $\approx 55$  km) horizontal resolution (Keup-Thiel et al. 2006). REMO was driven by the coupled atmosphere ocean general circulation model (AOGCM) ECHAM4/OPYC3 (Roeckner et al. 1996; Oberhuber 1993) data at T42 resolution. The AOGCM data describe the possible climate of 1961 until 2099 for current and possible future (IPCC B2) greenhouse gas concentrations. The Barents Region was laid in the centre of the regional climate model. We used the grid cell coordinates from the REMO model. As driving data for the spin-up period and 1901–1960 period, we used the CRU-point with the closest distance (in longitude–latitude degrees) to the REMO-grid point.

Global atmospheric  $\text{CO}_2$  concentrations for 1901–1960 were derived from a combination of ice-core measurements (1901–1950) and atmospheric observations (1950–1960; cf. Sitch et al. 2003). For the spinup period, the value 296 ppm (concentration in 1901) was used. For the period 1961–2000, we used the same  $\text{CO}_2$  concentration as the REMO model (IPCC B2).

### 2.3 JULES

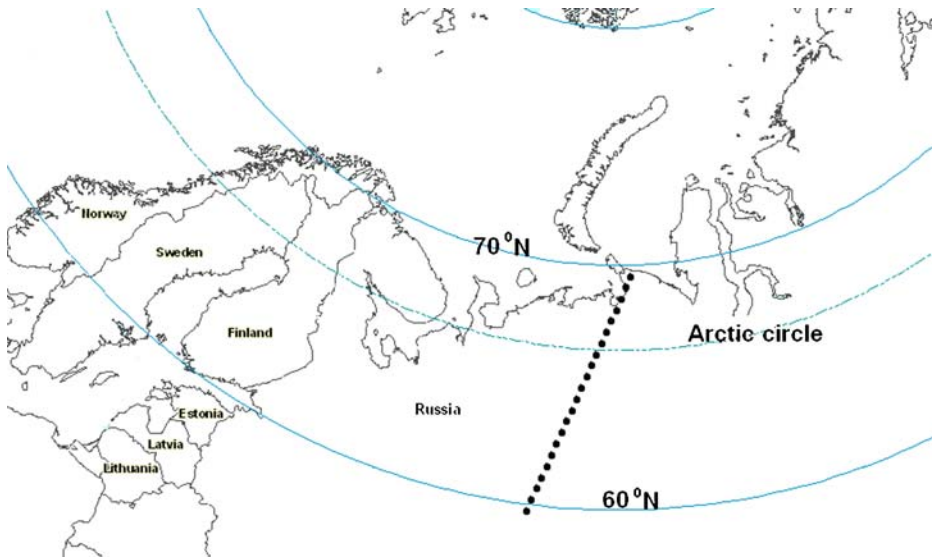
The JULES model has inherited its basic properties from the Hadley Centre climate model land surface scheme MOSES. It incorporates a fully mechanistic representation of the heat transfer and water flow in soil developed for large-scale applications (Cox et al. 1999). The soil hydrology module is based on the numerical solution of the Richards' equation. The soil temperature is calculated using the discretised heat diffusion equation. Both equations are solved for 4 soil layers with thicknesses from the surface of 0.1, 0.25, 0.65 and 2 m, respectively. Water phase changes are considered and soil thermal characteristics are functions of the water content (liquid and ice). The model includes a description of the insulating effect of snow on the soil surface. Evapotranspiration components from the canopy, vegetation and the bare soil are calculated as a function of the atmospheric specific humidity.

### 2.4 The transect in northern Russia

We studied a latitudinal transect in northern Russia which consists of 22 grid points and stretches from  $59.7^\circ\text{N}$  to  $69.5^\circ\text{N}$  (Fig. 1). It is situated in the Komi Republic and the Archangelsk Administrative District. This gradient covers the response of different ecosystems ranging from boreal needle leaved forest to tundra. The ecosystem model has been used in this region and shows that strong responses of vegetation, e.g. a northward shift of the forest biomes and an increase in shrub abundance, are expected for a future warmer climate (Wolf et al. 2008).

### 2.5 Iterative coupling approach

JULES and LPJ-GUESS are driven by the same climate input from the REMO model. For the control run, only the REMO data were used. We then ran three simulations, including



**Fig. 1** Barents Region and situation of the grid points

the output of JULES as additional input into LPJ-GUESS. Firstly, we used the JULES output on soil temperature (upper 10 cm) and controlled the GUESS soil respiration with this temperature estimate, instead of the intrinsic estimate of soil temperature. Additionally, photosynthesis was first started when the soil temperature increased above freezing ( $0^{\circ}\text{C}$ ), as water supply was assumed to be unavailable below this temperature. Secondly, we used the JULES estimations (upper four soil layers: 0.1, 0.25, 0.65 and 1 m thick) of unfrozen plant-available water as direct input into the LPJ-GUESS. We used thickness-weighted averages to scale to the two soil layers of GUESS (0.5 and 1.5 m thick). Finally, we ran a simulation that included both soil temperature and available soil water. The coupling was done for the 1981–2000 period, and results were analysed for this period. We ran four different simulations: (1) with JULES soil moisture; (2) with JULES soil temperature; (3) with JULES soil moisture and temperature and (4) the control run i.e. using LPJ-GUESS without estimations from JULES. The differences in soil temperature and soil moisture estimations between the two models are shown in Table 1.

## 2.6 Sensitivity of results to climate

There is some uncertainty in the climate model outputs of rainfall and air temperature. To test the land surface models' sensitivity to these climate uncertainties, the precipitation and air temperature were changed by fixed amounts to assess the effect that had on the modelled carbon fluxes.

Changes in precipitation were achieved by changing the input data to be 20 and 50% less or 20 and 50% more precipitation than REMO data (Table 2). For temperature sensitivity, we increased or decreased the temperature predicted by REMO by 1 and  $2^{\circ}\text{C}$ , which resulted in an additional four runs (Table 2). All the runs were compared to a control run with the unchanged REMO predictions on air temperature and precipitation. Overall, we ran nine simulations, one control and four sensitivity runs for air temperature and four sensitivity runs for precipitation.

**Table 1** Average soil temperature (upper layer, 0.5 m depth) and soil moisture (upper and lower layer, 1 m depth) as predicted by JULES and LPJ-GUESS

Month	Average soil temperature (°C)		Soil moisture (%)			
	JULES	LPJ-GUESS	JULES		LPJ-GUESS	
			upper layer	lower layer	upper layer	lower layer
Jan	-17.7	-16.3	0.01	0.00	0.57	0.73
Feb	-17.5	-16.6	0.01	0.00	0.57	0.73
Mar	-10.7	-11.8	0.01	0.00	0.54	0.73
Apr	-5.7	-6.5	0.01	0.01	0.46	0.73
May	0.2	-0.3	0.07	0.09	0.59	0.77
Jun	9.1	6.9	0.12	0.14	0.74	0.90
Jul	14.0	11.3	0.11	0.12	0.48	0.84
Aug	11.4	10.6	0.10	0.12	0.40	0.70
Sep	5.6	6.4	0.11	0.14	0.47	0.67
Oct	-1.0	0.1	0.11	0.10	0.57	0.71
Nov	-9.8	-7.8	0.06	0.03	0.57	0.72
Dec	-14.8	-13.2	0.04	0.03	0.57	0.72

## 2.7 Statistical analyses

The SPSS 13.0 statistical package was used for all analyses. Non-parametric tests were used in all cases, as the assumption of normal distribution might not be fulfilled. The tests for the effects of the different simulation runs were made as paired tests, as absolute values differed between grid points, e.g. along the latitudinal gradient. If the same data set was tested several times we corrected for multiple testing with the Dunn–Sidak method (Sokal and Rohlf 1995). The presented  $p$  values are the corrected values in such cases.

## 3 Results

### 3.1 Soil water content

Including JULES predictions of the content of unfrozen soil water decreases the vegetation carbon fluxes (e.g. the carbon uptake by vegetation) compared to the control run ( $p < 0.001$ , Wilcoxon Signed Ranks Tests,  $N=22$ , Fig. 2a). However this decrease is small, on average only 4% compared to the control run, except for the most northern grid point where the decrease is 23% of the control run.

Including the soil water predictions of JULES leads to a strong decrease of the soil carbon fluxes from the soil to the atmosphere (Fig. 2b). The average decrease in soil carbon emissions is 40% compared to the control run, a difference which is statistically significant ( $p < 0.001$ , Wilcoxon Signed Ranks Tests,  $N=22$ ). The importance of the feedback increases with increasing latitude from 32% at 59.7°N to almost 50% decrease in soil carbon fluxes at 69.5°N ( $r = -0.63$ ,  $p < 0.05$ , Spearman's Rho,  $N=22$ ).

### 3.2 Soil temperature

For the simulation run which included the JULES estimations of soil temperature (upper 10 cm), the vegetation carbon uptake is significantly lower than for the control run

**Table 2** Average air temperature and precipitation along the latitudinal transect and the values used for the study on uncertainty in air temperature and precipitation

Latitude (°N)	Temperature (°C) Control	Precipitation (mm year <sup>-1</sup> ) Control
69.49	-7.1	495
69.03	-6.9	547
68.58	-6.2	527
68.12	-5.5	567
67.66	-5.4	568
67.2	-4.8	616
66.74	-4.4	635
66.27	-4.0	649
65.81	-3.4	667
65.34	-3.1	676
64.87	-2.8	704
64.4	-2.6	718
63.93	-2.3	697
63.46	-2.0	721
62.99	-1.8	713
62.52	-1.5	715
62.04	-1.1	700
61.57	-0.6	689
61.09	-0.5	704
60.62	-0.2	723
60.14	0.2	751
59.66	0.5	747

( $p < 0.001$ , Wilcoxon Signed Ranks Tests,  $N=22$ , Fig. 2a). The decrease in vegetation carbon uptake is around 8% of the control run, except for the two most northern sites, where carbon uptake by vegetation is ca. 20% lower than the control.

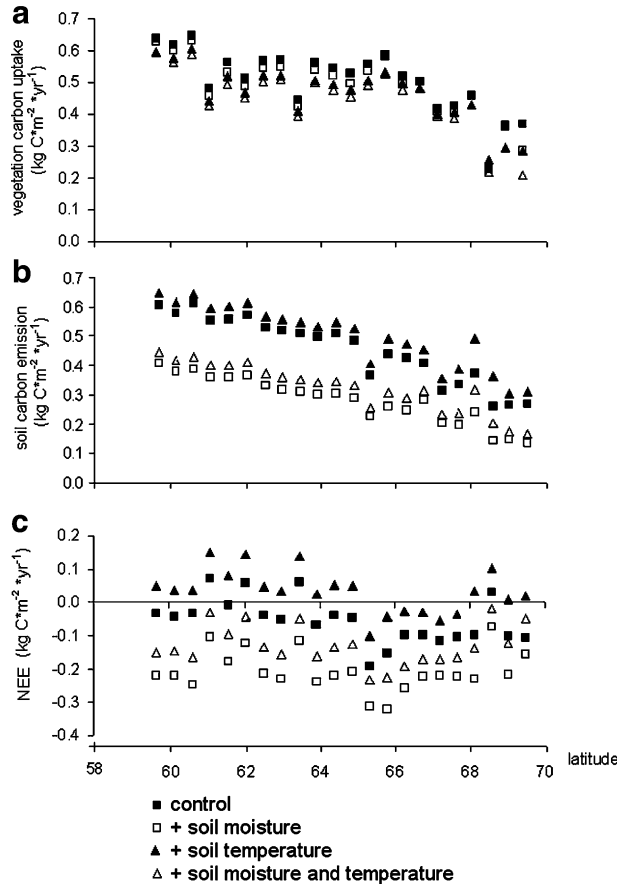
The soil carbon emissions are significantly influenced by including the soil temperature estimations from the JULES model ( $p < 0.001$ , Wilcoxon Signed Ranks Tests,  $N=22$ , Fig. 2b). There is an increase in carbon fluxes from soil to atmosphere of about 12% compared to the control run and the influence increases at higher latitudes ( $r=0.93$ ,  $p < 0.001$ , Spearman's Rho,  $N=22$ ).

### 3.3 Soil temperature and soil moisture

When both soil moisture and temperature estimates from the JULES model were used as input in the vegetation simulation, the carbon flux from atmosphere to vegetation is always smaller compared to the control run ( $p < 0.001$ , Wilcoxon Signed Ranks Tests,  $N=22$ , Fig. 2b). The changes are around 11% of the control run, except for the two most northern points, where the decrease in flux is around 20 and 40%.

The pattern is more complex for the carbon fluxes from soil to atmosphere, as the increase in flux due to the incorporation of JULES soil temperature estimates is overlaid with the decrease in soil carbon emissions due to the impact of soil moisture estimates from JULES (Figs. 2 and 3). The reduction in soil respiration is around 30% compared to the control run and there is no trend with latitude ( $p > 0.05$ ).

**Fig. 2** Impact of coupling JULES soil moisture and temperature to LPJ-GUESS on carbon dioxide fluxes. The control run uses the LPJ-GUESS estimations of soil moisture and temperature (*filled squares*). In the other runs JULES estimations of soil temperature (*filled triangles*) or soil moisture (*open squares*) or both (*open triangles*) were used as additional input in the LPJ-GUESS model. Each comparisons of treatment versus control is significant (for all  $p < 0.001$  Wilcoxon Signed Ranks Tests,  $N=22$ ). There are latitudinal trends in the response of soil carbon fluxes to changes in soil water content ( $r = -0.63$ ,  $p < 0.05$ , Spearman's Rho,  $N=22$ ) and soil temperature ( $r = 0.93$ ,  $p < 0.001$ , Spearman's Rho,  $N=22$ ), but no trend when soil moisture and soil temperature treatments were combined

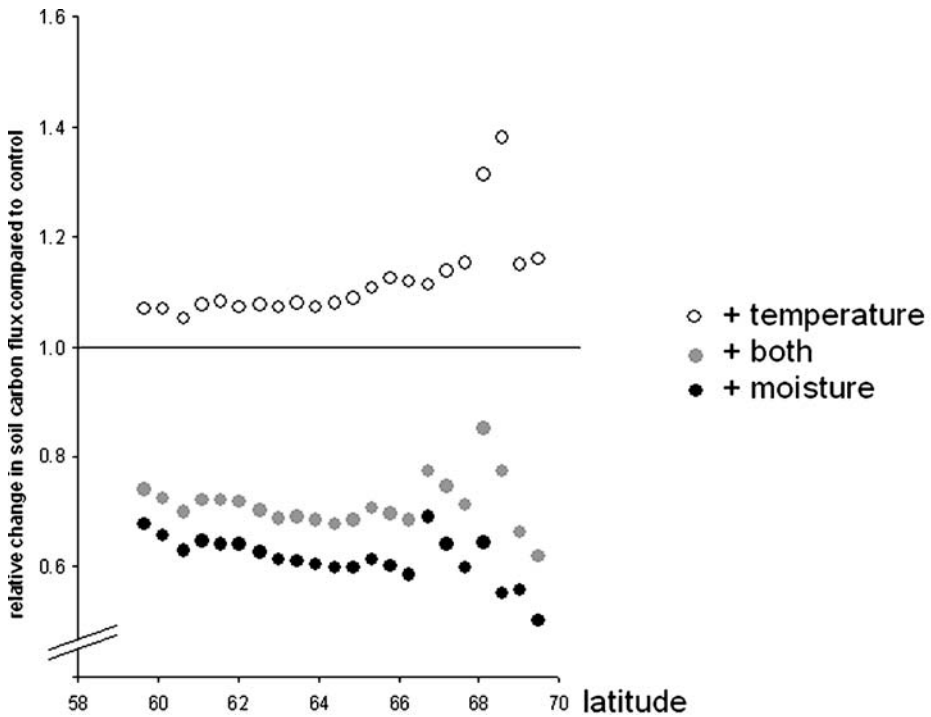


### 3.4 Air temperature

When we investigated the importance of uncertainties in temperature predictions, we increased or decreased the predicted temperature by 1 or 2°C. This results in a significant impact of on the vegetation carbon uptake ( $p = 0.003$ , Friedman Test,  $N=22$ ,  $df=4$ ). The differences are significant for the runs with a temperature lower than the predicted mean ( $T_{\text{REMO}} - 1^\circ\text{C}$ ,  $p < 0.01$ ,  $T_{\text{REMO}} - 2^\circ\text{C}$ ,  $p < 0.003$ , Dunn–Sidak correction for multiple testing,  $N=4$  tests), which results in a decrease in carbon uptake by vegetation. In contrast, if temperature was 1 or 2°C higher than the predicted mean, the increase in carbon uptake by vegetation is not significant ( $p > 0.05$ ). The effect of temperature is strongest in the more northern grid points (Fig. 4a) where any further increase in predicted temperature leads to an increase in carbon uptake by vegetation and a decrease in temperature leads to a decrease in uptake.

When we tested the effect of the assumption of a 1 or 2°C over and underestimation of air temperature, we found a significant effect on soil carbon emission ( $p < 0.0001$ , Friedman Test,  $N=22$ ,  $df=4$ ) where each temperature treatment is significantly different from the control run (Fig. 4b,  $p < 0.0003$  for each pair-wise comparison, considering correction for multiple testing, Dunn–Sidak,  $N=4$  tests). If the climate was increased by an additional 1 or 2°C, the soil carbon emissions increase (the effects were stronger in the 2°C temperature change compared to the 1°C). If the temperature was assumed to be 1 or 2°C lower than the





**Fig. 3** Changes in soil carbon emissions compared to the control run. Values larger than 1 indicate an increase in fluxes compared to the control, whereas values smaller than 1 indicate a decrease in soil carbon emissions. JULES estimations of soil temperature (*open circle*) or soil moisture (*black circle*) or both (*grey circle*) were used as additional input in the LPJ-GUESS model

predicted temperature, the soil carbon emission are lower than the control (Fig. 4b). The impact of uncertainties in the temperature is strongest in the higher latitudes (Fig. 4b).

### 3.5 Precipitation

To test the consequences of the uncertainties in precipitation predictions, we increased and decreased the predicted precipitation by 50%. This has only minor effects on the fluxes of carbon from the atmosphere to vegetation; less than 1% change in fluxes compared to the control run.

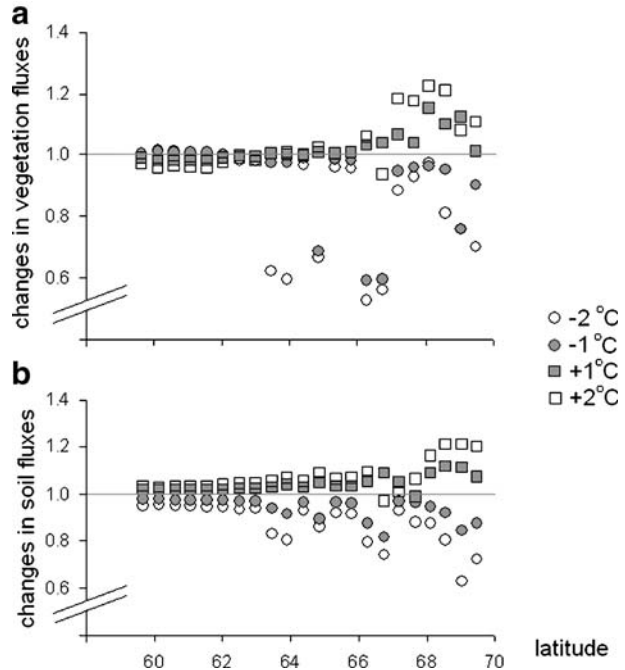
For the soil carbon emission, the effect of changes in the precipitation is small if the predicted precipitation is increased by 50%, which leads to an average 5% increase of soil carbon emissions. However, for a 50% decrease in precipitation, the average decrease of soil carbon emissions is around 11%. These results were significant ( $p < 0.0002$  for all tests, Dunn–Sidak correction for multiple testing,  $N=4$  tests).

## 4 Discussion

### 4.1 Soil carbon emissions

Soil carbon fluxes showed a clear response to the changes in assumptions about soil temperature. The soil carbon emission increased by ca. 12% when we used the JULES soil temperature estimates as input into the ecosystem model. Soil respiration is highly sensitive

**Fig. 4** Responses of vegetation carbon uptake and soil carbon emissions to uncertainties in the predictions of air temperature by the climate model REMO. We assumed an additional increase of temperature of 1 or 2°C (grey and open square), or that the temperature predictions were overestimated by 1 or 2°C (grey and open circles). For vegetation carbon fluxes, differences are significant for the runs with  $-1^{\circ}\text{C}$ ,  $p < 0.01$ , and  $-2^{\circ}\text{C}$ ,  $p < 0.003$ , but not for  $+1$  or  $+2^{\circ}\text{C}$ ,  $p > 0.05$  (Dunn-Sidak correction for multiple testing,  $N=4$  tests), for soil carbon fluxes, all temperature treatments are significant from the control run ( $p < 0.0001$ , Friedman Test,  $N=22$ ,  $df=4$ )



to soil temperature (e.g. Lloyd and Taylor 1994; Rustad et al. 2001). Therefore changes in the model formulation of soil temperature have a strong impact on the soil carbon emission. The importance of the model assumptions on soil temperature was also shown by McGuire et al. (2000a), who showed that keeping the soil temperature under snowpack at  $0^{\circ}\text{C}$  improved the model fit to observed data. However, the effect of soil moisture on soil carbon emission was even stronger, with a decrease in soil respiration by 40% when JULES soil moisture estimates were included in the LPJ-GUESS ecosystem model (Fig. 2c). This is in agreement with other modelling studies that show a tight coupling of soil respiration to soil moisture (McGuire et al. 2000b; Clein et al. 2000).

In this study, which covers a transect from boreal to arctic vegetation, we found that the importance of the two factors of soil temperature and moisture varies with latitude: soil temperature has a stronger correlation with latitude, showing a stronger impact in the northern grid points. However, soil moisture changes have a strong influence, that is greater than that of soil temperature, over the whole latitudinal range (Fig. 3). In this study, the importance of soil temperature and moisture was clearly shown. The impacts on soil carbon emissions could not be attributed to changes in NPP, which could be of high importance in some studies (Raich and Potter 1995), as vegetation carbon uptake was only little influenced. Neither were plant population dynamics important as suggested by Raich and Schlesinger (1992) as the time scale for such changes would be at least a few years, but substantial changes soil carbon emissions were observed already in the first year of coupling and remained important throughout the whole simulation period (1981–2000).

#### 4.2 Carbon fluxes between atmosphere and vegetation

In our study, the carbon fluxes from the vegetation showed little response to either soil temperature or soil moisture formulations. This is surprising, because the TEM 5.0 (Terrestrial

Ecosystem Model) produced a better fit of ecosystem carbon uptake to observed values when the effect of soil temperature on vegetation carbon uptake was considered (Zhuang et al. 2003) as the timing of the draw-down of atmospheric CO<sub>2</sub> at the start of the growing season was substantially improved. In our study, the limitation to photosynthesis in early spring, when air temperatures are already suitable for plant activity, but the soil might still be frozen and water unavailable, only reduces the annual photosynthesis by a little, usually less than 1% of total carbon uptake by vegetation. In addition, there is a small impact of soil temperature on root respiration that influences carbon uptake by vegetation. However, these responses were small and represent only a part of the total plant respiration. Only in the two most northern grid cells, was there a larger impact of soil temperature on carbon uptake by vegetation. In these grid cells, the growing season is very short, so even small changes in temperature regime might have a relatively larger impact than further south.

The impact of soil moisture on vegetation carbon fluxes did not show any changes along the considered gradient, except for the most northern grid cell. For instance, the drought stress which affected the growth of white spruce observed in Alaska (Barber et al. 2000) was not apparent in our study. This may be because the JULES model had a free-draining bottom boundary and no permafrost was modelled. This means that no over-saturation or paludification was considered, which could be important in Arctic ecosystems (Callaghan et al. 2005).

#### 4.3 Temperature and precipitation uncertainties

Ecosystems of northern latitudes are particularly sensitive to changes in air temperature (Lucht et al. 2002; ACIA 2005 and IPCC 2001), so uncertainties in the model predictions of air temperature should show the strongest consequences in the northern areas: this effect is shown in our tests (Fig. 4b). In contrast, precipitation increase or decrease showed little effect on vegetation and a small impact on soil carbon emissions. This is in agreement with field manipulation experiments, which showed only a small impact to additions of water on arctic vegetation (Press et al. 1998; Dormann and Woodin 2002). Our study also agrees with satellite observations showing that temperature and radiation limit plant productivity in high latitudes (Nemani et al. 2003) and modelling studies showing that responses in biogeochemical cycling and vegetation change at high latitudes are mainly sensitive to changes in air temperature and not precipitation (Lucht et al. 2002). The lack of response of vegetation to uncertainties in precipitation is also in line with the small response of vegetation to changes in soil moisture representation shown in our study.

Therefore, we can conclude that for this model version of LPJ-GUESS, uncertainties in precipitation of 50% will not greatly affect the predictions of carbon fluxes. However, uncertainties in air temperature predictions are likely to have a larger impact on the predictions of carbon uptake by vegetation and carbon emissions from soils in a future changing climate.

#### 4.4 Net ecosystem exchange

The changes in vegetation carbon uptake and soil carbon release influence net ecosystem carbon exchange (NEE, Fig. 2c) and hence predictions of whether the Arctic is a source or sink for carbon. NEE is the difference between vegetation carbon uptake and soil carbon release. In this study, both processes showed opposite but strong responses to uncertainties in the representation of soil moisture and soil temperature. Therefore, modelled NEE is also sensitive to the formulation of soil moisture and varied significantly between the different

simulations runs, from being a sink for CO<sub>2</sub> in the simulation with the JULES soil moisture estimations included in the ecosystem model, to being a possible source of carbon dioxide, if soil temperatures were used from JULES (Fig. 2c).

Long term changes in NEE are difficult to assess due to the high uncertainties related to the small fluxes. In addition to the uncertainties in model formulations, there are uncertainties related to the predictions in the precipitation and in air temperature which can alter the predictions about whether a given area is a source or sink of carbon. The importance of ecosystem respiration for NEE is in line with findings that respiration is highly important in carbon balance of European forests (Valentini et al. 2000).

In summary, NEE showed a strong response to climate signals; combined with the uncertainties in climate predictions and the uncertainties in model formulation, we conclude that we need further investigation and improved knowledge both from the experimental side, and also for the model formulations, in order to assess the net carbon fluxes of high latitude ecosystems.

#### 4.5 Additional factors

There are additional factors that influence the feedback of arctic systems, which are not explicitly considered in the LPJ-GUESS, such as methane emission (Christensen et al. 1999, 2003; Smith et al. 2004; Sitch et al. 2007), changes in canopy complexity from tundra to forest, which results in changes in ground heat fluxes, latent heat fluxes and sensible heat fluxes (Thompson et al. 2004; Beringer et al. 2005; Chapin et al. 2005), disturbance such as fire, extreme events (Marchand et al. 2005), erosion and permafrost degradation (Malmer et al. 2005). Another important factor is the possible significant increase in rate of net N mineralization of the upper organic soil horizon following soil warming (Rustad et al. 2001), which is not explicitly modelled in our study. Nitrogen is often considered to be a major limiting nutrient for plant production (Vitousek et al. 1997) and therefore an increase in the availability of inorganic N should increase plant carbon uptake (Callaghan et al. 2005). However, it is not certain that air warming will lead to soil warming because of increased interception of incoming radiation by an increased leaf area index (Callaghan et al. 2005) while Atkin and Cummins (1994) concluded that nutrient uptake was not increased greatly by the temperature treatments in experiments.

## 5 Conclusion

Soil moisture and soil temperature are important factors driving carbon fluxes, particularly soil carbon emissions. Therefore, ecosystem models should improve the formulation of soil characteristics to get a better representation of soil temperature and moisture. Furthermore, the dependence of soil carbon emissions on these factors needs to be reconsidered. The dependence of soil respiration on soil temperature uses the modified Arrhenius equation (Lloyd and Taylor 1994) and the use of two soil pools with different turnover rates have been assumed to be sufficient (Knorr et al. 2005). However, the relationship of soil processes to soil moisture need to be studied in more detail and incorporated into the ecosystem model.

Precipitation pattern might not have a strong influence on productivity of northern ecosystems because of high uncertainties. Even with an increase in air temperature, plant productivity might not become water-limited. However, soil carbon emissions are sensitive

to soil moisture and hence precipitation pattern and the ecosystem responses to these should be studied with improved ecosystem models.

It should be noted that even the most realistic of the soil models used in this study did not include permafrost and hence was unable to represent some of the important sub-surface processes of the Arctic region. The model does include a parameterisation of organic soils and a simple representation of the sub-grid scale patterns in hydrology which affect soil moisture (and hence soil temperature). A further study of how uncertainties in these representations affect the soil moisture and therefore the carbon fluxes from the soils would be useful. This study can be seen as a step towards more realistic analysis of carbon fluxes in the Arctic: Striving for a better understanding and representation of soil processes in ecosystem models will improve their reliability and predictions of future ecosystem changes and more accurate predictions of the future carbon sink source status of northern ecosystems which is critical for understanding feedbacks to global climate.

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