1 Oceanographic drivers of marine mammal and seabird habitat-use across shelf-

2 seas: A guide to key features and recommendations for future research and
3 conservation management

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# **10 ABSTRACT**

11 Mid-latitude ( $\sim$ 30-60°) seasonally stratifying shelf-seas support a high abundance and diversity of 12 marine predators such as marine mammals and seabirds. However, anthropogenic activities and 13 climate change impacts are driving changes in the distributions and population dynamics of these animals, with negative consequences for ecosystem functioning. Across mid-latitude shelf-seas, 14 15 marine mammals and seabirds are known to forage at a number of oceanographic habitats that structure the spatio-temporal distributions of prey. Knowledge of these and the bio-physical 16 mechanisms driving such associations are needed to improve marine management and policy. Here, 17 we provide a concise and easily accessible guide for both researchers and managers of marine systems 18 19 on the predominant oceanographic habitats that are favoured for foraging by marine mammals and 20 seabirds across mid-latitude shelf seas. We (1) identify and describe key discrete physical features 21 present across the continental shelf, working inshore from the shelf-edge to the shore line, (2) provide 22 an overview of findings relating to associations between these habitats and marine mammals and

23 seabirds, (3) identify areas for future research and (4) discuss the relevance of such information to conservation management. We show that oceanographic features preferentially foraged at by marine 24 25 mammals and seabirds include shelf-edge fronts, upwelling and tidal-mixing fronts, offshore banks and internal waves, regions of stratification, and topographically complex coastal areas subject to 26 strong tidal flow. Whilst associations were variable across taxa and through space and time, in the 27 majority of cases interactions between bathymetry and tidal currents appear to play a dominant role, 28 alongside patterns in seasonal stratification and shelf-edge upwelling. We suggest that the ecological 29 30 significance of these bio-physical structures stems from a capacity to alter the densities, distributions 31 (both horizontally and vertically) and behaviours of prey in a persistent and/or predictable manner 32 that increases accessibility for predators, and likely enhances foraging efficiency. Future 33 conservation management should aim to preserve and protect these habitats. This will require adaptive and holistic strategies that are specifically tailored to the characteristics of an oceanographic 34 35 feature, and where necessary evolve through space and time. Improved monitoring of animal movements and bio-physical conditions across shelf-seas would aid in achieving this. Areas for 36 future research include multi- disciplinary/trophic studies of the mechanisms linking oceanographic 37 38 habitats, prey and marine mammals and seabirds (which may also elucidate the importance of lesser studied features such as bottom fronts and Langmuir circulation cells), alongside a better 39 40 understanding of how predators perceive their environment and develop foraging strategies during 41 immature/juvenile stages. Estimates of the importance of bio-physical processes at a population level 42 should also be obtained. Such information is vital to ensuring the future health of these complex 43 ecosystems, and can be used to assess how anthropogenic activities and changes in the environment 44 will impact the functioning and spatio-temporal dynamics of these bio-physical features and their use 45 by marine predators.

# 46 KEYWORDS: Habitat selection · Foraging ecology · Bio-physical coupling · Conservation 47 management · Marine mammals · Oceanography · Seabirds

#### 48 1 Introduction

Mid-latitude ( $\sim$ 30-60°) seasonally stratifying shelf seas cover less than 8% of the world's oceans, yet 49 50 account for ~15% of marine global productivity (Muller-Karger et al., 2005; Simpson and Sharples, 51 2012). These regions support high abundances of species above the base of the food web, which 52 includes a diversity of marine predators such as marine mammals and seabirds. However, they are currently going through a period of rapid alteration, driven by the combined and cumulative effects 53 54 of a range of anthropogenic activities and impacts such as climate change, fisheries and the development of marine renewables (Walther et al., 2002; Frid et al., 2005; Witt et al., 2012; Avila et 55 56 al., 2018; Kroodsma et al., 2018). As a result, many populations of marine mammals and seabirds in shelf-seas have shifted in distribution (Bertrand et al., 2012; Hazen et al., 2013) or suffered severe 57 declines (Cury et al., 2011; McCauley et al., 2015; Paleczny et al., 2015), which has negatively 58 59 impacted the functioning of these systems as a whole (Heithaus *et al.*, 2008). Addressing this issue 60 represents a major environmental conservation challenge requiring response at the policy level 61 alongside informed management practices.

62 Marine mammals and seabirds meet their requirements for survival, growth and reproduction through 63 the exploitation of prey resources from their surrounding environment. Whilst typically these animals are highly mobile and capable of ranging vast distances (Block et al., 2011), foraging efforts are often 64 concentrated over localised spatio-temporal scales (Hastie et al., 2004; Sydeman et al., 2006; 65 Weimerskirch, 2007). Such heterogeneity in distributions is expected to match the organisation of 66 prey, but this has proved surprisingly challenging to demonstrate (Logerwell et al., 1998; Fauchald 67 and Erikstad, 2002; Gremillet et al., 2008; Torres et al., 2008), particularly at finer scales which may 68 be impacted by confounding factors (Schneider and Piatt, 1986; Hunt et al., 1992; Mehlum et al., 69 70 1999; Swartzman and Hunt, 2000; Vlietstra, 2005; Fauchald, 2009). Increasing evidence suggests the 71 behavioural patterns of marine predators (particularly those feeding on plankton and/or forage and 72 pelagic fish) are linked to bio-physical oceanographic processes that structure the accessibility and

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73 availability of these lower trophic level food sources (Cox *et al.*, 2013; Bertrand *et al.*, 2014;
74 Woodson and Litvin, 2015; McInnes *et al.*, 2017), and thus influence foraging efficiency/success
75 (Boyd *et al.*, 2016). Across shelf-sea environments, a number of discrete habitat features have been
76 identified as important locations that host enhanced foraging opportunities, including fronts, offshore
77 banks where internal waves propagate and tidally dependent island wakes (Hunt and Schneider, 1987;
78 Hunt *et al.*, 1999; Bost *et al.*, 2009; Bertrand *et al.*, 2014; Scales *et al.*, 2014a; Benjamins *et al.*, 2015).
79 Given the tight energy constraints of many marine mammals and seabirds (Cairns, 1988), these
80 structures can be thought of as critical habitat features.

81 Knowledge of the bio-physical processes that underlie links between oceanographic habitat features 82 and marine mammal and seabird foraging is vital to obtaining a comprehensive understanding of 83 marine ecosystem functioning. This will prove invaluable as we move towards implementing holistic management methods, that consider ecosystems in their entirety and aim to incorporate more 84 precautionary conservation measures (Arkema et al., 2006; Hooker et al., 2011). Early synthesises 85 86 and reviews have outlined the prominent bio-physical processes occurring across ocean environments and how these are linked to the spatio-temporal distributions of seabirds (e.g. Hunt, 1990, 1991, 1997; 87 88 Hunt et al., 1999), but no known equivalent review exists for marine mammals. Over the past 10-15 years, methodological and technological advances have substantially improved the way in which the 89 90 marine environment is studied, both in terms of how we collect data (Cooke et al., 2004; Hunt and Wilson, 2012; Brown et al., 2013; Waggitt and Scott, 2014; Photopoulou et al., 2015; Benoit-Bird 91 92 and Lawson, 2016; Macaulay et al., 2017) and quantitatively analyse it (Redfern et al., 2006; Wakefield et al., 2009; Brown et al., 2013; Carter et al., 2016; Bennison et al., 2017). As such, our 93 knowledge of links between oceanographic habitats and marine predators has substantially improved. 94 and a more mechanistic understanding of how these features aid marine mammal and seabird 95 foraging, and function as dynamic habitats is being attained. Although, more recent regional and/or 96 97 feature specific reviews have been published (e.g. fronts; Scales et al., 2014a, southern ocean fronts

98 and eddies; Acha *et al.*, 2004; Bost *et al.*, 2009, and tidal stream environments; Benjamins *et al.*, 99 2015), as of yet, a paper which encompasses links between both marine mammals and seabirds with 100 the full range of oceanographic habitat features that occur across mid-latitude, seasonally stratifying 101 shelf-seas as a whole is lacking. Such knowledge is still yet to be fully integrated into studies 102 examining behavioural patterns and habitat use by marine predators alongside marine management 103 strategies (Tremblay *et al.*, 2009; Fourcade *et al.*, 2018), and so it is particularly pertinent that this is 104 addressed, not least because the diverse human use of shelf-seas is accelerating impacts on these 105 ecosystems. We feel it important to synthesise the broad and diverse habitat features present in this 106 environment to facilitate conservation management and ecological research, and drive a policy 107 response to this crisis.

108 In light of this, we provide a concise and easily accessible guide for both researchers and managers 109 of marine systems on the predominant oceanographic habitats that are favoured for foraging by 110 marine mammals and seabirds across mid-latitude shelf seas. We identify and describe key discrete 111 physical features present across the continental shelf, working inshore from the shelf-edge to the 112 shore line (Figures 1 & 2, Table 1), and discuss links to marine mammals and seabirds. We then 113 highlight the key characteristics of these features that make them attractive as foraging habitats. We 114 identify areas where knowledge is lacking and make recommendations for the direction of future 115 research. Finally, we discuss how these insights can be used to improve the conservation management 116 of shelf-sea environments. Our aim is to provide a concise overview, in a format that is broken down 117 into feature specific sections accessible to non-oceanographers. In doing so we hope to encourage 118 both researchers and conservation managers of marine systems to move towards the identification, 119 bio-physical characterisation and incorporation of discrete oceanographic habitat structures that 120 promote prey availability into future studies and management strategies.

## 121 2 The shelf edge

The shelf edge marks the transition zone from the comparatively shallow (usually less than 200m) waters of the continental shelf to the deep abyssal plains (where depths exceed 2000m) of the open ocean (Simpson and Sharples, 2012). This region is relatively narrow, with a typical lateral distance for around 50km, and so generally characterised by a steep sloping profile (e.g. the European continental shelf-edge). Along the shelf edge, dependent upon geographical location (e.g. bordering a major eastern boundary current), shelf-edge fronts and wind-driven upwelling fronts support high levels of primary and secondary productivity which attract a diversity of marine mammals and seabirds.

# 130 2.1 Shelf-edge fronts

131 Shelf-edge fronts (also shelf-break and shelf-slope fronts) occur at the interface between on-shelf and 132 open-ocean waters (Figure 1), and are marked by strong gradients in salinity, and sometimes 133 temperature. As currents, pushed onto the shelf via tidal forcing, are interrupted by the steep sloping topographic profiles of these regions, upwelling pushes the surface mixed layer above the critical 134 depth for phytoplankton growth (Fournier *et al.*, 1979), whilst simultaneously facilitating exchange 135 with the nutrient rich waters of the open-ocean's bottom boundary layer (Springer et al., 1996; Ryan 136 et al., 1999). High levels of primary productivity are typically sustained, sometimes perennially 137 (Fournier et al., 1979), attracting planktivorous grazers alongside large numbers of pelagic fish 138 139 (Podesta et al., 1993; Sabatés and Olivar, 1996; Springer et al., 1996; Genin, 2004; Greer et al., 2015). 140 Dependent upon the lateral extent and topography of the adjacent continental shelf, these features 141 may be far from land, and so relatively inaccessible to those foragers constrained to land-based 142 colonies (e.g. breeding seabirds and some seals). Links to marine predators are dominated by those 143 taxa capable of performing far-ranging foraging trips (e.g. black petrel *Procellaria parkinsoni*, fork-144 tailed storm petrel Oceanodroma furcate, northern fulmar Fulmarus glacialis and short-tailed 145 albatross Phoebastria albatrus; Schneider, 1982; Stone et al., 1995; Piatt et al., 2006; Freeman et al., 146 2010) or that are not restricted to a central location for breeding (e.g. Cuvier's beaked whale Ziphius 147 *cavirostris*, Risso's dolphin *Grampus griseus*, sperm whale *Physeter microcephalus* and spotted 148 dolphin *Stenella attenuata*; Baumgartner, 1997; Waring *et al.*, 2001; Pinedo *et al.*, 2002; Azzellino 149 *et al.*, 2008; Scott and Chivers, 2009). Where the shelf edge is nearer to land (e.g. the Skagerrak), 150 shorter ranging breeding species may forage at these features (e.g. little auk *Alle alle*; Skov and 151 Durinck, 1998). Shelf-edge fronts may also be important to non-breeding individuals in the late 152 summer, autumn and winter, when constraints to a central land-based breeding location no longer 153 apply (e.g. Cory's shearwater *Calonectris borealis* and little auk; Haney and McGillivary, 1985a; 154 Brown, 1988).

# 155 2.2 Wind-driven upwelling fronts

156 Along the major eastern boundary currents of western North America (the California current), Peru 157 (the Humboldt current) and west Africa (the Benguela, Canary and Somali currents), strong cross 158 winds in combination with Coriolis forcing and Ekman transport form intense upwelling systems, 159 which sustain some of the highest levels of primary and secondary productivity globally (Longhurst 160 *et al.*, 1995). In other regions, similarly structured upwelling systems may occur (e.g. the southern 161 shelf of Australia and along the eastern boundary of the Labrador Current; Kinsella *et al.*, 1987; 162 Kampf *et al.*, 2004), albeit on a smaller and less impressive spatio-temporal scale.

Along, or immediately inshore of the shelf-edge, upwelling fronts mark where these systems meet on-shelf coastal waters. Strong convergent flows accumulate and retain the phytoplankton biomass and small nekton generated by adjacent upwelling systems (Bjorkstedt *et al.*, 2002), which attracts lef large numbers of pelagic and forage fish (Ainley *et al.*, 2005; Reese *et al.*, 2011; Watson *et al.*, 2018). Due to the typically narrow extent of adjacent shelves (e.g. western Africa and western America's), upwelling fronts are often proximate to land. As such, the prey aggregating effects of these features are exploited by a diverse range of marine predators (Bourne and Clark, 1984; Forney and Barlow, 1998; Hoefer, 2000; Camphuysen and van der Meer, 2005; Croll *et al.*, 2005; Tynan *et al.*, 2005; 171 Ainley *et al.*, 2009) that includes those individuals constrained to land-based colonies or a shoreward
172 distribution (e.g. lactating New Zealand fur seals *Arctocephalus forsteri* and Northern elephant seals
173 *Mirounga angustirostris* alongside numerous breeding seabirds such as Cape gannet *Morus capensis*,
174 common guillemot *Uria aalge*, Humboldt penguin *Spheniscus humboldti*, kelp gull *Larus*175 *dominicanus*, Peruvian booby *Sula variegata*, rhinoceros auklet *Cerorhinca monocerata* and a
176 number of phalarope species; Briggs *et al.*, 1984; Croll, 1990; Weichler *et al.*, 2004; Ainley *et al.*,
177 2005; Crocker *et al.*, 2006; Baylis *et al.*, 2008; Sabarros *et al.*, 2014).

178 The intensities of upwelling systems and their associated fronts can vary seasonally and/or inter-179 annually with climatic conditions and wind patterns (Kinsella *et al.*, 1987; Bograd *et al.*, 2009a), 180 which can substantially impact the structuring of surrounding ecosystems with concomitant 181 consequences for marine mammals and seabirds (Schneider and Methven, 1988; Schneider, 1994; 182 McGowan *et al.*, 1998; Abraham and Sydeman, 2004; Wolf *et al.*, 2009; Black *et al.*, 2011; Woodson 183 and Litvin, 2015). For example, in years when decreased upwelling intensity reduces the availability 184 of high quality foraging habitats around frontal zones (e.g. with the El Nino Southern Oscillation; 185 ENSO), breeding seabirds along the west coast of the America's display signs of reduced body 186 condition (e.g. common guillemots; Croll, 1990), whilst others (e.g. Cassin's Auklet *Ptychoramphus* 187 *aleuticus*, Humboldt penguins and marbled murrelet *Brachyramphus marmoratus*) extend their 188 foraging trips, which may result in reduced reproductive success (Becker and Beissinger, 2003; 189 Hennicke and Culik, 2005; Bertram *et al.*, 2017).

#### 190 3 The mid-shelf: from the shelf-edge to near-shore coastal waters

191 The mid-shelf extends from the shelf-edge to near-shore coastal waters (Figure 1) with topographic 192 depths typically ranging from around 50m to 200m. In mid-latitude, temperate zones this region 193 stratifies seasonally between late spring and autumn when increased solar irradiation heats surface 194 waters sufficiently to increase buoyancy levels and overcome tidal and wind-driven mixing (Pingree, 195 1975; Pingree *et al.*, 1976; Simpson and Sharples, 2012). A two-layer system is formed, characterised
196 by a surface mixed layer of warm, nutrient deficient water and a bottom boundary layer of dense,
197 cold nutrient rich water (Figure 1). This structuring underlies a number of oceanographic processes
198 that appear important to marine mammals and seabirds, particularly in areas where the spatial extent
199 of the continental on-shelf region is large and tidal ranges considerable (e.g. the Canadian, European,
200 northeast USA continental shelf and the eastern Bering Sea Shelf).

#### 201 3.1 The annual spring bloom

202 The development/onset of stratification in the spring drives a significant annual phytoplankton bloom (Pingree et al., 1976; Sambrotto et al., 1986; Sharples et al., 2006). The timing of this bloom varies 203 204 annually as a result of climatic fluctuations in air temperature/solar irradiation and wind stress (Sharples *et al.*, 2006), which can lead to a temporal mismatch between fish spawning and plankton 205 206 production (match-mismatch hypothesis; Cushing, 1975). This can have bottom-up impacts at higher trophic levels by influencing fish recruitment (Beaugrand and Kirby, 2010; Sigler et al., 2016) and 207 208 food availability (Durant et al., 2007), and has been shown to effect the breeding success of a number 209 of seabirds including Atlantic puffin Fratercula arctica, black-legged kittiwake Rissa tridactyla, 210 common guillemot and rhinoceros auklet off the coast of British Columbia, and across the North and 211 Norwegian Seas (Durant et al., 2006; Scott et al., 2006; Borstad et al., 2011). Such impacts 212 sometimes occur at a lag of 1-2 years to underlying shifts in environmental conditions (Zador et al., 213 2013). Changes in prey availability have also been linked to spatial variability in the distribution of 214 the spring bloom alongside the oceanographic conditions within which it occurs (Table 2). For 215 example, across the eastern Bering Sea, in years when the spring bloom occurs in warmer offshore 216 waters (due to earlier sea ice retreat; Hunt and Stabeno, 2002), changes in the abundance, 217 composition, distribution and survival of predominant plankton and juvenile fish species results in 218 shifts in the diets and distributions of several seabird populations (Springer et al., 2007; Renner et al.,

219 2016; Hunt *et al.*, 2018), which can result in demographic impacts (Satterthwaite *et al.*, 2012; Zador 220 *et al.*, 2013).

# 221 3.2 Vertical interfaces in offshore stratified regions (the pycnocline) and sub-surface productivity

222 Following the spring bloom, productivity within the mid-shelf region redistributes and is maintained 223 by a number of oceanographic processes (Richardson et al., 2000; Weston et al., 2005). An important 224 feature is the vertical interface between the low nutrient, warm surface mixed layer and the cool, 225 dense, high nutrient bottom boundary layer, where steep vertical gradients in density form a 226 pycnocline, which can alternatively be referred to as the thermocline (when vertical density gradients are temperature driven) or the halocline (when vertical density gradients are driven by changes in 227 228 salinity). This structure acts as a barrier between surface and bottom boundary waters by inhibiting 229 the vertical transport of nutrients and plankton (Stepputtis et al., 2011). In some instances, the 230 pcynocline may be composed of both a thermocline and halocline. Alternatively, the effects of 231 vertical changes in temperature and salinity can cancel each other out, resulting in no pcynocline. 232 Across shelf-seas, offshore seasonal summer stratification is predominantly thermally driven (with a 233 temperature driven pycnocline; Simpson and Sharples, 2012), although in regions subject to high 234 levels of freshwater input (e.g. the Skagerrak between the North and Baltic Seas; Skov and Durinck, 235 2000) saline gradients may also be important. The majority of studies investigating interactions 236 between marine predators and the pycnocline have focused on links with temperature delineated 237 thermoclines (e.g. Takahashi et al., 2008; Kokubun et al., 2010; Pelletier et al., 2012; Nordstrom et 238 al., 2013; ven Eeden et al., 2016).

High levels of sub-surface primary productivity often concentrate around the pycnocline, and can account for over 50% of water column productivity (Weston *et al.*, 2005), alongside ~30% of total annual productivity (Richardson and Christoffersen, 1991). This is maintained through the summer works by two sources of episodic mixing events, each of which results in an influx of nutrients from

the bottom boundary layer that enhance productivity. First, increased tidal currents during spring
tides generate turbulent dissipation (due to friction with the sea bed), that may extend up through the
bottom boundary layer temporarily eroding the base of the pycnocline (Sharples, 1999, 2008;
Sharples *et al.*, 2001; Allen *et al.*, 2004). Second, wave and wind driven surface mixing (due to
changes in wave/wind direction and/or velocity with prevailing weather conditions) may partially
break down vertical stratification (Sharples and Tett, 1994; Rippeth *et al.*, 2005; Williams *et al.*,
additionally aid in the retention of small organisms such as phytoplankton (Franks, 1995; Durham *et al.*, 2009; Cheriton *et al.*, 2010), whilst a synchronous accumulation of zooplankton (McManus *et al.*,
2005) can result in a propagation of food supply across multiple trophic levels.

Sub-surface productivity at and around the pycnocline has been linked to foraging by a number of marine predators, such as little auk, northern fulmar, northern gannet *Morus bassanus* and grey seal *Halichoerus grypus* (Skov and Durinck, 2000; Scott *et al.*, 2010). In diving species, individuals may repetitively descend to the pycnocline (e.g. African penguin *Spheniscus demersus*, northern fur seal *Callorhinus ursinus*, northern right whale *Eubalaena glacialis*, rhinoceros auklet and thick-billed murre *Uria lomvia*; Baumgartner and Mate, 2003; Matsumoto *et al.*, 2008; Takahashi *et al.*, 2008; Kuhn, 2011; ven Eeden *et al.*, 2016), where peaks in prey density (Hansen *et al.*, 2001; Baumgartner and Mate, 2003) increase foraging efficiency (Pelletier *et al.*, 2012). In years when a pycnocline is absent or highly dispersed, foraging success tends to decrease, with concomitant consequences for seabird breeding success (Ropert-Coudert *et al.*, 2009a). Alternatively, individuals (e.g. thick-billed murres in the southeastern Bering Sea) may expand the range of habitats foraged at (both in terms of horizontal extent and dive depth; Kokubun *et al.*, 2010).

265 The depth of the pycnocline is also an important determinant of foraging habitat suitability (Hunt *et al.*, 1993; Skov and Durinck, 2000; Nordstrom *et al.*, 2013). Increased light attenuation with depth
267 means productivity around shallower pcynoclines is likely enhanced compared with deeper 11

peynoclines (Skov and Durinck, 2000). Moreover, for near-surface feeders (e.g. northern fulmar) a shallower pycnocline may make prey available at more accessible depths (Skov and Durinck, 2000), and for those that dive from the surface (e.g. least auklet *Aethia pusilla*), reduce foraging energetic costs (Hunt *et al.*, 1990; Haney, 1991; Skov and Durinck, 2000; Langton *et al.*, 2011). In thermally stratified waters, exothermic prey may redistribute themselves near the surface in an attempt to avoid unfavourable cool bottom boundary waters below the pycnoline (e.g. mackerel *Scomber scombrus*; Grégoire, 2006).

# 275 3.3 Internal waves and offshore banks

276 Internal waves form within stratified regions when tidal currents are interrupted by areas of abrupt, 277 raised and/or uneven topography (Figure 1; Mann and Lazier, 2006), and commonly occur in proximity to the shelf edge (Bertrand et al., 2014) and around offshore banks (Palmer et al., 2013), 278 279 reefs and rock pinnacles (Moum and Nash, 2000). Resultant locally induced upwelling causes an 280 oscillation in the pycnocline that can exceed an amplitude of 30m and approach ~50% of local water 281 depth (Witman et al., 1993; Palmer et al., 2013). Nutrient fluxes across the pycnocline sustain 282 exceptionally high levels of sub-surface productivity (Richardson *et al.*, 2000; Tweddle *et al.*, 2013), 283 whilst the simultaneous creation of a number of convergent (aggregating) and divergent (dispersing) 284 zones (Figure 1) can alter the vertical distributions of plankton and small nekton (Lennert-Cody and 285 Franks, 1999; McManus et al., 2005; Bertrand et al., 2008), forcing large aggregations of prey to the surface (Embling *et al.*, 2013) that are foraged at by a range of marine predators (Moore and Lien, 286 2007; Stevick et al., 2008; Scott et al., 2013; Bertrand et al., 2014). These features appear to be 287 especially important to those taxa that near-surface feed on plankton and/or forage fish such as black-288 legged kittiwake, humpback whale Megaptera novaeangliae and several species of petrel and 289 290 shearwater (Haney, 1987; Stevick et al., 2008; Hazen et al., 2009; Embling et al., 2012). The 291 generation of internal waves is tidally mediated (Pineda et al., 2015), and patterns in the occurrence 292 of surface prey aggregations alongside marine mammal and seabird foraging regularly reflect this 293 (Hazen *et al.*, 2009; Embling *et al.*, 2012). Further complexities in the shape of a topographic 294 structure (e.g. steep-sided crests and mounts) may enhance the density of surface aggregations by 295 concentrating tidal currents whilst simultaneously moderating the passage of internal waves to 296 increase upwelling flows and surface convergence (Stevick *et al.*, 2008).

#### 297 4 Tidal-mixing fronts

Tidal-mixing fronts mark the transition zones between the seasonally stratifying waters of the midshelf and mixing coastal waters (Simpson and Hunter, 1974; Pingree and Griffiths, 1978; Schumacher *et al.*, 1979), and are marked by strong horizontal surface temperature gradients (Miller, 2009). Their positions are dependent upon the ability of tidal currents and wind stress to overcome the buoyant effects of surface heat fluxes and mix the entire water column (Fearnhead, 1975; Franks, 1992a; Acha *et al.*, 2004). This is a function of water column depth, and so the position of a tidal-mixing front can be roughly predicted from the ratio of total water depth (*h*) to tidal velocity (*u*) -  $h/u^3$  (the Simpson-Hunter parameter; Simpson and Hunter, 1974; Simpson and Sharples, 2012). Once established, variation in a tidal-mixing front's position occurs (Figure 2, Table 2), in response to changes in the strength of tidal currents with the spring-neap cycle (Sharples and Simpson, 1996; Simpson and Sharples, 2012), as well as from variation in heat flux and wind-driven mixing (Kachel *et al.*, 2002; Nahas *et al.*, 2005; Pisoni *et al.*, 2015).

Tidal-mixing fronts are often associated with elevated and persistent primary productivity that has the potential to propagate across multiple trophic levels (Coyle and Cooney, 1993; Munk *et al.*, 1995; Gregory Lough and Manning, 2001). Where the pycnocline of stratified offshore waters shallows to meet inshore mixing waters, increased light exposure supplemented with runoff nutrients (of coastal waters) alongside those mixed up from the bottom boundary layer results in productivity levels several orders of magnitude higher than in surrounding waters (Pingree *et al.*, 1975; Simpson *et al.*, 1979; Franks, 1992a). Additional convergent flows (Pingree *et al.*, 1974) may redistribute the 317 horizontal and vertical distributions of weak or passively swimming organisms (e.g. plankton grazers
318 attracted to the high productivity levels of the front; Coyle *et al.*, 1998), resulting in near-surface
319 retention and accumulation (Franks, 1992b; Epstein and Beardsley, 2001).

320 A diverse range of marine predators forage around tidal-mixing fronts (Haney and McGillivary, 321 1985b; Begg and Reid, 1997; Goold, 1998; Hunt et al., 1999; Weir and O'Brien, 2000). Associations 322 are particularly prominent in colonial seabirds, and land-based breeding sites are often located in 323 proximity to these features (Hunt, 1997). Large numbers of near-surface feeding planktivores, such 324 as least auklet and short-tailed shearwater Puffinus tenuirostris, forage at and around tidal-mixing 325 fronts in concordance with patches of increased zooplankton abundance (Hunt *et al.*, 1996; Jahncke 326 et al., 2005), which are often concentrated near the sea's surface (Harrison et al., 1990; Hunt and 327 Harrison, 1990; Russell et al., 1999). These features may also attract large cetacean species including 328 a number of lunge-feeding rorquals (e.g. blue whale *Balaenoptera musculus*, fin whale *Balaenoptera* 329 physalus and humpback whale; Doniol-Valcroze et al., 2007; Dalla Rosa et al., 2012). Piscivores, 330 such as black-legged kittiwake, common dolphin Delphinus delphis, common guillemot, Magellanic 331 penguin Spheniscus magellanicus and northern gannet, also frequently forage at tidal-mixing fronts 332 (Kinder et al., 1983; Durazo et al., 1998; Goold, 1998; Boersma et al., 2009; Scales et al., 2014b; 333 Cox et al., 2016, 2017), likely because the aggregating effects of these features on plankton 334 predictably attract high densities of forage and pelagic fish (Hansen et al., 2001; Alemany et al., 335 2009; Brigolin et al., 2018). Across the southeastern Bering Sea, individuals present at these features 336 have been directly linked to high density patches of fish prey (Decker and Hunt, 1996; Kokubun et 337 al., 2008), where capture rates were increased (Vlietstra et al., 2005). For both piscivores and 338 planktivores, fronts with strong surface flow gradients may be particularly attractive, possibly due to 339 additional aggregative effects on small biomass (Schneider *et al.*, 1987). Reduced productivity at 340 frontal zones alongside geographical shifts in typical locations (in years of abnormal oceanographic 341 conditions; e.g. ENSO events) have been linked to increases in short-tailed shearwater mortality rates342 (Napp and Hunt, 2001).

## 343 5 Near-shore coastal waters and estuaries

On the shoreward side of a tidal-mixing front, turbulence generated through friction between tidal currents and the seabed extends the entire water column (due to shallow depths), and prevents thermal stratification (Simpson and Sharples, 2012). As such, the majority of this region remains permanently mixed throughout the year. Concentrated patches of primary productivity are generally limited to regions of fresh water influence (ROFIs) around estuarine systems, where salinity driven stratification may occur. Other notable areas of interest to marine mammals and seabirds, such as those associated with tidally driven turbulence around topographic structures, likely function by mechanically altering the behaviours and distributions of zooplankton and fish prey, as indicated by periodicity in their use (Zamon, 2002, 2003). Unlike the mid-shelf, where the seasonal development of thermal stratification splays a dominant role in the formation of foraging habitat, features occurring in near-shore coastal waters may persist throughout the year and, in some cases, are targeted perennially by marine predators (Skov and Prins, 2001).

# 356 5.1 Channels, headland and island wakes, nearshore reefs and bays

In near-shore coastal regions, marine mammals and seabirds frequently forage within tidally active areas (Nol and Gaskin, 1987; Marubini *et al.*, 2009; Anderwald *et al.*, 2012; Benjamins *et al.*, 2015; Warwick-Evans *et al.*, 2016; Waggitt *et al.*, 2018), resulting in distinct regularities in their distributions and behaviours that coincide with particular tidal phases (Becker *et al.*, 1993; Hunt *et al.*, 1998; Irons, 1998; Isojunno *et al.*, 2012; De Boer *et al.*, 2014; Ijsseldijk *et al.*, 2015). Specifically, areas such as narrow channels, headlands, islands, reefs and bays often function as periodic foraging hotspots, where interactions between strong tidal currents (often exceeding  $1.5\text{ms}^{-1}$ ) and complex topography create prosperous foraging opportunities for marine predators (Cairns and Schneider, 365 1990; Coyle *et al.*, 1992; Zamon, 2003; Benjamins *et al.*, 2015; Couperus *et al.*, 2016). In some 366 instances, several of these features may occur in close proximity to one another (Bailey and 367 Thompson, 2010).

#### 368 5.1.1 Channels and narrow passes

369 Channels, corridors and narrow passes are typical of estuaries, fjords and island groups. During 370 strong tidal flows, these features may act as bottlenecks creating predictable and exploitable 371 concentrations of zooplankton and fish prey advected from adjacent areas (Zamon, 2001, 2002; 372 Couperus *et al.*, 2016), which are exploited by a number of marine mammals and seabirds (Thompson et al., 1991; Lescrauwaet et al., 2000; Zamon, 2001; Holm and Burger, 2002; Hastie et al., 2004, 373 374 2016; Ladd et al., 2005; Bailey and Thompson, 2010). Rapid currents and turbulence along these passages (Nimmo Smith et al., 1999) may additionally disorientate fish and break down shoal 375 376 cohesion (Liao, 2007), increasing catchability (Zamon, 2001, 2003; Crook and Davoren, 2014). The 377 fine scale foraging distributions of several piscivorous alcids, cetaceans and pinnipeds (e.g. Atlantic puffin, common guillemot, harbour porpoise *Phocoena phocoena* and harbour seal *Phoca vitulina*) 378 across channels and narrow passes are known to concentrate in central and/or narrow areas where 379 380 current flows are maximised (Pierpoint, 2008; Hastie et al., 2016; Waggitt et al., 2016a). However, 381 it should be noted that due to the energetics of navigating turbulent flows (Wilson *et al.*, 2001; Heath 382 and Gilchrist, 2010), some individuals may forage around the periphery of these currents where 383 speeds are reduced (Pierpoint, 2008; Wilson et al., 2013; Cole et al., 2018), or avoid such areas 384 completely (Wilson et al., 2001; Embling et al., 2010; Waggitt et al., 2016b).

In some cases, the steep sides of a channel, corridor or narrow pass may additionally provide a barrier suitable for prey herding (Heimlich-Boran, 1988), and cetaceans that employ complex group foraging strategies (Simila and Ugarte, 1993; Fertl and Wilson, 1997; Duffy-Echevarria *et al.*, 2008) are strategies (because the second strategies (Heimlich-Boran, 1988; Hastie *et al.*, 2004; Bailey and 389 Thompson, 2010). Where the edge of a channel or pass acts as a barrier to tidal currents (e.g. in the 390 presence of shallow banks and/or meanders), resultant upwelling and current circulation can force 391 advected zooplankton, and sometimes small nekton (e.g. forage fish), into dense aggregations 392 (Simard et al., 1986; Lavoie et al., 2000; Davies et al., 2013) that may be driven towards the surface 393 (Simard et al., 2002), making these features additionally attractive to both bulk-feeding baleen whales 394 (Cotté and Simard, 2005) and surface feeding birds (e.g. a number of auklet species alongside 395 Bonaparte's Larus Philadelphia and Mew Larus canus gull; Vermeer et al., 1987; Hunt et al., 1998). 396 In some instances, ephemeral features, known as Langmuir circulation cells, form convergent zones 397 at the sea's surface (over scales of a few to a couple of hundred metres; Barstow, 1983). These can 398 further entrain plankton and small nekton (Hamner and Schneider, 1986), which near-surface/surface 399 foraging seabirds (e.g. northern fulmars and prions *Pachyptila spp*) have been observed exploiting 400 (Goss et al., 1997; Ladd et al., 2005). Where fast currents pour down into a channel, or where 401 previous upwelled waters descend (Hunt et al., 1998), downwelling structures may form (Hunt et al., 402 1998; Waggitt *et al.*, 2016a). Whilst these features have been linked to the foraging distributions of 403 two benthic/demersal feeders (black guillemot *Cepphus grille* and European shag *Phalacrocorax* 404 aristotelis; Waggitt et al., 2016a) alongside one upper-water column feeder (least auklet; Hunt et al., 405 1998), the exact mechanisms driving these interactions are unclear, but appear to be site and species 406 specific (Hunt et al., 1998; Waggitt et al., 2017).

# 407 5.1.2 Headland and island wakes

When headland and island features interrupt the passage of strong tidal current flows, a leeward wake
(or eddy) may form. At the interface with non-wake waters, shear induced hydrographic fronts
(Wolanski and Hamner, 1988; Johnston and Read, 2007) may accumulate zooplankton which become
retained within the calm waters of the wake (Alldredge and Hamner, 1980). In addition, turbulent
flows around these structures may disorientate fish prey that use the wake to forage, or seek refuge
from adjacent strong tidal currents (Liao, 2007; Robinson *et al.*, 2007; Tarrade *et al.*, 2008). For

414 example, in the Bay of Fundy (Canada), harbour porpoise, fin and minke *Balaenoptera acutorostrata* 415 whale exploit dense patches of euphasiids *Meganyctiphanes norvegica* and herring *Clupea harengus* 416 along the edge of an island wake during flood tides (Johnston *et al.*, 2005b, 2005a; Ingram *et al.*, 417 2007). Bottlenose dolphins in the Moray Firth (UK) concentrate foraging activity along the surface 418 signatures of hydrographic fronts, that form during specific tidal conditions in proximity to a headland 419 on the edge of a deep, steep-sided channel (Bailey and Thompson, 2010), while across Glacier Bay 420 and Icy Strait in southeastern Alaska, humpback whale distributions are disproportionately distributed 421 within the leeward waters of several headlands (Chenoweth *et al.*, 2011). Where fish actively avoid 422 turbulent flows (Nichol and Somerton, 2002), hydrographic fronts may act as a barrier to their 423 movements, and so some predators (e.g. bottlenose dolphin and killer whale) may use these features 424 for prey herding (Heimlich-Boran, 1988; Benjamins *et al.*, 2015). However, as of yet this has not 425 been directly explored.

# 426 5.1.3 Nearshore reefs, banks and ridges

Where current flows run across nearshore topographically complex structures such as reefs, banks and ridges, under specific tidal conditions friction can generate shear instabilities, turbulence and upwelling (Coyle *et al.*, 1992; Jones *et al.*, 2014). Peaks in common guillemot, harbour porpoise, phalaropes and thick-billed murre occurrence, corresponding to the locations and times at which these hydrographic features manifest, likely reflect changes in plankton and fish prey distributions that aid capture (Brown and Gaskin, 1986; Coyle *et al.*, 1992; Skov and Thomsen, 2008; Jones *et al.*, 2014). Further offshore where tidal currents are less pronounced (but still within boundaries of near-coastal hydrographic features such as reefs and ridges may also act to trap plankton and small biomass during downward migration, which may be particularly attractive to planktivorous species foraging upon euphausiids, such as the short-tailed shearwater (Hunt *et al.*, 1996).

437 5.1.4 Bays

Where the curvature of a headland or series of small islands results in the formation of a bay, complex tidal circulation patterns (enforced by the curved profile of the bay) may accumulate plankton and small nekton through advection and retention (Gomez-Gutierrez and Robinson, 2006; Rogachev *et al.*, 2008). In bays characterised by the presence of steep topographic barriers and ledges, interactions with these circulation patterns can generate localised upwelling. This can force accumulated biomass into dense surface aggregations, which appear to provide important foraging opportunities for a number of planktivorous species that either surface feed (e.g. black-legged kittiwake and red-necked phalarope *Phalaropus lobatus*; Drew *et al.*, 2013; Thorne and Read, 2013) and/or bulk feed (e.g. bowhead whale *Balaena mysticetus* and North Atlantic right whale; Jiang *et al.*, 2007; Rogachev *et al.*, 2008).

## 448 5.2 Regions of freshwater influence (ROFIs): Estuarine plume and tidal intrusion fronts

449 Within and around estuarine systems, typical circulation patterns, forcing dense water below less 450 dense water, promote the two-layer stratification of outflowing nutrient rich freshwater and intruding 451 saline waters (Simpson and Sharples, 2012). Where this stratification meets coastal mixing waters, 452 high horizontal gradients in salinity and density mark the position of either a tidal intrusion front 453 (dense saline coastal water intruding into the estuary) or a plume front (brackish water discharging 454 out of the estuary; Simpson and Nunes, 1981; Lewis, 1984). Stratification increases stability in the 455 water column, and allows plankton to redistribute and settle at or above the salinity driven pycnocline (halocline) where waters are nutrient rich and light exposure increased. Resultant productivity levels 456 are high (Cloern, 1991), and attract large numbers of zooplankton and forage fish (Govoni et al., 457 458 1989; Kingsford and Suthers, 1994; Kaltenberg et al., 2010; Phillips et al., 2017), which may be concentrated at the surface signature of the front due to additional convergent flows (Govoni *et al.*, 459 460 1989). A number of piscivores, such as black and red throated diver Gavia stellata/arctica, bottlenose 461 dolphin, common guillemot, little penguin, northern fur seal and sooty shearwater Ardenna grisea, 462 forage at estuarine plume and tidal intrusion fronts (Skov and Prins, 2001; Mendes et al., 2002; 463 Pelland et al., 2014; Zamon et al., 2014; Kowalczyk et al., 2015a; Phillips et al., 2017). The location, occurrence and strength of these features may be tidally mediated and/or dependent upon local current 464 465 patterns, rainfall and wind events (Sharples and Simpson, 1993; Kingsford and Suthers, 1994; Choi and Wilkin, 2007; Schlacher et al., 2008). Such variability may further impact the abundance and 466 distribution of zooplankton and forage fish, alongside the behaviours of marine mammals and 467 seabirds that forage at these feature (Mendes et al., 2002; Schlacher et al., 2008; Kowalczyk et al., 468 469 2015b; Lin et al., 2015; Phillips et al., 2018). For example, bottlenose dolphin foraging at a tidal 470 intrusion front in the Moray Firth (UK), did so only during flood tide conditions when the front was 471 most pronounced (Mendes et al., 2002), whilst sooty shearwater and common guillemots have been 472 shown to spatially track the boundary of the Columbia River plume (USA) as it evolves through time 473 (Phillips et al., 2018). In Port Phillip Bay (Australia), little penguin breeding success has been linked 474 to the occurrence and intensity of rainfall events, and subsequent changes in the dynamics of a local 475 estuarine plume front and prey availability (Kowalczyk et al., 2015b).

# 476 6 Oceanographic habitats and the facilitation of foraging

It is evident that bio-physical processes strongly influence habitat choice by a diversity of marine
mammals and seabirds that feed upon plankton and/or forage and pelagic fish, and across shelf-seas
a range of oceanographic features are favoured for foraging. Such associations appear attributable to
predictable increases in prey accessibility and availability that stem from changes in the abundance
and density, depth distribution, behaviour and/or patch persistence of prey, which together likely
facilitate trophic transfer and enhance foraging efficiency (Pelletier *et al.*, 2012; Boyd *et al.*, 2016).
For example, notable declines in the reproductive success of some marine predator populations have
mirrored a change/reduction in the availability of prey enhancing oceanographic habitats following
environmental changes with prevailing climatic and weather conditions (Hennicke and Culik, 2005;
Scott *et al.*, 2006; Boersma and Rebstock, 2009; Ropert-Coudert *et al.*, 2009a; Wolf *et al.*, 2009;
Borstad *et al.*, 2011). The importance of difference aspects of prey availability (e.g. prey abundance 20

versus depth distribution, versus predictability) will depend upon species specific foraging strategies
(e.g. surface feeder versus diver, available search time, differences in prey type; Hunt *et al.*, 1999;
Langton *et al.*, 2011), alongside individual energetic requirements (Hennicke and Culik, 2005;
Goldbogen *et al.*, 2011). For central place foragers constrained to land (e.g. many breeding seabirds
and seals), some features may be inaccessible due to their proximity and limitations on an individuals
foraging range (e.g. shelf-edge fronts; Thaxter *et al.*, 2012).

# 494 6.1 Increased abundance and density of prey

Sites where prey are abundant in dense aggregations allow individuals to reduce foraging effort while maintaining sufficient yields (Enstipp *et al.*, 2007; Benoit-Bird *et al.*, 2013; Goldbogen *et al.*, 2015; Thaxter *et al.*, 2016). Behavioural changes alongside increases in prey capture rates at a number of oceanographic features (Vlietstra *et al.*, 2005; Rogachev *et al.*, 2008; Cox *et al.*, 2016) reflect the presence of abundant and densely concentrated prey (Decker and Hunt, 1996; Vlietstra *et al.*, 2005; Stevick *et al.*, 2008), suggesting these habitats can aid individuals in maximising foraging efficiency (Ropert-Coudert *et al.*, 2009a; Goldbogen *et al.*, 2011; Pelletier *et al.*, 2012). This typically occurs in tandem with increases in primary and secondary productivity (e.g. around shelf-edge fronts, upwelling fronts and tidal-mixing fronts; Decker and Hunt, 1996; Ainley *et al.*, 2005; Jahncke *et al.*, 2005) or specific flow characteristics (e.g. convergent zones at offshore banks subject to the passage of internal waves; Embling *et al.*, 2012, 2013), and may be particularly important for those individuals with especially high energetic needs (e.g. chick provisioning seabirds; Hennicke and Culik, 2005) and/or whose foraging strategies are particularly costly (Green *et al.*, 2009; Goldbogen *et al.*, 2011).

# 508 6.2 Depth distribution of prey

509 Depth distribution plays a key role in prey accessibility, particularly for those taxa that feed at or near 510 the surface and/or have limited dive capabilities (Embling *et al.*, 2012; Boyd *et al.*, 2015). Surface 511 convergent zones at shelf-edge fronts, upwelling fronts and tidal-mixing fronts are frequently used 512 by these foragers, as are localised upwelling structures related to interactions between topography 513 and tidal currents (e.g. at offshore banks and around the coast). In such areas, individuals have been 514 directly linked to shallow prey aggregations (Russell *et al.*, 1999; Stevick *et al.*, 2008; Embling *et al.*, 515 2012). For air-breathing diving predators, changes in the depth distributions of their prey may reduce 516 the energetic costs of capture by either allowing individuals to concentrate search activity within a 517 restricted proportion of the water column, or reducing overall dive depths from the surface (Ropert-518 Coudert *et al.*, 2009b; Benoit-Bird *et al.*, 2011). For example, the foraging efficiencies of some diving 519 seabirds appear to be tied to the presence of vertical prey aggregating features such as the pycnocline 520 (Ropert-Coudert *et al.*, 2009a; Pelletier *et al.*, 2012).

## 521 6.3 Prey behaviour

522 Bio-physical conditions and processes may elicit a change in prey behaviour that further alters 523 densities and depth distributions to increase vulnerability to predation by marine mammals and 524 seabirds. The breakdown of fish school cohesion likely makes individual prey items easier to catch 525 (Crook and Davoren, 2014; Kilian *et al.*, 2015; Hastie *et al.*, 2016), as may the formation of prey 526 concentrations in areas of reduced current speeds (e.g. around island wakes) as fish attempt to avoid 527 such disruptions (Liao, 2007; Robinson *et al.*, 2007; Benjamins *et al.*, 2015). Alternatively, the 528 temperature preferences of some prey mean they may redistribute themselves to aggregate in warmer, 529 near surface waters (Grégoire, 2006).

#### 530 6.4 Persistence and predictability

The oceanographic features favoured for foraging by marine mammals and seabirds typically occur in a persistent and/or predictable manner. If individuals can learn and remember the locations at which encountering accessible prey is more probable (Hunt *et al.*, 1999; Gende and Sigler, 2006; Weimerskirch, 2007; Davoren, 2013; Regular *et al.*, 2013; Grecian *et al.*, 2018), concentrated search for (Hamer *et al.*, 2009; Pettex *et al.*, 2010; Dragon *et al.*, 2012; Patrick *et al.*, 2014) can increase 536 foraging efficiency (Stephens and Krebs, 1986). Targeted search patterns have been shown to 537 coincide with the occurrence of a number of oceanographic habitats that are repetitively visited 538 (Bailey and Thompson, 2010; Sabarros *et al.*, 2014; Scales *et al.*, 2014b). Moreover, these behaviours 539 have been shown to develop as individuals mature, suggesting they are beneficial (Grecian *et al.*, 540 2018). In some cases, the scales across which targeted search effort occur are tied to those of 541 corresponding physical processes. For example, gannets foraging around tidal-mixing fronts restrict 542 search behaviours over scales of between 2km to 10km (Hamer *et al.*, 2009; Scales *et al.*, 2014b) 543 reflective of those over which the locations of these features vary with tidal- and wind- driven events 544 (Nahas *et al.*, 2005; Pisoni *et al.*, 2015). Bottlenose dolphins foraging around topographically 545 controlled tidal structures display highly localised search behaviours over 100's of metres, which 546 mirrors the similarly scaled predictability of these features (Bailey and Thompson, 2010).

# 547 7 Future research directions

548 Whilst associations between marine predators and bio-physical processes have been documented 549 across numerous studies, there are a number of areas in which improvements are still necessary. 550 Concurrent measurements of sub-surface oceanography, low- to mid- trophic level prey distributions 551 and marine mammal and seabird behaviours around many discrete physical features are lacking (e.g. 552 tidal stream environments), or limited to a specific set of locations and sites (e.g. the Bering Sea, 553 British Isles and Canadian Continental Shelf). Moreover, fine-scale three-dimensional measurements of marine mammal and seabird movements are rarely integrated, which would allow for estimates of 554 prey capture attempts and energetic expenditure to be calculated and used to evaluate foraging 555 effort/efficiency (Viviant et al., 2010; Watanabe and Takahashi, 2013; Richard et al., 2016). 556 Achieving this would greatly increase our understanding of how physical habitats impact prey 557 558 availability for marine predators and facilitate trophic transfer (Embling et al., 2012), whilst also 559 elucidating site and species specificity (Waggitt et al., 2017) alongside the drivers of spatio-temporal 560 variability in marine mammal and seabird distributions (Certain et al., 2007). Such efforts may also

561 reveal the importance of other, lesser studied structures. For example, to our knowledge, the 562 ecological significance of bottom fronts (which may be important to deeper diving predators) is yet 563 to be investigated (Hill et al., 2008). In addition, despite some evidence suggesting Langmuir 564 circulation cells are exploited by surface feeding seabirds in near-coastal regions (Goss et al., 1997; 565 Ladd et al., 2005), the importance of these features further offshore is yet to be explored (Barstow, 566 1983). Distinguishing the way in which individuals perceive their environment via knowledge 567 transfer (Machovsky-Capuska et al., 2014), learning and memory (Regular et al., 2013; Grecian et 568 al., 2018), sight (Bodey et al., 2014; Tremblay et al., 2014; Bairos-Novak et al., 2015) and smell 569 (Savoca and Nevitt, 2014) would also be beneficial, as would an increased knowledge of 570 immature/juvenile foraging behaviours and how these develop through time (de Grissac et al., 2017; 571 Votier et al., 2017; Grecian et al., 2018). Further inter-taxa research (e.g. marine mammals versus 572 seabirds, planktivores versus piscivores, and surface feeders versus divers) would compliment this, 573 and provide additional insight of the selection pressures that have shaped the evolution of the at-sea 574 behaviours of these animals, and driven the necessary adaptations required for foraging in dynamic 575 waters. Such investigations may also be useful for assessing the potential of these taxa to adapt in 576 response to climate change. Finally, there is a distinct lack of studies determining the importance of oceanographic habitat features at a population level, which should be addressed. Future research 577 578 should aim to fill these gaps if we are to improve our understanding of habitat use by marine mammals However, achieving this will require novel methodological 579 and seabirds across shelf-seas. 580 techniques. Adaptive survey designs (Embling et al., 2012; Suberg et al., 2014; Waggitt and Scott, 581 2014; Waggitt et al., 2016a; Benjamins et al., 2017) that incorporate active and passive acoustics 582 (Williamson et al., 2015; Benoit-Bird and Lawson, 2016; Macaulay et al., 2017; Malinka et al., 2018) 583 alongside underwater videography (Machovsky-Capuska et al., 2011; Crook and Davoren, 2014) may 584 prove particularly useful, as will animal borne biologging via the attachment of accelerometers 585 (Viviant et al., 2010; Watanabe and Takahashi, 2013), cameras (Votier et al., 2013; Watanabe and

586 Takahashi, 2013; Tremblay et al., 2014), GPS loggers (Yoda et al., 2014), oceanographic sensors (Charrassin *et al.*, 2008) and satellite relay systems (e.g. the Argos satellite system; Photopoulou *et* 587 588 al., 2015; CLS, 2016; Cox et al., 2018). Moreover, outputs from remote-sensing and oceanographic 589 modelling can be used to initially identify discrete features of interest (Scales et al., 2014a; Waggitt 590 et al., 2016a), and/or provide data over extended areas/time-spans/retrospectively. This may require 591 novel processing and analytical routines, that can characterise and link bio-physical processes to 592 information on animal behaviour and movement (d'Ovidio et al., 2004; Miller, 2009; Embling et al., 593 2012; Boyd et al., 2014; Bayle et al., 2015; Mattei et al., 2018; Pirotta et al., 2018), with particular 594 attention paid towards the spatio-temporal scales of investigations (Figure 2, Table 1; Mannocci et 595 al., 2017; Scales et al., 2017). Ideally, multiple approaches should be combined and integrated with 596 demographic data, which would yield an overview of ecosystem dynamics unprecedented in detail (Boyd et al., 2015; Barbraud et al., 2017; Carroll et al., 2017), and can later be used to force individual 597 and population based models to determine the adaptability of these environments to future change 598 (e.g. anthropogenic or climatically driven; Boyd et al., 2016; Barbraud et al., 2017). 599

# 600 8 Applications to conservation management

Over the last 20-30 years, marine management and policy has started to evolve from single species based protocols and strategies (typically tailored to a specific fishery/sector), to more holistic methods that consider ecosystems in their entirety, and incorporate more precautionary conservation measures (e.g. the European Marine Strategy Framework Directive 2008/56/EC; Pikitch *et al.*, 2004; Arkema *et al.*, 2006; Curtin and Prellezo, 2010). Key to this is the preservation of core shelf-sea habitats, such as those oceanographic features identified here as favourable for foraging by marine mammals and seabirds (Hooker and Gerber, 2004; Taylor *et al.*, 2005; Heithaus *et al.*, 2008; Game *et al.*, 2009; Dickey-Collas *et al.*, 2017; Sherley *et al.*, 2017). However, achieving adequate protection is complicated by the diverse nature of such structures, which occur over a range of spatio-temporal scales (Figure 2 and Table 1), and function via a variety of bio-physical mechanisms that may impact 611 prey availability in different ways (e.g. through inducing changes in depth distribution, persistent, 612 predictability, abundance and/or behaviour; see section 6). A sophisticated understanding of these 613 dynamics can be used to effectively implement appropriate and customised conservation management 614 strategies (Authier *et al.*, 2017).

615 For example, initial implementations of holistic management have focused on fixed area based 616 protection through the designation of marine protected areas (MPAs; Hyrenbach, 2000; Hooker and 617 Gerber, 2004). Here, the aim is to reduce overlap with spatially explicit threats (e.g. commercial 618 fisheries, marine renewables and maritime traffic) that may cause direct mortality and/or disrupt 619 accessibility to favoured habitats (e.g. through prey depletion or displacement; Pichegru *et al.*, 2010; 620 Gormley *et al.*, 2012). However, while persistent and predictable bathymetrically tied tidal features 621 (e.g. offshore banks, channels and island wakes) may be particularly well suited to such measures, to 622 accommodate individuals relying on habitats that are variable through space and time (e.g. upwelling and tidal-mixing fronts), adaptive approaches are required. A recent solution to this is dynamic ocean 623 624 management (DOM), defined as management that is adjustable through space and time in response 625 to the shifting nature of the ocean and its users (Hobday et al., 2014; Lewison et al., 2015; Maxwell 626 et al., 2015). Such methods can thus allow for geographical changes in habitat locations with, for example, prevailing environmental conditions (e.g. position in spring-neap cycle alongside short term 627 628 weather events, seasonal trends and longer term climatic changes; Nahas et al., 2005; Bograd et al., 629 2009a; Pisoni et al., 2015). However, implementing DOM requires multidisciplinary and novel 630 monitoring approaches, so as boundaries are effectively designated and modified (Hazen *et al.*, 2018). 631 Moreover, attention needs to be paid towards the spatio-temporal scales over which data is acquired 632 and protection implemented, which should reflect the characteristics and spatio-temporal variability 633 of an oceanographic feature (Figure 2, Table 2; Mannocci et al., 2017; Scales et al., 2017). Satellite 634 remote-sensing can aid in this by allowing oceanographic features to be tracked both instantaneously 635 in near real time and over longer periods, although this is only applicable where a distinct surface 636 signature is presented (Miller, 2009; Scales *et al.*, 2014a). Sensors deployed on diving animals 637 autonomous/unmanned survey vehicles and/or moorings may compliment such information 638 (Charrassin *et al.*, 2008; Suberg *et al.*, 2014; Photopoulou *et al.*, 2015), as can outputs from 639 oceanographic and statistical modelling (Brodie *et al.*, 2018; Mattei *et al.*, 2018). Where stationary 640 management is preferred, long-term time series data may be used to develop buffer zones that extend 641 the boundaries of a fixed MPA, such that fluctuations in the spatio-temporal occurrence of a habitat 642 feature is captured (e.g. the entire spring-neap/weather dependent range of a tidal-mixing front; Nahas 643 *et al.*, 2005; Grantham *et al.*, 2011; Pisoni *et al.*, 2015). Such strategies should still be evaluated at 644 regular intervals to assess potential geographical shifts in response to climate change (Queiros *et al.*, 645 2016).

646 A shift towards management strategies than consider habitat hydrology in addition to geography is 647 also beneficial to ensuring that the functioning of oceanographic features favoured for foraging is 648 maintained (Dickey-Collas et al., 2017). For example, anthropogenically generated structural 649 changes to the marine environment (e.g. marine renewable energy installations; MREIs) may alter 650 the bio-physical and spatio-temporal characteristics of oceanographic environments (e.g. 651 interruptions to near coastal current regimes may alter and/or manifest tidally driven features such as 652 hydrographic fronts, wakes and localised upwelling systems; Brostrom, 2008; Shields et al., 2011; 653 Benjamins et al., 2015; Fraser et al., 2018). This may cause shifts in the geographical locations at 654 which these structures occur and impact prey availability (e.g. depth distribution and predictability; 655 Becker and Beissinger, 2003), and thus should be carefully considered during the planning stages of development projects. Impact assessments at sites where oceanographic features favoured for 656 657 foraging are present need to consider how variability in the dynamics of such structures influence 658 patterns in area use by marine mammals and seabirds, and thus the validity of associated evaluations 659 (Benjamins et al., 2017; Cox et al., 2017). Climate change impacts will likely also alter the 660 functioning of these habitats. Increases in the frequency and intensity of extreme climatic (e.g. ENSO

661 associated changes in upwelling; Cai *et al.*, 2014) and weather (e.g. storm induced turbulent mixing; 662 Young et al., 2011) events may again, modify the dynamics and geographical locations of critical 663 oceanographic features (Hazen et al., 2013; Sherley et al., 2017), such as upwelling fronts, the pycnocline and tidal-mixing fronts (Ropert-Coudert et al., 2009a; Pisoni et al., 2015). For example, 664 decreases in the reproductive outputs of a number of seabirds have been linked to changes in the 665 666 availability of oceanographically generated foraging habitats following irregularities in prevailing 667 climatic and weather conditions (Hennicke and Culik, 2005; Durant et al., 2006; Scott et al., 2006; 668 Ropert-Coudert et al., 2009a). Minimising the impacts of this will require adaptive, innovative and 669 precautionary management strategies that minimise cumulative stressors (Field and Francis, 2006; 670 Lester et al., 2010; Sherley et al., 2017), particularly since our understanding of how ecosystems will 671 cope and respond to such alterations is largely unknown. Again, MPAs and DOM informed by 672 studies on the dynamics of oceanographic features favoured for foraging will likely play an important 673 role in this (Halpern et al., 2010; Grantham et al., 2011; Briscoe et al., 2016; Dickey-Collas et al., 674 2017).

#### 675 9 Conclusions

A range of oceanographic features are favoured for foraging by marine mammals and seabirds across shelf-seas in mid-latitude temperate zones. Whilst associations are diverse and variable in nature (both between sites and across species), in the majority of cases intricate interactions between bathymetry and tidal currents play a dominant role, alongside patterns in seasonal stratification and shelf-edge upwelling. The attractiveness of a favoured oceanographic feature appears to stem from persistent and/or predictable increases in prey accessibility which facilitates foraging. Changes in prey abundance and density, behaviour and depth distribution are fundamental to this, and in a number of cases have been shown to improve foraging efficiencies. However, our knowledge of interactions between marine predators and oceanographic features favoured for foraging is far from complete, and future work should aim to further our understanding of the functional mechanisms linking bio-

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686 physical processes, prey and marine mammals and seabirds. In many instances, detailed and 687 simultaneous three-dimensional measurements of sub-surface oceanography, prey densities and 688 distributions alongside marine predator behaviours (across three-dimensions) are lacking, and this 689 should be addressed. Such efforts may also elucidate the importance of lesser studied features such 690 as bottom fronts and Langmuir circulation cells. A better understanding of how predators perceive 691 their environment and develop foraging strategies during immature/juvenile stages would also be 692 beneficial, as would comparative inter-taxa research and estimates of the importance of 693 oceanographic habitat features at a population level. Knowledge of the bio-physical processes that 694 underlie habitat use by marine mammals and seabirds across shelf-seas should be used to inform 695 future conservation management and policy. This will require improvements in the monitoring of 696 oceanographic conditions such that adaptive strategies can be implemented which, where necessary, 697 can evolve through space and time in response to the dynamic nature of the ocean. This would aid 698 the preservation and protection of oceanographic features that facilitate trophic transfer, and are thus 699 critical to the functioning of shelf-sea environments. Such holistic approaches are vital to ensuring 700 the future health of these complex ecosystems.

## 701 Acknowledgements

702 This work was funded by a PhD studentship to SLC by the Natural Environment Research Council
703 (NERC; NE/J500380/1). We thank GL Hunt and one anonymous reviewer for comments on earlier
704 versions of this manuscript.

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# Figures



Increasing distance offshore (typically remaining within 200nm of major landmasses)

Figure 1. Cross shelf schematic giving an overview of the typical oceanographic structure of thermally stratified shelf-sea environments at midlatitudes during the summer months. Adapted from Simpson and Sharples (2012). Black directional arrows show (1) turbulent mixing around the seabed, offshore banks and islands (circular arrows), and (2) convergent and divergent upwelling and downwelling currents associated with the passage of internal waves (angled arrows).

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Figure 2. Overview of the typical spatio-temporal scales oceanographic structures across shelf-sea environments function over. Further details can be found in Table 1.

706 Table 1. Details of the generalised spatio-temporal scales over which oceanographic features favoured for foraging across shelf-seas function. Columns 707 from left to right list a features name, the spatial and temporal extent its expanse covers respectively, and finally the spatio-temporal scales over which 708 variation may occur. References are provided for specific examples of features functioning in this manner. Note that regional and site specific differences 709 in physical forcing mean there is variation around these generalisations in spatio-temporal scale, which should be assessed.

BIO-PHYSICAL HABITAT FEATURE	SPATIAL EXTENT	TEMPORAL EXTENT	SPATIO-TEMPORAL VARIATION
Shelf-edge fronts	Lateral extent of 10's - ~100km. Length may exceed 100's km's (e.g. Celtic shelf break and along the edge of the continental shelf of the Bering Sea; Pingree <i>et al.</i> , 1981; Springer <i>et al.</i> , 1996).	Typically persist perennially, although strength and associated productivity may vary seasonally (Fournier <i>et al.</i> , 1979; Ryan <i>et al.</i> , 1999).	Geographical location may shift 10's km's with prevailing meteorological conditions, seasonal changes and climatic fluctuations (Linder and Gawarkiewicz, 1998). Intensity may also vary over similar temporal scales (e.g. the strength of frontal interfaces as determined by horizontal gradients in sea surface temperature), and inter-annually with impacts on associated productivity (Fournier <i>et al.</i> , 1979).
Wind-driven upwelling fronts	Lateral extent of 10's to ~100km. Length may exceed 100's km's (e.g. along the California and Humbolt Currents; Acha <i>et al.</i> , 2004; Letelier <i>et al.</i> , 2009; Kahru <i>et al.</i> , 2012).	Generally display some seasonality in occurrence in response to upwelling intensity with current flows wind events (Kampf <i>et al.</i> , 2004; Bograd <i>et al.</i> , 2009b; Letelier <i>et al.</i> , 2009). Once established, may persist for several months.	Geographical location may shift 10-100's km's seasonally and inter-annually with climatic fluctuations in upwelling intensity, which may also impact frontal intensity and productivity levels (Bograd <i>et al.</i> , 2009b; Letelier <i>et al.</i> , 2009). Surface convergences and instabilities within these zones may be more ephemeral in occurrence, and linked to local meteorological conditions (Capet <i>et al.</i> , 2008).
Spring bloom	Can extend across stratified section of the shelf- sea, encompassing areas exceeding 100's km <sup>2</sup> (e.g. Celtic & North Seas; Pingree <i>et al.</i> , 1976; Holligan <i>et al.</i> , 1989).	Seasonally occurring in spring for a period lasting no more than a month (typically around 1-2 weeks; Pingree <i>et al.</i> , 1976; Townsend and Spinrad, 1986; Mills <i>et al.</i> , 1994).	Initial date may vary with the spring-neap cycle alongside changes in climatic conditions (Hunt and Stabeno, 2002; Mann and Lazier, 2006; Sharples <i>et al.</i> , 2006), which can also impact magnitude of productivity (Sambrotto <i>et al.</i> , 1986). Regional variation in initial start date also occurs (e.g. latitudinally; Henson <i>et al.</i> , 2009). Smaller scaled shifts in geographical occurrence may also occur inter-annually (e.g. across the Bering Sea; Hunt and Stabeno, 2002).
Vertical interfaces in stratified regions (i.e. the pycnocline)	Vertical extent of 10 cm's to a few metres (Simpson and Sharples, 2012). Horizontally extends across offshore stratified section of shelf-sea, which may encompass 100's km <sup>2</sup> (Pingree, 1975; Holligan <i>et al.</i> , 1989).	In offshore waters may be persistent perennially, although depth and inshore extent varies seasonally, and is most prominent during spring, summer and early autumn (Pingree, 1975).	Inshore extent and depth can vary inter-annually, seasonally, and on short time-scales (days to weeks), with climatic variation, weather events and tidal currents, as can the intensity of the pycnocline (i.e. gradient of change in density; Cairns and LaFond, 1966; Skov and Durinck, 2000; Ropert-Coudert <i>et al.</i> , 2009; Kokubun <i>et al.</i> , 2010). As the pycnocline approaches tidal-mixing fronts at the boundaries of coastal mixing waters, its depth shallows (Pingree, 1975). Around offshore banks, reduced depths are observed with internal wave passage and localised upwelling (Stevick <i>et al.</i> , 2008; Embling <i>et al.</i> , 2012).
Sub-surface productivity	May occur over a larger area exceeding 10's km's (Weston <i>et al.</i> , 2005), or locally in concentrated patches of 100's m's to km's (Scott	Can occur across a season, or ephemerally for a few days/weeks (Richardson and Christoffersen, 1991; Sharples <i>et al.</i> , 2001;	Productivity and entrainment may be highest when the water column stabilises, following a period of tidal and wave induced vertical mixing (McManus <i>et al.</i> , 2005; Cheriton <i>et al.</i> , 2007;

	<i>et al.</i> , 2010). Productivity is generally vertically concentrated within a few m's of the pycnocline (Sharples <i>et al.</i> , 2001).	Sharples, 2008).	Durham <i>et al.</i> , 2009). May be particularly elevated around offshore banks where internal waves propagate or there is localised upwelling (Lennert-Cody and Franks, 1999; Richardson <i>et al.</i> , 2000; Embling <i>et al.</i> , 2012).
Offshore banks & internal waves	Spatially predictable and tied to topographic features generally occurring over $1-10$ 's km <sup>2</sup> (e.g. Jones bank; (Palmer <i>et al.</i> , 2013).	Appear to be seasonal features linked to thermal stratification between late spring and early autumn. May function intermittently with specific tidal conditions (Embling <i>et al.</i> , 2012, 2013; Palmer <i>et al.</i> , 2013)	Closely tied to bathymetric structures. Propagation of internal waves may vary with spring-neap modulation and storms (Embling <i>et al.</i> , 2013; Palmer <i>et al.</i> , 2013). Those associated with the shelf edge (e.g. within the Humboldt current; Bertrand <i>et al.</i> , 2008), may be influenced by changes in upwelling intensity.
Tidal-mixing fronts	Small lateral extent of typically 100's m's to 10's km (Schumacher <i>et al.</i> , 1979; Decker and Hunt, 1996; Durazo <i>et al.</i> , 1998). Length may exceed 10's km's (e.g. tidal mixing fronts around the Bristish Isles; Fearnhead, 1975; Pingree and Griffiths, 1978; Miller, 2009).	Seasonally occurring from late spring to early autumn (e.g. Fearnhead, 1975; Pingree and Griffiths, 1978; Kachel <i>et al.</i> , 2002; Acha <i>et al.</i> , 2004).	Locations are coarsely predictable and typically tied to a ratio of total water depth ( <i>h</i> ) and tidal velocity ( <i>u</i> ) - $h/u^3$ (Simpson and Hunter, 1974). Inshore/offshore shifts may occur over scales of 10's km's, in response to changes in current strength with the spring-neap cycle, alongside surface induced mixing during wind events and decreased solar irradiance (Kachel <i>et al.</i> , 2002; Nahas <i>et al.</i> , 2005; Pisoni <i>et al.</i> , 2015). This may follow seasonal patterns (Hill and Simpson, 1989). Small scale variation (100's m's to km's) in the surface signature of a front may occur throughout the tidal cycle and due to prevailing wind conditions (Durazo <i>et al.</i> , 1998).
Channels, headland & island wakes, nearshore reefs & bays	Predictable, topographically tied localised features, occurring over 100's m's to a few km's (Zamon, 2002; Johnston and Read, 2007; Bailey and Thompson, 2010; Jones <i>et al.</i> , 2014). Some channels, bays and passes may extend across 10's km's (e.g. Academy Bay, Sea of Okhotsk, and passes between Aleutian Islands; Ladd <i>et al.</i> , 2005; Rogachev <i>et al.</i> , 2008).	May occur perennially, although only at specific times in the tidal cycle (Zamon, 2003; Johnston and Read, 2007; Bailey and Thompson, 2010; Waggitt <i>et al.</i> , 2016).	Functionality may vary with strength and direction of water currents across diurnal and bi-weekly tidal cycles (Johnston and Read, 2007; Bailey and Thompson, 2010; Jones <i>et al.</i> , 2014). Local wind patterns may temporarily intensify surface convergences (e.g. occurrence of Langmuir circulation cells; Goss <i>et al.</i> , 1997; Ladd <i>et al.</i> , 2005).
ROFI's: estuarine plume & tidal intrusion fronts	May be localised over 100's m's (e.g. Moray Firth fronts; Mendes <i>et al.</i> , 2002) or larger encompassing 10's to 100's of km's (e.g. Columbine Estuarine & Mississippi River Plume Fronts; Govoni <i>et al.</i> , 1989; Phillips <i>et al.</i> , 2018).	Occur perennially, sometimes periodically with the tidal cycle (Simpson and Nunes, 1981; Mendes <i>et al.</i> , 2002; Phillips <i>et al.</i> , 2018).	Occurrence may be linked to specific phases of the tidal cycle (e.g. Mendes <i>et al.</i> , 2002). Geographical location and strength may vary with tidal cycle and experience spatial variation with spring-neap changes in current strength. Local weather conditions and climate will also impact geographical location, occurrence and strength (e.g. rainfall and wind; Kowalczyk <i>et al.</i> , 2015b; Phillips <i>et al.</i> , 2018). Productivity blooms may follow such cycles (Cloern, 1991).