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**Geoarchaeological and Paleoenvironmental context of the human settlement
in the Eastern Tandilia Range, Argentina**

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Abstract

The Tandilia range is a discontinuous series of hills and valleys located at the central area of the Pampean region (Argentina). The process of human settlement reached in this zone the highest population densities and reoccupation levels of Southern South America, developing under the strongly changing climatic conditions that characterized the Pleistocene-Holocene transition. This paper analyzes the geoarchaeological and paleoenvironmental record of the Pleistocene-Holocene transition (c. 9400-10,500 ¹⁴C BP) recovered from cave and rockshelter sedimentary successions of the Tandilia range. The analysis of several lines of evidence (stratigraphic, geoarchaeological, archaeological, zooarchaeological, and paleobotanical), suggests that a cold arid to semi-arid climate characterized the late Pleistocene (c. 10,500 BP), and changed abruptly to warmer conditions towards the beginning of the Post-Glacial Period. As a consequence of this climatic amelioration, the populations inhabiting the Tandilia range would have been benefited by a higher availability of resources, which could explain the higher levels of site density and reoccupation found in the area during this period.

Keywords: *early human settlement, pampean region, Argentina, paleoindian, Pleistocene/Holocene*

1. Introduction

Although the process of human occupation of the Pampean region began c. 12,100 ^{14}C BP (Politis, et al., 2004; Steele and Politis, 2009; Prates, et al., 2011), most archaeological evidence is found between c. 11,000 and 9600 ^{14}C BP. During this period, a wide variety of pampean sub-environments became fully occupied. The highest human population densities were registered in the eastern Tandilia range (Mazzanti, 2003; Flegenheimer, et al., 2006; Mazzanti, et al., in press) and, in a lesser extent, in the contiguous Interserrana Plain (Martínez, 1999; Politis, et al., 2004) (Figure 1).

The region studied in this paper comprises the eastern Tandilia range (Figures 1 and 2), an area characterized by hills, isolated rocky outcrops, and extensive table-form massifs with interior valleys. Given their geomorphological characteristics, these hills would have constituted a very suitable environment for human settlement, particularly during the periods of cold and dry climate that characterized the Pleistocene-Holocene transition. The most important benefits offered would have consisted in a wide visual domain over the surrounding landscapes, the presence of abundant freshwater sources, and the proximity to the Atlantic Ocean, with the consequent increased biodiversity, which included flora and fauna both of continental and coastal origin. In this context, the hydrological catchments would have acted as a nucleus for attracting human groups, as it occurred in the most important river of the region, the Quequén Grande river (Martínez, 1999; Gutierrez and Martínez, 2008; Martínez *et al.*, this volume). Palynological evidence also suggested that these micro-environmental conditions would have favoured the development of extensive grasslands (Prieto, 1996; Quatrocchio et al., 2008) inhabited by herbivorous animals, that eventually became important for hunting and settlement for specialized or domestic tasks.

The earliest occupation levels registered in the caves and rockshelters, dated between c. 10,500 and 9600 ^{14}C BP, were recurrently found over a notable stratigraphic discordance that indicates a change in depositional and post-depositional processes (Martínez, 2007). The end of the Pleistocene was also characterized by a continental-scale massive extinction of megamammals (Martin, 1984; Politis, et al., 1995; Haynes, 2006). In the Tandilia range and surrounding areas, this extinction occurred at the end of the Pleistocene and affected most of the big herbivorous species utilized by humans, leading to a decrease in the diversity of potential preys, as is reflected in the archaeological records at Cueva Tixi and Cueva El Abra. However, human subsistence was not severely affected, since populations maintained a generalist pattern that prevailed throughout the early Holocene (Quintana, 2001).

The high concentration of archeological sites observed in the Tandilia microregion enhances the importance of this area for the study of the processes of early settlement in the

Pampean region. Besides the nine sites presented in this work (Figure 2), a total of eight rockshelter and open-air sites have been already studied in the adjacent central-eastern sector of the Tandilia range (Flegenheimer et al., 2006; Prates et al., 2011). As a whole, this diversity of sites demonstrates the functional variability and the types of settlements that characterized early human populations in this region. These earliest settlements would have developed under fluctuating paleoclimatic conditions (Politis et al., 2004; Martínez, 2006; Grill et al., 2007) similar to those described for southern Patagonia between 12,000 and 11,000 ^{14}C BP (Miotti and Salemme, 1999, 2003; Borrero, 2003).

In this paper, the stratigraphic and paleoenvironmental context of nine sites attributable to the initial human population of the eastern Tandilia range are discussed. The archaeological sites studied include Cueva Tixi, Abrigo Los Pinos, Cueva El Abra, Cueva La Brava, Cueva Burucuyá, Lobería I Sitio 1, and Amalia Sitio 2, as well as two small rockshelters, El Mirador and Abrigo La Grieta, where excavations have been recently initiated (Figure 2). The main objective is to analyze the paleoenvironmental conditions prevailing in the region during the settlement process, through the analysis of several types of proxy records, which include geoarchaeological, zooarchaeological, archaeological, paleobotanical (charcoals), and microfossil (diatom and phytolith) data. The studied records have been recovered from stratified sediments located at the base of the sequences of caves and rockshelters. These circumstances optimized the discovery of cultural contexts with good integrity and favoured the use of several proxy data.

Although research is still in progress, the information provided in the present contribution becomes of great value in order to address questions of local and regional interest regarding the dynamics of human population in the Pampean region: What was the geographical significance of the Tandilia range in the process of human settlement? Why did human occupation develop so intensively after c. 10,500 ^{14}C BP? Moreover, this contextual framework of the social dynamics of the area poses more general questions related to the initial population of Southern South America, particularly the Atlantic region, which lead to the question: How were the prevailing paleoenvironmental conditions and landscape during the earliest human occupation of the eastern Tandilia range?

2. Material and Methods

Allostratigraphic units of caves and rockshelters were defined based on their sedimentary characteristics (texture, color, types of limits and structures) and pedogenetic features. Particle size analysis was performed by sieving and pipetting. Organic carbon content was determined by the Walkley and Black method (Walkley and Black, 1934), and pH was measured with a pH-meter in a

suspension with a soil/water ratio of 1:2.5. Radiocarbon dates were performed both through the ^{14}C conventional and Accelerated Mass Spectrometer (AMS) methods (Table 2).

Diatom analyses were conducted on the Loberia I, Cueva Tixi and Cueva El Abra sequences. In each site, sediment samples were collected following the defined allostratigraphic units. Samples were processed through standard methodologies (see Hassan et al., 2009) and permanent drop slides were made with Naphrax®. On each slide at least 300 diatom valves were counted in random transects using a magnification of 1000X. The species were characterized according to their ecological requirements, and particularly moisture preferences, following McMillan and Rushforth (1985); Van Dam et al. (1994); Johansen (1999); Lowe et al. (2006) and Poulíčková y Hašler (2007). Constrained incremental sum of squares cluster analysis (single linkage, Euclidian distance) was performed on diatom percentages to define diatom zones in the fossil sequences using the computer program TILIA CONISS (Grimm, 2004).

Detailed phytolith analysis were conducted in a representative profile of Loberia 1 site 1, and then compared with selected samples from the lower section of Cueva Tixi archaeological site, in a chronological lapse between 10.375 ± 90 and 7.700 yr B.P. The samples were processed using standard wet oxidation and heavy flotation techniques (with sodium polytungstate solution, specific gravity 2,345) following Zucol, et al. (2010a and b); while morphotype assignation and counting were made according to the descriptive scheme of Zucol et al. (2010c). The phytolith diagram was performed using POLPAL, Numerical Analysis programs (Walanus and Nalepka, 2002), and comparative cluster analysis were conducted using PAST (Paleontological statistics software, Hammer, et al., 2001).

The archaeological charcoals were identified by examining freshly fractured surfaces (cross, radial and tangential sections) using reflected light. They were identified using standard wood key, wood atlases and descriptions by Metcalfe and Chalk (1950), Tortorelli (1956), Tuset (1963) and InsideWood database (2004-onwards). The terminology for the wood features follows the IAWA Lists of Features Suitable for Hardwood Identification (IAWA Committee, 1989). Charcoal remains were compared to extant species, particularly present-day taxa from the Eastern edge of the Tandilia Range (Buenos Aires, Argentina). The collected remains were deposited in the Colección Científica del Laboratorio de Arqueología de la Facultad de Humanidades, Universidad Nacional de Mar del Plata.

3. Study Area and Stratigraphic Context

The climate is humid-temperate with a mean annual precipitation of 979 mm (for the last 20 years) and mean temperatures of 20°C in January and 7°C in July (Romanelli et al., in press). The

potential evapotranspiration values estimated for the same period by the Thornthwaite method, ranged from 696 to 766 mm/year, with an average of 740 mm/year (Kruse et al., 1997).

Two mountain ranges are present in the southern Pampean region: Ventania in the southwest (maximum height: 1,243 masl), and Tandilia in the northeast (525 masl; Figure 1). The two ranges are separated by the Interserrana Plain, an extensive plain filled by continental Cenozoic deposits, with a maximum altitude of 200 masl (Figure 1). The eastern Tandilia range constitutes a group of table-like hills of flat summit, composed of lower-Paleozoic sub-horizontal quartzite strata (Balcarce Formation, Dalla Salda and Iñiguez, 1979). The hills that form this portion of the range present maximum elevations of c. 300 masl, they are elongated in an east-west direction, and end in the Atlantic Ocean. They are separated by wide valleys and undulated plains (Figures 1 and 2). The late Pleistocene-Holocene sediments that cover the intermountain and plain areas are sometimes present in the horizontal summits of the main ranges. These deposits are composed of loess or reworked silt of aeolian and volcanoclastic origin (Teruggi, 1957; Tricart, 1973).

Most of the analyzed archaeological sites of eastern Tandilia are located in caves and rockshelters developed on quartzite cliffs near to the ranges summits. In some cases, however, these are located in low quartzite rocks outcropping in the Interserrana Plain (Figure 2). The origin of these cavities has been attributed to dissolution processes of the siliceous cement on quartzites, occurred during wet periods during the Cenozoic (see Martínez, 2007, and references therein). Although caves and rockshelters are known as highly efficient sediment traps (Colcutt, 1979), in Tandilia only a few of them present sedimentary matrices. The infilling material can be both endogenous, derived from ceiling and wall by dissolution, physical weathering and collapse; or exogenous sediments, mainly composed by primary or reworked loess blown in by the wind or carried by running water (Table 1). Dissolution of quartzite can be very important in the endogenous processes since quartz sandy sediments constitute a significant percentage of the stratigraphic sequences of some caves (i.e.; up to 90% in Los Pinos and Cueva Tixi; Martínez et al., 1999).

The development of the late Pleistocene-Holocene sedimentary matrices of the caves and rockshelters is very variable, ranging from 0.40 to 2 m (Figure 3). These deposits are usually graded beds, indicating a decrease in energy from the late Pleistocene to the late Holocene.

Early human occupations in eastern Tandilia sites are dated between 9570 and 10,672 ¹⁴C BP (see Table 2). These archaeological levels are located immediately above erosional or depositional stratigraphic unconformities (Figures 3 and 4). These discordances separate the coarser and stone-rich Pleistocene sediments from the sandy-clayed Holocene silts (Figure 3). The presence of these coarse angular rock fragments has been related to mechanical weathering processes linked

to the cold and arid climatic conditions (e.g. Tonni and Fidalgo, 1978; Prado et al., 1987; Alberdi et al., 1989; Bonadonna et al., 1999; Martínez, 2001) that characterized the Last Glacial Maximum (IOS 2; Martínez, 2007). There are also distinct colour changes between sediments located above (brown-gray, 10 YR 5 / 2) and below the unconformities (very pale brown, 10 YR 7 / 4) (Martínez et al., 1999; Martínez and Osterrieth, 2003).

Holocene sediments show typical signs of humid conditions, like pedogenetical modifications (i.e. Los Pinos and Cueva Burucuyá) and diatomaceous levels (i.e. Cueva Tixi, Cueva La Brava, Cueva Burucuyá). Moreover, in most of the profiles the organic matter content was higher in Holocene than in late Pleistocene sediments (Table 1). pH varied between 3 and 5 (the only exception was Cueva Tixi, where it ranged between 7 and 8; Martínez et al., 1999; Martínez, and Osterrieth, 2001), which may explain the low preservation of bones in these sites.

4. Archaeological Context

This section summarizes the results of the archaeological studies performed on the nine sites mentioned above, which are representative of the functional variability of the local system and provides information about the early settlement of hunter-gatherer groups in the Tandilia range valleys. Three of the analyzed sites showed evidence of residential occupation (Cueva Tixi, Cueva El Abra, and Abrigo Los Pinos), particularly through a repeated use of their interior chambers for a variety of activities related to residential settlement. The archaeological records from the other four sites (Cueva La Brava, Cueva Burucuyá, Amalia Sitio 2, and Lobería I Sitio 1) indicated only occasional utilization, and presented lower amounts of archaeological materials (Mazzanti, 2003; Mazzanti et al., 2010). The two most recently excavated sites (Cueva El Mirador and Abrigo La Grieta), preliminarily assigned to the Pleistocene-Holocene transition because of their stratigraphic characteristics and the nature of the archaeological materials assemblages, evidenced a more specific set of activities that were complementary to the functions of the surrounding sites.

Six sites were situated on the same hill, with five of these located in the valley of Arroyo La Vigilancia, at a very short distance from each other (Figure 2). This proximity has allowed addressing the mobility patterns of domestic groups, in terms of the repeated use of these strategic local sites on the landscape. The archaeological evidence analyzed includes lithic artefacts and mineral pigments collected in all sites, materials from bowl-shaped combustion structure found in three sites (Cueva La Brava, Cueva Burucuyá, and Abrigo Los Pinos), and an abundant and diverse assemblage of well preserved faunal remains found in Cueva Tixi (Mazzanti, 1997, 1999, 2003; Quintana, 2001).

Studies of lithic technology integrated information about the diversity of transported stones used in various stages of artefacts production (Valverde, 2002, 2003, 2004). Some patterns for the procurement of raw material and the mobility strategies, that characterized the social circuits of interactions among these ancient populations, can now be proposed (Mazzanti, 2003). The results are in general agreement with the modalities of stone material procurement proposed by Bayón and Flegenheimer (2004) for the Pampean region, which involve carrying and possible visits to procurement source areas (Tandilia, Ventania, and the Atlantic coast) by highly mobile groups. Some of the artefacts found in the Tandilia sites were made with a red silicified limestone, a material that could have originated in Uruguay (Flegenheimer et al., 2003, Bonomo and Blasi, 2010). The presence of artefacts made from this type of stone could reflect the extensive social interaction space within which these initial human settlement groups developed.

Artefacts recovered at sites with evidence of residential activities indicated a strong tendency toward the exploitation of local stones (medium- and coarse-grained orthoquartzites from the Balcarce Formation. The only exception was Cueva Tixi, where allocthonous stones predominated (fine-grained orthoquartzites from the Sierras Bayas Group). These patterns reflect important aspects of the technological decisions made by these early groups and provide evidence about their mobility patterns within the mountain range system. Rocks from Barker (Sierras Bayas Group) (Figure 1) were exploited to obtain fine orthoquartzites and mineral pigments (Porto López and Mazzanti, 2010). These raw materials were transported in a variety of forms (nodules, cores, flakes, tools) to sites located more than 70 km away. Movements to the Atlantic coast were also inferred around 11,000 ¹⁴C BP, based on the presence of basalt cobbles and silica stones collected from beaches about 60 km away. To the west, outcrops of andesites, silicas, ftanites, silicified dolomite stone, and other types of rocks are found. Towards the south, in the Ventania mountain range, sources of rhyolites and meta-quartzites are found (Bayón et al., 1999). All these stone types were identified in very low proportions in assemblages from the nine sites, mainly in waste flake deposits. Their presence indicates a profound level of knowledge on stone sources, both within and beyond the Tandilia range system. More ephemeral or re-occupied sites also indicate that small, very mobile groups developed multiple, diverse settlements with high levels of functional variability and predominant use of local resources.

5. Faunal Remains

The zooarchaeological record from eastern Tandilia comprises materials from the Cueva El Abra and Cueva Tixi sites, although only the later contains abundant and well-preserved remains (Quintana, 2004; Quintana et al., 2003). The Cueva Tixi record consists of 4,149 bone specimens

identified to various taxonomic levels, including 33 species-level assignments (this number is approximate since the small bird remains were not identified) (Table 3). Nine of these species showed evidence of human activity, such as fractures, cut marks, or manufacturing of bone artefacts. Most of the species not utilized by humans were transported to the cave by birds during periods without human occupation. These birds created a deposit of pellets, from which 2,423 microvertebrate bones were recovered (caviomorph and cricetid rodents, marsupials, and birds). These bones showed a good preservation, with moderate fragmentation and little evidence of alteration by digestion.

The sample from Cueva El Abra is poor and only consumption of the deer *Ozotoceros bezoarticus* could be identified. Indetermined fragments of ungulate bones were also present, indicating an intense alteration of the skeleton. The occupants of Cueva Tixi developed a generalized subsistence strategy (Quintana and Mazzanti, 2001). Resources exploited included high-rank taxa such as guanaco (*Lama guanicoe*) and deer (*Ozotoceros bezoarticus*); low-rank taxa such as vizcachas (*Lagostomus maximus*), coypos (*Myocastor coypus*), ñandú (*Rhea* sp.) eggs, and three species of armadillos (*Chaetophractus villosus*, *Dasypus hibrydus*, and *Zaedyus pichiyi*); and intermediate-rank taxa such as the ñandú and a giant extinct armadillo (*Eutatus seguini*) (Table 3) (Quintana and Mazzanti, 2001).

Caviidae rodents, lizards, and medium-sized birds were available in the landscape during the late Pleistocene (and were recovered from Paleoindian levels), but were only incorporated into the subsistence system during the Late Holocene. In contrast with sites on the Interserrana Plain, in these mountainous sites megamammals were not recorded, although their presence in natural Late-Pleistocene strata confirms their presence in the area. Three species recorded for this period, *Eutatus seguini*, *Dusicyon avus* (a large fox), and *Galea tixiensis* (a wild cavy) became extinct during the late Pleistocene, the Middle Holocene, and the Late Holocene, respectively. A fourth species, *Monodelphys* sp., possibly became extinct during the Late Holocene (Goin, 2001). The faunal assemblages from these late Pleistocene levels contained some species considered to have value as climate indicators. However, there are also non-analogous associations present that demonstrate the co-existence of species from arid environments (*Lestodelphys halli*, *Zaedyus pichiyi*, and *Dolichotis patagona*) and sub-tropical environments (*Cavia aperea*, *Monodelphys* sp., *Dasypus hybridus*, and *Oxymycterus* sp.) (Quintana, 2001; Goin, 2001).

A third Paleoindian site, Cueva Burucuyá, lacked faunal remains associated with the early human occupation. Underlying the archaeological level, dated between to 10,000 – 10,672 BP, a lower stratum presented a markedly uneven upper surface that contains deep ichnites from megamammals, which cover the entire excavated area (Quintana et al., 2001) (Figure 4). This

assemblage of footprints forms the base of the stratum in which the archaeological materials were found.

6. Palaeobotanical Record

This section presents an overview of the results of the palaeobotanical studies carried out in the studied sites. As these studies are currently in progress, analysis will focus only on the sequences where the Pleistocene-Holocene transition has been analyzed in detail. Stratigraphic diagrams showing the distribution of the paleobotanical indicators along the complete sequences are presented, in order to provide an historical framework for the discussion of the changes inferred for the Pleistocene-Holocene transition.

6.1. Diatoms

Diatom communities respond quickly to environmental changes and are composed of a large number of species with different tolerances to physical and chemical variables. Consequently, they are well-established indicators for both monitoring present changes in water condition, and tracking past environmental changes through the analysis of the assemblages preserved in sediments (see Smol and Stoermer, 2010). In Argentina, continental diatoms have been used to infer paleoenvironments in southern (Zárate et al., 2000; Hassan et al., 2004), and western Pampas (González and Maidana, 1998), and Patagonia (Markgraf et al., 2003; Mayr et al., 2005).

Diatom studies were carried out in three of the sites studied in the present contribution: Lobería I, Cueva Tixi and Cueva El Abra. The Lobería I sequence was dated between c. 9800 to 160 ¹⁴C BP. A total of 35 diatom species were identified, from which only 13 reached relative abundances >2% in at least one sample and were plotted (Figure 5). Cluster analysis allowed dividing the sequence into two three diatom zones (Figure 5), from which the first one (Zone I, units 7 to 5E) comprised the Pleistocene-Holocene transition. The four samples grouped into this zone were dominated by different diatom assemblages, which could be indicating fluctuating environmental conditions for this moment. Unit 7 was dominated by *Diadismus gallica*, accompanied by *Nitzschia hantzschiana*. *D. gallica* is frequently confined to the surface of caves, mosses, and neutral to slightly alkaline soils, being able to grow under low light intensity (Johansen, 1999). Unit 6B was dominated by *Melosira* sp., together with *Diadismus contenta* and *N. hantzschiana*. Unit 6 A was characterized by a high abundance of *N. hantzschiana*. Unit 5E was dominated by *Pinnularia borealis* and *P. subcapitata*. Those taxa inhabit environments with higher water content, in some cases attached to moss surfaces, and are absent in strictly aerial substrates.

Eight samples, corresponding to levels B to F, were studied in Cueva Tixi. These levels were dated to between c. 10000 and 700 ¹⁴C BP (Martínez, 2007). A total of 35 diatom taxa were identified, from which only 13 reached frequencies $\geq 2\%$ in at least one sample (Figure 6A). Three diatom zones were defined according to cluster analysis. Zone I, located at the base of the sequence, comprised the Pleistocene-Holocene transition. This sample was dominated by *Denticula elegans* and *Nitzschia hantzschiana*, and indicates low water availability in the environment, as is dominated by taxa highly tolerant to dissection.

Five sediment samples were studied in Cueva El Abra, corresponding to units 4 to 1, ranging from c. 9800 to 960 ¹⁴C BP (Martínez, 2007). Twenty four diatom taxa were identified in the remaining sedimentary units, from which only 11 reached frequencies $\geq 2\%$ in at least one level and were plotted (Figure 6B). Sedimentary levels from the Pleistocene-Holocene transition (unit 4) did not contain diatom frustules, possibly indicating a lack of water availability into the cave.

Overall, diatom assemblages were dominated mainly by aerophilic diatoms (Figure 7). These taxa are present in a variety of sub-aerial environments, such as mosses, wood, wet rocks, as well as lichen phycobionts. Their development is mainly limited by their ability to survive to dry moments of different longitude and periodicity (Johansen, 1999). Their dominance in the studied caves indicates that although they have been wet during the Holocene, the moisture was mostly restricted to cave walls and other wet substrates. Many of the identified taxa, such as *Diadsmis gallica*, *D. contenta*, *Frankophila similioides*, *Nitzschia amphibia* and *Pinnularia subcapitata*, have been registered alive in modern environments of the Tandilia range. They inhabit caves with high moisture levels, where water infiltrates through the walls allowing the development of mosses (Hassan, unpublished data). This modern information is, however, scarce and further study on the distributional patterns of the modern diatom flora inhabiting these caves is needed in order to adjust the paleoenvironmental inferences, particularly in relation to moisture reconstructions.

6.2. Phytoliths

Phytolith studies are important archaeobotanical resources, especially in early archaeological records, where it is difficult to find well preserved botanical macroremains. Recent phytolith studies carried out in the Pampean region allowed advancing both in paleoenvironmental and human plant management knowledge (Osterrieth, et al., 2000; Zucol, et al., 2002; Zucol, et al., 2007, Zucol, et al., 2008; Colobig and Zucol, 2008; Zucol, et al., 2010c, Mazzanti, et al., 2010; Colobig, 2011). These siliceous microremains are very resistant to destruction and representative of local processes, constituting excellent proxies for analyzing the discrepancies between the archaeological settlements and the surrounding areas from a paleoecological perspective.

The Lobería I profile showed a very complete phytolith record, which allowed dividing the sequence into five phytolith zones (Figure 8). The Pleistocene-Holocene transition (samples 6B and 7) was included in zone I, and characterized by the presence of elongated elements with smooth (Figure 9.10), wavy (Figure 9. 11) or serrate (Figure 9.18) contour associated with short (Fig. 9.16) and elongated truncated cones (Figure 9. 17); point-shaped phytoliths were more abundant in sample 7. These dominant elements were accompanied by low abundances of bilobate (Figure 9.14, 15) and polylobate phytoliths, other truncated cones types, saddle, globular and festucoid boat-shaped elements. Abundant articulated epidermal phytoliths were observed (Figure 9. 12), more of them with dicot affinities (Figures 9.9). Sponje spicules and chrysophycean stomatocysts were scarce.

In Cueva Tixi, the phytolith assemblage of the basal section (unit F), corresponding to the Pleistocene-Holocene transition, was characterized by the dominance of elongated phytoliths of smooth and wavy contour (Figures 9.3), together with short prismatic, truncated cones (Figures 9.6 and 9.7) and bilobate elements. Elongated phytoliths with serrate and denticulate contour, with concave or straight ends, point shaped, polyhedral bulliforms, fan-shaped, rondel, globular (Figure 9. 13) and bilobate (Figure 9.5) were also present, although in lower abundances. Articulated phytoliths with different epidermal trichome elements were also abundant, while sponje spicules and stomatocysts were scarce. The comparative phytolith assemblage's analysis from lower sections of the studied sites (Figure 10) can establish that the sample E1 of Cueva Tixi show similarities with the sample 7 from Lobería I sequence, chronologically located in different times.

6.3. Charcoal

Archaeological sites of the Eastern Tandilia Range have scarce and very fragmentary anthracological records. However, these materials appear as small twigs and sprigs without the diagnostic features that allow a good systematic identification, are suitable for radiocarbon dating.

Archaeological charcoal can be employed to identify trees, shrubs and herbs in the area around human settlements (Bonomo et al., 2011 and references therein). Also, concentrated charcoal is usually a sign of human selection and used of forest resources (Figueiral, 2005). The archaeological charcoal used as paleoenvironmental indicator has been discussed for varied authors (see Marconetto, 2010) because the charcoals are the result of selective practices performed by humans rather than a faithful representation of environment (Piqué i Huerta, 1999; Marconetto 2005, 2010). However, is possible generate palaeoecological and palaeoenvironmental data on the basis of anatomical features of the archaeological charcoal instead of focusing on the floristic associations of the remains (Marconetto, 2010).

Of all the anthracological material recovered, samples from Cueva El Abra were the only that allowed a systematic determination. These pieces were assigned to *Salix* sp. (Salicaceae), recognized by semi-ring porous to diffuse porous, predominantly solitary vessels and rarely in radial multiples. The vessels are hexagonal in outline. Apotracheal axial parenchyma is scarce. Rays are very narrow and short, uniseriate and biseriate and 7–27 cells high. Rays are homogeneous composed of upright or square cells. *Salix* sp. was used as fuel because these charcoals were found concentrated and might indicate selection of resources by hunter-gatherer groups. This combustion structure has been dated by ^{14}C (AMS) to 9834 ± 65 B.P.

7. Discussion

The sedimentary sequences analyzed in this paper show a record of early occupations that cover a span of c. 1000 y (9600–10,500 ^{14}C BP). Various lines of evidence (stratigraphic, sedimentological, geoarchaeological, diatoms, phytoliths, wood charcoal, faunal, and archaeological) indicate that the post-glacial climatic amelioration took place very quickly around 10,500 BP. The late Pleistocene-Holocene transition represented one of the most distinctive and recurrent sedimentary records found in the caves and rockshelters of the eastern Tandilia range. This transition is stratigraphically represented by an erosional unconformity, which is marked by a change in grain size, colour, and compaction in relation to the overlying and underlying sediments. This unconformity corresponds to a change in the sedimentary dynamics, and is located immediately below the first early occupation layers (c. 9600–10,500 AP). The underlying late Pleistocene sediments reflect colder and drier paleoclimatic conditions, while the overlying deposits correspond to the post-glacial climatic amelioration. Specifically, the Holocene deposits have smaller grain sizes, show pedogenetic modifications, and have higher amounts of organic matter. The identification of this boundary could represent a useful criterion for locating new sites, since it separates the archaeologically sterile deposits from the deposits reflecting human occupation. Zárata (1997) identified this boundary at Cerro El Sombrero (in the central-east Tandilia range); although in this case the underlying deposits consisted of a truncated paleosol that, according to the author, would have developed prior to the Late Glacial Maximum. Similar deposits were also identified by Martínez et al. (1997) at the site of Abrigo Los Pinos, beneath the deposit underlying the occupation layer (U4) (Figure 3). These were assigned tentatively to the last interglacial (OIS 5e, c. 120 ky BP).

Paleoclimatic information regarding the Pleistocene-Holocene transition in the Pampean region is abundant and is derived from a wide range of proxy evidence. Based upon information from sites located in the adjacent Interserrana Plain, Politis, et al. (2004), Martínez (2001), and Grill

et al. (2007) postulated that initial population took place under arid to semi-arid climatic conditions. The human populations that would have lived in this area c. 11,000-10,000 BP would therefore have done so under colder, dry sub-humid conditions (Quattrocchio et al., 2008; Tonni and Cione, 1994; Prieto, 1996; Tonni et al., 1999, 2003). Between c. 10,500 and 9000 BP a change in these dry conditions towards higher rainfall levels and temperatures has been indicated by numerous proxy records from the Pampas region (Tonni and Fidalgo, 1978; Prado et al., 1987; Alberdi et al., 1989; Tonni y Cione, 1994; Prieto, 1996; Zárate et al., 1995, 2000; Zarate, 1997; Borrero et al., 1998; Bonadonna et al., 1999; Tonni et al., 1999; Martinez, 1999, 2001, 2007; Prieto et al., 2004; Grill et al., 2007; Gutierrez y Martínez, 2008; Gutierrez et al., 2010).

Phytoliths and diatoms were the most abundant botanical remains recovered from the studied profiles. The analysis of phytoliths contributed to the knowledge of the paleoclimatic conditions prevailing in the Pleistocene-Holocene transition, as well as to the discussion of some aspects of plant management in these early times where other botanical records are scarce. The sensibility of phytoliths to microenvironmental changes may explain not only the variations of elements produced by local and current vegetation inside the caves, but also those produced by anthropic remains. In the Loberia I site 1, the presence of pooid and danthonioid elements in the basal and middle zones (I and II) associated with the alternation of panicoid/chloridoid of the zone I suggest that climatic conditions were dry and temperate during the Pleistocene-Holocene transition. On the other hand, the Cueva Tixi sequence shows an increase of pooid (including festucoid) and chloridoid phytoliths in the basal sample (F). Cueva Tixi profile indicates colder conditions to the early Holocene (zone III) than Loberia I site 1. The presence of cold weather and dry conditions has also been reported for the early Holocene-Pleistocene by regional studies of other proxies (Prado, et al., 1987; Bonadonna, et al., 1995; Tonni, et al., 1999).

Variations in the phytolith analysis among the two sites can be attributed to the position of each cave in the landscape and its differential human use. For instance, one question that needs further discussion is whether phytoliths observed in Loberia I site 1 were the result of anthropic management or a consequence of the microclimate of the site and its altitudinal context. In this sense, an important variable to consider is the position in the landscape of each of the caves: while Lobería is located in low landscapes (98 m in height), Cueva Tixi is situated in a higher and more exposed (176 masl). Hence, these differences in altitude and exposition may be influencing the variability of the botanical microremains inside the sites. The presence of globular elements in early times may indicate human activity, but must be considered that are also present in the soil parent material. Recent evidence shows that globular elements are present along the complete sequences

inside the sites, and they are not coincident with megathermic elements, suggesting a possible introduction by human groups (Colobig, 2011).

Diatom assemblages were present in Lobería I and Cueva Tixi successions since late Pleistocene (before c. 10,000 ¹⁴C BP), while in Cueva El Abra sediments younger than c. 2900 ¹⁴C BP did not contain diatom valves. In the Loberia I sequence, dominant diatoms from unit 7 (late Pleistocene) indicated the lowest moisture availability conditions from all the sequence, as signaled by the aerial diatom *Diadesmis gallica*. Overall, the conditions in this cave were drier during the late Pleistocene and were followed by a climatic amelioration since the Pleistocene-Holocene transition. The dominance of *Denticula elegans* and *Nitzschia hantzschiana* at the basal levels (Unit F, figure 4) of Cueva Tixi (before c. 10,000 ¹⁴C BP) also indicates low moisture environmental conditions for the Pleistocene-Holocene transition, as these are very tolerant to long dry periods. The absence of diatoms at the basal levels of Cueva el Abra (late Pleistocene to mid Holocene) and their appearance through the late Holocene, are suggesting that this site also evolved from dry conditions during the Pleistocene-Holocene transition to wetter conditions through late Holocene. Hence, although local constraints cannot be discarded, the coincidence in the inferences made from these three sites could be signaling dry conditions for the late Pleistocene in this region. The possible climatic amelioration occurred during the Holocene is also supported by Cueva tixi record, where an increase in diatom assemblages indicating wetter conditions was recorded through middle Holocene (Unit C).

Hearts are a common feature in early settlements of most of the sites. Charcoal fragments recorded in Cueva El Abra, dated 9800 BP were identified as *Salix* sp. This taxon was part of a riparian landscape with gallery forests and is characteristic of the few forest communities that occur in the area. *Salix* sp. is very common across rivers and streams and the flood pulse is a key-factor in the establishment of the *Salix* forests (Liotta, 2001). This species lives in subhumid and semiarid regions under a temperate to warm-temperate climate. Furthermore, the soil should be moist, sandy, and with good drainage. The Cueva El Abra site is located at the foot of the Arroyo La Vigilancia valley, which contains a small river course that lacks flow during the dry seasons. This river reflects the basic drainage characteristics of the eastern Tandilia range, where no larger, permanent rivers exist. The gallery forests that characterize the vegetation in these valleys would have required a precipitation regimen sufficient to maintain permanent river flow, an unlikely scenario given the drier climatic conditions at the end of the Pleistocene. However, the presence of *Salix* charcoal supports the idea of a wetter rainfall regimen, which would have been established around 10,000 BP in these mountains. This is in agreement with the recording of new mammalian taxa that make up non-analog associations. The presence of species characteristic of subtropical environments, such as

Cavia aperea, *Monodelphys* sp., *Dasypus hybridus*, and *Oxymycterus* sp., could be considered as components of an emerging, expanded biological richness that developed with the arrival of new environmental conditions, which may have characterized a large ecotonal area.

The archaeological sites discussed here comprised part of the same mobility system, and also reflect social processes of territorial consolidation that involved other sites of similar ages in the adjacent regions of the Tandilia range and the Interserrana Plain. A strong relationship with the Paso Otero 5 site was especially noticed, as a consequence of the types of stone utilized (orthoquartzites from the Sierras Bayas Group, basalts from the Atlantic coast, and silicified limestones from southern Uruguay and/or the Pampas). The Paso Otero 5 site also shows signs of reutilization by very mobile groups during the period 10,440- 9400 BP, and is considered as a locus of specialized hunting and/or scavenging activities related to megamammals (Martínez and Gutiérrez, 2011).

The archaeological evidence found in the studied sites indicates that the early population of the area occurred at the end of the Pleistocene, in coincidence with the period of major climatic and environmental changes, such as the process of deglaciation that originated the extinction of part of the fauna. Hunter-gatherer societies were repeatedly present in the eastern mountainous environment, but as part of much broader mobility ranges that included the adjacent Interserrana Plain and the Atlantic Ocean coast. The search for resources, especially stone for production of tools, involved an area of social interaction that included territories as far away as the Río de la Plata river basin in present-day Uruguay (Flegenheimer et al., 2003)

Although records of occupations earlier than those discussed here do exist, both on the Interserrana Plain (Arroyo Seco, Site 2) and in the Tandilia range (Cerro La China sites 1 and 2 and Cerro El Sombrero) (Flegenheimer, 1987; Zárate and Flegenheimer, 1991) (Figure 1), they come from different stratigraphic contexts and not all are specifically from cave sites. The particular stratigraphic situation found in all nine of the sites discussed in this work, located in the same micro-region, reflect the magnitude of post-glacial climate change. This is represented by the stratigraphic unconformity and the development of new conditions in terms of sedimentary dynamics around 10,500 BP. The change to wetter conditions would also have generated a very advantageous pattern of resource availability in the mountainous landscapes (water, vegetation, and fauna), thereby attracting human occupation. This is manifested by the high density of sites represented in the same valley, and in the recording of recurrent episodes of use of certain strategic temporary sites.

8. Conclusion

The late Pleistocene-Holocene boundary is represented in the stratigraphical sequences of caves and rockshelters from eastern Tandilia range by a clear discordance. It is dated between 10,500 and 9600 BP, in coincidence with the initial human settlement. The identification of this unconformity could serve as a useful criterion for locating new occupations, since it separates the archaeologically sterile deposits from those that contain human occupations. Despite the local micro-environmental conditions recorded in caves and rockshelters, a regional generalized trend was inferred from the different evidence analyzed. In this sense, new paleobotanical records also indicate an abrupt climatic change from more arid late Pleistocene conditions to post-glacial climatic amelioration occurred towards c. 10,500 BP.

The Tandilia mountain range represented an optimal environment that favoured early population of the Pampas region. For the micro-region of the eastern Tandilia range, human groups were interested in occupying particular mountainous loci during the process of territorial consolidation of the Pampas. The environmental change towards warmer and wetter climatic conditions led to a wider range of natural resources, which included a concentrated degree of biodiversity in flora and fauna from both continental and coastal zones, as well as the abundance of fresh water. This in turn appears to have led to both higher density and greater intensity of human occupation in the eastern Tandilia range.

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FIGURE CAPTIONS

Figure 1. Location of the study area in Buenos Aires province. Some archaeological sites used as reference: 1. Cerro La China. 2. Cerro El Sombrero. 3. Arroyo Seco. 4. Paso Otero 5. Inferred paleocoast at c. 11,000 BP which correspond to c. -70 masl (after Guilderson *et al.*, 2000) is indicated on Atlantic ocean.

Figure 2. Location of sites and archaeological localities in Eastern Tandilia Range.

Figure 3. Stratigraphical profiles of the sites (modified of Martínez and Osterrieth, 2003 and Mazzanti *et al.*, 2010).

Figure 4. Stratigraphic sequence of Burucuyá cave composed of late Pleistocene to Holocene sandy-clayed silts, diatomaceous silts and a diatomite layer. A remarkable late Pleistocene-Holocene boundary related to footprints of mammals (ichnites) can be observed.

Figure 5. Relative frequency diagram of diatom composition and moisture groups for the sequence from Loberia I site.

Figure 6. Relative frequency diagram of diatom composition and moisture groups from: A) Cueva Tixi, and B) Cueva El Abra sites.

Figure 7. Dominant diatom species: 1) *Achnanthes coarctata*, 2) *Amphora pediculus*; 3) *Caloneis hyalina*; 4) *Cocconeis placentula*; 5) *Denticula elegans*; 6) *Diadismis contenta*; 7) *D. gallica*; 8) *Eunotia* sp.; 9) *Frankophila similioides*; 10) *Hantzschia amphyoaxis*; 11) *Hyalodiscus* sp.; 12) *Luthicola mutica*; 13) *Melosira* sp.; 14) *Melosira dickiei*; 15) *Navicula atomus*; 16) *N. fragilaroides*; 17) *Navicula* cf. *minima*; 18) *Nitzschia hantzschiana*; 19) *N. inconspicua*; 20) *N. terrestris*; 21) *Orthoseira dendroteres*; 22) *O. roeseana*; 23) *Pinnularia borealis*; 24) *P. subcapitata*. Scale: 5 μm .

Figure 8. Phytolith Diagram of Archaeological locality Loberia I, Site 1.

Figure 9. Different non articulated and articulated phytoliths observed in the studied profile samples. 1-2 Saddle. 3, 11 Elongated elements with wavy contour. 4, 6-7, 16-17 Truncated cones. 5, 14-15 Bilobates. 8 Dicot affinity element. 9, 12 Articulated phytoliths. 10 Elongated element with smooth contour. 13 Globular. 18 Elongated element with serrate contour. Scale bar in 11: 20 μm .

Figure 10. Cluster analysis of Archaeological locality Loberia I, Site 1 and Cueva Tixi sites using correlation coefficient (cophenetic correlation coefficient: 0.8424).

Table 1. Field descriptions and laboratory data for Loberia I profile. Notes: AU= Allostratigraphic Unit; SH=Soil Horizon; AC= Archaeological Component; C=Color (dry); OM=Organic Matter; P=Phosphorus. ▲=Archaeological Materials.

Table 2: Radiocarbon dates from the studied sites.

Table 3. Faunal record of paleoindian levels of Cueva Tixi and Cueva El Abra. † Extinct species. (1) A *Monodelphis* species different to *M. dimidiata* (see Goin 2001). (2) A species undetermined of *Canis*, may be *Canis avus* young specimens. (3) A species undetermined of *Canis*. (4) At least 4 species. In parentheses are the numbers of armadillo plates or bird eggshells.

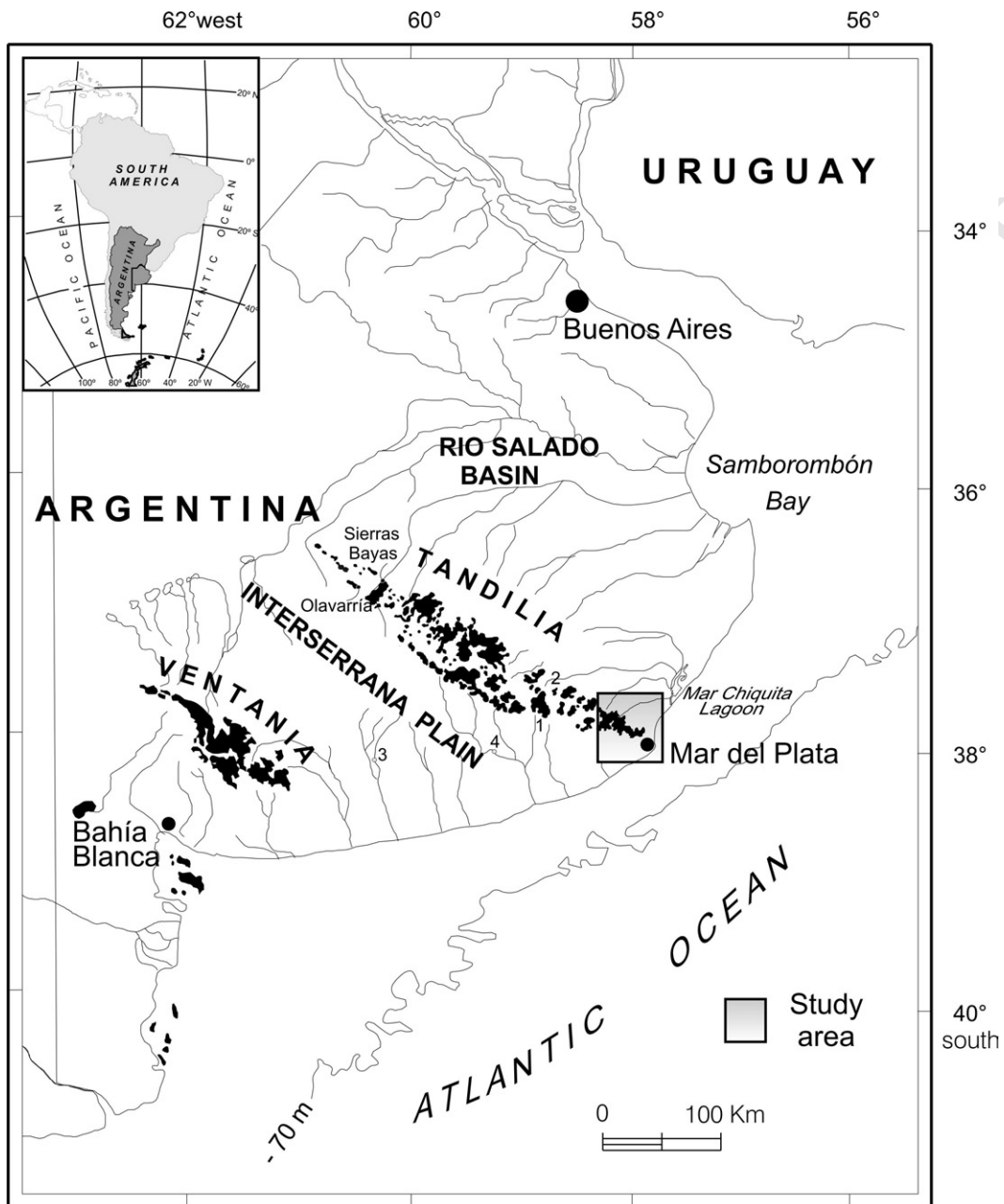
AU	Depth (cm)	SH	AC	Sand %	Silt %	Clay %	C (10YR)	pH	OM	N total (g kg ⁻¹)	P (ppm)	
1	6.5	AC		32	49	19	3/2	3,98	9.96	6,66	1389.6	
2	14.5	2AC	▲	32	49	19	3/3	3,93	8.94	4,56	1361.1	
3	15	3C	▲	32	43	25	3/1.5	5,39	10.83	6,99	1591.8	
4	25	4C	▲	38	40	22	3/3	4,15	6.89	3,78	679.9	
5	a	27	5Cb	▲	39	40	21	3.5/3	4,19	5.88	3,64	729.9
	b	34			36	43	21	4/2	4,05	6,53	3,23	553.7
	c	40			37	46	17	4/2.5	4,04	5.86	3,10	394.8
	d	48			37	48	15	4/2	4,08	4.88	3,30	539.2
	e	55			43	43	14	4/2.5	3,88	5.00	2,56	528.5
6	a	94		▲	40	50	10	4/3	3,70	4.74	2,26	967.5
	b	104		▲	45	43	12	4/2	4,04	3.97	2,26	1157.9
7	180			55	38	7	7/3	4,17	6.69	2,28	1208.7	

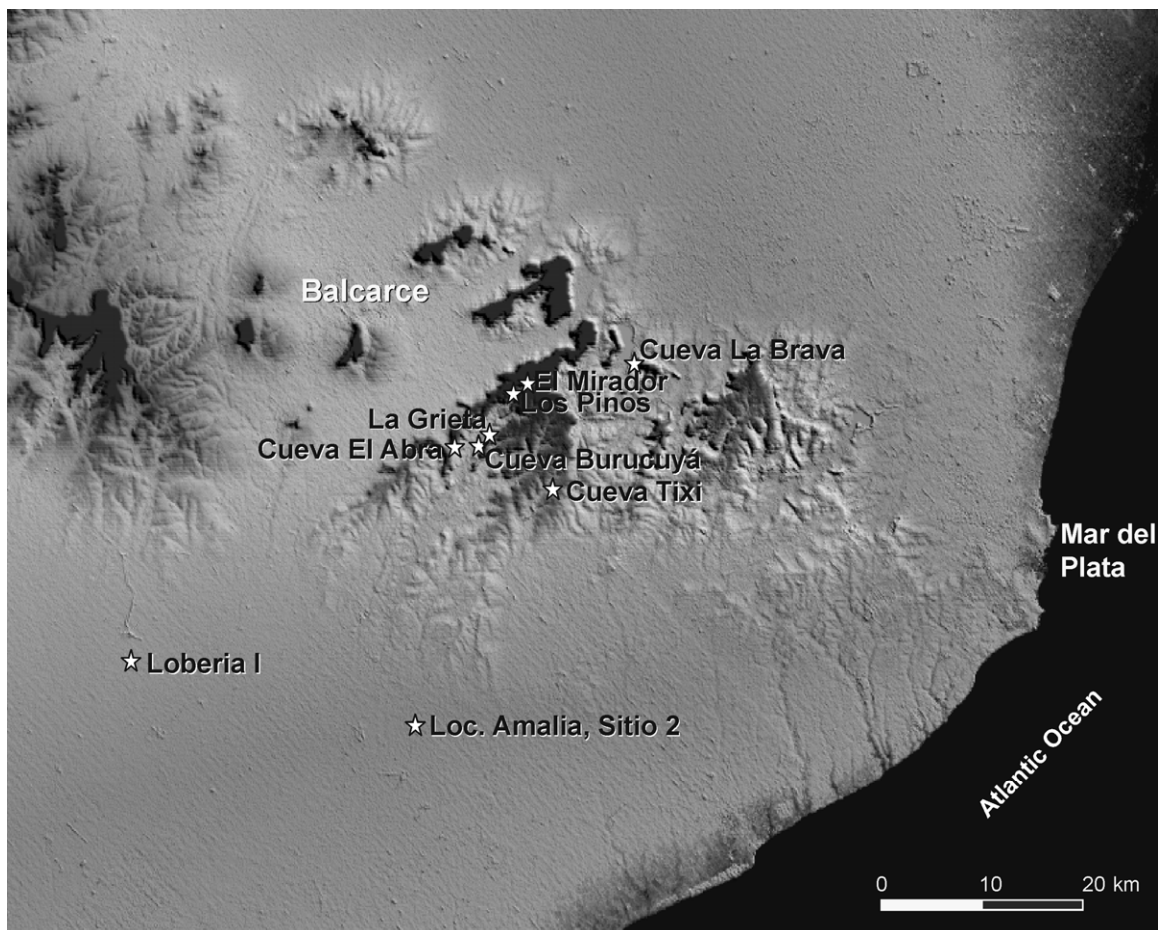
Table 2

Site	¹⁴ C Age (yr BP) Charcoal	Lab
Cueva Tixi	10,045 ± 95 (AMS)	AA-12131
	10,375 ± 90 (AMS)	AA-12130
Abrigo Los Pinos	9570 ± 150	LP-630
	10,465 ± 65 (AMS)	AA-24045
	10,415 ± 70 (AMS)	AA-24046
Cueva Burucuyá	10,000 ± 120	LP-863
	10,672 ± 56 (AMS)	AA-94640
Cueva El Abra	9834 ± 65 (AMS)	AA-38098
	10,270 ± 200 (AMS)	AA-94641
Cueva La Brava	9670 ± 120	LP-550
	10,178 ± 54 (AMS)	AA-94639
Amalia Sitio 2	10,425 ± 75 (AMS)	AA-35499
Lobería I Sitio 1	9787 ± 81 (AMS)	AA-81063

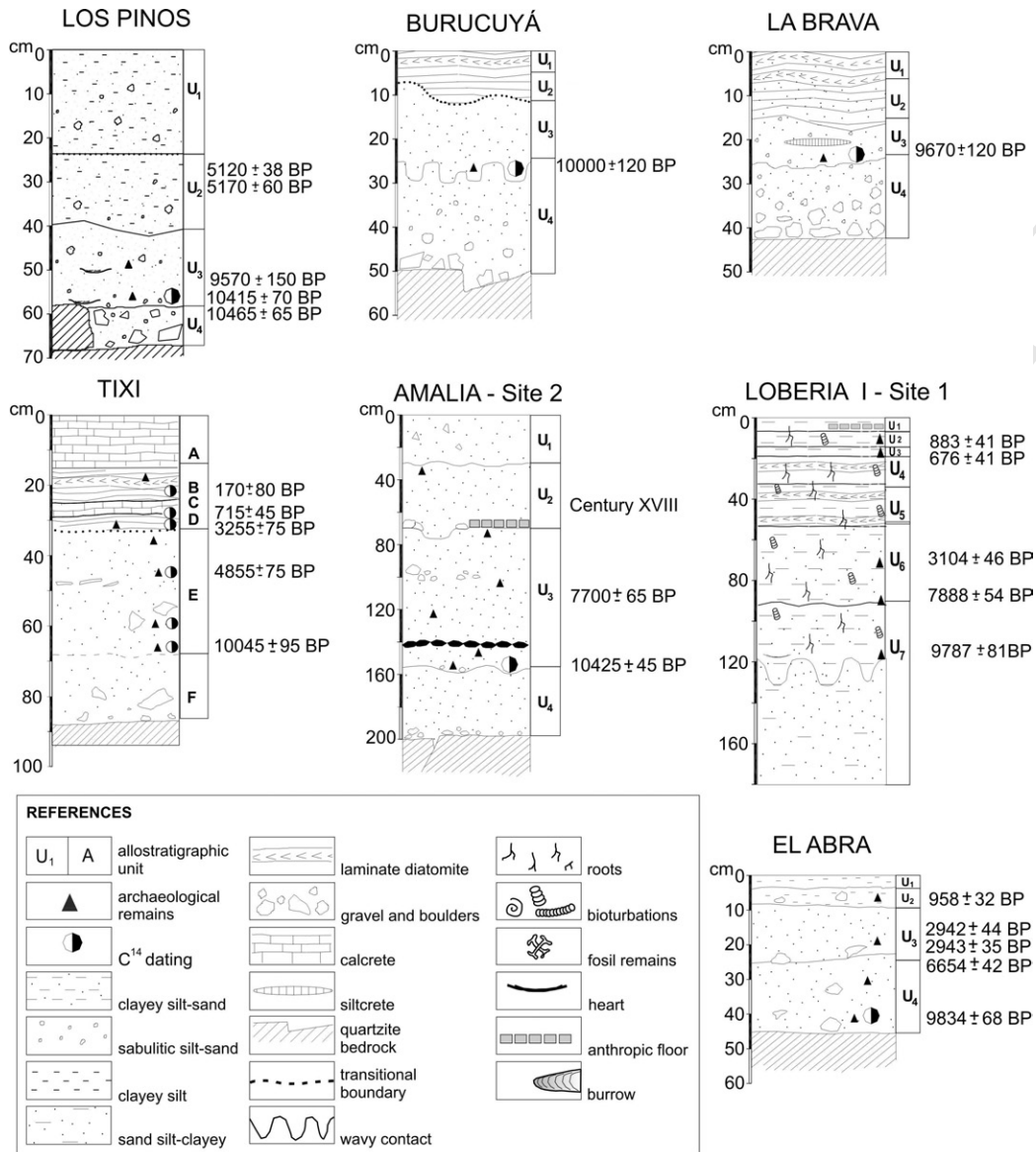
	Cueva Tixi NISP	Cueva El Abra NISP
<i>Lestodelphys halli</i>	14	
†(1) <i>Monodelphis</i> sp.	5	
<i>Thylamys</i> sp.	6	
Didelphidae indet.	3	
<i>Akodon azarae</i>	41	
<i>Calomys</i> sp.	3	
<i>Holochilus brasiliensis</i>	3	1
<i>Necromys obscurus</i>	27	
<i>Oxymycterus rutilans</i>	9	
<i>Reithrodon auritus</i>	464	
<i>Cavia aperea</i>	17	
<i>Ctenomys talarum</i>	190	8
<i>Dolichotis patagonum</i>	4	3
† <i>Galea tixiensis</i>	41	
<i>Lagostomus maximus</i>	31	12
<i>Myocastor coypus</i>	4	
† <i>Canis avus</i>	8	
(2) <i>Canis</i> sp. A	1	2
(3) <i>Canis</i> sp. B	19	
<i>Lycalopex gimnocercus</i>	1	
<i>Puma concolor</i>	2	
<i>Chaetophractus villosus</i>	6 (412)	96
<i>Dasypus hybridus</i>	6	
† <i>Eutatus seguini</i>	8 (20)	
<i>Zaedyus pichiy</i>	232	13
<i>Lama guanicoe</i>	51	
<i>Ozotoceros bezoarticus</i>	23	1
<i>Tupinambis</i> cf. <i>merianae</i>	5	
<i>Bothrops</i> sp.	4	
Colubridae indet.	1	1
Amphibia indet.	1	
<i>Rhea americana</i>	7 (31)	
(4) Aves indet	(24)	1

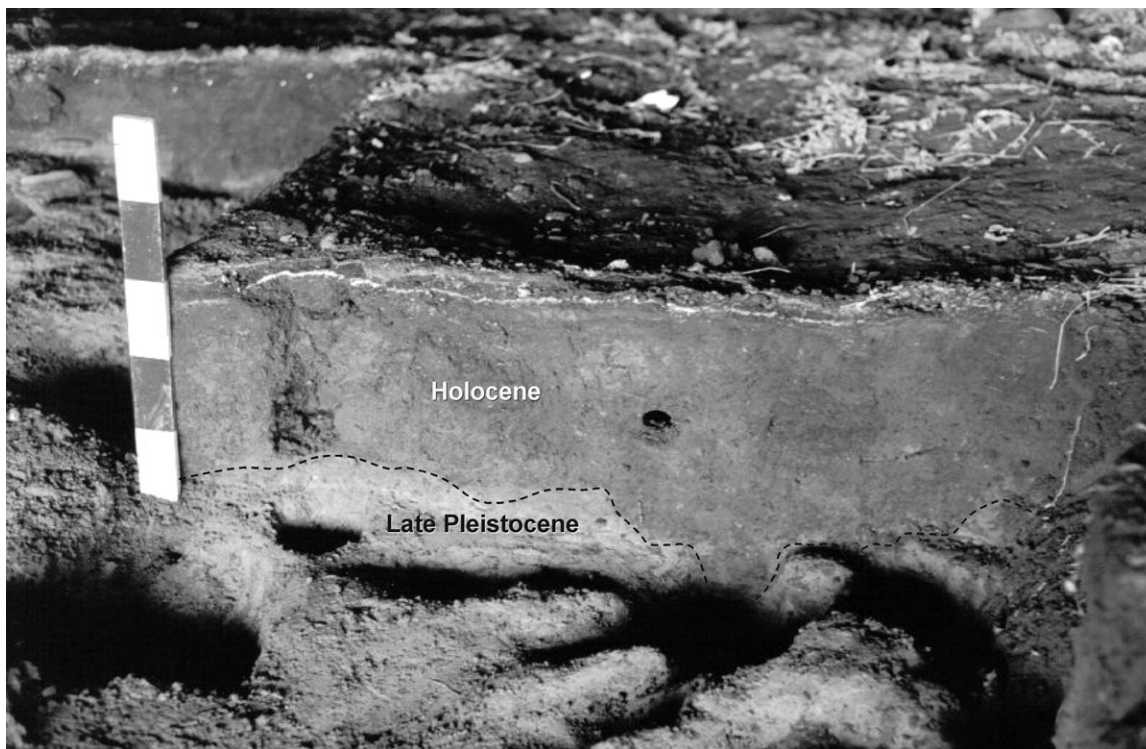
Table 3



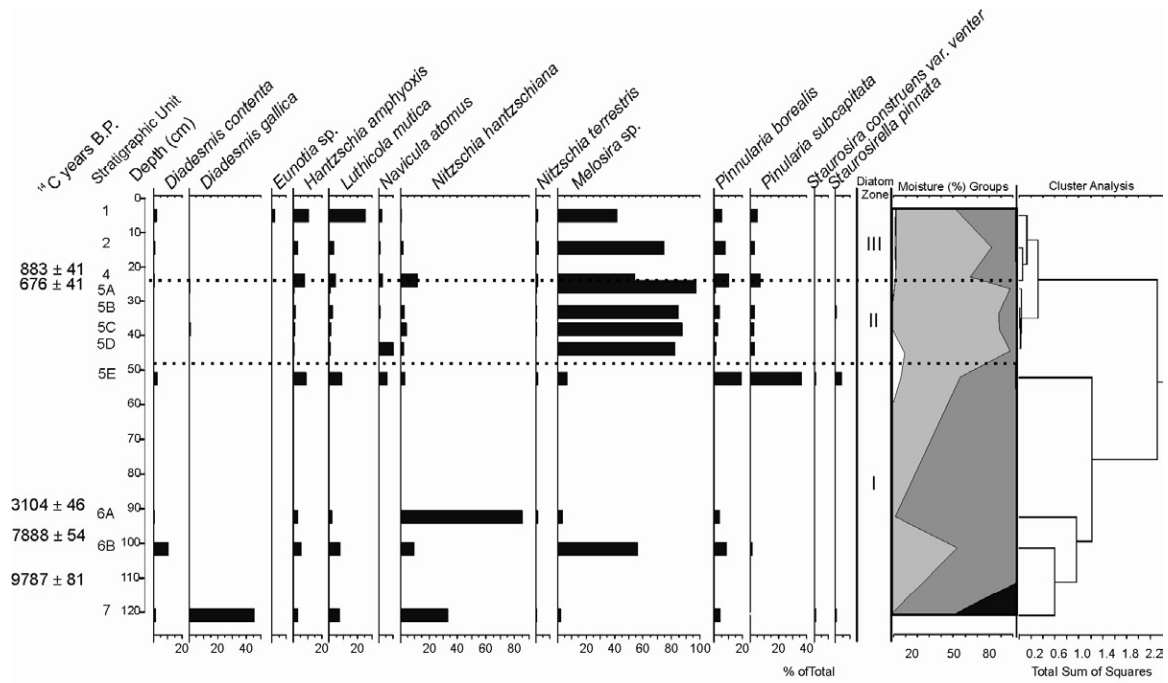


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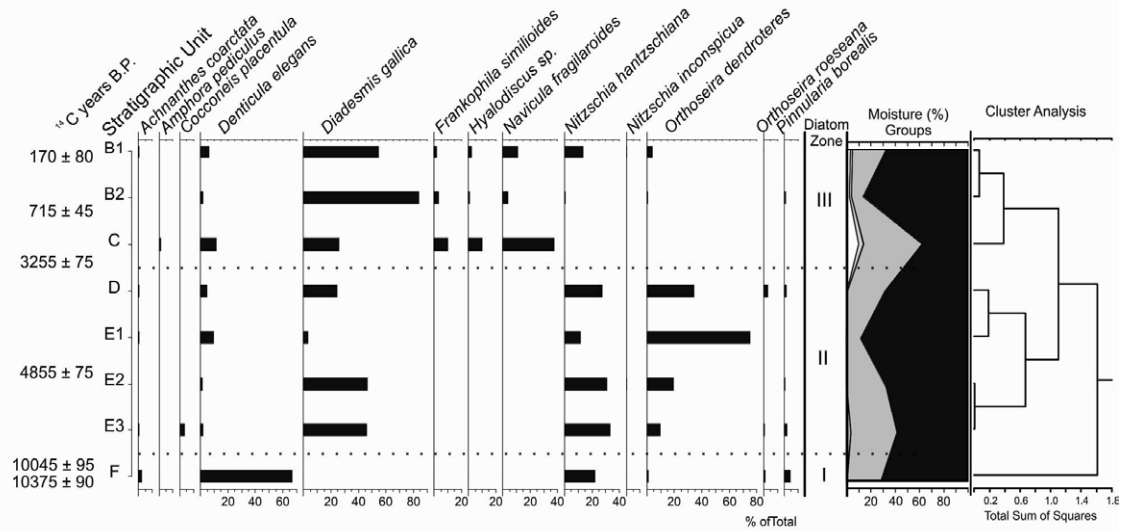
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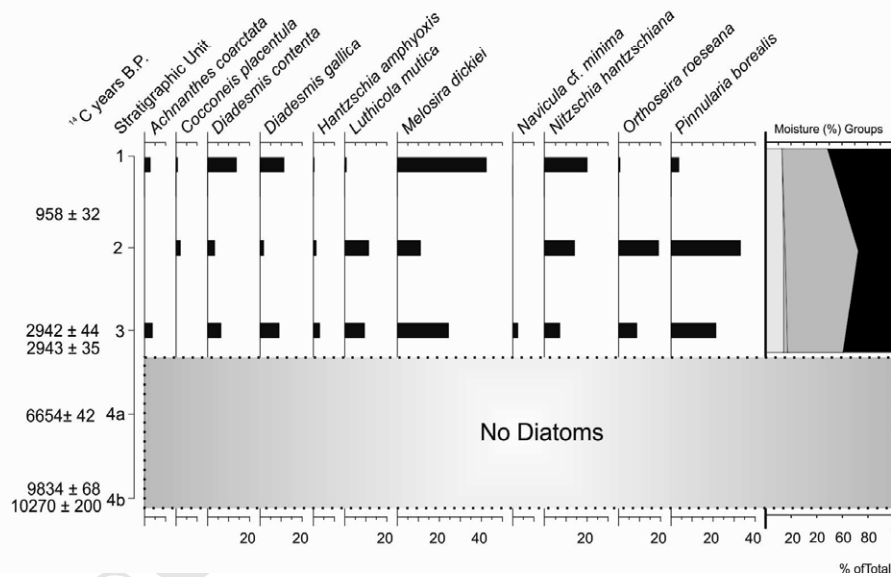
Classification according to moisture requirements (Van Dam et al., 1994)

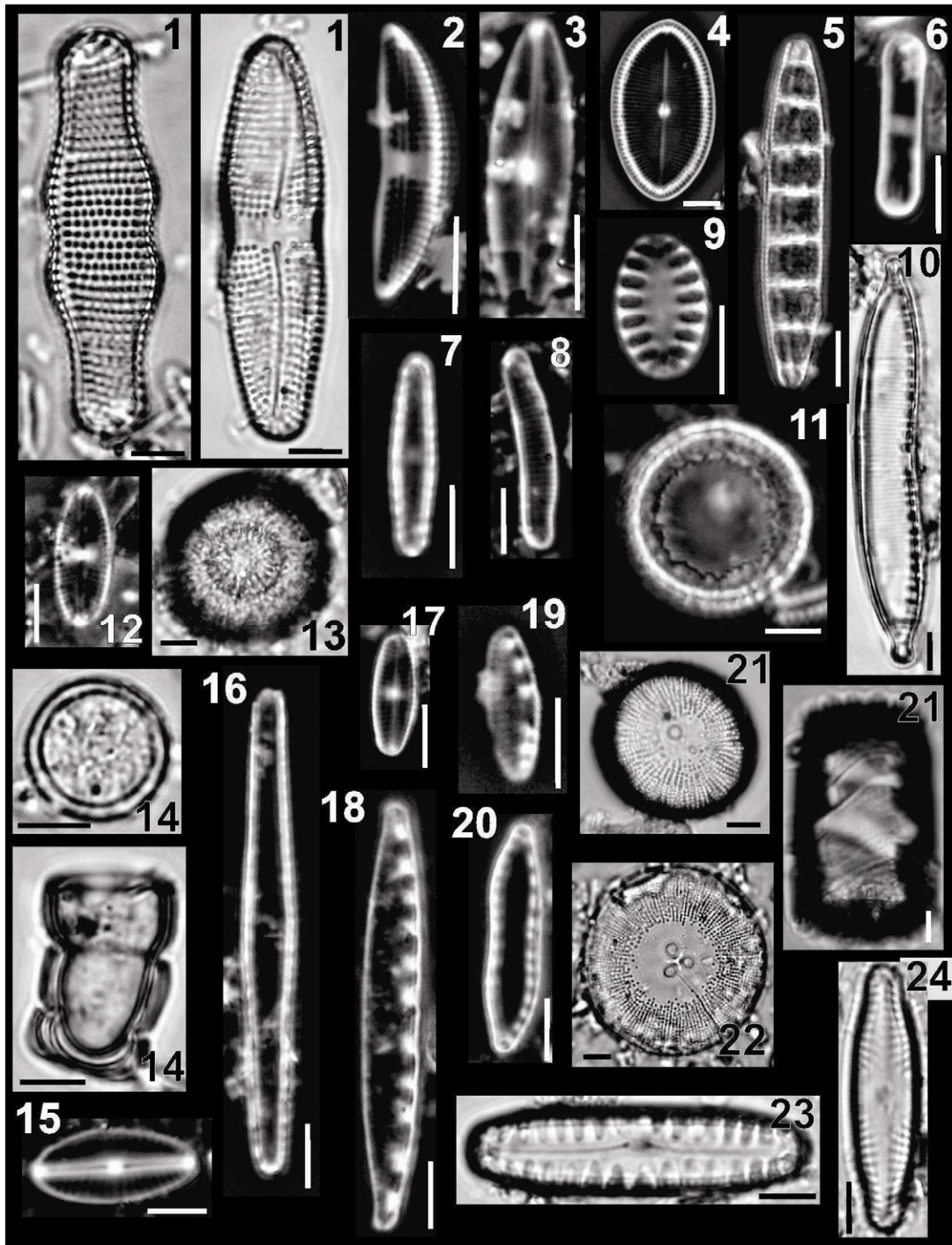
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|--|---|
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| <input type="checkbox"/> Never, or only very rarely, occurring outside water bodies. | <input type="checkbox"/> Mainly occurring on wet and moist or temporarily dry places. |
| <input type="checkbox"/> Mainly occurring in water bodies, sometimes on wet places. | <input type="checkbox"/> Nearly exclusively occurring outside water bodies. |

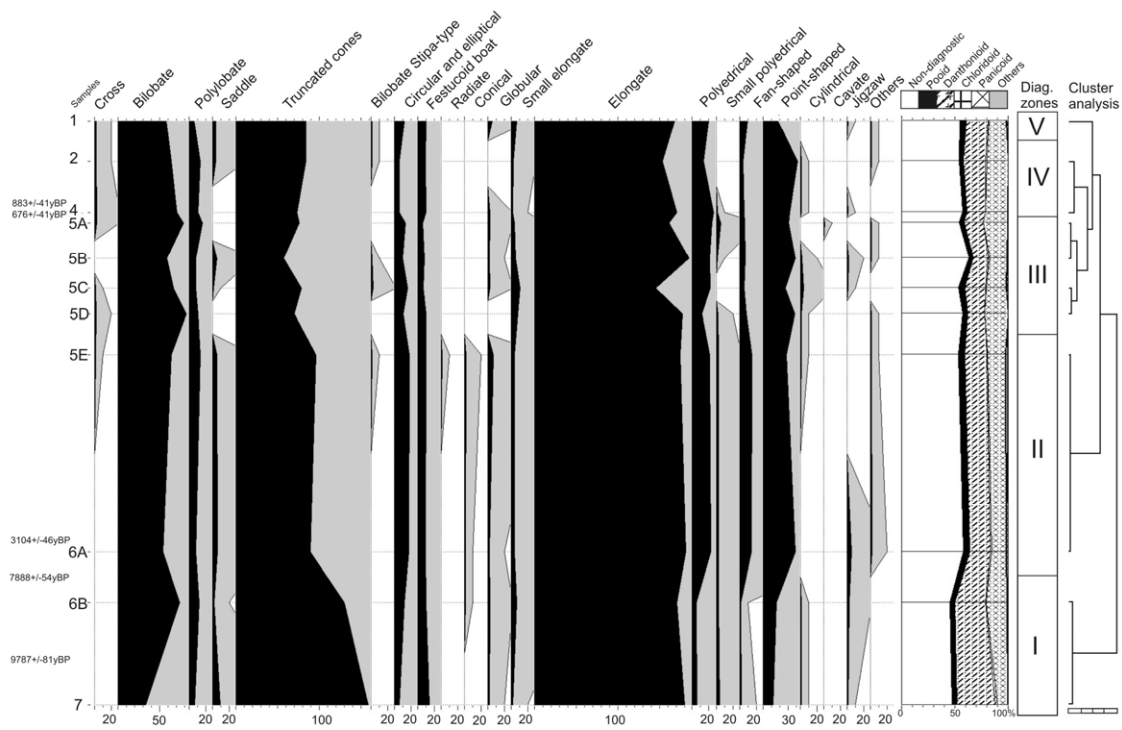
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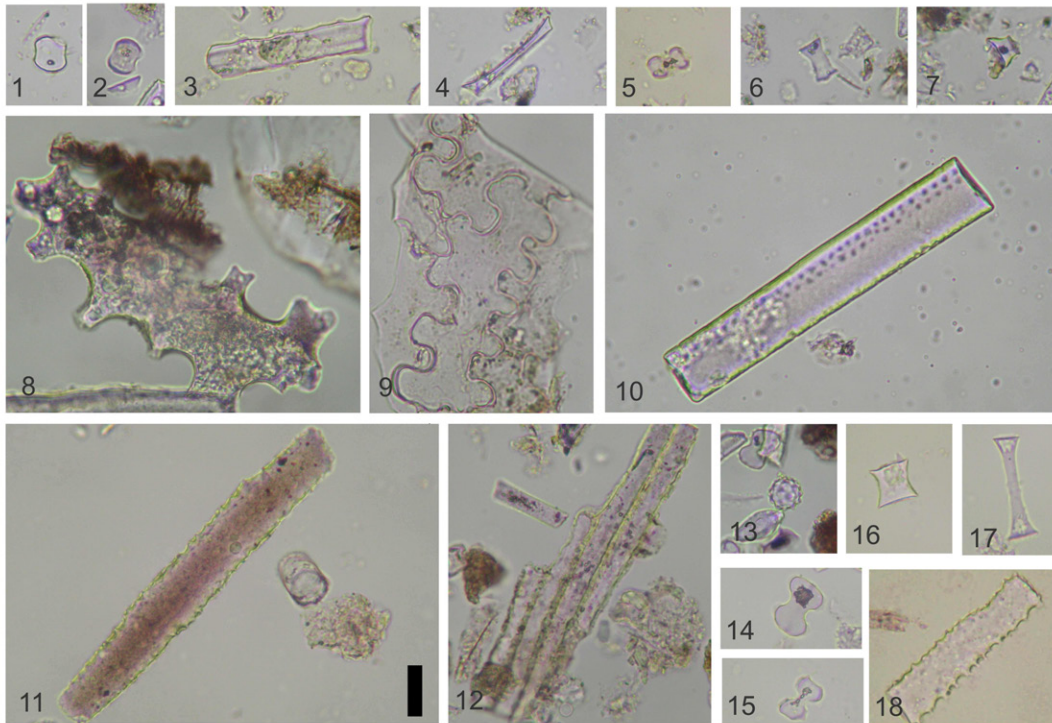


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