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DETERMINATION OF THE TURBULENT TEMPERATURE–HUMIDITY CORRELATION FROM SCINTILLOMETRIC MEASUREMENTS

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Abstract. We report on the investigation and successful application of the bichromatic correlation of optical and microwave signals for determining the area-averaged correlation of temperature–humidity fluctuations. The additional technical effort is marginal compared to the common ‘two-wavelength method’, which has (in contrast) the restriction that only two of the three relevant meteorological structure parameters can be deduced. Therefore, in the past, it was often assumed that the turbulent humidity and temperature fluctuations are perfectly positively or negatively correlated. However, as shown in this study, over non-homogeneous terrain when the flow conditions are not ideal, this assumption is questionable. The measurements were analysed statistically, and were compared to *in situ* measurements of the Bowen ratio Bo and the correlation of temperature–humidity fluctuations using eddy-covariance techniques. The latter is in good agreement to that derived by scintillometry. We found that the correlation is not ± 1 but as low as -0.6 for Bo smaller than -2 , and up to 0.8 for Bo larger than 1 .

Keywords: Area-averaged fluxes, Bichromatic correlation, Correlation of temperature–humidity fluctuations, Electromagnetic waves, Heterogeneous land surface, Microwave and optical scintillometer.

1. Introduction

Many applications in meteorology and the atmospheric sciences demand continuous measurements of turbulent surface heat fluxes that are representative of extended areas. Over heterogeneous terrain, local point measurements, using profile or eddy-covariance techniques, are not suitable, and it has been found that scintillometry appears to be a reliable method for measuring spatially averaged fluxes (e.g. Beyrich et al., 2002; Meijninger et al., 2002a, 2002b, and references therein).

In order to measure sensible and latent heat fluxes using scintillometry, several investigations have been carried out to date by applying the so-called

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two-wavelength method (e.g. Kohsiek and Herben, 1983; Hill et al., 1988; Andreas, 1989; Hill, 1997; Green et al., 2000, 2001; Meijninger et al., 2002a). With this method the path-averaged refractive index structure parameter C_n^2 is measured at two wavelengths λ_i simultaneously. The parameter C_n^2 can be expressed as

$$C_{n_i}^2 = A_i^2 C_T^2 + 2A_i B_i C_{Tq} + B_i^2 C_q^2, \quad (1)$$

with the three unknowns, the temperature structure parameter C_T^2 , the (specific) humidity structure parameter C_q^2 and the temperature–humidity structure parameter C_{Tq} . The constants $A_i = A(\lambda_i, P, T, q)$ and $B_i = B(\lambda_i, P, T, q)$ depend on the mean temperature (T), the air pressure (P), the specific humidity (q) and the electromagnetic wavelength (λ_i). With measurements of C_n^2 at only two wavelengths, it is impossible to find all three meteorological structure parameters, which is a serious disadvantage of the two-wavelength method. Thus it can only be applied by assuming that the three parameters are not independent, and it is often assumed that $C_{Tq} = \pm(C_T^2 C_q^2)^{1/2}$, implying that the correlation coefficient

$$r_{Tq} = \frac{C_{Tq}}{\sqrt{C_T^2 C_q^2}} \quad (2)$$

between temperature and humidity fluctuations is assumed to be ± 1 (e.g. Kohsiek and Herben, 1983; Hill et al., 1988; Andreas, 1989; Hill, 1997; Green et al., 2000, 2001; Meijninger et al., 2002a). In fact, this assumption is theoretically demanded for a flow that strictly obeys the Monin-Obukhov similarity theory (MOST) (Hill, 1989). For non-ideal flow conditions, the assumption of $C_{Tq} = \pm\sqrt{C_T^2 C_q^2}$ is not always justifiable, and, especially over non-homogeneous terrain, the strict validity of MOST and hence the perfect temperature–humidity correlation is rather questionable (Andreas, 1987; Andreas et al., 1998; De Bruin et al., 1993). Therefore it would be a considerable advantage to know also the parameter C_{Tq} independently. Together with more accurate determinations of C_T^2 and C_q^2 should be feasible, leading to better heat flux estimates. Furthermore, it is thought that scalar–scalar correlations are especially sensitive to violations of MOST (Hill, 1989; Andreas et al., 1998). Thus, measurements of C_{Tq} , together with C_T^2 and C_q^2 , would also be well suited to answering fundamental questions about the MOST validity and applicability.

In order to determine all three meteorological structure parameters in (1) three independent C_n^2 measurements should be available; this demand could be met by the ‘three-wavelength method’ (Phelps and Pond, 1971; Hill et al., 1988; Andreas, 1990; Hill, 1997). Unfortunately, it is not easy to find three independent wavelengths in available electromagnetic spectral regions. In

addition, Andreas (1990) showed theoretically that the accuracy of C_{Tq} by this method is poor, and it would be a rather costly experimental exercise, since at least three scintillometers are needed. These may be reasons why the three-wavelength method never has been experimentally applied.

As suggested by Lüdi (2002), another possibility of measuring path-averaged values of C_{Tq} is by cross-correlating two electromagnetic signals at different wavelengths that pass through the same volume of air. Thereby the effects of refractive dispersion on the bichromatic correlation have to be taken into account. Because only the powers (or the intensities) of the two signals (and not the electromagnetic fields) must be correlated, the additional technical effort is marginal compared to the common two-wavelength method.

The cross-correlation between two signals at different wavelengths has already been discussed (Hill, 1988; Hill and Lataitis, 1989) in order to determine the inner scale of turbulence. However, to our knowledge, the determination of C_{Tq} through this technique has not been previously analysed and tested.

In this paper, we present measurements collected from 15 April 2003 to 29 May 2003 in Lindenberg (Germany), carried out in clear air conditions (with the propagation path free of hydrometeors) with two scintillometers at wavelengths of 940 nm and 3.2 mm. With the correlation of the two signals, we infer therefrom three quantities, the two monochromatic variances and the bichromatic covariance, by which C_T^2 , C_q^2 and C_{Tq} can be extracted. The detailed description of our experimental set-up is outlined in Section 2, and the theoretical derivation of the measurable quantities is given in Section 3. A sensitivity analysis of the measurable quantities is given in Section 4. The measured correlation of path-averaged temperature–humidity fluctuations is discussed in Section 5, followed by conclusions in Section 6.

As far as we know, the current investigation presents the first path-averaged measurements of turbulent temperature–humidity correlations.

2. Description of the Experiment

The study region is a heterogeneous landscape around the Meteorological Observatory Lindenberg of the Deutscher Wetterdienst (DWD) 65 km south-east of Berlin, Germany. The land use in the area is composed of forest (42%), agricultural fields (41%), lakes (6.5%), meadows (5%) and villages (3.5%), and is quite typical of large parts of northern central Europe south of the Baltic Sea. A more detailed description of the landscape (including maps) is given in Beyrich et al. (2002).

As mentioned in Section 1, two scintillometers, one in the optical to near-infrared region (wavelength $\lambda = 940$ nm) and a second in the millimeter

region ($\lambda = 3.2$ mm), are operated over a path length of 4.7 km in a north-south orientation. The microwave and optical scintillometers have antenna diameters of 0.4 and 0.15 m, respectively, which are identical for transmitters and receivers. Both scintillometers have diverging beams with full power beam widths of approximately 1 degree, and have been installed on two permanent towers. Many technical details about the optical instruments can be found in Moene et al. (2005). The receivers (north end of path) are mounted at the top of a 30-m tower and the transmitters are installed at a height of 50 m on the 99-m meteorological tower near Falkenberg. The effective height of the propagation path is 45 m above ground (Beyrich et al., 2002).

As mentioned above, both electromagnetic beams are required to cross the same air volume and therefore the two beams should be as close to each other as possible. If both beams are focused by the same antenna dishes, we obtain the best possible configuration. However, this set-up usually cannot be realised, and an acceptable and realistic set-up is shown in Figure 1. The distance between the optical axis of both transmitters should be as small as possible, and the same is true for the receivers; additionally, the beams should intersect in the middle of the path. In our set-up this latter configuration was realised only approximately, as the mounting possibilities on the 99-m tower were critical. The relative positions of the transmitters and receivers with respect to each other are $\mathbf{d}_T = (y, z) = (0.25, -1)$ m and $\mathbf{d}_R = (y, z) = (0, 0.5)$ m (cf. Figure 1). As will be shown in the next section, the performance of this set-up is not much worse than the optimal configuration.

For the analysis presented here, measurement periods of 10 min were used. This choice is a compromise between stationarity and statistical significance (Caccia et al., 1987); if the period is too long, the assumption of atmospheric stationarity may be violated, and if the period is too short the statistical weight of the estimations is very poor.

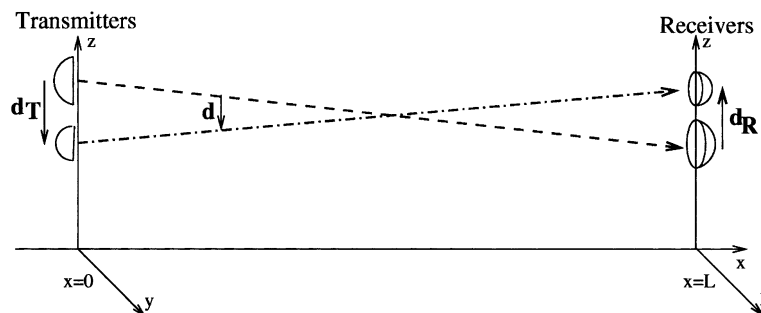


Figure 1. Schematic view of the measurement configuration. The microwave and optical beams are shown as dashed-dotted lines, respectively. The relative positions of the transmitters and receivers are $\mathbf{d}_T = (y, z) = (0.25, -1)$ m and $\mathbf{d}_R = (y, z) = (0, 0.5)$ m.

The data shown herein were obtained during clear air conditions. Thus, any data obtained during precipitation or during fog (leading to a low optical signal) were rejected. In addition, according to the saturation criteria of Ochs and Hill (1982) and Wang et al. (1978), data that may be contaminated by saturation of scintillation were eliminated. In this study also, measured vertical gradients of specific humidity and temperature are used (cf. Section 5). As several *in situ* meteorological stations are available we also rejected measurements when the gradients derived by different stations were not equal to within a factor of 2.

Before computing the variances and the covariance of the measured signals, we bandpass filtered them. The optical signal was split into two parts, which were differently filtered; one part of the optical signal was used to determine the signal variance and the other part was used to determine the cross-covariance between optical and microwave signals. As the latter part of the optical signal was filtered the same as the microwave signal it is not discussed explicitly. The upper frequency cut-offs of the signals are determined by the receiving hardware: the received microwave signal was low-pass filtered at $f=20$ Hz and sampled at $f=40$ Hz (Nyquist theorem). The respective frequencies for the optical signal are higher by a factor of 20. According to the theoretical model (Section 3 and Figure 3), scintillations at microwave wavelengths dominate for $0.04 \lesssim f/v \lesssim 10\text{m}^{-1}$ and scintillations at optical wavelengths for $0.1 \lesssim f/v \lesssim 100\text{m}^{-1}$, with f the temporal frequency and v the wind velocity component transverse to the propagation path. With v equal to several m s^{-1} , signal fluctuations due to turbulence are expected at $\lambda_1 = 940$ nm only in the range of $0.2 < f < 400$ Hz and at $\lambda_2 = 3.2$ mm for $0.06 < f < 20$ Hz. These ranges hence reveal the bandpass filtering of the signals. The suppression of low-frequency fluctuations eliminates gain drifts of the instruments and changes of the atmospheric opacity. Furthermore, as shown in Appendix B, absorption-induced fluctuations are also suppressed. The filtered signals are used thereafter to compute the 10-min averaged variances and the covariance. These values will be compared with the theoretical expressions, to be derived in the following section.

3. Theoretical Background

3.1. THE COVARIANCE FUNCTION OF POWER FLUCTUATIONS

The signals are collected by receivers with extended apertures, where the collected power P of a receiver is the integrated intensity I over the aperture area of the incident electromagnetic wave. Thus the covariance function B_P of the normalised power P detected by two receivers, with collecting circular apertures \sum_{R_1} and \sum_{R_2} of radius R_{R_1} and R_{R_2} , is given by

$$\begin{aligned}
B_P(k_1, k_2) &\doteq \frac{\langle P'_1(k_1)P'_2(k_2) \rangle}{\langle P_1(k_1) \rangle \langle P_2(k_2) \rangle} \\
&= \left(\frac{1}{\pi^2 R_{R1}^2 R_{R2}^2} \right) \int_{\Sigma_{R1}} \int_{\Sigma_{R2}} \frac{\langle I'_1(k_1, \mathbf{r}_1) I'_2(k_2, \mathbf{r}_2) \rangle}{\langle I_1(k_1) \rangle \langle I_2(k_2) \rangle} d^2 \mathbf{r}_1 d^2 \mathbf{r}_2 \quad (3)
\end{aligned}$$

with the electromagnetic wavenumbers k_1 and k_2 of the measured powers P_1 and P_2 , respectively and $d^2 \mathbf{r}$ is the area element of integration. The power and intensity fluctuations are defined as $P' = P - \langle P \rangle$ and $I' = I - \langle I \rangle$, with $\langle \dots \rangle$ representing the ensemble average (here given by the average over the sample period). The vectors \mathbf{r}_1 and \mathbf{r}_2 describe the receiving apertures in the plane $x = L$, with L the length of the propagation path. The intensity I , which, in our case, originates from an extended source, can be treated in the same way as the receivers (reciprocity theorem), i.e. as a sum of intensities J from point sources at positions ρ (e.g. Ochs and Wang, 1978). The normalised intensity covariance function (integrand in Equation (3)) is then given by

$$\frac{\langle I'_1(k_1, \mathbf{r}_1) I'_2(k_2, \mathbf{r}_2) \rangle}{\langle I_1(k_1) \rangle \langle I_2(k_2) \rangle} = \frac{\left\langle \int_{\Sigma_{T1}} d^2 \rho_1 J_1(k_1, \rho_1, \mathbf{r}_1) \int_{\Sigma_{T2}} d^2 \rho_2 J_2(k_2, \rho_2, \mathbf{r}_2) \right\rangle}{\left\langle \int_{\Sigma_{T1}} d^2 \rho_1 J_1(k_1, \rho_1) \right\rangle \left\langle \int_{\Sigma_{T2}} d^2 \rho_2 J_2(k_2, \rho_2) \right\rangle}, \quad (4)$$

In our application, the sources $T1$ and $T2$ are circular transmitters with radius R_{T1} and R_{T2} , respectively. A theoretical examination with the well-known weak fluctuation theory is appropriate because in the mm range, scintillation is weak. Also at the optical wavelength scintillation is mostly weak, as both the receiver and the transmitter have large apertures. Nevertheless, as noted in Wang et al. (1978), as well as in Ochs and Hill (1982), strong scintillation with saturation effects can occur. Therefore, data from the optical scintillometer that do not satisfy the saturation criteria of Ochs and Hill (1982) and Wang et al. (1978) (for our set-up: $C_n^2 \leq 1.4 \times 10^{-14} \text{ m}^{-2/3}$) have to be rejected as the applied theory is not appropriate.

Interchanging the ensemble averages with the integrals (due to the linearity of integration and ensemble average operators), using $J \propto e^{2\chi}$ (with χ the log-amplitude, cf. Tatarskii, 1971) in (4) and presuming $4\langle \chi_1(k_1, \rho_1, \mathbf{r}_1) \chi_2(k_2, \rho_2, \mathbf{r}_2) \rangle \ll 1$ (due to the weak fluctuation approximation) we obtain:

$$\begin{aligned}
B_P(k_1, k_2) &= \left(\frac{4}{\pi^4 R_{R1}^2 R_{R2}^2 R_{T1}^2 R_{T2}^2} \right) \\
&\times \int_{\Sigma_{R1}} \int_{\Sigma_{R2}} \int_{\Sigma_{T1}} \int_{\Sigma_{T2}} d^2 \mathbf{r}_1 d^2 \mathbf{r}_2 d^2 \rho_1 d^2 \rho_2 B_\chi(k_1, k_2), \quad (5)
\end{aligned}$$

$$B_\chi(k_1, k_2) = \langle \chi_1(k_1, \rho_1, \mathbf{r}_1) \chi_2(k_2, \rho_2, \mathbf{r}_2) \rangle. \quad (6)$$

The log-amplitude correlation function $B_\chi(k_1, k_2)$ of two electromagnetic waves, neglecting dispersion effects, propagating through a turbulent medium, is well-known (e.g. Ishimaru, 1978, p. 401). From the latter, the extension of taking into account the effects of dispersion is straightforward. It is found for spherical wave propagation that (Hill and Lataitis, 1989)

$$B_\chi(k_1, k_2) = 2\pi k_1 k_2 \int_0^L dx \int_{-\infty}^{\infty} d^2 \boldsymbol{\kappa} H \cdot e^{j\boldsymbol{\kappa} \mathbf{D}} \Phi_{n_1 n_2}(\boldsymbol{\kappa}), \quad (7)$$

$$H = \sin\left(\frac{\gamma(L-x)\kappa^2}{2k_1}\right) \sin\left(\frac{\gamma(L-x)\kappa^2}{2k_2}\right), \quad (8)$$

where $\mathbf{D} = \boldsymbol{\rho}'' + \gamma(\mathbf{r}'' - \boldsymbol{\rho}'')$, $\boldsymbol{\rho}'' = \boldsymbol{\rho}_1 - \boldsymbol{\rho}_2$, $\mathbf{r}'' = \mathbf{r}_1 - \mathbf{r}_2$. Furthermore we have set $\gamma = x/L$ and $\boldsymbol{\kappa} = (\kappa_\gamma, \kappa_z)$ is the two-dimensional spatial wavenumber. The refractive index cospectrum $\Phi_{n_1 n_2}(\boldsymbol{\kappa})$ is assumed to be locally homogeneous and isotropic. Note, that absorption-induced fluctuations are not taken into account in (7) because they are negligible as shown in Appendix B. Due to the fact that the two transmitters T_1 and T_2 are centred at different locations in the ($x=0$)-plane it is convenient to decompose $\boldsymbol{\rho}_i$ for each transmitter by $\bar{\boldsymbol{\rho}}_i + \boldsymbol{\rho}'_i$, with $\bar{\boldsymbol{\rho}}_i$ the centre position of the transmitter T_i . Then the integrals over the transmitter areas Σ_{T_1} and Σ_{T_2} can be solved by

$$\int_{\Sigma_{T_i}} d^2 \boldsymbol{\rho}'_i e^{\pm j\boldsymbol{\kappa} \boldsymbol{\rho}'_i (1-\gamma)} = 2\pi R_{T_i}^2 \frac{J_1(\boldsymbol{\kappa}(1-\gamma)R_{T_i})}{\boldsymbol{\kappa}(1-\gamma)R_{T_i}}, \quad i = 1, 2,$$

where J_1 is the Bessel function of the first kind of order one. For the same reason we also decompose $\mathbf{r}_i = \bar{\mathbf{r}}_i + \mathbf{r}'_i$, allowing to solve the integrals over the receiving apertures Σ_{R_1} and Σ_{R_2} . The assumption of isotropy allows us to integrate (7) over direction, i.e.

$$\int_{-\infty}^{\infty} \int_{-\infty}^{\infty} d^2 \boldsymbol{\kappa} e^{j\boldsymbol{\kappa} \mathbf{d}} = 2\pi \int_0^{\infty} d\kappa \kappa J_0(\kappa|\mathbf{d}|), \quad (9a)$$

$$\mathbf{d} = (1-\gamma)\mathbf{d}_T + \gamma\mathbf{d}_R, \quad (9b)$$

where J_0 is the Bessel function of the first kind of order zero, $\mathbf{d}_T = \bar{\boldsymbol{\rho}}_1 - \bar{\boldsymbol{\rho}}_2$ is the distance between the transmitters, and $\mathbf{d}_R = \bar{\mathbf{r}}_1 - \bar{\mathbf{r}}_2$ the separation of the receivers (cf. Figure 1).

The refractive index cospectrum $\Phi_{n_1 n_2}(\boldsymbol{\kappa})$ is related to the power spectra of temperature and specific humidity fluctuations, $\Phi_T(\boldsymbol{\kappa})$ and $\Phi_q(\boldsymbol{\kappa})$, and the temperature–humidity cospectrum $\Phi_{Tq}(\boldsymbol{\kappa})$, by (Hill et al., 1980)

$$\Phi_{n_1 n_2}(\boldsymbol{\kappa}) = A_1 A_2 \Phi_T(\boldsymbol{\kappa}) + (A_1 B_2 + A_2 B_1) \Phi_{Tq}(\boldsymbol{\kappa}) + B_1 B_2 \Phi_q(\boldsymbol{\kappa}). \quad (10)$$

The constants A_i and B_i are the same as in Equation (1). In the inertial subrange of turbulence it is generally accepted that all of the spectra in (10) are proportional to $0.033\kappa^{-11/3}$ (three-dimensional Kolmogorov spectrum) with the respective structure and cross-structure parameters as the coefficients. Thus

$$C_{n_1 n_2} = A_1 A_2 C_T^2 + (A_1 B_2 + A_2 B_1) C_{Tq} + B_1 B_2 C_q^2. \quad (11)$$

For $\lambda_1 = \lambda_2$, (11) coincides with (1). Note that the dissipation subrange of turbulence is suppressed due to the use of extended receivers and transmitters. Also the large-scale range (production or input range) of turbulence is not observable due to the high-pass filtering effect of the expression H (cf. Equation (8)). Thus, effects of inhomogeneity scales larger than the Fresnel size $\sqrt{\lambda L}$ of the mm-wave scintillometer (i.e. larger than approximately 4 m) are suppressed. As the inertial subrange of turbulence extends to considerably larger scales than only 4 m (particularly at 45 m above ground) the adoption of the Kolmogorov spectrum ($\propto \kappa^{-11/3}$) is reasonable and justified.

With the refractive cospectrum $\Phi_{n_1 n_2}(\kappa) = 0.033 C_{n_1 n_2} \kappa^{-11/3}$ the expression for B_P finally becomes:

$$B_P(\mathbf{d}) = 8.448\pi^2 k_1 k_2 \int_0^L dx C_{n_1 n_2} W(\mathbf{d}, x), \quad (12)$$

where the weighting function W is given by

$$W(\mathbf{d}, x) = \int_0^\infty d\kappa \kappa^{-8/3} H \cdot F \cdot J_0(\kappa|\mathbf{d}|),$$

$$F = \frac{J_1(\kappa(1-\gamma)R_{T1})J_1(\kappa(1-\gamma)R_{T2})J_1(\kappa\gamma R_{R1})J_1(\kappa\gamma R_{R2})}{\kappa^4 R_{R1} R_{R2} R_{T1} R_{T2} (1-\gamma)^2 \gamma^2},$$

$$\gamma = x/L.$$

Equation (12) describes the covariance function of the power fluctuations for two extended transmitters with radius R_{T1} and R_{T2} separated by \mathbf{d}_T , each observed by individual receivers, separated by \mathbf{d}_R , collecting the signals through apertures of radius R_{R1} and R_{R2} , respectively. The distances \mathbf{d}_T and \mathbf{d}_R are related to \mathbf{d} by (9). The expression F describes the averaging effects due to the extended transmitter and receiver apertures. Let us mention that the aperture averaging effects are significant only for the optical signal. At microwaves, they are negligible since the Fresnel zone ($\sqrt{\lambda L}$) is much larger (by a factor of more than 10) than the aperture radius.

3.2. WEIGHTING FUNCTIONS

Equation (12) is the basic equation in our analysis to infer meteorological structure parameters from scintillation. With our measuring set-up (cf. Sections 1 and 2) we measure three signals, namely the cross-covariance of the detected power of the two signals at different wavelengths and the two monochromatic variances ($\lambda = \lambda_2$ for $\mathbf{d} = 0$ in Equation (12)), yielding path averaged $C_{n_1 n_2}$, $C_{n_1}^2$ and $C_{n_2}^2$. The relative weighting of these latter parameters along the path is described by the weighting function $W(\mathbf{d}, x)$ which depends on \mathbf{d} (Equation (9)). In Figure 2 the function $W(0, x)$, normalised to $W(0, x = L/2)$, for $C_{n_1}^2$ and $C_{n_2}^2$ are shown as dashed-dotted and solid lines, respectively. The parameters for computing the curves are as in our experiment; $L = 4.7$ km, $\lambda_1 = 940$ nm, $\lambda_2 = 3.2$ mm, $R_{R1} = R_{T1} = 0.075$ m and $R_{R2} = R_{T2} = 0.2$ m. The significant reduction of the weighting function for $C_{n_1}^2$ close to the receiver and transmitter is due to the fact that the ratio of aperture radius and Fresnel zone ($\sqrt{\lambda L}$) is larger than 1.

The normalised weighting function for $C_{n_1 n_2}$ should also be similar to the former ones, in particular it should be also maximal in the middle of the path. In addition, in order to have a strong signal, the integral of $W(\mathbf{d}, x)$ should be as large as possible. These demands are best met by minimising the distance \mathbf{d} . This means that \mathbf{d}_T and \mathbf{d}_R should be as small as possible, and the two beams should cross each other in the middle of the path (where $\mathbf{d} = 0$, cf. Figure 1). Computed weighting functions of $C_{n_1 n_2}$ are shown as dotted and

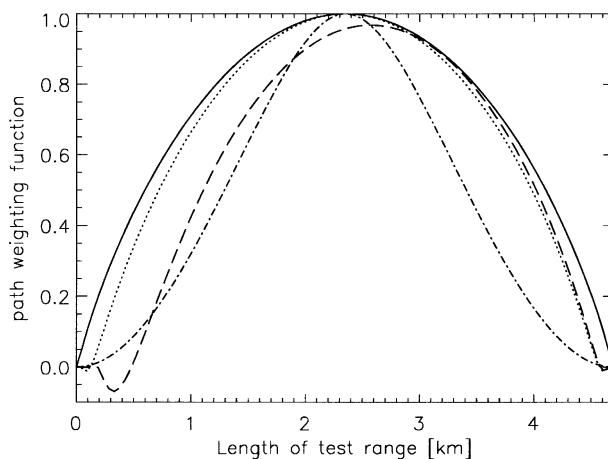


Figure 2. Normalised weighting functions $W(\mathbf{d}, x)$ for $C_{n_1 n_2}$ for different separations of the receivers and transmitters are shown for relative positions of $\mathbf{d}_T = (0.25, -1)$ m and $\mathbf{d}_R = (0, 0.5)$ m (dashed, actual set-up) and $\mathbf{d}_T = -\mathbf{d}_R = (0, 0.5)$ m (dotted). Normalised weighting functions $W(0, x)$ for $C_{n_1}^2$ ($\lambda = 940$ nm) and $C_{n_2}^2$ ($\lambda = 3.2$ mm) are shown as dashed-dotted and solid lines, respectively.

dashed lines in Figure 2, The dotted curve results from $\mathbf{d}_T=(0, 0.5)$ and $\mathbf{d}_R=(0, -0.5)$ (metres) and shows the best possible realisation that could be achieved with our receivers and transmitters. The dashed curve results from $\mathbf{d}_T=(0.25, -1.0)$ and $\mathbf{d}_R=(0, 0.5)$, representing the actual set-up. The weighting functions of $C_{n_1n_2}$ in Figure 2 are again normalised to $W(\mathbf{d}=0, x=L/2)$, i.e. normalised to the maximum of the dotted curve. The dotted curve deviates only slightly from the solid curve, thus $C_{n_1n_2}$ would almost be equally weighted with this configuration as $C_{n_2}^2$. As mentioned above, in our experiment it was not possible to mount the transmitters perfectly, and the two beams do not intersect in the middle of the path. Therefore, the maximum is slightly reduced and shifted toward the receivers. Nevertheless the maximum of the weighting function for our realisation (dashed curve) is still close to the middle of the path and all three parameters $C_{n_1}^2$, $C_{n_2}^2$ and $C_{n_1n_2}$ are averaged basically over the same air volume.

Replacing $C_{n_1}^2$, $C_{n_2}^2$ and $C_{n_1n_2}$ by representative mean values along the path x , Equation (12) can also be written as

$$B_P = \int_0^\infty d\kappa G(\kappa) C_{n_1n_2}, \quad (13a)$$

$$G(\kappa) = 8.448\pi^2 k_1 k_2 \int_0^L dx \kappa^{-8/3} H \cdot F \cdot J_0(\kappa|\mathbf{d}|). \quad (13b)$$

The spectral filter functions $G(\kappa)$ for the optical, the microwave and the covariance signals are given as dashed, solid and dotted lines, respectively in Figure 3, showing the relative contribution of inhomogeneity scales to the

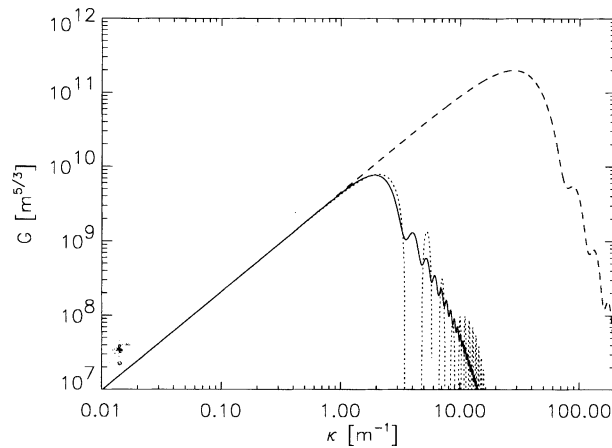


Figure 3. Spectral filter functions for $C_{n_1}^2$ (dashed), $C_{n_2}^2$ (solid) and $C_{n_1n_2}$ (dotted). The refractivities n_1 and n_2 stand for $\lambda_1=940$ nm and $\lambda_2=3.2$ mm, respectively.

measured signals. At the optical wavelength $G(\kappa)$ extends to higher spatial frequencies κ ; therefore the temporal frequency bandpass (before the detection of the signal) was chosen to be higher than at the other two signals (cf. Section 2). The filters $G(\kappa)$ for the microwave and covariance signals are almost equal. Therefore, we filtered the corresponding measured signals equally (cf. Section 2).

3.3. METEOROLOGICAL STRUCTURE PARAMETERS FROM $C_{n_1}^2$, $C_{n_2}^2$ AND $C_{n_1n_2}$

From the measured refractive structure parameters $C_{n_1}^2$, $C_{n_2}^2$ and $C_{n_1n_2}$ the meteorological structure parameters, C_T^2 , C_q^2 and C_{Tq} can be inferred. From (1) and (11) we obtain

$$C_p = MC_s \tag{14}$$

with

$$M = \begin{pmatrix} A_1^2 & 2A_1B_1 & B_1^2 \\ A_2^2 & 2A_2B_2 & B_2^2 \\ A_1A_2 & (A_1B_2 + A_2B_1) & B_1B_2 \end{pmatrix},$$

where C_s and C_p are vectors with components (C_T^2, C_{Tq}, C_q^2) and $(C_{n_1}^2, C_{n_2}^2, C_{n_1n_2})$, respectively, Equation (14) can be inverted if

$$\det[M] = -A_1^3B_2^3 \left(1 - \frac{A_2B_1}{A_1B_2}\right)^3 \neq 0, \tag{15}$$

requiring $(A_2B_1)/(A_1B_2) \neq 1$. In the optical region we determined the wavelength-dependent parameters A_i and B_i with expressions given in Andreas (1988) and in the mm-wavelength region we used the Millimeter-Wave Propagation Model (MPM93) of Liebe et al. (1993). MPM93, also known as the ‘‘Liebe Model’’, was recently experimentally validated at 94 GHz and good agreement between the modelled data and the observations was found for the clear atmosphere (Martin et al., 2000). For the present wavelengths and typical meteorological conditions ($T=20$ °C, $P=1000$ hPa and 90% relative humidity) the ratio $(A_2B_1)/(A_1B_2) = -0.01$. The inverse matrix M^{-1} is given by

$$M^{-1} = \frac{1}{(A_1B_2 - A_2B_1)^2} \begin{pmatrix} B_2^2 & B_1^2 & -2B_1B_2 \\ -A_2B_2 & -A_1B_1 & (A_1B_2 + A_2B_1) \\ A_2^2 & A_1^2 & -2A_1A_2 \end{pmatrix}. \tag{16}$$

Consequently, there exists an unambiguous solution for C_T^2, C_q^2 and C_{Tq} , found from measurements of $C_{n_1}^2, C_{n_2}^2$ and $C_{n_1n_2}$ given by

$$C_s = M^{-1}C_p \tag{17}$$

4. Sensitivity Analysis

In this section we consider how measurement uncertainties of $C_{n_1}^2$, $C_{n_2}^2$ and $C_{n_1 n_2}$ affect the parameters C_T^2 , C_q^2 and C_{Tq} . As already mentioned in Section 1, the “three-wavelength method” (Phelps and Pond, 1971; Andreas, 1990; Hill et al., 1988; Hill, 1997) is the only alternative for measuring path-averaged structure parameters, C_q^2 , C_T^2 and C_{Tq} . In Andreas (1990) a detailed sensitivity analysis of the “three-wavelength method” was carried out. In order to compare the two methods, we performed an analogous sensitivity analysis and use (as far as possible) the same nomenclature.

From (17) it is seen that uncertainties in the measured C_{pj} affect each of the meteorological structure parameters. With the notation already introduced above the absolute uncertainty of C_{si} can be expressed as

$$dC_{si} = \sum_{j=1}^3 \frac{\partial C_{si}}{\partial C_{pj}} dC_{pj}. \quad (18)$$

However, it is better to consider the relative uncertainties since the C_p values can range over several orders of magnitude. Therefore (18) is modified to (Andreas, 1990)

$$\frac{dC_{si}}{C_{si}} = \sum_{j=1}^3 S_{ij} \frac{dC_{pj}}{C_{pj}}, \quad (19)$$

$$S_{ij} = \frac{C_{pj}}{C_{si}} \frac{\partial C_{si}}{\partial C_{pj}}, \quad (20)$$

where dC_{si}/C_{si} is the relative uncertainty of the computed meteorological structure parameter, and dC_{pj}/C_{pj} is the relative uncertainty of the measured refractivity structure parameter. In the following the non-dimensional sensitivity coefficients S_{ij} will be referred to as S_{Tj} , S_{qj} and S_{Tqj} ($j=1, \dots, 3$). As shown by Andreas (1990) the sensitivity coefficients can be expressed as functions of the Bowen ratio Bo and the temperature–humidity correlation coefficient r_{Tq} , defined in (2). The Bowen ratio Bo is an important parameter to understand the sensitivity S ; Bo is defined as the ratio of the sensible to the latent heat fluxes. Since Bo has the same sign as C_{Tq} and $Bo^2 = (c_p \langle wT \rangle)^2 / (L_v \langle wq \rangle)^2 = (c_p^2 C_T^2) / (L_v^2 C_q^2)$ it follows (Andreas, 1990)

$$Bo = \frac{\text{sgn}[C_{Tq}]}{K} \sqrt{\frac{C_T^2}{C_q^2}}, \quad (21)$$

where w , T and q are, respectively, the vertical wind velocity, temperature and specific humidity fluctuations. Further, $K = L_v/c_p$, where c_p is the specific heat of air at constant pressure, and L_v is the latent heat of vaporisation or sublimation,

Strictly speaking, the assumption of $\langle wT \rangle^2 / \langle wq \rangle^2 = C_T^2 / C_q^2$ leading to (21) is only absolutely correct when Monin-Obukhov similarity is valid. With the computed sensitivity functions we know what the relative uncertainties in C_T^2 , C_q^2 and C_{Tq} will be for given relative uncertainties in the measured refractive index structure parameters. The derivations of the sensitivity functions are straightforward (see also Andreas, 1990), and their explicit expressions are given in Appendix A. They are shown as a function of Bo in Figures 4 and 5. Figure 4 shows the curves for S_T and S_q , respectively, with $|r_{Tq}| = 1$, where the curves are computed for typical meteorological conditions; $T = 20$ °C, $P = 1000$ hPa and relative humidity of 90%, as also chosen by Andreas (1990). The sensitivity coefficients depend only weakly on the meteorological conditions.

If the $|S|$ values are larger than 1, the resulting error in the meteorological structure parameter is accordingly larger than the measured refractivity structure parameters. If the S value is close to zero, the respective measure of C_{pj} is not sensitive to the meteorological structure parameter C_{si} under consideration. Thus, in general, optimum measurements can be made by having two S values close to 0 and the other value close to 1 (or -1).

From the sensitivity functions S_{Tj} (cf. Figure 4, upper panel) it is seen that C_T^2 is well measurable when $|Bo| \gtrsim 0.05$. In this range (except for $-0.2 \lesssim Bo \lesssim -0.05$) S_{T1} (from the optical scintillometer) is close to 1 whereas the other two S_{Tj} are close to 0. Only for $|Bo| \lesssim 0.05$ is C_T^2 scarcely measurable. This was also found with the “three-wavelength method” (Andreas, 1990).

From the sensitivity functions S_{qj} (cf. Figure 4, lower panel) it is seen that the information is primarily encountered in the microwave signal, as S_{q2} is approximately 1, whereas the others are mostly close to 0. The structure parameter C_q^2 can be well measured since $|Bo| \lesssim 3$; otherwise larger errors in C_q^2 have to be taken into account. In this respect the “three-wavelength method” is worse as it is more limited to the range $|Bo| < 0.5$ (Andreas, 1990).

From the sensitivity functions S_{Tqj} (cf. Figure 5, upper panel) it is seen that C_{Tq} is much more difficult to measure than the other two meteorological structure parameters. According to S_{Tq1} and S_{Tq2} for $|Bo| < 1$ even small errors in $C_{n_1}^2$ and $C_{n_2}^2$ result in large errors in C_{Tq} . Measurement errors in $C_{n_1 n_2}$ are not so critical because $|S_{Tq3}|$ is small. However, our method can be well used in the two windows around $1 < |Bo| < 10$ with $|r_{Tq}| = 1$. The lower panel of Figure 5 shows that these two ranges become narrower when $|r_{Tq}|$ is smaller. We also tested other wavelength combinations, in particular the combination with an infrared signal from the window of $\lambda = 7.8\text{--}19$ μm . When we pair a visible-to-near infrared signal with the latter, the resulting measurability of C_{Tq} is completely useless. A little better, but still more-or-less useless, and much worse than our combination, is the combination of a millimeter signal with an infrared signal from the above mentioned window. This shows that even though our sensitivity for measuring C_{Tq} is not exhilarating, it cannot simply be improved by choosing other wavelength combinations.

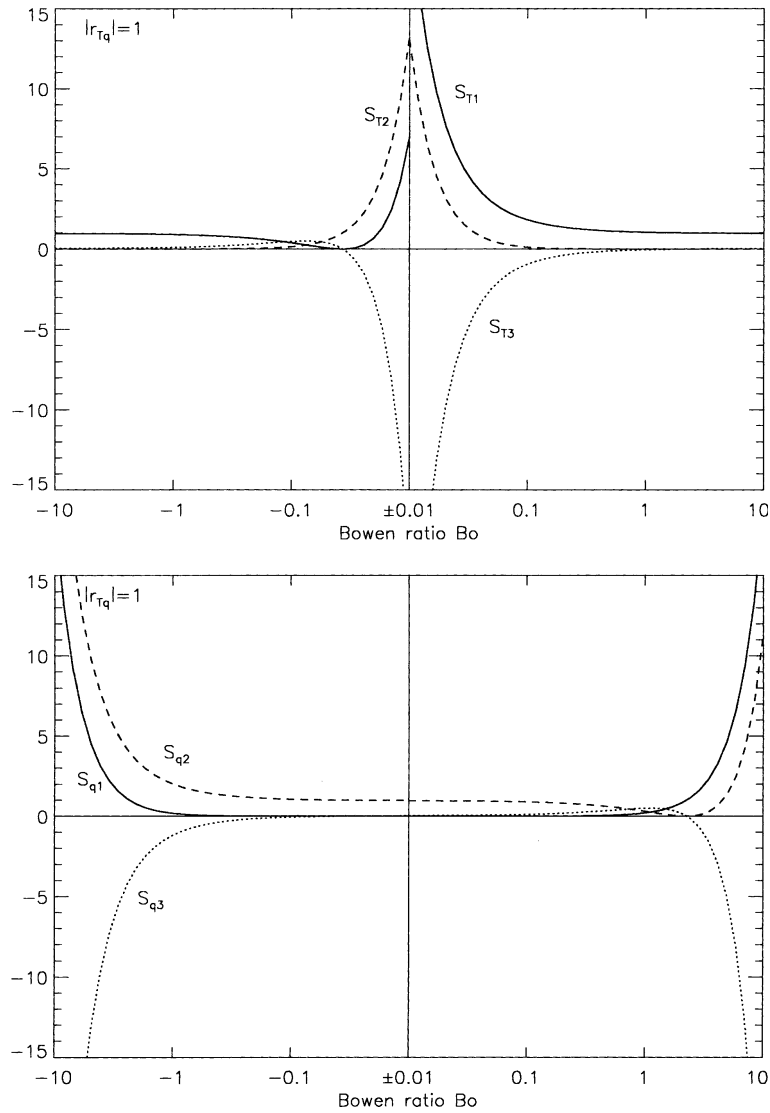


Figure 4. The sensitivity coefficients S_T (upper panel) and S_q (lower panel) for the two wavelengths $\lambda_1=940$ nm, $\lambda_2=3.2$ mm and for $|r_{Tq}|=1$. Environment conditions are $T = 20$ °C, relative humidity of 90% and $P=1000$ hPa.

5. Measurement Results and Discussion

In order to determine C_{Tq} , we find from the results of the previous section that uncertainties in $C_{n_1 n_2}$ measurements are rather uncritical, but errors in $C_{n_1}^2$ and $C_{n_2}^2$ values should be very small. Due to the fact that these structure parameters are averaged over a long path length of several kilometres,

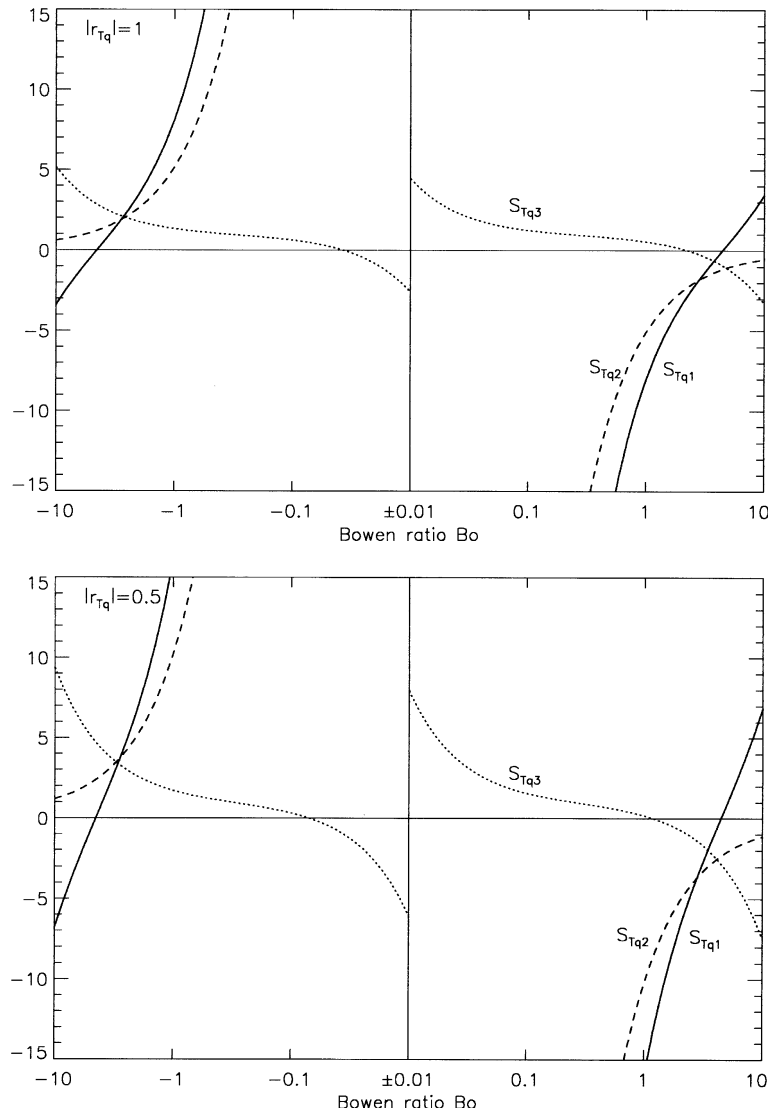


Figure 5. The sensitivity coefficients S_{Tq} for $|r_{Tq}|=1$ (upper panel) and for $|r_{Tq}|=0.5$ (lower panel). All other parameters are the same as in Figure 4.

already with rather short averaging times of 10 min the statistical errors of $C_{n_1}^2$ and $C_{n_2}^2$ are small. Similarly as specified by Scintec for the commercial BLS900 instrument, we found experimentally that our optical scintillometer reaches the noise limit at $C_n^2 \approx 5 \times 10^{-17} \text{ m}^{-2/3}$. With typical daytime values of $C_n^2 \approx 5 \times 10^{-15} \text{ m}^{-2/3}$ (cf. Figure 6) this corresponds to a measurement uncertainty of approximately 1%. Even smaller are the statistical measurement errors in $C_{n_2}^2$; the noise limit of the microwave scintillometer was found

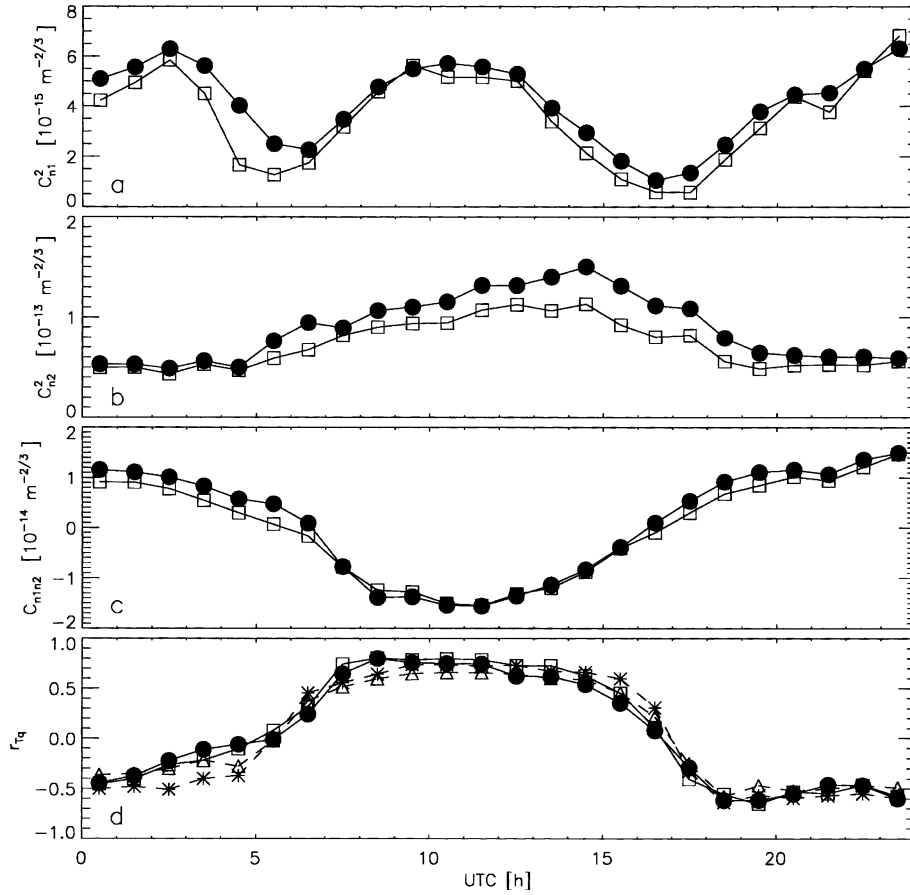


Figure 6. Median (squares) and mean (circles) diurnal cycle of $C_{n_1}^2$ (panel a), $C_{n_2}^2$ (panel b), $C_{n_1 n_2}$ (panel c) and the correlation coefficient r_{Tq} (panel d). The dashed curves are median (asterisk) and mean (triangle) values derived by eddy-covariance techniques. The local mid-day is at 1100 UTC.

to be considerably lower than $10^{-15} \text{ m}^{-2/3}$. As typical measurement values are around $10^{-13} \text{ m}^{-2/3}$ (cf. Figure 6) the errors are smaller than 1%. The uncertainty of $C_{n_1 n_2}$, derived from the cross-covariance, is larger. It also can be estimated experimentally: the cross-covariance between the two signals should disappear when they are correlated with a time lag. As measurements are not perfect the time-lagged cross-covariance does not absolutely disappear but reaches the noise limit, corresponding to $C_{n_1 n_2} \approx 3 \times 10^{-15} \text{ m}^{-2/3}$. With typical $C_{n_1 n_2}$ values (cf. Figure 6) around midnight and midday ($\pm 1.5 \times 10^{-14} \text{ m}^{-2/3}$) this corresponds approximately to statistical errors of 20%. Let us note that this value is in good agreement with the theoretical expectation of the signal-to-noise ratio of the bichromatic correlation (bichromatic

covariance B_P divided by the square root of the two monochromatic variances): the bichromatic correlation was computed numerically and was found to be $0.04r_{n_1n_2}$, where $r_{n_1n_2} = C_{n_1n_2}/(C_{n_1}^2 C_{n_2}^2)^{-1/2}$ with values between -1 and $+1$. With the above mentioned signal-to-noise ratio of $C_{n_1}^2$ and $C_{n_2}^2$ being at least 100 and 200, respectively the detection limit of the bichromatic correlation is then $(1/2)(1/100 + 1/200) \approx 0.007$. Thus, this limit is at least five times smaller than the bichromatic correlation (since $|r_{n_1n_2}|$ is not vanishing), which corresponds well to the empirically found error of $C_{n_1n_2}$ of 20%.

For the quantification of the temperature–humidity correlation the parameter r_{Tq} , defined in (2), is best suited. This parameter and the measured refractivity structure parameters are shown in Figure 6 as mean (circles) and median (squares) diurnal cycles for the whole measuring campaign (15 April 2003 to 29 May 2003). Panel *a* shows $C_{n_1}^2$ from the optical scintillometer; as expected (e.g. Wesely and Alcaraz, 1973), the optical scintillation is high around midday (1100 UTC) and disappears almost around sunrise (0600 UTC) and sunset (1700 UTC). During the night $C_{n_1}^2$ is again large. The refractivity structure parameter for microwaves has a less distinct diurnal cycle (Figure 6, panel *b*). In general $C_{n_2}^2$ is larger during daytime. Compared to the optical structure parameter, $C_{n_2}^2$ is larger by more than one order of magnitude, caused by atmospheric humidity fluctuations. For both structure parameters the mean values (circles) are slightly larger than the median values (squares) because these structure parameters obey a log-normal rather than a normal distribution (e.g. Lüdi and Magun, 2005 and references therein).

One of our key parameters, the cross-structure parameter $C_{n_1n_2}$, derived from cross-correlating the optical and microwave signals, is shown in panel *c*. Obviously $C_{n_1n_2}$ again has a characteristic diurnal cycle, being negative during daytime and positive during nighttime.

From these refractive parameters and mean meteorological quantities (measured at 40 m height, next to the transmitters), the meteorological structure parameters and r_{Tq} (Equation (2)) can be derived (cf. Section 3.3). From panel *d* (Figure 6) it is seen that the correlation between humidity and temperature fluctuations also has a clear diurnal cycle: the $T - q$ correlation is high and positive during daytime and negative during nighttime. The anti-correlation during nighttime seems to be less pronounced than the positive daytime correlation. From 19 May 2003 to 18 June 2003 *in situ* fluctuation measurements were performed at the 50 m level of the meteorological tower using a sonic anemometer-thermometer and an infrared hygrometer. Therefrom, r_{Tq} has also been determined by applying the eddy-covariance method, and is shown as dashed curves in Figure 6d. The triangles and the asterisks indicate the mean and median values, respectively. Obviously, the *in situ* measured correlation coefficients are very similar to those derived by scintillometry, and the small differences are probably due to the different

measurement periods. Of particular interest is also the fact that again the nighttime anti-correlation is less pronounced than the daytime positive correlation. Similar diurnal cycles of r_{Tq} were also observed by other *in situ* investigations using eddy-covariance instruments (e.g. Andreas et al., 1998).

The determined values of r_{Tq} of all usable 10-min samples are shown in Figure 7 as a function of Bo . As r_{Tq} is not directly measured, but computed from C_{Tq} , C_q^2 and C_T^2 that are contaminated with errors, some unphysical values of $|r_{Tq}|$ larger than 1 are seen in Figure 7. The Bowen ratio can be determined by measuring temperature and humidity at two heights, and approximating this “gradient- Bo ” by (Andreas et al., 1998)

$$Bo = \frac{c_p(T_2 - T_1)}{L_v(q_2 - q_1)}. \quad (22)$$

For determining the temperature and humidity gradients necessary for the application of (22) we used data from psychrometer measurements at 2, 10, and 40 m above the ground. We derived Bo from the measurements at 2 and 40 m. For $|Bo| > 0.2$, we used only those values that were equal within a factor of 2 with Bo derived from the other two height combinations.

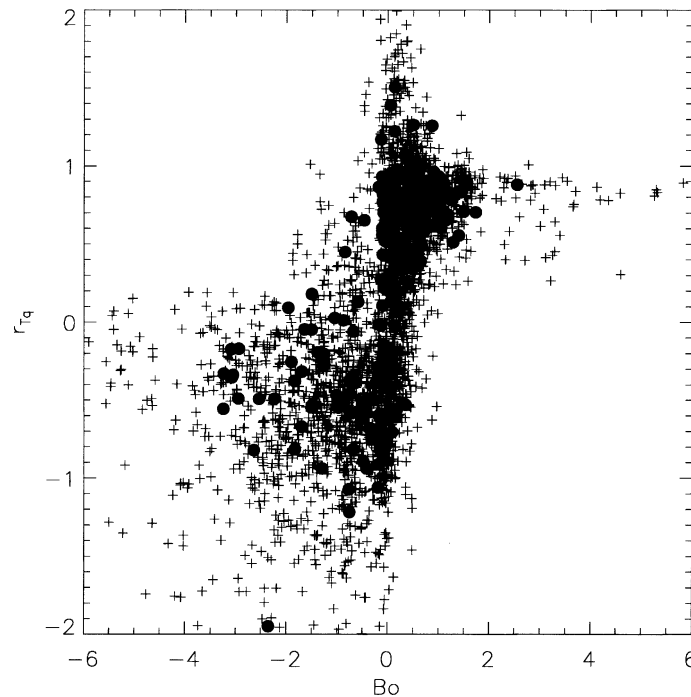


Figure 7. Determined correlation coefficients r_{Tq} versus the Bowen ratio Bo . For most points (crosses) Bo was determined by Equation (22). During the period from 19 to 29 May 2003, Bo was also derived from eddy-covariance measurements of the sensible and latent heat fluxes at 50 m height (circles).

As mentioned above, starting on May 19, 2003, direct wind, temperature and humidity fluctuation measurements were performed at the 50-m level of the meteorological tower. From these measurements, turbulent fluxes of sensible and latent heat have been determined using the eddy-covariance method, and the Bo ratio was directly calculated from the flux data (bold dots in Figure 7). A comparison between the latter Bo and the “gradient- Bo ” (Equation (22)) showed that the gradient- Bo is reasonably accurate. As expected from the foregoing sensitivity analysis in Section 4, the scatter of r_{Tq} for individual measurements is large, especially for $|Bo| \lesssim 1$. Some scatter is likely also due to the fact that the point measure of Bo is not representative for the whole propagation path. Nevertheless, a clear relation between r_{Tq} and Bo can be well observed: the correlation coefficient r_{Tq} is close to zero for small $|Bo|$ and rises toward 1 for larger Bo . For negative Bo the coefficient r_{Tq} becomes also negative. This finding is substantiated in Figure 8 where we divided all measurement values into classes of Bo ranges and evaluated the data in each class statistically. For $|Bo| < 1$ the class were chosen in steps of 0.2; for $|Bo| \geq 1$ the classes were chosen in steps of 1.0. The mean values (circles in Figure 8) of r_{Tq} for each class rise steadily from -0.6 for $Bo < -2$ to zero at small $|Bo|$ and to $r_{Tq} = 0.8$ for $Bo > 2$. The vertical bars indicate the statistical errors of the mean values (standard deviation divided by the square root of the number of measurements), which are in the order of 0.05. (cf. Table I). The median values

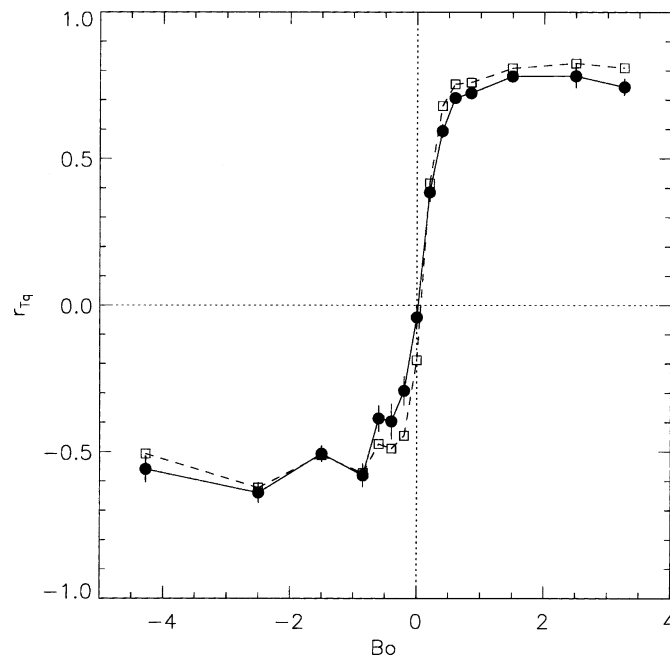


Figure 8. Mean values (circles), their statistical errors (vertical bars) and median values (squares) of r_{Tq} for Bo classes. The numerical values are summarised in Table I.

TABLE I

Summary of the statistically averaged r_{Tq} values in Bo classes. $\overline{r_{Tq}}$ and $\langle r_{Tq} \rangle$ are the mean and median values of r_{Tq} , respectively, S is the standard deviation, $S_{\bar{r}} (= S/\sqrt{n})$ is the mean error of the mean value and n is the number of measurements.

Bo	$\overline{r_{Tq}}$	$\langle r_{Tq} \rangle$	S	$S_{\bar{r}}$	n
< -3	-0.559	-0.507	0.495	0.045	123
-3.0-2.0	-0.639	-0.623	0.479	0.035	191
-2.0-1.0	-0.507	-0.513	0.506	0.027	340
-1.0-0.7	-0.581	-0.573	0.519	0.039	176
-0.7-0.5	-0.387	-0.474	0.525	0.044	141
-0.5-0.3	-0.397	-0.489	0.649	0.059	118
-0.3-0.1	-0.292	-0.446	0.678	0.049	188
-0.1-0.1	-0.041	-0.188	0.756	0.040	357
0.1-0.3	0.385	0.416	0.600	0.032	356
0.3-0.5	0.593	0.679	0.481	0.023	435
0.5-0.7	0.707	0.754	0.302	0.017	288
0.7-1.0	0.723	0.759	0.275	0.019	199
1.0-2.0	0.781	0.808	0.164	0.012	188
2.0-3.0	0.781	0.825	0.179	0.040	20
≥ 3	0.744	0.809	0.191	0.029	45

(squares) are also shown in Figure 8. The latter ones and the mean values are basically equal to within the statistical errors, indicating that r_{Tq} seems to be normally distributed. This was also confirmed by the fact that the higher moments (particularly the skewness) were close to zero for each sample population of the Bo -Classes. Of particular interest is the fact that the three r_{Tq} values at $Bo > 1$ are equal within the error bars, namely $r_{Tq} = 0.78 \pm 0.01$, $r_{Tq} = 0.78 \pm 0.04$, and $r_{Tq} = 0.74 \pm 0.03$, indicating that r_{Tq} reaches an upper limit. Also for negative Bo the anti-correlation reaches a minimal value: the two r_{Tq} values at $Bo < -2$ are also equal within the error bars, i.e. $r_{Tq} = -0.64 \pm 0.04$ and $r_{Tq} = -0.56 \pm 0.05$. Obviously, as already mentioned above, the negative correlation is less pronounced. All measurement points in Figure 8, including the statistical errors, are summarised in Table I.

Even though there are quite large uncertainties in the behaviour and absolute values of r_{Tq} , there do not exist many investigations into the correlation of temperature-humidity fluctuations. In general, it is believed that when sensible and latent heat fluxes have the same direction (i.e. $Bo > 0$), the correlation between humidity and temperature fluctuations is positive. On the other hand, when sensible and latent heat fluxes have an opposite flow-direction (i.e. $Bo < 0$), humidity and temperature fluctuations are anti-correlated. In a flow that strictly obeys MOST, the correlation coefficient

between any two conservative scalars should be ± 1 (Hill, 1989). However, this is not always true, and MOST can be violated (Hill, 1989; Andreas et al., 1998). It is thought that scalar–scalar correlations, such as r_{Tq} , are especially sensitive indicators of deviations from MOST (Hill, 1989). Over the metre-scale heterogeneous landscape “Sevilleta”, Andreas et al. (1998) measured with an *in situ* eddy-covariance instrument values of r_{Tq} being 0.8 or less with a mean value of 0.76 for daytime measurements. Consequently it was concluded that violations of MOST are due to the metre-scale heterogeneity of the Sevilleta. The heterogeneity at our test site is higher than at the Sevilleta, as the dominant heterogeneity scales are in the order of hundreds of metres. Nevertheless our measured values of r_{Tq} are similar. The mean value of r_{Tq} of 0.8 for large Bo is slightly larger than their 0.76.

From other investigations, there exist several hypotheses to explain the non-perfect temperature–humidity correlation observed elsewhere (cf. Priestley and Hill, 1985; De Bruin et al., 1993; Andreas et al., 1998). In our case, the explanation of Andreas et al. (1998) for the Sevilleta landscape appears to be most plausible, as the heterogeneity leads to distributed heat and moisture sources that cannot produce temperature and humidity fluctuations with perfect correlation.

Finally, as mentioned in the introduction, let us close this section by noting that measured r_{Tq} (or C_{Tq}) lead to better estimates of C_T^2 and C_q^2 , yielding more accurate heat flux estimates. For daytime non-stable stratification ($r_{Tq} > 0$, $Bo > 0$), C_T^2 and C_q^2 , was determined by the conventional two wavelength method (assuming $r_{Tq} = 1$) and was compared to the respective structure parameters determined by our method. The more r_{Tq} deviates from 1 the worse is the estimation by the two-wavelength method. Assuming $r_{Tq} = 1$ instead of $r_{Tq} = 0.8$ leads to an overestimation of C_q^2 by a factor of around 1.16 and an underestimation of C_T^2 , by a factor of 1.03. For values of r_{Tq} between 0.4 and 0.6 the conventional method leads to a C_q^2 overestimation and C_T^2 underestimation by factors of 1.25 and 1.1, respectively. For r_{Tq} smaller than 0.4 the overestimation of C_q^2 remains approximately around a factor of 1.25 but the underestimation of C_T^2 is raised to a factor of 1.2.

The fact that C_q^2 is notably overestimated by the two-wavelength method, even when the deviation from perfect positive r_{Tq} correlation is only small ($r_{Tq} \approx 0.8$), can perhaps partly explain the conclusions of Beyrich et al. (2004) – that is, that conventionally derived scintillometric latent heat fluxes in Lindenberg have been found to be significantly larger than respective *in situ* measurements.

6. Conclusion

We investigated theoretically and experimentally the measurability of C_{Tq} and r_{Tq} by cross-correlating the detected signals of an optical and a

microwave scintillometer. For the first time (as far as we know), it was possible to measure path-averaged correlations of temperature–humidity fluctuations. Compared to the traditional two-wavelength method the necessary instrumental modifications are marginal. Hence, in comparison to the suggested three (or more) wavelength method by Andreas (1990), which also measures path-averaged C_{Tq} values, our method is less expensive and simpler in set-up. However, unlike the preliminary hopes mentioned in Lüdi (2002), the sensitivity analysis showed that our method does no better in measuring C_{Tq} than the three-wavelength method. Therefore, the errors in individual measurements can be large. The correlation of temperature–humidity fluctuations derived by scintillometry was compared to that derived by eddy-covariance, and good agreement was found. Our results are similar to those found over the (metre-scale) heterogeneous Sevilleta by Andreas et al. (1998), as r_{Tq} fell to -0.6 for Bo smaller than -2 and rose to a value of 0.8 for $Bo > 1$. The statistical errors in the averaged r_{Tq} in each Bo class are approximately 0.05 , which is quite small.

The fact that r_{Tq} is smaller than 1 shows that our observations violate MOST. However, the quite large temperature–humidity (anti-)correlation for $|Bo| > 1$ also indicates that MOST is not seriously violated, as also concluded for the Sevilleta data (Andreas et al., 1998). Therefore, in order to infer heat fluxes from scintillometer observations made over heterogeneous landscapes, the similarity functions of MOST might still be applicable. However, when the common two-wavelength method is used, we suggest that non-perfect correlation of temperature–humidity fluctuations should be presumed. For Lindenberg and for similar landscapes (as encountered in large parts of northern central Europe south of the Baltic Sea) the function found in Figure 8 (and Table I) may serve as a reference function.

Appendix A. Sensitivity Functions

The explicit form of the sensitivity functions shown in Figures 4 and 5 are given below:

$$S_{T1} = \frac{(A_1^2 + 2 \frac{A_1 B_1 |r_{Tq}|}{KBo} + \frac{B_1^2}{K^2 Bo^2}) B_2^2}{(A_1 B_2 - A_2 B_1)^2}, \quad (A1)$$

$$S_{T2} = \frac{(A_2^2 + 2 \frac{A_2 B_2 |r_{Tq}|}{KBo} + \frac{B_2^2}{K^2 Bo^2}) B_1^2}{(A_1 B_2 - A_2 B_1)^2}, \quad (A2)$$

$$S_{T3} = -2 \frac{(A_1 A_2 + \frac{(A_1 B_2 + A_2 B_1) |r_{Tq}|}{KB_0} + \frac{B_1 B_2}{K^2 B_0^2}) B_1 B_2}{(A_1 B_2 - A_2 B_1)^2}, \quad (\text{A3})$$

$$S_{Tq1} = - \frac{(\frac{A_1^2 B_0 K}{|r_{Tq}|} + 2 A_1 B_1 + \frac{B_1}{|r_{Tq}| B_0 K}) A_2 B_2}{(A_1 B_2 - A_2 B_1)^2}, \quad (\text{A4})$$

$$S_{Tq2} = - \frac{(\frac{A_2^2 B_0 K}{|r_{Tq}|} + 2 A_2 B_2 + \frac{B_2}{|r_{Tq}| B_0 K}) A_1 B_1}{(A_1 B_2 - A_2 B_1)^2}, \quad (\text{A5})$$

$$S_{Tq3} = \frac{(\frac{A_1 A_2 B_0 K}{|r_{Tq}|} + A_1 B_2 + A_2 B_1 + \frac{B_1 B_2}{|r_{Tq}| B_0 K}) (A_1 B_2 + A_2 B_1)}{(A_1 B_2 - A_2 B_1)^2}, \quad (\text{A6})$$

$$S_{q1} = \frac{(A_1^2 K^2 B_0^2 + 2 A_1 B_1 |r_{Tq}| K B_0 + B_1^2) A_2^2}{(A_1 B_2 - A_2 B_1)^2}, \quad (\text{A7})$$

$$S_{q2} = \frac{(A_2^2 K^2 B_0^2 + 2 A_2 B_2 |r_{Tq}| K B_0 + B_2^2) A_1^2}{(A_1 B_2 - A_2 B_1)^2}, \quad (\text{A8})$$

$$S_{q3} = -2 \frac{(A_1 A_2 K^2 B_0^2 + (A_1 B_2 + A_2 B_1) |r_{Tq}| K B_0 + B_1 B_2) A_1 A_2}{(A_1 B_2 - A_2 B_1)^2}. \quad (\text{A9})$$

B. Absorption-induced Fluctuations

Sensitive to possible absorption-induced fluctuations are the mm-wave variance, yielding $C_{n_2}^2$, and the cross covariance, yielding $C_{n_1 n_2}$. In order to estimate the influence of absorption on scintillation at 94 GHz, the analysis of Niveen et al. (1998) and Hill et al. (1980) are useful. The variance of the detected signal can be regarded as the sum of three variances, namely σ_R^2 due to the real part n of the refractive-index fluctuations; σ_I^2 , due to the imaginary part m of the refractive-index fluctuations; and σ_{IR} , due to the covariance of the real and imaginary parts of the refractive-index fluctuations. (In Niveen et al. (1998) the latter was not considered.) They are given by

$$\sigma_R^2 = 8.448 \pi^2 k_1 k_2 \int_0^L dx \int_0^\infty d\kappa \kappa^{-8/3} \cdot F \cdot A, \quad (\text{B1})$$

$$\sigma_I^2 = 8.448 \pi^2 k_1 k_2 \int_0^L dx \int_0^\infty d\kappa \kappa^{-8/3} \cdot F \cdot B, \quad (\text{B2})$$

$$\sigma_{IR} = 8.448\pi^2 k_1 k_2 \int_0^L dx \int_0^\infty d\kappa \kappa^{-8/3} \cdot F \cdot C, \quad (\text{B3})$$

with

$$A = \left(\sin\left(\frac{\gamma(L-x)\kappa^2}{2k_1}\right) \right)^2 C_n^2, \quad (\text{B4})$$

$$B = \left(\cos\left(\frac{\gamma(L-x)\kappa^2}{2k_1}\right) \right)^2 C_m^2, \quad (\text{B5})$$

$$C = 2 \left(\cos\left(\frac{\gamma(L-x)\kappa^2}{2k_1}\right) \right) \left(\sin\left(\frac{\gamma(L-x)\kappa^2}{2k_1}\right) \right) C_{nm}, \quad (\text{B6})$$

where $k_1 = k_2$ is the wavenumber of the mm wave. All parameters are defined as in Section 3. Neglecting temperature fluctuations, the computation (with MPM93 and the *in situ* meteorological data) of the ratios C_m^2/C_n^2 and C_{nm}/C_n^2 are found to be smaller than 4×10^{-7} and 6×10^{-4} , respectively. The boundaries of the integration over the spatial wave number κ , in (B.1) to (B.3) can be narrowed by temporal filtering of the measured signal with an appropriate band-pass filter (cf. Section 2). In our experiment the integration boundaries are approximately at $\kappa \cong 0.04$ and 20 m^{-1} , respectively, and with these integration limits it is found that $\sigma_I^2/\sigma_R^2 \lesssim 0.0003$ and $\sigma_{IR}/\sigma_R^2 \lesssim 0.001$. Thus, absorption effects at the mm-wave variance can be safely neglected. The same was found by Wouter Meijninger (personal communication, 2004) who also computed the absorption effects at 94 GHz, following Nieveen et al. (1998). The small influence of absorption at 94 GHz is due to the fact that this frequency is in an atmospheric absorption window.

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References

- Andreas, E. L.: 1987, 'On the Kolmogorov constants for the temperature-humidity Cospectrum and the Refractive Index Spectrum', *J. Atmos. Sci.* **44**, 2399–2406.
- Andreas, E. L.: 1988, 'Estimating C_n^2 over Snow and Sea Ice from Meteorological Data', *J. Opt. Soc. Amer. A* **5**, 481–495.
- Andreas, E. L.: 1989, 'Two-wavelength Method of Measuring Path-averaged Turbulent Surface Heat Fluxes', *J. Atmos. Ocean. Tech.* **6**, 280–292.
- Andreas, E. L.: 1990, 'Three-wavelength Method of Measuring Path-averaged Turbulent Heat Fluxes', *J. Atmos. Ocean. Tech.* **7**, 801–814.
- Andreas, E. L., Hill, R. J., Gosz, J. R., Moore, D. I., Otto, D. W. and Sarma, A. D.: 1998, 'Statistics of Surface-layer Turbulence over Terrain with Metre-scale Heterogeneity', *Boundary-Layer Meteorol.* **86**, 379–408.
- Beyrich, F., Bange, J., Bernhofer, C., Foken, T., Hennemuth, B., Huneke, S., Kohsiek, W., Leps, J.-P., Lohse, H., Lüdi, A., Mauder, M. and Meijninger, W. M. L.: 2004, 'Energy and Water Vapour Fluxes over a Heterogeneous Land Surface: The LITFASS-2003 Experiment', in *Proceedings of the 16th Symposium on Boundary Layers and Turbulence*, Portland, Maine U.S.A.
- Beyrich, F., De Bruin, H. A. R., Meijninger, W. M. L., Schipper, W. and Lohse, H.: 2002, 'Results from One-year Continuous Operation of a Large Aperture Scintillometer over a Heterogeneous Land Surface', *Boundary-Layer Meteorol.* **105**, 85–97.
- Caccia, J. L., Azouit, M. and Vernin, J.: 1987, 'Wind and C_n^2 Profiling by Single-star Scintillation Analysis', *Appl. Opt.* **26**, 1288–1294.
- De Bruin, H. A. R., Kohsiek, W. and Den Hurk, B. J. J. M.: 1993, 'A Verification of Some Methods to Determine the Fluxes of Momentum, Sensible Heat and Water Vapour using Standard Deviation and Structure Parameter of Scalar Meteorological Quantities', *Boundary-Layer Meteorol.* **63**, 231–257.
- Green, A. E., Astill, M. S., McAneney, K. J. and Nieveen, J. P.: 2001, 'Path-averaged Surface Fluxes Determined from Infrared and Microwave Scintillometers', *Agric. For. Meteorol.* **109**, 233–247.
- Green, A. E., Green, S. R., Astill, M. S. and Caspari, H. W.: 2000, 'Estimating Latent Heat Flux from a Vineyard Using Scintillometry', *J. Terr. Atmos. Ocean. Sci.* **11**(2), 525–542.
- Hill, R. J.: 1988, 'Comparison of Scintillation Methods for Measuring the Inner Scale of Turbulence', *Appl. Opt.* **27**, 2187–2193.
- Hill, R. J.: 1989, 'Implications of Monin-Obukhov Similarity Theory for Scalar Quantities', *J. Atmos. Sci.* **46**, 2236–2244.
- Hill, R. J.: 1997, 'Algorithms for Obtaining Atmospheric Surface-layer Fluxes from Scintillation Measurements', *J. Atmos. Ocean. Tech.* **14**, 456–467.
- Hill, R. J., Bohlander, R. A., Clifford, S. F., McMillan, R. W., Priestley, J. T. and Schoenfeld, W. P.: 1988, 'Turbulence-induced Millimeter-wave Scintillation Compared with Micrometeorological Measurements', *IEEE Trans. Geosci. Remote Sens.* **26**, 330–342.
- Hill, R. J., Clifford, S. F. and Lawrence, R. S.: 1980, 'Refractive-index and Absorption Fluctuations in the Infrared Caused by Temperature, Humidity, and Pressure Fluctuations', *J. Opt. Soc. Amer.* **70**, 1192–1205.
- Hill, R. J. and Lataitis, R. J.: 1989, 'Effect of Refractive Dispersion on the Bichromatic Correlation of Irradiances for Atmospheric Scintillation', *Appl. Opt.* **28**, 4121–4125.
- Ishimaru, A.: 1978, 'Wave Propagation and Scattering in Random Media,' Vol. 2, Academic Press, New York, San Francisco, 339 pp.
- Kohsiek, W. and Herben, A. J.: 1983, 'Evaporation Derived from Optical and Radio-wave Scintillation', *Appl. Opt.* **22**, 2566–2569.

- Liebe, H. J., Hufford, G. A., and Cotton, M. G.: 1993, 'Propagation Modeling of Moist Air and Suspended Water/Ice Particles at Frequencies Below 1000 GHz', in: *AGARD 52nd Specialists' Meeting of the Electromagnetic Wave Propagation Panel*. Mallorca, Spain.
- Lüdi, A.: 2002, 'Determination of the Structure Constants of Temperature, Humidity and temperature-humidity by Scintillometry at two Wavelengths', IAP Research Report 2002-13, Institute of Applied Physics, University of Bern, Bern, Switzerland. (in German), 8 pp.
- Lüdi, A. and Magun, A.: 2005, 'Refractivity Structure Constant and Length Scales from mm Wave Propagation in the Stably Stratified Troposphere', *J. Atmos. Sol. Terr. Phys.* **67**, 435–447.
- Martin, L., Mätzler, C., and Kämpfer, N.: 2000, 'Transmission Measurements at 94 GHz to Validate the Liebe Model', in *AP 2000 Millennium Conference on Antennas and Propagation, SP-444, ESA Publication Division*. Davos, Switzerland.
- Meijninger, W. M. L., Green, A. E., Hartogensis, O. K., Kohsiek, W., Hoedjes, J. C. B., Zurbier, R. M. and De Bruin, H. A. R.: 2002a, 'Determination of Area-averaged Water Vapour Fluxes with Large Aperture and Radio Wave Scintillometers over a Heterogeneous Surface Flevoland Field Experiment', *Boundary-Layer Meteorol.* **105**, 63–83.
- Meijninger, W. M. L., Hartogensis, O. K., Kohsiek, W., Hoedjes, J. C. B., Zurbier, R. M. and De Bruin, H. A. R.: 2002b, 'Determination of Area-averaged Sensible Heat Fluxes with a Large Aperture Scintillometer over a Heterogeneous Surface – Flevoland Field Experiment', *Boundary-Layer Meteorol.* **105**, 37–62.
- Moene, A. F., Meijninger, W. M. L., Hartogensis, O. K., Heusinkveld, B. G., and de Bruin, H. A. R.: 2005, 'The Effects of Finite Accuracy in the Manufacturing of large Aperture Scintillometers', Internal report 2005/1, Wageningen University, Meteorology and Air Quality Group, Wageningen, The Netherlands, 19 pp.
- Nieveen, J. P., Green, A. E. and Kohsiek, W.: 1998, 'Using a Large-aperture Scintillometer to Measure Absorption and Refractive Index Fluctuations', *Boundary-Layer Meteorol.* **87**, 101–116.
- Ochs, G. R. and Hill, R. J.: 1982, 'A Study of Factors Influencing the Calibration of Optical C_n^2 Meters', NOAA Tech. Memo. ERL WPL-106, NOAA, 24 pp.
- Ochs, G. R. and Wang, T. I.: 1978, 'Finite Aperture Optical Scintillometer for Profiling Wind and C_n^2 ', *Appl. Opt.* **17**, 3774–3778.
- Phelps, G. T. and Pond, S.: 1971, 'Spectra of the Temperature and Humidity Fluctuations and the Turbulent Fluxes of Momentum, Moisture and Sensible Heat over the Ocean', *J. Atmos. Sci.* **28**, 918–928.
- Priestley, J. T. and Hill, R. J.: 1985, 'Measuring High-frequency Humidity, Temperature and Radio Refractive Index in the Surface Layer', *J. Atmos. Ocean. Tech.* **60**, 233–251.
- Tatarskii, V. I.: 1971, 'The Effects of the Turbulent Atmosphere on Wave Propagation', Isreal Program for Scientific Translation Ltd., Jerusalem, 472 pp.
- Wang, T. I., Ochs, G. R. and Clifford, S. F.: 1978, 'A Saturation-resistant Optical Scintillometer to Measure C_n^2 ', *J. Opt. Soc. Amer.* **68**, 334–338.
- Wesely, M. L. and Alcaraz, E. C.: 1973, 'Diurnal Cycles of the Refractive Index Structure Function Coefficient', *J. Geophys. Res.* **78**, 6224–6232.