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SECONDARY CIRCULATION IN NATURAL STREAMS

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ABSTRACT

Secondary circulation which is sometimes referred to as secondary flow, secondary current or transverse current is an important phenomenon in natural streams and plays an important role in many natural processes in streams such as stream channel meander, bank erosion, bed scour, resuspension, and movement of sediment. Secondary circulation is that component of flow which is not in the main flow direction and is small as compared to the main flow velocity. A computerized data collection system for secondary circulation data acquisition in natural streams was developed and utilized in the field. The system includes an electromagnetic current meter, a micro-computer, an interface between the computer and the current meter, and a specially designed supporting structure. Secondary circulation data was collected in the Sangamon River near Mahomet, Illinois, utilizing the data collection system. A mathematical model for secondary circulation based on an existing model has been developed and tested against the data collected in the field. Model results generally reproduce similar secondary circulation patterns as observed from the field data but over-estimate the magnitudes of the velocities.

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i

KEYWORDS - computer; data acquisition; model; secondary circulation; secondary current; streams; transverse current; velocity

CONTENTS

	Page
Abstract	i
Introduction	1
Objectives	1
Acknowledgments	2
Review of Literature	2
Historical Background	3
Laboratory Measurements	4
Field Measurements	4
Indirect Techniques	5
Mathematical Modeling	6
Mathematical Model	8
Mathematical Derivations	8
Coordinate Transformation	9
Coordinate System	9
Scaling Factors	11
Derivation of the Equation for the $\mathtt{V}_{ig \xi}$ Velocity Component	14
Derivation of the Equation for the V $_\eta$ Velocity Component	20
Approximation of Shear Stresses	20
Computer Model	22
Phase I Analysis	25
Phase II Analysis	27
Grid System Construction	27
Solving for (Y,Z) Given (ξ,η)	28
Computation for $ abla_{\xi}$	31
Computation for $ extsf{V}_\eta$	
Field Data Collection	33
Field Data Collection Site	33
Instrumentation	35
Three-Dimensional Flow Measurement Setup	35
Marsh-McBirney 527 Electromagnetic Current Meter	35
Interface and Computer	37
Support System for the Current Meter	37
Rotating Bucket Current Meter	38
Suspended Sediment Samplers	38

Bed Load Sampler
Sonar Stage Meter
Data Collection Procedures
Secondary Circulation
Data Reduction40
Sampling Period41
Stream Discharge Measurement47
Suspended Sediment Concentrations and Particle Size Sampling $\dots 48$
Bed Load Measurement
Bed Material Sampling49
Stage Measurement
Summary of Field Data Collected50
Secondary Circulation50
Stream Discharge53
Suspended Sediment Concentration
Bed Load
Bed Material Characteristics59
Stage
Model Results
Comparison of Computed and Measured Velocities
Shear Stress Distribution68
Summary and Conclusions
References
Appendicesprinted in a separate volume
Appendix A. Computer Program
Appendix B. Measured Velocities
Appendix C. Computed Velocities

iii

INTRODUCTION

The main component of flow in any stream, river, or canal might be expected to be parallel to the channel bed in the downstream direction. However, because of a variety of reasons including the resistance flow due to stream banks, a rotational or cellular movement of flow will occur in the main channel. This persistent flow component that is not in the main flow direction has been termed secondary circulation, secondary flow, secondary current, or transverse current. This type of flow, which is always present in any open channel flow, is completely different from induced flow circulations in some pockets found on the side of open channels.

The direction of secondary flow is strongly associated with the primary velocity distribution. Areas of high velocities near the bed or bank are associated with outward secondary flow, while areas of low velocities are associated with inward secondary flow. This phenomenon has significant implications with respect to bank erosion and bed scour, since areas of high velocities are much more susceptible to erosion and scour.

Considerable agreement exists among researchers that secondary circulation is an important phenomenon in natural rivers and is partially responsible for the development of meander patterns, bank erosion, bed scour, resuspension and movement of sediment, and an increase in boundary shear and flow resistance (24, 26, 29, 37, 38, 39, 41, 44, 45, 47, 53, 54). For example, the existence of strong secondary circulation near a certain region close to the bank with a component of velocity directed toward the bank will indicate the location of potential bank erosion areas.

Objectives

The main objectives of this proposed research were:

- 1) To determine the magnitude and nature of secondary circulation in natural streams by direct measurement in the field.
- 2) To develop and verify a mathematical model for predicting secondary circulation in natural streams.
- 3) To determine the effects of secondary circulation on bank erosion by quantifying its effects on velocity and shear stress distribution, and on dispersion and transport of sediment in natural streams.

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REVIEW OF LITERATURE

It is important to have a clear understanding of what the term secondary circulation means before proceeding further. When researchers talk about secondary flow, they are actually describing the three-dimensional nature of open channel flow. Ciray (18) defined the secondary velocity as follows:

> When the magnitude of the vector composed by any two components of the local velocity vector in a three dimensional flow is small compared with the magnitude of the third component, the latter forms the main flow whereas the remaining two form the secondary.

Although this definition appears to be adequate to some extent, it lacks completeness. Perhaps more attention should be paid to the individual components of velocity than to the resultant of two vector quantities. In this line of thought, Chiu et al. (13) defined the secondary flow as the two components of flow which are in the transverse direction not coincident with the primary flow. Because of the presence of snags, obstructions, trees, etc., on the bed of a river, the velocity vector in a direction other than the

primary direction of the flow may be of greater magnitude than the primary velocity vector. Considering all these factors, Bhowmik (7) proposed the following definition:

In an open channel flow, any persistent component of velocity which is not parallel to or tangent to the center line of the channel and also not parallel to and in the same general downstream direction of the overall bed slope should be termed as secondary velocity.

Secondary current will be taken as the flow which is parallel to and also in the same direction as the secondary velocity. The definition stated above will ensure that the primary flow velocity is interpreted as the velocity in the downstream direction that coincides with the general direction of flow.

Historical Background

Historically, late nineteenth century hydraulic engineers, noting that the maximum velocity is below the water surface rather than at the water surface, offered the explanation that the frictional resistance at the airwater interface is responsible for this phenomenon. However, many researchers working at the time did not accept this idea, and they tried and were successful in correctly identifying the mechanisms of secondary circulation. They also indicated that these mechanisms might be responsible for the lower flow velocity at the air-water interface. These researchers did not identify secondary circulation by name, but they recognized its presence in concept. They include Bazin and Barcy in 1865, Gerrard in 1878, Francis in 1878, Wood in 1879, Stearns in 1883, and Cunningham in 1882 (7, 19).

In 1909 Gibson (30) analyzed the then-existing literature and offered his thoughts on why the filament of maximum velocity is below the water surface. He said that this phenomenon is "... due to the action of the sides of the channel in producing transverse currents inwards along the surface and outwards along the bed of the stream, thus distributing a layer of slowly moving water over the central part of the stream." He also indicated that the presence of curves in the stream's alignment can and will modify this phenomenon. He showed that the secondary flow velocity is about 5 percent of the main stream velocity.

It is noteworthy that almost 100 years after the initial postulations about the effect of wall roughness on the transverse flow phenomenon that were

made by Wood (59), Stearns (55), Francis (25), and others, Perkins (49) showed that the mean streamwise vorticity in turbulent flow arises both from mean flow skewing and from the inhomogeneity of anisotropic wall turbulence.

In the early twentieth century many others also worked on the transverse flow phenomenon. In 1917, Helstrom estimated the secondary flow velocity to be about 11 percent of the main stream velocity (32). In 1930 Nikuradse published results showing that secondary flow exists in noncircular conduits (27).

Laboratory Measurements

Once the concept of secondary flow was recognized, researchers tried to measure the secondary circulation, especially in laboratory channels under ideal flow conditions. The simplest method is of course to use dye or neutrally buoyant particles to measure qualitatively the existence and the direction of secondary currents (3, 58). Another fairly sophisticated instrument used to measure secondary current in laboratory experiments is the hot-wire anemometer. Researchers such as Brundrett and Baines (8), Gessner and Jones (28), Hoagland (33), Tracy (57), and Liggett et al. (42) utilized hot wire anemometry systems to measure components of secondary currents. Launder and Ying (40) collected turbulence-induced secondary current data by using a DISA hot-wire probe. Muller (43) reported the results of laboratory experiments where a Laser Doppler Anemometer was used to measure the vertical and axial component of flow velocity. The presence and absence of secondary flow were interpolated on the basis of the increase or decrease of the turbulence and momentum exchange. All of Muller's experiments were conducted for supercritical flow ranges, i.e., the Froude number was more than unity. However, all these research projects were confined to laboratory experimentations.

Field Measurements

While various researchers have measured secondary circulation in laboratory experiments, field instrumentations and measurements are rare. Only a very limited number of attempts have been made to measure secondary circulation in the field.

Kanwisher and Lawson (36) have described an electromagnetic flow sensor for measuring flow in natural open channels, which can be modified to measure

both the main stream and lateral flow velocities. This electromagnetic flow sensor has the potential of being used in the field to measure secondary flow velocities.

One of the most serious and comprehensive attempts that was made by any group of researchers to measure secondary currents in gravel bed rivers was that made by Bathurst et al. (4, 5). Their measurements were made at four bends on the River Severn in Wales, United Kingdom. Direct measurements of secondary velocities were made with an electromagnetic flow meter that had the capability of measuring two mutually perpendicular components of velocity simultaneously. The measuring head was 2 inches (50 mm) in diameter, and the instrument was accurate to within 0.33 fps (10 mm/s). All the measurements were confined to bends. The instrument worked out fairly well. The primary velocity data were collected with an Ott C-31 current meter. Isovels were developed, and their shapes were compared with the shape and pattern of secondary cells that could be sketched on the basis of the shape of the primary isovels. The data indicated that the peak values of shear stress were associated with the core of the maximum velocity.

Thorne et al. (56) conducted other detailed bend-flow hydraulics measurements on the Fall River, Rocky Mountain National Park, Wyoming. Primary and lateral velocity components were measured using a Marsh-McBirney Model 511 two-component electromagnetic current meter. The current meter simultaneously measures two mutually perpendicular velocity components with an accuracy of 0.01 ft/sec (3 mm/s). Much more accurate data were recorded, and secondary cells were often found to exist at outer and inner banks. The device used by Thorne et al. could not record the orientation of the sensor probe, which remained a difficulty in the data reduction. Demissie et al. (20, 21) developed a measuring system in which the current meter is supported by a self-standing system and data are transmitted to a computer on the stream bank for processing and storage. The flow velocity components are measured by a Marsh-McBirney current meter, Model 527, which can also measure the bearing of the sensor from the magnetic north in a clockwise direction.

Indirect Techniques

Field data collected by Bathurst et al. (4, 5) indicated that it is possible to draw secondary flow cells on the basis of the primary velocity isovels. Such an attempt was made by Bhowmik (7) for the field data collected

from the Kaskaskia River in Illinois. Fig. 1 shows the isovels at a section where the secondary flow cells were sketched. Data shown are from the same cross section but for two flow conditions: low and high flows. If it is assumed that the isovels are a set of flexible membranes held in place by fluids in between them, then a bulging or deformity in their shape indicates the presence of some force acting normal to the face of the membranes. Thus, if the membranes bulge inward, it indicates the presence of a force from the outside to the inside and vice versa. The approximate locations of the secondary cells in Fig. 1 were drawn on the basis of this technique. Data from any other isovels can similarly be used to draw the patterns of secondary cells. Brundrett and Baines (8) and Liggett et al. (42) expressed some doubt as to the validity of sketching secondary flow cells on the basis of primary velocity isovels. Their conclusions were based on laboratory experiments. However, even though it is not a precise technique, it provides a good qualitative picture of secondary circulation in natural channels.

Recently Ikeda (35) indicated a technique of sketching the patterns of secondary current cells in sand bed laboratory channels on the basis of the nature and existence of bed undulations or topography. In a deformable channel such as a sand bed channel, the bed profile in the lateral direction may partially reflect the magnitude and direction of the secondary current cells near the bed. Only qualitative information is obtained by using isovels or bed topography.

Mathematical Modeling

A few researchers have worked on the development of mathematical models related to secondary circulation in open channels. However, most of their work has been confined to cases with flow around bends. A strong secondary current is generated in a bend because of the imbalance of horizontal fluid forces; thus it is easier to quantify this current for flow around bends than for straight reaches.

Einstein and Li (23) have shown theoretically that uniform flow in straight channels can exist only in the case of laminar flow. For turbulent flow, they have shown that secondary currents will be generated and that the nonexistence of parallel isovels results from the presence of secondary flows, especially near the frictional boundary. Some other researchers who have



Figure 1. Isovels and secondary flow cells in a natural river (7)

worked on the mathematical formulation of secondary currents are Ibragimov et al. (34), Ananian (2), Ciray (18), Prasad (50), Tracy (57), Liggett et al. (42), and Odgaard (46).

One of the most significant contributions to the mathematical modeling of secondary circulation was made by Chiu and his associates (10-17). Much of Chiu's work related to secondary circulation around channel bends. He also worked in the analyses of secondary currents in straight open channels, and has indicated that whenever the three components of shearing force are not in equilibrium everywhere in a channel, the presence of secondary current is inevitable.

By introducing a coordinate transformation function, Chiu et al. (12) developed a model which appears to have considerable potential for application to natural open channel flow problems. A curvilinear orthogonal coordinate system which is constructed by computed isovels and their orthogonal trajectories forms the basis for the model. The model was used to compute secondary flows on the basis of some hydraulic data collected from the Rio Grande Canal (12, 13). Later, different functions for approximating isovels and shear stresses were used to improve the model (14-17). Comparisons of model results with measured data on the East Fork River (Wyoming) showed reasonable agreement. Field data such as cross-sectional area, discharge, depth, water surface profile, and primary velocity distribution patterns are needed to use the model.

MATHEMATICAL MODEL

Mathematical Derivations

This part of the report discusses the derivation of equations used in a new mathematical model for calculating secondary velocity components. The model is based primarily on Chiu's work (10-17). Detailed derivations for some functions are presented here so the mathematical model can be clearly understood.

In general, the model consists of the momentum and the continuity equations. Once the isovels for a cross section are approximated by mathematical functions, the momentum and the continuity equations are then transformed into a curvilinear coordinate system, which is based on the simulated isovels and their orthogonal trajectories. After the transformation

of these equations, one of the secondary velocity components is computed directly from the momentum equation, and the other component is derived by solving the continuity equation. Shear stresses are assumed to vary quadratically from the boundary line to the point of maximum velocity in the new coordinate system. Finally the secondary velocity components are transformed back to the cartesian coordinate system by following inverse transformation rules.

Coordinate Transformation

Coordinate System

The curvilinear coordinate system consists of two components ξ and η as shown in Fig. 2, where a constant value of ξ is an equal velocity line (isovel) and η is the coordinate perpendicular to the ξ lines. Assuming that the primary velocity distribution follows the logarithmic rule, which is reasonable for most natural channels, the velocity distribution can be represented by:

$$u = \frac{u}{k} \ln \frac{\xi}{\xi o}$$
(1)

in which

u = the primary velocity

k = universal constant

 ξ = transformed coordinate

 ξ_0 = coefficient characterizing the velocity distribution of the primary flow

 $u_* = \sqrt{gRS_e} = mean shear velocity, where$

g = gravitational acceleration

R = hydraulic radius

S_e = energy slope

Eq. 1 defines the primary velocity distribution in a direction vertical to ξ curves in the channel cross section. Since u* and k are constant at a given cross section, Eq. 1 states that for a given ξ all the velocities on this curve will have the same magnitude, i.e., a constant ξ is an isovel curve. The curve $\xi = \xi_0$ is the line where primary velocities vanish. The coordinates of ξ and η in terms of Y and Z coordinates are given as (15):



Figure 2. Curvilinear coordinate system

$$\xi = Y(1-Z)^{\beta_i} \exp(\beta_i Z - Y + 1)$$

$$\beta_{i} \left[\frac{D + \delta_{y} + \varepsilon}{B_{i} + \delta_{i}} \right]^{2}$$

$$\eta = \pm \frac{1}{Z} \quad (|1 - Y|) \qquad \exp \left[Z + \beta_{i} \left(\frac{D + \delta_{y} + \varepsilon}{B_{i} + \delta_{i}} \right) Y \right]^{2} \qquad (3)$$

(2)

in which

$$Y = \frac{y + \delta y}{D + \delta y + \varepsilon}$$
(4)

$$Z = \frac{|z|}{B_{i} + \delta_{i}}$$
(5)

 β_i , δ_i , and δ_y are parameters determined from the distribution of the primary velocity; i = 1 or 2 for either the left or right half of a cross section; ε = the distance from the water surface to the point of maximum velocity; B_i = the width of either the left or right half of a cross section; and D is the water depth (Fig. 2). The point of maximum velocity is assumed to be either below the water surface (as shown in Fig. 2 when $\varepsilon > 0$), at the water surface (when $\varepsilon = 0$), or above the water surface (when $\varepsilon < 0$).

Scaling Factors

Eqs. 2 and 3 define a transformation relation between the new curvilinear coordinate system and the (Y, Z) coordinate system. The scaling factors are derived by referring to Fig. 3 and assuming that \mathbf{j} and \mathbf{k} are unit vectors in the Y and Z directions and tangent to ξ and η at point p; the directions of \mathbf{j} and \mathbf{k} are in the increasing direction of ξ , η ; ΔS_{η} and ΔS_{ξ} are the incremental distances along the positive ξ and η directions; and \mathbf{r} is the position vector to p from a fixed origin.



Figure 3. Relationship between the (Y-Z) coordinate system and the curvilinear coordinate system $(\xi - \eta)$.

The derivatives of \vec{r} with respect to ξ and η are derived as follows:

$$\frac{\partial \mathbf{r}}{\Delta \xi} = \lim_{\Delta \xi \to 0} \left(\frac{\Delta \mathbf{r}}{\Delta \xi} \right)_{\eta = \mathbf{c}}$$
$$= \lim_{\Delta \xi \to 0} \left(\frac{\Delta \mathbf{r}}{\Delta s \eta} \cdot \frac{\Delta s \eta}{\Delta \xi} \right)_{\eta = \mathbf{c}}$$
$$= \left(\frac{\partial s \eta}{\partial \xi} \right)_{\mathbf{p}} \cdot \left(\frac{\partial \mathbf{r}}{\partial s \eta} \right)_{\mathbf{p}}$$
$$= \left(\frac{\partial s \eta}{\partial \xi} \right)_{\mathbf{p}} \cdot \mathbf{k}$$

Set
$$h_{\xi} = \frac{\partial s_{\eta}}{\partial \xi}$$

Then

$$h_{\xi} \cdot \overline{k} = \frac{\partial \overline{r}}{\partial \xi}$$

Similarly

$$h_{\eta} \cdot \overline{j} = \frac{\partial \overline{r}}{\partial \eta}$$

On the other hand, the derivations of \vec{r} with respect to ξ and η can be written in terms of Y and Z as:

$$\frac{\partial \vec{r}}{\partial \xi} + \frac{\partial \vec{r}}{\partial \eta} = \left(\frac{\partial Y}{\partial \xi} + \frac{\partial Y}{\partial \eta}\right)\vec{j} + \left(\frac{\partial Z}{\partial \xi} + \frac{\partial Z}{\partial \eta}\right)\vec{k}$$

By using the above relations the following equations are derived:

$$h_{\xi}^{2} = \left(\frac{\partial Y}{\partial \xi}\right)^{2} + \left(\frac{\partial Z}{\partial \xi}\right)^{2}$$
(6)

and

$$h_{\eta}^{2} = \left(\frac{\partial Y}{\partial \eta}\right)^{2} + \left(\frac{\partial Z}{\partial \eta}\right)^{2}$$
(7)

The differentials of ξ and η with respect to Y and Z can be obtained from Eqs. 2 and 3 directly. They are:

$$\frac{\partial \xi}{\partial Y} = (1-Y)(1-Z)^{\beta_{i}} \exp (\beta_{i}Z - Y + 1)$$
(8)

$$\frac{\partial \xi}{\partial Z} = \beta_{i} (1-Z)^{\beta_{i}} Y \exp \left(\beta_{i} Z - Y + 1\right) \left(-\frac{Z}{1-Z}\right)$$
(9)

For
$$Y < 1$$
 $\frac{\partial \eta}{\partial Y} = \frac{1}{Z} C_i (1-Y)^{C_i} \exp (Z + C_i Y) (\frac{-Y}{1-Y})$ (10)

$$\frac{\partial \eta}{\partial Z} = \frac{1}{Z} (1-Y)^{C_{i}} \exp \left(Z + C_{i}Y\right) \left(1 - \frac{1}{Z}\right)$$
(11)

For
$$Y > 1$$
 $\frac{\partial \eta}{\partial Y} = \frac{1}{Z} C_1 (Y-1)^C i \exp (Z + C_1 Y) (\frac{-Y}{Y-1})$ (12)

$$\frac{\partial \eta}{\partial Z} = \frac{1}{Z} (Y-1)^{C_{i}} \exp (Z + C_{i}Y) (\frac{1-Z}{Z})$$
(13)

where

$$C_{i} = \beta_{i} \left[\frac{D + \delta_{y} + \varepsilon}{B_{i} + \delta_{i}} \right]^{2}$$
(14)

The differentials of Y and Z with respect to ξ and η can be obtained from equations 8 to 13 as follows:

$$\frac{\partial Y}{\partial \xi} = \frac{Y(1-Z)^2(1-Y)}{[(1-Y)^2(1-Z)^2 + \beta_i C_i Y^2 Z^2]} \cdot \frac{1}{\xi}$$
(15)

$$\frac{\partial Z}{\partial \xi} = \frac{-c_{i} Y^{2} Z (1-Z)}{[(1-Y)^{2}(1-Z)^{2} + \beta_{i} c_{i} Y^{2} Z^{2}]} \frac{1}{\xi}$$
(16)

$$\frac{\partial Y}{\partial \eta} = \frac{-\beta_{i} Y Z^{2} (1-Y)}{[(1-Y)^{2}(1-Z)^{2} + \beta_{i} C_{i} Y^{2} Z^{2}]} \cdot \frac{1}{\eta}$$
(17)

$$\frac{\partial Z}{\partial \eta} = \frac{-Z (1-Y)^2 (1-Z)}{[(1-Y)^2 (1-Z)^2 + \beta_i C_i Y^2 Z^2]} \cdot \frac{1}{\eta}$$
(18)

The scaling factors can now be derived from Eqs. 6 and 7, and Eqs. 15 to 18. As given by Chiu and Lin (16) the relations for ${\rm h}_\xi$ and ${\rm h}_\eta$ are:

$$h_{\xi} = \frac{(D + \delta_{y} + \varepsilon)(1 - Z) Y}{\xi \sqrt{(1 - Y)^{2}(1 - Z)^{2} + C_{i}^{2} Y^{2} Z^{2}}}$$
(19)

$$h_{\eta} = \frac{(B_{i} + \delta_{i})}{(D + \delta_{v} + \varepsilon)} \cdot \frac{\xi}{\eta} \cdot \frac{Z (1 - Y)}{Y (1 - Z)} h_{\xi}$$
(20)

Derivation of the Equation for the $\mathbf{V}_{\boldsymbol{\xi}}$ Velocity Component

The momentum equation in the curvilinear coordinate system is derived on the assumption that the control volume in Fig. 4 represents a volume of water in the curvilinear coordinate system.



Figure 4. A control volume in the curvilinear coordinate system

Since each $\xi = \xi_j$ line in Fig. 4 represents an equal primary velocity line, there is no change of primary velocity along this ξ line, i.e.,

$$\frac{\partial \mathbf{u}_{\mathbf{j}}}{\partial \eta} \bigg|_{\substack{\xi = \xi_{\mathbf{j}}}} = 0$$

where u_j is the primary velocity on the j^{th} isovel. The longitudinal component of the momentum equation in cartesian coordinates is given as :

$$\frac{\partial u}{\partial t} + \frac{\partial u}{\partial x} + \frac{\partial u}{\partial y} + \frac{\partial u}{\partial z} = -\frac{1}{\rho} \left[\frac{\partial \tau_{xx}}{\partial x} + \frac{\partial \tau_{yx}}{\partial y} + \frac{\partial \tau_{zx}}{\partial z} \right]$$
(21)

in which

x, y, z are the coordinates

- t = the time
- u = the primary velocity in x direction

v = the velocity components in y direction

- w = the velocity component in z direction
- ρ = the water density
- p = the hydrostatic pressure
- g = gravitational acceleration
- H = distance from the channel bottom to a datum
- τ_{ix} = is the shear stress on the ith face acting in the x direction in which i = x, y, or z.

This equation can be written in x, Y, Z coordinates directly using the transformation equations given in Eqs. 4 and 5. Since the partial derivatives are as shown in the following equation:

$$\partial Y = \partial y$$

$$\partial Z = \begin{cases} \partial z & \text{if } z > 0 \\ -\partial z & \text{if } z < 0 \end{cases}$$

Eq. 22 will change in sign only for z < 0.

The transformation relationship between Y, Z coordinates and ξ , η coordinates can be written in a general form as follows:

$$\vec{e}_{\xi} = h_{\xi} \frac{\partial \xi}{\partial Y} \vec{j} + h_{\xi} \frac{\partial \xi}{\partial Z} \vec{k}$$
(23)
$$\vec{e}_{\eta} = h_{\eta} \frac{\partial \eta}{\partial Y} \vec{j} + h_{\eta} \frac{\partial \eta}{\partial Z} \vec{k}$$

(22)

in which $\overline{e_{\xi}}$ and $\overline{e_{\eta}}$ are unit vectors in the ξ and η directions, respectively, and \overline{j} and \overline{k} are unit vectors in the Y and Z directions, respectively.

Using Eq. 23, the transformation relationships for the velocity components between the two coordinate systems can be written as:

$$V_{\xi} = h_{\xi} \frac{\partial \xi}{\partial Y} v + h_{\xi} \frac{\partial \xi}{\partial Z} w$$
(24)

(25)

$$V_{\eta} = h_{\eta} \frac{\partial \eta}{\partial Y} v + h_{\eta} \frac{\partial \eta}{\partial Z} w$$

where V_{ξ} and V_{η} are velocity components in the ξ and η directions, respectively. Since ξ and η are orthogonal to each other as are the Y and Z coordinates, the following relations exist:

$$h_{\eta} \frac{\partial \eta}{\partial Y} = -h_{\xi} \frac{\partial \xi}{\partial Z}$$

$$h_{\xi} \frac{\partial \xi}{\partial Y} = h_{\eta} \frac{\partial \eta}{\partial Z}$$

These can be proved by referring to Fig. 5.

$$h_{\eta} \frac{\partial \eta}{\partial Y} = \cos (\eta, Y) = \sin (Y, \xi)$$
$$h_{\xi} \frac{\partial \xi}{\partial Z} = \cos (\xi, Z) = \cos (Z, \eta) = -\sin (Y, \xi)$$



Figure 5. Definition sketch for derivatives between Y, Z and ξ , η

Therefore,

$$h_{\eta} \frac{\partial \eta}{\partial Y} = - h_{\xi} \frac{\partial \xi}{\partial Z}$$

By definition,

 $angle(Z,\eta) + angle(\eta,Y) = angle(\eta,Y) + angle(Y,\xi)$

Thus,

angle $(Z, \eta) = angle(Y, \xi)$.

Similarly,

$$h_{\xi} \frac{\partial \xi}{\partial Y} = \cos (\xi, Y)$$

$$h_{\eta} \frac{\partial \eta}{\partial Z} = \cos (\eta, Z)$$

Therefore

$$h_{\xi} \frac{\partial \xi}{\partial Y} = h_{\eta} \frac{\partial \eta}{\partial Z}$$

Now the left hand side (LHS) of the momentum equation (Eg. 21) can be rewritten as:

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} + w \frac{\partial u}{\partial z} = \frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + \left[v\left(\frac{\partial u}{\partial \xi}\frac{\partial \xi}{\partial Y} + \frac{\partial u}{\partial \eta}\frac{\partial \eta}{\partial Y}\right) + w\left(\frac{\partial u}{\partial \xi}\frac{\partial \xi}{\partial Z} + \frac{\partial u}{\partial \eta}\frac{\partial \eta}{\partial Z}\right)\right]$$

$$= \frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + \left[v\left(\frac{\partial u}{\partial \xi}\frac{\partial \xi}{\partial Y}\right) + w\left(\frac{\partial u}{\partial \xi}\frac{\partial \xi}{\partial Z}\right)\right]$$

$$= \frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + \left(v\frac{\partial \xi}{\partial Y} + w\frac{\partial \xi}{\partial Z}\right)\frac{\partial u}{\partial \xi}$$

$$= \frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + \frac{V_{\xi}}{h_{\xi}}\frac{\partial u}{\partial \xi} \qquad (26)$$

where only one velocity component, V_{ξ} , is retained. The transformation of the right hand side (RHS) of the momentum equation into ξ and η coordinates can also proceed as follows:

The shear stresses in the x direction can be written as

$$\begin{aligned} \tau_{\mathbf{Y}\mathbf{x}} &= (\tau_{\xi\mathbf{x}} \cdot \mathbf{\hat{e}}_{\xi} + \tau_{\eta\mathbf{x}} \cdot \mathbf{\hat{e}}_{\eta}) \cdot \mathbf{\hat{j}} \\ \tau_{\mathbf{Z}\mathbf{x}} &= (\tau_{\xi\mathbf{x}} \cdot \mathbf{\hat{e}}_{\xi} + \tau_{\eta\mathbf{x}} \cdot \mathbf{\hat{e}}_{\eta}) \cdot \mathbf{\hat{k}} \end{aligned}$$
(27)

Since there is no velocity change in the η direction on each ξ curve,

$$\tau_{\eta \mathrm{X}} = \mathrm{o} \tag{28}$$

Substituting this equation, as well as Eq. 23, into Eq. 27 results in the following equations:

$$\tau_{Yx} = \tau_{\xi x} \cdot h_{\xi} \frac{\partial \xi}{\partial Y}$$
(29)
$$\tau_{Zx} = \tau_{\xi x} \cdot h_{\xi} \frac{\partial \xi}{\partial Z}$$

Using the chain rule for derivatives, the following relations can be derived.

$$\frac{\partial \tau_{\mathbf{Y}\mathbf{x}}}{\partial \mathbf{Y}} = \frac{\partial \tau_{\mathbf{Y}\mathbf{x}}}{\partial \xi} \frac{\partial \xi}{\partial \mathbf{Y}} + \frac{\partial \tau_{\mathbf{Y}\mathbf{x}}}{\partial \eta} \frac{\partial \eta}{\partial \mathbf{Y}}$$

$$\frac{\partial \tau_{Zx}}{\partial Z} = \frac{\partial \tau_{Zx}}{\partial \xi} \frac{\partial \xi}{\partial Z} + \frac{\partial \tau_{Zx}}{\partial \eta} \frac{\partial \eta}{\partial Z}$$

Substituting Eq. 29 into Eq. 30 gives:

$$\frac{\partial \tau_{\mathbf{Y}\mathbf{x}}}{\partial \mathbf{Y}} = \frac{\partial}{\partial \xi} \left(\tau_{\xi\mathbf{x}} \cdot \mathbf{h}_{\xi} \right) \left(\frac{\partial \xi}{\partial \mathbf{Y}} \right)^2 + \frac{\partial}{\partial \eta} \left(\tau_{\xi\mathbf{x}} \cdot \mathbf{h}_{\xi} \right) \left(\frac{\partial \xi}{\partial \mathbf{Y}} \cdot \frac{\partial \eta}{\partial \mathbf{Y}} \right)$$

$$\frac{\partial \tau_{\mathbf{Z}\mathbf{x}}}{\partial \mathbf{Z}} = \frac{\partial}{\partial \xi} \left(\tau_{\xi\mathbf{x}} \mathbf{h}_{\xi} \right) \cdot \left(\frac{\partial \xi}{\partial \mathbf{Z}} \right)^2 + \frac{\partial}{\partial \eta} \left(\tau_{\xi\mathbf{x}} \cdot \mathbf{h}_{\xi} \right) \left(\frac{\partial \xi}{\partial \mathbf{Z}} \cdot \frac{\partial \eta}{\partial \mathbf{Z}} \right)^2$$

Summation of the two terms results in the following:

$$\frac{\partial \tau_{\mathbf{Z}\mathbf{x}}}{\partial \mathbf{Z}} + \frac{\partial \tau_{\mathbf{Y}\mathbf{x}}}{\partial \mathbf{Y}} = \frac{\partial}{\partial \xi} \left(\tau_{\xi\mathbf{x}} \cdot \mathbf{h}_{\xi} \right) \left[\left(\frac{\partial \xi}{\partial \mathbf{Y}} \right)^2 + \left(\frac{\partial \xi}{\partial \mathbf{Z}} \right)^2 \right] + \frac{\partial}{\partial \eta} \left(\tau_{\xi\mathbf{x}} \cdot \mathbf{h}_{\xi} \right) \left[\left(\frac{\partial \xi}{\partial \mathbf{Y}} \frac{\partial \eta}{\partial \mathbf{Y}} + \frac{\partial \xi}{\partial \mathbf{Z}} \right)^2 \right]$$
(32)

From Eq. 25 and the equations derived from Fig. 5, the following relations can be derived:

$$\left(\frac{\partial \xi}{\partial Y}\right)^2 + \left(\frac{\partial \xi}{\partial Z}\right)^2 = \left(\frac{1}{h_{\xi}}\right)^2$$

$$\frac{\partial \xi}{\partial Y} \frac{\partial \eta}{\partial Y} + \frac{\partial \xi}{\partial Z} \frac{\partial \eta}{\partial Z} = -\frac{1}{h_{\xi}} \frac{1}{h_{\eta}} \cos(\xi, Y) \sin(\xi, Y) + \frac{1}{h_{\xi}} \frac{1}{h_{\eta}} \cos(\eta, Z) \sin(\xi, Y) = 0$$

Substituting the above relations into Eq. 32 results in the following equation:

(30)

(31)

$$\frac{\partial \tau_{Yx}}{\partial Y} + \frac{\partial \tau_{Zx}}{\partial Z} = \frac{1}{h_{\xi}2} \frac{\partial}{\partial \xi} (\tau_{\xi x} h_{\xi})$$
$$= \frac{\tau_{\xi x}}{h_{\xi}2} \frac{\partial}{\partial \xi} (h_{\xi}) + \frac{1}{h_{\xi}} \frac{\partial}{\partial \xi} (\tau_{\xi x})$$

The RHS of the momentum equation can now be rewritten as:

$$-\frac{1}{\rho} \left[\frac{\partial}{\partial x} \left(\mathbf{p} + \rho \mathbf{g} \mathbf{H} \right) - \left(\frac{\partial \tau_{\mathbf{X}\mathbf{X}}}{\partial \mathbf{x}} + \frac{\partial \tau_{\mathbf{Y}\mathbf{X}}}{\partial \mathbf{Y}} + \frac{\partial \tau_{\mathbf{Z}\mathbf{X}}}{\partial \mathbf{Z}} \right) \right]$$
$$= -\frac{1}{\rho} \left[\frac{\partial}{\partial \mathbf{x}} \left(\mathbf{p} + \rho \mathbf{g} \mathbf{H} \right) - \frac{\partial}{\partial \mathbf{x}} \left(\tau_{\mathbf{X}\mathbf{X}} \right) - \frac{\tau_{\xi\mathbf{x}}}{h_{\xi^2}} \frac{\partial}{\partial \xi} \mathbf{h}_{\xi} - \frac{1}{h_{\xi}} \frac{\partial}{\partial \xi} \left(\tau_{\xi\mathbf{X}} \right) \right]$$

An equation to further simplify the term involving $h\xi^2$ is derived as follows.

Since $h_{\xi} \frac{\partial \xi}{\partial Y} = h_{\eta} \frac{\partial \eta}{\partial Z}$, differentiating both sides by ξ and using the same relationship again results in

$$\frac{1}{h_{\xi}} \frac{\partial h_{\xi}}{\partial \xi} = \frac{1}{h_{\eta}} \frac{\partial h_{\eta}}{\partial \xi}$$

This equation is substituted into the transformed RHS equation above; then

$$RHS = -\frac{1}{\rho} \left[\frac{\partial}{\partial x} (p + \rho g H) - \frac{\partial}{\partial x} (\tau_{XX}) - \frac{\tau_{\xi X}}{h_{\xi} h_{\eta}} \frac{h_{\eta}}{\partial \xi} - \frac{1}{h_{\xi}} \frac{\partial}{\partial \xi} (\tau_{\xi X}) \right]$$

Equating the transformed LHS and RHS equations results in the following momentum equation.

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + \frac{\nabla_{\xi}}{h_{\xi}} \frac{\partial u}{\partial \xi} = -\frac{1}{\rho} [\frac{\partial}{\partial x} (p + \rho g H) - \frac{\partial}{\partial x} (\tau_{xx}) - \frac{\tau_{\xi X}}{h_{\xi} h_{\eta}} \frac{\partial h_{\eta}}{\partial \xi} - \frac{1}{h_{\xi}} \frac{\partial}{\partial \xi} (\tau_{\xi x})]$$

or

$$\nabla_{\xi} \left(\frac{\rho}{h_{\xi}} \frac{\partial u}{\partial \xi} \right) = -\rho \frac{\partial u}{\partial t} - \rho u \frac{\partial u}{\partial x} - \frac{\partial}{\partial x} \left(p + \rho g H \right) + \frac{\partial}{\partial x} \left(\tau_{xx} \right) + \frac{\tau_{\xix}}{h_{\xi} h_{\eta}} \frac{\partial h_{\eta}}{\partial \xi} + \frac{1}{h_{\xi}} \frac{\partial}{\partial \xi} \tau_{\xix} \right)]$$

or

$$\nabla_{\xi} = \left(\frac{\rho}{h_{\xi}} \frac{\partial u}{\partial \xi}\right) \left[-\rho \frac{\partial u}{\partial t} - \rho \frac{\partial u}{\partial x} - \frac{\partial}{\partial x}\left(p + \rho gH\right) + \frac{\partial}{\partial x}(\tau_{xx}) + \frac{\tau_{\xi x}}{h_{\xi}h_{\eta}} \frac{\partial h_{\eta}}{\partial \xi} + \frac{1}{h_{\xi}} \frac{\partial}{\partial \xi}(\tau_{\xi x})\right] (33)$$

where the velocity component in the ξ direction is isolated from the velocity component in the η direction.

Derivation of the Equation for the \mathtt{V}_η Velocity Component

The continuity equation in the x, Y, Z coordinate system is

$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial Y} + \frac{\partial w}{\partial Z} = 0$$

Applying the transformation rule between the two coordinate systems on $\partial v / \partial Y$ and $\partial w / \partial z$ results in the following relation:

$$\frac{\partial \mathbf{v}}{\partial \mathbf{Y}} + \frac{\partial \mathbf{w}}{\partial \mathbf{Z}} = \frac{1}{\mathbf{h}_{\xi} \mathbf{h}_{\eta}} \left[\frac{\partial}{\partial \xi} \left(\mathbf{h}_{\eta} \mathbf{V}_{\xi} \right) + \frac{\partial}{\partial \eta} \left(\mathbf{h}_{\xi} \mathbf{V}_{\eta} \right) \right]$$

The continuity equation in the curvilinear coordinate system is therefore:

$$\frac{\partial \mathbf{u}}{\partial \mathbf{x}} + \frac{1}{\mathbf{h}_{\xi} \mathbf{h}_{\eta}} \left[\frac{\partial}{\partial \xi} (\mathbf{h}_{\eta} \mathbf{V}_{\xi}) + \frac{\partial}{\partial \eta} (\mathbf{h}_{\xi} \mathbf{V}_{\eta}) \right] = 0$$

or

$$\frac{\partial}{\partial \eta} (h_{\xi} V_{\eta}) = - h_{\xi} h_{\eta} \frac{\partial u}{\partial x} - \frac{\partial}{\partial \xi} (h_{\eta} V_{\xi})$$

Integrating both sides of this equation results in:

$$h_{\xi} V_{\eta} = -\int_{\eta^{\star}}^{\eta} [(h_{\xi} h_{\eta} \frac{\partial u}{\partial x} + \frac{\partial}{\partial \xi} (h_{\eta} V_{\xi})] d\eta + C$$

in which

 $\eta \star$ = the η value on the water surface

C = an integration constant

The integration constant can be determined by using the boundary condition on the water surface. The continuity equation in the x, ξ , η coordinate system can finally be written as:

$$\nabla_{\eta} = -\frac{1}{h_{\xi}} \int_{\eta^{\star}}^{\eta} \left[\left(h_{\xi} h_{\eta} \frac{\partial u}{\partial x} + \frac{\partial}{\partial \xi} \left(h_{\xi} \nabla_{\xi} \right) \right] d\eta + \nabla_{\eta^{\star}}$$
(34)

where V_{η}^* is the velocity V_{η} at point (ξ , η^*). The velocity component in the η direction is therefore computed from Eq. 34.

Approximation of Shear Stresses

To compute the velocity component V_{ξ} from Eq. 33, the shear stresses are needed. However, the shear stresses are also dependent on the velocity components. Therefore an additional equation is needed to solve the shear stresses and velocities. The additional equation is provided from the distribution of shear stress in the new coordinate system. Chiu and Hsiung (14) assumed that the shear stress varies quadratically in the ξ direction as follows:

$$\tau_{\xi x}(\xi, \eta) = \alpha_{0} + \alpha_{1} [\xi_{\max}(\eta) - \xi] + \alpha_{2} [\xi_{\max}(\eta) - \xi]^{2}$$
(35)

in which α_0 , α_1 , and α_2 are coefficients and $\xi_{\max}(\eta)$ is the maximum ξ value on an η curve. The coefficients α_0 , α_1 , and α_2 are determined from boundary conditions. As developed by Chiu, these are:

a) $\varepsilon < 0$ (the point of maximum velocity is below the water surface)

$$\alpha_{0} = 0$$
 (36)

$$\alpha_1 = f_1(\xi_0', \eta) \tag{37}$$

in which

• .

 $\xi_0' = \xi$ curve that approximates the channel boundary and

$$f_{1}(\xi,\eta) = \frac{-h_{\xi}F_{0} - \alpha_{2}\left\{2 - \frac{1}{h\eta}\frac{\partial h\eta}{\partial \xi}\left[\xi_{\max}(\eta) - \xi\right]\right\}(\xi_{\max}(\eta) - \xi)}{1 - \frac{1}{h_{\eta}}\frac{\partial h\eta}{\partial \xi}\left[\xi_{\max}(\eta) - \xi\right]}$$

where

$$F_{o} = \rho \left(\frac{\partial u}{\partial t} + \frac{\partial u}{\partial x}\right) + \frac{\partial}{\partial x} \left(p + \rho g H\right) - \frac{\partial \tau_{xx}}{\partial x}$$
(38)

 ξ_{\max} $(\eta) = 1$ for η curves that intersect with the stream banks or bed For η curves that intersect with the water surface, it can be reasonably assumed that $\tau_{k\eta} = 0$ if the wind effects are ignored; then

$$\alpha_1 = -\alpha_2 \left[1 - \xi_{\mathsf{D}}(\eta)\right] \tag{39}$$

where $\xi_{D(\eta)}$ is the ξ value of this η curve at the water surface.

$$\alpha_2 = \frac{-1}{\left[\frac{\xi_{\text{max}}(\eta) - \xi_{\text{o}}' f_2(\eta)}{\left[\frac{\xi_{\text{max}}(\eta) - \xi_{\text{max}}(\eta) - \xi_{\text{max}}(\eta) - \xi_{\text{max}}(\eta) - \xi_{\text{max}}(\eta) - \xi_{\text{max}}(\eta)\right]}}\right]$$
(40)

in which

$$f_{2}(\eta) = \frac{\xi_{\max}(\eta) - \xi_{o'}}{1 - \frac{1}{h_{\eta}} \frac{\partial h\eta}{\partial \xi} \left[\xi_{\max}(\eta) - \xi_{o'}\right]}$$
(41)

and the overbar means the average value over the wetted perimeter, which is computed as

$$\tilde{A} = \frac{1}{W.P} \int_{W.P.} A h_{\eta} (\xi_0', \eta) d\eta$$
(42)

and W.P. = total length of wetted perimeter.

- b) $\varepsilon = o$ (the point of maximum velocity is on the water surface). For this case, Eqs. 36, 37, and 40 can still be used.
- c) $\varepsilon > o$ (the point of maximum velocity is above the water surface). In this case the maximum value of $\xi(\eta)$ at the water surface is

 $\xi_{\max}(\eta) = \xi_0(\eta)$

and Eqs. 36, 37, and 40 are still valid.

Computer Model

The mathematical equations for computing the secondary velocities were presented in the preceding section. The equations, which were established in cartesian (x, y, z) coordinates, were first translated into the cross-sectionwise (x, Y, Z) coordinates, and were then transformed into the curvilinear (x, ξ , η) coordinates. Several parameters were used in these translation and transformation processes. The functions of these parameters were to relate the channel geometry and primary velocity distribution pattern to the governing equations.

After the equations in the curvilinear coordinate system are established, the computations are carried out in a grid system constructed on the (ξ, η) coordinates. The secondary velocities in the (x, y, z) coordinate system are obtained by using inverse transformations.

The overall computational procedures are hence separated into two phases. Phase I determines the best fitted coefficients ξ_i , β_i , k, δ_y , ξ_0' , and ξ_0 for approximating the primary velocity distribution and the channel geometry. Phase II uses these parameters to construct a computational grid system and carries out the computations for secondary circulation. In order to make a clear presentation of the model, two flow charts, shown in Figs. 6 and 7, are used to illustrate the computational procedures.



Figure 6. Flow chart of phase I analysis





Phase I Analysis

The Phase I analysis performs the following functions: 1) analyzes field data and determines parameters; 2) derives parameters for Phase II computations.

The channel cross section and isovels of the measured primary velocities are plotted first. Then determinations are made of the mean elevation of the channel bottom, H; the hydraulic radius, R; and the distance from the water surface to the point of maximum velocity, ε (Fig. 8). The axis which passes through the point of maximum velocity divides the channel cross section into two computational halves, as defined by subscripts i = 1 or 2 previously. The water surface width of each half, B_i, and the water depth D are then determined. Before proceeding to the steps described in Fig. 6, additional data need to be determined. These include $(y_b, z_b)_i$, and the corresponding slopes S_i for each bank. This information is needed for the following computations.



Figure. 8. Definition sketch for field parameters

On the assumption that the ξ_0' curve approximates the channel geometry, this curve should pass through (y_b, z_b) . At such a point, the slope of the ξ_0' curve (i.e. boundary) is

$$S_{\xi} = \frac{\partial y}{\partial \xi} = \frac{\partial \xi \, \partial z}{\partial \xi / \partial y} = \frac{(\partial \xi / \partial z) (\partial z / \partial \xi)}{(\partial \xi / \partial Y) (\partial Y / \partial y)}$$
$$= \beta_{i} \frac{D + \delta_{y} + \varepsilon}{B_{i} + \delta_{i}} \frac{Y \, Z}{(1 - Z) (1 - Y)}$$
(43)

Equation 43 can be rewritten as

$$\beta_{i} = (S_{\xi}) \frac{B_{i} + \delta_{i}}{D + \delta_{y} + \varepsilon} \frac{(B_{i} + \delta_{i} - z_{b})(D + \varepsilon - y_{b})}{z_{b}(Y_{b} + \delta_{y})}$$
(44)

On the other hand, the ξ_0 ' curve can also be described, by using Eq. 2, as

$$\xi_{0}' = \frac{y_{b} + \delta_{y}}{D + \delta_{y} + \varepsilon} (1 - \frac{|z_{b}|}{B_{1} + \delta_{1}})^{\beta_{1}} \exp\left(\frac{\beta_{1}|z_{b}|}{B_{1} + \delta_{1}} - \frac{y_{b} + \delta_{y}}{D + \delta_{y} + \varepsilon} + 1\right)$$
(45)

or as

$$\beta_{i} = \frac{\ln(\delta_{y}) - \ln(y_{b} + \delta_{y}) + y_{b} / (D + \delta_{y} + \varepsilon)}{z_{b} / (B_{i} + \delta_{i}) + \ln(B_{i} + \delta_{i} - |z_{b}|) - \ln(B_{i} + \delta_{i})}$$
(46)

This β_i value is the same β_i value described by Eq. 44. Moreover, the boundary condition at the origin gives the ξ_0' value as

$$\xi_{0}' = \frac{\delta_{y}}{D + \delta_{y} + \varepsilon} \exp \left(-\frac{\delta_{y}}{D + \delta_{y} + \varepsilon} + 1\right)$$
(47)

which is derived from Eq. 45 by setting $y_b = z_b = o$. Equating Eqs. 44 and 46 gives an equation which contains two unknowns: δ_i and δ_y . By assuming a δ_y value, δ_i and β_i values can be computed. Therefore an iterative procedure to determine the δ_i and β_i values is used to generate the closest computed velocity distribution to the measured primary velocities in a given channel.

Since the boundary curve ξ_0' and all the other ξ curves have to fit a logarithmic velocity distribution as described in Eq. 1, the following iterative steps depicted in Fig. 6 are used to determine the coefficients needed for Phase II.

- (1) assume trial $\delta_{\rm V}$ values
- (2) calculate $\delta_{\rm i}$ from Eq. 44 and Eq. 46
- (3) compute β_i

- (4) compute ξ values at every data point where u is measured
- (5) perform linear regression on u/u* and ln ξ
- (6) determine k and ξ_0 from intercept and slopes of step (5)
- (7) repeat step (1) to step (6)

The ξ_0 value obtained in step 6 is the line which defines zero velocity, while the ξ_0' line is the line which approximates the boundary. The best fitted coefficients are those coefficients which generate a primary velocity distribution that has the least standard deviation in comparison with the measured velocities.

The measured hydraulics parameters, i.e., D, B_i, ε , H, R, and the derived parameters δ_i , β_i , k, ξ_0 , ξ_0' , and τ_* , are then transmitted to Phase II for velocity and shear stress computations.

Phase II Analysis

The Phase II analysis involves solving the equations discussed in the last section for field variables. The parameters, determined from field data in Phase I, relate these equations to local hydraulic conditions. The characteristics of these equations require building the computations on a grid system, similar to that of a finite difference method. Using the ξ and η curves to form the grid, field data parameters are calculated at the grid nodal points. Also many variables (e.g., h_{ξ} and h_{η}) require the corresponding Y, Z coordinate values in their calculations. Consequently the Y, Z coordinates at each grid point have to be retrieved.

Grid System Construction

The grid system is constructed by drawing ξ curves and η curves with predetermined $\Delta \xi$ and $\Delta \eta$ intervals, between the maximum ξ and η values and the minimum ξ and η values. These maximum and minimum ξ and η values can be derived from Eqs. 2 and 3 once the parameters which depend on field data are determined through Phase I analysis, but the $\Delta \xi$ and $\Delta \eta$ values have to be determined by trial and error. The $\Delta \xi$ and $\Delta \eta$ values, once determined, are used in the numerical evaluation of the velocity components V_{ξ} and V_{η} , and significantly affect the accuracy and cost of computations. From Eq. 34 it can be seen that a smaller $\Delta \eta$ value will give greater accuracy in the numerical integration. On the other hand, a larger $\Delta \eta$ value will result in a shorter computation time. It is therefore desirable to have the $\Delta \eta$ value as

large as possible and to keep the computational error within an allowable range.

The $\Delta\eta$ values, however, depend on $\Delta\xi$ (Eq. 34) values. The $\Delta\xi$ value becomes available in the numerical evaluation of derivatives with respect to ξ . Finer $\Delta\xi$ values are needed near the channel boundary areas, since the primary velocity gradient is the highest at this location. The derivatives with respect to ξ , when evaluated by a numerical method, produce error. This error is the sum of two errors, truncation and rounding off. The round-off error is negligible on a large machine such as the CDC-6400 that carries numbers to 5 digits beyond the decimal point; therefore the $\Delta\xi$ value is selected to minimize the truncation error. The computation of $\Delta h_{\xi}/h_{\eta}\xi$ (in Eq. 33) is used for checking this $\Delta\xi$ selection. Two examples of a grid system generated using the procedures above are illustrated in Fig. 9.

Solving for (Y,Z) Given (ξ,η)

The values of ξ and η can be computed directly from given Y, Z coordinates by using Eqs. 2 and 3. However, obtaining the values of Y and Z for given ξ and η coordinates requires solution of the two nonlinear equations. The "modified Regula Falsi" iteration method (51) is used to solve these non-linear equations. Since the variations of ξ and η curves are different in various regions, the following four algorithms are used to solve the equations:

- (1) Solve Y for fixed Z from ξ curve
- (2) Solve Z for fixed Y from ξ curve
- (3) Solve Y for fixed Z from η curve
- (4) Solve Z for fixed Y from η curve

The algorithms are illustrated in Fig. 10. With these four algorithms, one can solve for Y (or Z) for an assumed value of Z (or Y) on a selected curve of ξ or η . The choice of the ξ (or η) curve for solutions depends on the tangent values at two consecutive points on that curve. If the tangent values are very close to each other on one curve, then the other curve is selected. The algorithm converges quickly after the first (Y, Z) coordinate point is found, which then becomes the initial trial value for the next (Y, Z) point.



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Figure 9. Grid systems for computational model





Figure 10. Algorithms for solving Y or Z from ξ and η

Computation for $V\xi$

The V_{ξ} component is computed from Eq. 33 on every grid point. In general, the normal stress term $\partial \hat{\tau}_{xx} / \partial x$ is small and is included in the hydrostatic pressure variation term. If the flow is steady and uniform, then

$$\frac{\partial}{\partial \mathbf{x}} (\rho \mathbf{g} \mathbf{H}) = \rho \mathbf{g} \mathbf{S}$$
(48)

where S is the slope of the channel bottom. Since the primary velocity is assumed to have a logarithmic distribution, and the shear stress a quadratic distribution, the velocity and shear stress derivatives can be written as:

$$\frac{\partial \mathbf{u}}{\partial \xi} = \frac{\partial}{\partial \xi} \left(\frac{\mathbf{u}_{\star}}{\mathbf{k}} \, \ln \frac{\xi}{\xi_0} \right) = \frac{\mathbf{u}_{\star}}{\xi \mathbf{k}} \tag{49}$$

and

$$\frac{\partial \tau_{\xi \mathbf{x}}}{\partial \xi} = -\alpha_1 - 2\alpha_2(\xi_{\max} - \xi)$$
(50)

The derivative of h_{η} with respect to ξ can be computed in Y and Z coordinates by using this relationship:

$$\frac{1}{h_{\eta}} \frac{\partial h_{\eta}}{\partial \xi} = \frac{\partial}{\partial \xi} \quad (\ln h_{\eta})$$

Therefore

$$\frac{1}{h_{\eta}} \frac{\partial h_{\eta}}{\partial \xi} = - \frac{\beta_{i} C_{i} Y^{2} Z^{2} (1-Z)^{2} + C_{i} Y^{2} (1-Y)^{2} (1-Z)^{2}}{\xi [(1-Y)^{2} (1-Z)^{2} + \beta_{i} C_{i} Y^{2} Z^{2}]^{2}}$$
(51)

With these terms known and $\partial u \ \partial x$ determined from field measurement, the V_{ξ} component at each node is computed in a straightforward manner.

Computation for $V\eta$

Eq. 34 can be rewritten as:

$$V_{\eta} = -\frac{1}{h_{\xi}} \int_{\eta^{\star}}^{\eta} \left[h_{\xi}h_{\eta} \frac{\partial u}{\partial x} + \frac{\partial}{\partial \xi} \left(h_{\xi}V_{\xi}\right)\right] + V_{\eta^{\star}}$$
(52)

The following steps are needed to solve Eq. 52.

(1) Determine the boundary value ${\rm V}_{\eta\star}$

$$V_{z} = V_{\xi} (h_{\xi} \frac{\partial \xi}{\partial Z})^{-1} \text{ or } V_{z} \text{ from field data}$$

$$V_{\eta \star} = h_{\eta} \quad V_{\xi} \frac{\partial \eta}{\partial \xi}$$
(53)

(2) By using Leibnitz's rule, transform Eq. 52 as follows:

$$V_{\eta} = -\frac{1}{h_{\xi}} \int_{\eta^{\star}}^{\eta} h_{\xi} h_{\eta} \frac{\partial u}{\partial x} d\eta - \frac{1}{h_{\xi}} \frac{\partial}{\partial \xi} \int_{\eta^{\star}}^{\eta} (h_{\eta} V_{\xi}) d\eta + V_{\eta^{\star}}$$
(54)

or =
$$\sum \frac{-1}{h_{\xi}} \int_{\eta_{i}}^{\eta_{i+1}} h_{\xi} h_{\eta} \frac{\partial u}{\partial x} d\eta - \sum \frac{1}{h_{\xi}} \frac{\partial}{\partial \xi} \Psi + V_{\eta \star}$$
 (55)

where $\Psi = \int_{\eta \star}^{\eta} (h_{\eta} \nabla_{\xi}) d\eta$

(3) The numerical integration for each integral on the right-hand side is evaluated as:

$$\int_{\eta_{i}}^{\eta_{i+1}} A h_{\eta} d\eta = \int_{-1}^{+1} (N_{1}A_{1} + N_{2}A_{2}) (N_{1}h\eta_{1} + N_{2}h\eta_{2}) \left[\frac{dN}{dr} \eta_{1} + \frac{dN}{dr} \eta_{2} \eta_{2}\right] dr$$
$$= \frac{1}{2} (\eta_{2} - \eta_{1}) \prod_{m=1}^{M} (N_{1}A_{1} + N_{2}A_{2}) (N_{i}h_{\eta_{1}} + N_{2}h_{\eta_{2}}) W_{m}$$
(56)

where $N_1 = \frac{1-r}{2}$, $N_2 = \frac{1+r}{2}$, r is the reference parameter ranging from -1 to +1, M is the Gaussian point, and W_m is the weight coefficient.

(4) The derivative of Ψ with respect to ξ is calculated by:

$$\frac{\partial}{\partial \xi} \Psi_{n} = \frac{[\Psi_{n}]_{i+1} - [\Psi_{n}]_{i-1}}{2\Delta \xi}$$
(57)

which is the central difference approximation.

The whole computation is carried out on an ξ curve, and the values of Ψ on neighboring ξ curves are also required.

A complete listing of the computer program for the model is provided in Appendix A.
FIELD DATA COLLECTION

The field data collection component of the project consists of development of the instrumentation needed to measure secondary circulation in natural channels, and the actual data collection. A computerized field data collection system was developed for the project and used successfully in the field. In addition to the acquisition of secondary circulation data, other measurements and sampling for flow discharges, suspended sediment concentrations, bed loads, and bed material were performed during the field data collection periods. This section of the report describes the instruments used to collect the data and the data collection procedures, and presents a summary of the data collected.

Field Data Collection Site

Three sets of field data were collected for the project. Data were collected in May and June 1984 and in August 1985. All the data were collected in the Sangamon River near Mahomet (near the junction of Sections 19, 20, 29, and 30, Township 20N., Range 7E.; river mile 180.6 above the Illinois River; and about 3.8 miles downstream of the Highway 47 bridge, Mahomet, IL). The Sangamon River near Mahomet, among other sites considered, was found to be the best site in terms of water depth requirements and site accessibility. The location of the study site is shown in Fig. 11. The stream channel at the study site is nearly a straight reach for about 500 feet, and then follows a gradual bend. At the upstream end of the site there is a bridge. The bridge piers did not obstruct the flow during the field trips; thus there were no eddies from the bridge piers. This was felt to be an important consideration in all secondary circulation measurements.

The channel width at this reach is almost constant (approximately 75 feet) and the cross section geometry is almost trapezoidal, which is an ideal case for the application of the mathematical model. Water depth at the study site varied from 2 to 4 feet, which made the data collection program relatively easy. The site is easily accessible by car, which facilitated the transport of equipment to the study site.



Figure 11. Location of study site

<u>Instrumentation</u>

Three-Dimensional Flow Measurement Setup

The instrumentation for measuring secondary circulation includes a Marsh-McBirney (Model 527) electromagnetic current meter, a Commodore (Model 8032) micro-computer with tape drive, an interface between the current meter and the computer, and a rigid support system for the current meter (Fig. 12). The support system includes a yoke which allows the current meter to be rotated 90 degrees so that both the lateral and vertical components of the secondary velocities may be measured.

Using the setup shown in Fig. 12, a continuous velocity reading from the current meter can be digitized and transmitted to the computer. Either the time history or an average and standard deviation of the velocity for a given period can be calculated and stored on cassette tapes.

Marsh-McBirney 527 Electromagnetic Current Meter

The Marsh-McBirney 527 Electromagnetic Water Current Meter consists of a transducer probe with a geomagnetic compass, a cable, and a signal processor. The probe is a 4-inch-diameter sphere on a 1-inch-diameter rod. The current meter measures the flow of water in a plane normal to the longitudinal axis of the probe. The current meter operates on the Faraday Principle of Electromagnetic Induction, which, simply stated, says: "a conductor moving in a magnetic field (generated from within the probe) produces a voltage that is proportional to its velocity." The electrodes placed on the wall of the probe detect the voltages caused by water flowing past the probe in a plane normal to the probe's axes. The flow sensing volume around the probe is a sphere 12 inches in diameter (3 probe diameters).

The two components of velocity together with the compass reading can be directly monitored on the three panel meters on the signal processor. The panel meter has three selectable full-scale ranges of ± 2 , ± 5 , and ± 10 ft/sec. The outputs are within 2% of full scale over the velocity range of the instrument, and the compass accuracy is $\pm 10^{\circ}$ at tilt angles of up to 25° . Also provided on the signal processor is a voltage output jack, which makes the velocity and the compass readings available to external data logging equipment; i.e., computers or any other data logging devices.



Figure 12. Field instrumentation setup

Interface and Computer

Since the Marsh-McBirney electromagnetic current meter output is in an analog signal, and it is almost impossible to visually average the velocity readings from the panel meters, it was decided to construct a data logger system based on a micro-computer. The initial step is to build an interface between the current meter and the computer. An interface which can receive the current meter output at selected frequencies and average the values for specified time periods was built at the Illinois State Water Survey Electronic Shop. The interface transfers the voltage outputs to the computer as digital signals.

The computer used for this system is a Commodore micro-computer (Model 8032) with a cassette tape drive. A computer program was developed to control the sampling procedure; to calculate the means, standard deviations of the velocity components, and the angle; and to store the data on cassette tapes. Additional information such as on site descriptions, coordinates and rotation of the current meter, time, and the gain selected for amplification of the signal is entered into the computer before sampling is initiated. The data stored on the tapes include the descriptive information and the actual data.

Once in the office, the data stored in the tapes is transmitted from the micro-computer to a main-frame computer (CYBER) by using a communication program and a modem. Further analysis of the data is carried out on the CYBER computer.

Support System for the Current Meter

The support system for the current meter was an important consideration in the development of the data acquisition system. The current meter has to be placed at a known depth in the stream, and the orientation of the current meter has to remain the same during the sampling period to obtain consistent data. To achieve these objectives, a rigid support system which could be moved from place to place had to be built. This was accomplished by modifying an aluminum stepladder and building a supporting platform, pole, and yoke as shown in Fig. 12. The yoke was designed to allow a 90° rotation of the current meter so that secondary currents both in the transverse and vertical directions could be measured. The yoke is supported by a pole which can be raised or lowered at regular intervals to change the depth of the measuring point. The pole is supported by the platform, which is attached to the

ladder. Additional bracings were used to reduce the vibration of the current meter generated by the stream current. A complete cross-sectional velocity distribution is measured by moving the whole support system along the channel width, lowering and raising the pole, and rotating the current meter.

Rotating Bucket Current Meter

This instrument, a standard Price-type meter, has a rotor with six coneshaped cups mounted on a vertical stainless steel shaft. The meter has been calibrated by the manufacturer to a standard rating table where the number of rotations of the rotor over a specified time corresponds to a specific velocity.

The current meter is attached to a 30-lb weight to keep it from drifting in the current and is suspended into the stream from a bridge deck with a cable/winch/crane assembly. This meter was used to measure flow velocity in the Sangamon River at the bridge site shown in Fig. 11. These velocity data were used to determine the discharge at the test reach.

Suspended Sediment Samplers

Two types of suspended sediment samplers were used for this investigation: the US DH-59 and DH-48, which are standard samplers designed for the United States Geological Survey. These samplers are designed to accumulate a water-sediment sample from a stream vertical at such a rate that the velocity in the nozzle at the point of intake is as close as possible to the stream velocity. These types of samplers, properly used, will withdraw a sample of the sediment water mixture that represents an average of the sediment concentration through a stream vertical weighted for velocity so that proper sediment load calculations can be made.

Both samplers use standard 1-pint glass milk bottles to collect the sediment-water samples. The DH-59 sampler was designed to be suspended by a hand-held rope or a cable/winch/crane assembly. This sampler was used to determine sediment concentrations of the stream at the bridge site (Fig. 11). The DH-48 sampler was used to sample for sediment concentration at sites away from the bridge. It was designed to be suspended from a wading rod and lowered into the stream by hand.

Bed Load Sampler

A Helley-Smith bed load sampler was used to measure the component of the total sediment load that moves in close proximity to the streambed. This sampler was designed for sampling coarse material where the diameter of the bed load material is above 0.25 mm. The sampler was designed to be lowered to the streambed, left in contact with the bed for a certain period of time, and then retrieved from the water. Samples are retained by a mesh bag (0.25-mm mesh openings) open to the flow through a 3-inch-square opening. The quantity of material trapped by the sampler over a certain period of time indicates the amount of bed load moving through the 3-inch-square area, and by extrapolation (to the total bed width of the stream) the total quantity of bed load carried by the stream can be calculated.

Sonar Stage Meter

An electronic sonar distance meter was used to measure the stage or height of the stream. Readings were made throughout the data collection period in order to assess the changes over time in stage. The sonar meter is placed at a known datum on the bridge and transmits the sonar signal downward to the stream water surface. The water surface reflects the signal back to the meter and the distance is calculated on the basis of the lag time between transmission and return. The meter is calibrated for the effect of air temperature on the speed of sound. The instrument used was an Exact Technologies Corporation Dimensional Measurement Computer. It has a resolution of 1 inch over a range of up to 50 feet.

Data Collection Procedures

Secondary Circulation

Secondary circulation data were collected from cross section to cross section within the study reach. Measurements are taken at several depths for each vertical. Normally ten or more verticals per cross section were selected for data collection. Generally data collection at one cross section lasted one full day.

The data collection procedure for secondary circulation is as follows: First the support platform and pole are moved into the stream and firmly set at the data collection site. Then the yoke with the current meter is securely attached to the pole on the platform. The cable from the current meter is

then connected to the signal processor, which is located on the stream bank. The output from the signal processor is first transferred to the interface and then to the computer by cables between the devices. After all the connections between the devices are completed, power is turned on to activate the system. The computer program which controls the data collection and storage is then loaded into the computer and run. The computer prompts with header lines for inputting general information such as location, date, and time. It then asks for coordinates and gain (or amplification) being used. Once the above procedures are completed, sampling is initiated and terminated by pressing designated keys on the computer keyboard. Finally, the data are displayed on the screen for inspection and then transferred to cassette tapes for further analysis.

Data Reduction

All the cross sections where measurements were taken were aligned to be normal to the stream banks. However, this selection does not guarantee that the cross sections are normal to the main flow direction. Furthermore, it is very difficult to keep the orientation of the current meter consistent for all measurements. Thus the measured values of secondary flows might not represent the true values. The measured velocities were adjusted to reflect the true longitudinal and secondary currents by using the orientation of the current meter with respect to the magnetic north, which is measured simultaneously with the velocities, and by using one of the following two criteria: 1) The discharges measured by the two current meters, the electromagnetic and the rotating bucket mechanical current meter, should be equal; and 2) the continuity of secondary flows must be satisfied between downstream and upstream cross sections.

The method employed to check the continuity of secondary discharge is that proposed by Dietrich and Smith (22), and also used by Thorne et al. (56). In this method, the net secondary discharge is determined by integrating the secondary velocity over the depth. The continuity of discharge requires that the net secondary discharge at a cross section be equal to that of the section immediately upstream or downstream. If the two differ, the section is reoriented to change the secondary velocities until an agreement is reached.

The measured velocity components V_x and V_y , shown in Fig. 13, are first converted to resultant velocities from which the downstream and cross-stream velocities are computed.

The magnitude of the resultant velocity = $\sqrt{v_x^2 + v_v^2}$

The angle of the resultant velocity with respect to magnetic north (θ) is determined as:

$$\theta = \begin{cases} \tan^{-1}(V_{x}/V_{y}) & \text{if } V_{y} > 0 \\ \tan^{-1}(V_{x}/V_{y}) + \pi & \text{if } V_{y} < 0 \end{cases}$$
(58)

The angle of the main downstream velocity with respect to the magnetic north (θ_N) is given by:

$$\theta_{\rm N} = \theta + \theta_{\rm c} \tag{59}$$

in which

 $V_{\rm v}$, $V_{\rm x}$ = measured velocity

 $\boldsymbol{\theta}_{\rm C}$ = bearing of sensor with respect to magnetic north

 $\boldsymbol{\theta}_{N}$ = bearing of left bank with respect to magnetic north

 θ = bearing of resultant flow with respect to magnetic north

The downstream and cross stream components of the velocity are then given by:

downstream velocity = (resultant) (cos θ) cross stream velocity = (resultant) (sin θ) (60)

Sampling Period

Two major factors are considered before selecting the sampling period: 1) the sampling period should be long enough that the velocity fluctuations will average out and a true mean velocity will be obtained; and 2) the sampling duration should be short enough that the measurements for at least one cross section can be completed within a day.

Theoretically the true mean velocity is the average of the velocity measurements for an infinite time period. In practice, however, a period of averaging time is determined which gives a mean velocity very close to the long-time average with specified tolerance limits.

One method of estimating the averaging period is to determine the mean velocities for several averaging periods and then select the averaging period which gives a mean within a specified range from the true mean. This will require taking long period samples for each of the sampling periods. However,





an equivalent result is obtained from one long period sample as follows. A set of sampling periods τ_i , i = 1,2,3 ...M are first selected. Then the sample is subdivided into blocks of N_i samples for each τ_i . The mean, \bar{u} , and the standard error, S_m, which is the standard deviation in the mean, for each averaging period are then computed by the following equations (48)

$$\overline{u} = \frac{1}{T_m} \int_0^{T_m} u \, dt \tag{61}$$

$$\frac{\begin{bmatrix} N_{i} \\ \Sigma \\ j=i \end{bmatrix}}{\begin{bmatrix} N_{i} \\ N_{i} \end{bmatrix}}$$

(62)

where

 $v_{j} = \frac{1}{\tau} \int_{0}^{t_{i}} u_{i} dt = time average velocity for the period of \tau_{i}$ $N_{i} = the number of averages with an average period of \tau_{i}$ $T_{m} = the sampling period of the whole sample$ $\overline{u} = the long-term average or the true mean$

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u = velocity at time t

To generalize the standard error for variable magnitudes of velocity, it is convenient to non-dimensionalize the standard error by the mean velocity. The normalized error, S_{m_i} , is therefore given by:

$$\overline{S}_{m_{i}} = \frac{S_{m_{i}}}{\overline{u}}$$
(63)

Two sets of velocity measurements are shown in Fig. 14 to illustrate how the sampling periods were selected. The first measurement was taken near the center of the channel (42 feet from the left bank) and 1.25 feet from the water surface, while the second sample was taken close to the stream bank (26 feet from the left bank) and 1 foot from the water surface.

The total sampling time, Tm, for both samples was 5 minutes. The sampling periods, τ_i , were 14, 27, 50, 99, and 150 seconds. The relationship between the sampling period and the normalized standard error for both samples is shown in Fig. 15. As shown in the figure the normalized standard error near the bank is greater than at the center of the channel. This is probably because of greater secondary currents and turbulence intensity near the banks



TIME, seconds

Figure 14a. Velocity fluctuations at a point (distance = 42 ft, depth = 1.25 ft), cross section 9, August 9, 1985



Figure 14b. Velocity fluctuations at a point (distance = 26 ft, depth = 1.00 ft), cross section 7, August 7, 1985



Figure 15. Relationship between sampling period and the normalized standard error for two sets of velocity measurements

than at the center of the channel. However, the normalized error for both samples decreases with increasing sampling period. The error decreases sharply for sampling periods more than 60 seconds and gradually thereafter. It was decided to use 60 seconds as the sampling period for an efficient and reliable data collection procedure.

Stream Discharge Measurement

The discharges of the Sangamon River at the study site were measured several times during each field trip. Two types of instrumentation were used to measure discharge of the Sangamon River: 1) the electromagnetic twodimensional current meter, and 2) the rotating bucket mechanical current meter. The electromagnetic meter was employed primarily to measure secondary circulation. However, since the meter also measures the flow in the downstream direction, the data generated were used to compute the stream discharge. The rotating bucket mechanical meter was used to measure stream velocity and to compute stream discharge for a comparative analysis between results of the two methods.

The stream discharge measurement techniques used for the rotating bucket current meter were those recommended by the United States Geological Survey (9) and the American Society for Testing and Materials, Standard Practice for Open-Channel Flow Measurement of Water by Velocity-Area Method, Designation: D 3858-79 (1). Stream discharges are determined by subdividing a cross section of the stream into partial sections; the average velocity of the flow at each partial section is measured and multiplied by the flow area of the partial section. The sum of the individual partial section discharges equals the total stream discharge.

Discharge measurements performed at the bridge site used the rotating bucket current meter. The stream cross section under the bridge was divided into 7 partial sections, each approximately 10 feet wide. The velocity meter, suspended by a cable/winch/crane assembly, was lowered into the stream at the midpoint of each partial section. A depth gage built into the winch read the total depth of the stream at the midpoint of the partial section. This depth was recorded and later used to calculate the flow area of the partial section. Velocity measurements were then made at the vertical in the midpoint of the partial section. The meter was positioned below the water surface at 0.2 and 0.8 of the total depth (for total depths greater than 2.0 ft) or at 0.6 of the

total depth (for total depths less than 2.0 ft). The velocity of the stream at the measured points was recorded and used to calculate the average velocity of the partial section. Stream discharges were then calculated by multiplying the mean velocity for each partial section by the respective cross-sectional area. Total discharge for the section was obtained by summing all the individual partial discharges.

Suspended Sediment Concentrations and Particle Size Sampling

Sediment sampling was performed using the US DH-59 and US DH-48 samplers as described in "Instrumentation." The sediment sampling was performed to determine the concentration of suspended sediment and the total sediment load in the stream. The methods used were developed by the USGS (31) and are proposed practices of the American Society for Testing and Materials Committee D-19 (water) and Subcommittee D-19.07 (sediment) (1).

The sampling methodology employed the equal width increment method where the stream cross section is divided into three or four partial sections and the midpoint of each partial section is sampled. At each midpoint a clean sample bottle is inserted into the sampler and the sampler is carefully lowered from the stream surface to the bottom and raised back to the surface at a constant rate of 1 foot per second. The sample bottle is then capped, identified, and noted in the field book. This procedure is repeated at each partial section midpoint until the entire width is sampled. The purpose of using equal width increment and depth integrated techniques is to quantify the total amount of sediment carried in suspension by a stream where the concentration of sediment varies across the width and depth.

Suspended sediment concentration sampling was performed at the study site on most of the days of data collection. The sample bottles were identified and labeled in the field and then brought into the office for logging and preparation for delivery to the Inter-Survey Geotechnical Laboratory for analysis.

Suspended sediment particle size analysis was performed on three samples obtained by dip sampling. Dip sampling was used instead of the methods outlined above for suspended sediment concentration because of the low concentration of suspended sediment, which made it necessary to collect a large quantity of samples for the laboratory analysis.

Bed Load Measurement

The coarse fraction of the sediment load (>0.25 mm) moving in contact with or within 3 inches of the streambed was sampled using a Helley-Smith bed load sampler. This sampler was lowered to the streambed at the thalweg and left there for a period of five minutes. The sampler allows the streamflow near the bed to pass through the intake opening and into a fine mesh. The sample in the mesh bag at the time of retrieval was then examined for quantity and types of sediment (inorganic, wood, leaves, etc.). Field observations of the sample were then entered into the field notebook. Only one bed load measurement was collected as part of this investigation because of low flow conditions during sampling periods, at which time no appreciable bed load movement was observed.

Bed Material Sampling

The streambed sediment was sampled in order to characterize the stream bed and to determine the distribution of sediment particle sizes that are available for transport as bed load and suspended load. The streambed was sampled across its width 100 ft downstream of the bridge sampling site. Three discrete samples were obtained from the bed at three locations: the midpoint, and the left and right quarter points. The bed samples were obtained using a shallow bucket 10 inches across and 3 inches deep. The bucket was lowered through the water column by hand and then tilted so that one edge could cut a sample from the top inch of the streambed. The bucket and sample were brought to the water surface carefully to avoid agitating or dispersing the material. The excess water was drained away from the sample and the sample was placed into whirl-pack sampling bags. The samples averaged approximately 500 grams each. All sample bags were labeled to identify date, time, and sampling location. The samples were then brought into the office for logging and preparation and delivered to the Inter-Survey Geotechnical Laboratory for analysis.

Stage Measurement

The relative change in the stream surface elevation was monitored daily at the bridge site using a sonar distance meter. At a minimum, readings were made at the beginning and end of the work day. The sonar readings were

adjusted to reflect possible temperature effects on the speed of sound. Relevant data were recorded in the field notebook for later analysis.

Summary of Field Data Collected

A summary of all the field data collected for the project is presented in this section. A complete list of all the data is included in Appendix B. All the cross sections where data were collected are shown in Fig. 16. Table 1 shows the dates when the data were collected at each cross section.

Secondary Circulation

The velocity measurements used to construct secondary circulation were first processed separately to illustrate the distribution of each component in a cross section. Fig. 17 shows the distribution of the longitudinal, transverse, and vertical components of the velocity at cross section 7. It should be noted that different scales are used for the three components shown in the figure, so that the distributions of the lateral and vertical components, which are much smaller than the longitudinal component, can be shown effectively. The distribution of the longitudinal component (Fig. 17a) indicates that the higher velocities are located close to the center of the channel and the lower velocities are near the banks, as should be the case.

Date Month/day/year	Cross section number where data were collected
06/27/84	7
06/28/84	6
06/29/84	5
07/09/84	3
07/10/84	4
07/12/84	2
07/13/84	1
08/05/85	5
08/06/85	6
08/07/85	7
08/08/85	8
08/09/85	9

Table 1. Dates of Field Data Collection



Figure 16. Location of cross sections at the study site





Figure 17. Distribution of measured longitudinal, transverse and vertical components of velocity at cross section 7, June 27, 1984

The distribution of the transverse component shown in Fig. 17b is, however, highly variable in magnitude and direction. The vertical components for this particular cross section (Fig. 17c) show some variation in magnitude but not as much as the lateral components. Fig. 17 is just an example to illustrate the type of data collected and how they were analyzed. Similar plots for all the data collected in the field are given in Appendix B.

The longitudinal velocity was further processed to generate isovels as shown in Fig. 18. Since the model used for this project utilizes the isovels as the basis for coordinate transformation, it was felt important to generate the isovels for measured data so that they could be compared with the model. Isovel plots for all the measurements are also shown in Appendix B.

The transverse and vertical components of the velocity were combined to generate the secondary current vectors as shown in Fig. 19. For this particular case, the measured values indicate primarily upward and lateral movements of the secondary currents. The downward currents were either missed in the measuring process or are located close to the banks. As with the previous plots, the rest of the secondary current plots are included in Appendix B.

Stream Discharge

The discharge of the Sangamon River at the study site was measured at the bridge cross section using the rotating bucket current meter for the purposes of monitoring the changes in the daily streamflows. Stream discharges were also calculated on the basis of the secondary current measurements at the different cross sections. Table 2 presents a summary of the discharge measurements, as determined by both the rotating bucket and electromagnetic current meters. As shown in the table there are some differences in the stream discharges calculated from the two sets of measurements. Considering that the discharges were not measured at the same cross sections, and considering the errors in stream flow measurements, the agreement between the two sets of measurements was satisfactory for most cases.



Figure 18. Isovels of the longitudinal velocity at cross section 7, June 27, 1984



DISTANCE, ft

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i. Leu

Figure 19. Measured secondary current vectors at cross section 7, June 27, 1984

	<u>Water discha</u>	Suspended sediment		
<u>Date</u>	<u>Rotating bucket</u>	<u>Electromagnetic</u>	(ppm)	(tons/day)
6/27/84		139		
6/28/84	·	140	205	77
6/29/84		119		
7/09/84	71	48		
7/10/84	65	58	114	20
7/12/84	36			
8/05/85	126	75	85	29
8/06/85	100	115	94	25
8/07/85	73	66	48	10
8/08/85		115		
8/09/85	80	59	52	11
8/14/85	70		103	20

Table 2. Water Discharge and Sediment Load Measurements During Data Collection Period

The highest and lowest discharges measured in the 1984 data collection period were 140 cfs on June 28, and 36 cfs on July 12, respectively. For the 1985 data the highest discharge was 126 cfs on August 5 and the lowest was 59 cfs on August 9. The discharges measured during the field data collection periods are all below the mean streamflow and generally fall in the range of flows exceeded from 40 to 70 percent of the time as indicated in the flow duration curve shown in Fig. 20. The flow duration curve was developed from historical streamflow records of the Sangamon River gaging station at Mahomet, which is 4 miles upstream of the study site. Data collection was not conducted during higher flows because of logistic requirements of deep water measurements, which require using a boat to move the instrumentation from station to station. For water depths less than 4 feet the instrumentation is moved from station to station by wading in the stream channel and physically carrying the platform.



Figure 20. Flow duration curve for the Sangamon River at Mahomet, Illinois

Suspended Sediment Concentration

A total of 48 suspended sediment concentration measurements were made during the data collection field trips. Several of the measurements were detailed cross-sectional measurements, while some were representative samples for calculating sediment load. The concentration measurements were averaged for the whole cross section and the suspended sediment loads were calculated from the average concentrations. These results are shown in Table 2. The suspended sediment concentrations during the 1984 data collection periods are generally higher than those in 1985. This is probably the influence of a major flood event in 1984 which preceded the data collection period. The suspended sediment load in the Sangamon River during the data collection period ranged from a low of 10 tons per day on August 7, 1985, to a high of 77 tons per day on June 28, 1984 (Table 2).

One detailed suspended sediment concentration measurement was performed on August 8, 1985, to investigate the distribution of sediment in the stream channel. A total of 21 point samples were collected at seven verticals. Fig. 21 shows the point concentration values and the equal concentration lines generated from the data. As can be seen in the figure, the suspended sediment concentration is higher towards the channel bottom and there are some regions of higher concentration which might be associated with secondary currents. However, there is no significant variation of concentration at that particular time of data collection, and more intensive data collection similar to that done here will be needed at different flows to conclusively associate concentration distribution with secondary currents.

Bed Load

Only one bed load measurement was made at the study site. This measurement, which was made on August 6, 1985, yielded a bed load sample of 4.5 grams for a sampling period of five minutes. The sample consisted of inorganic fine gravel and coarse sand. The bed load flow area was estimated to be the central 50 feet of the stream channel, which had a total width of 69 feet. Bed load was calculated by multiplying the sampler intake width by the bed load flow area width and then extending the rate per unit time to a 24hour period to determine total load per day. The bed load was estimated to be 0.3 tons per day. In comparison, the suspended load was 25 tons per day.



Figure 21. Distribution of suspended sediment concentration in parts per million, cross section 8, August 8, 1985

Therefore the bed load represents approximately 1.2 percent of the total sediment load (bed and suspended) on August 6, 1985, at this location.

Bed Material Characteristics

The Sangamon River's bed is composed primarily of sand and some gravel. Sand and gravel bars are common in the river. Riffles in the river are composed of gravels up to 3 inches in diameter and coarse sand. Pools in the river occur between the riffles and are underlaid by sand and silt.

Table 3 summarizes the laboratory and statistical analyses of bed material samples obtained at the study site in 1984 and 1985. In 1984 two composite samples obtained by mixing four discrete samples taken at equal distances across the stream were collected at cross sections 2 and 4. In 1985 three separate samples were collected at cross section 8, one at mid-channel and one each at the one-fourth point of the stream width from the stream bank. In table 3 the individual sample results of the 1985 samples are averaged in order to provide a composite average for comparison with the 1984 samples. Sand predominates in the composite samples, and gravel makes up the bulk of the remaining sample with less than 5 percent silt and clay. All the composite samples are descriptively classified as "gravelly sand with silt."

The bed material varies in composition with respect to particle size distribution across the width and length of the river. The individual samples obtained on August 14, 1985, show a trend of increasing average particle size from the west to the east as shown in table 3. The longitudinal variation of particle sizes can also be seen in table 3 by comparing the composite samples obtained at the two cross sections in 1984. The gravel concentration was higher at cross section 2 than at cross section 4, and the sand concentrations were lower at cross section 2 than at cross section 4 (Fig. 16). The data provided here are intended to provide a general description of the channel bed material characteristics at the study site for use in comparing and contrasting different study sites. This was not a detailed channel morphologic investigation.

Stage

The stage (height) of the river was measured during the 1985 data collection period to assess changes over time in the stage and water

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Characteristics
Table 3.

	<u>Description</u>	gravelly sand	gravelly sand	silty sand	gravelly sand	sandy gravel	gravelly sand
Silt	<u>and clay</u>	4.73	1.86	3.61	4.68	4.32	4.20
	<u> </u>	58.44	81.19	93.40	69.17	37.38	66.65
	<u>%Gravel</u>	36.83	16.95	2.99	26.15	58.30	29.15
Standard	<u>deviation</u>	.25	.41	.53	.33	.21	.29
	<u>Median</u>	1.02	.73	.51	.68	2.85	.31
	<u>Mean</u>	1.41	.86	.50	.86	2.06	1.17
	<u>Location</u>	x-s #2*	×7# S-X	west $1/4$	midpoint	east $1/4$	composit
	<u>Date</u> **	7/10/84	7/12/84	8/14/85	8/14/85	8/14/85	8/14/85

* x-s = cross section

west and east 1/4 refer to a point one-fourth of the stream width from the stream bank

**All 1985 samples obtained from a cross section 100' downstream of the bridge site.

All 1984 samples were composited prior to laboratory analysis from 4 discrete samples obtained at equal distances across the stream width.

discharge. Figure 22 summarizes the stage measurements over the period August 5-14, 1985. The highest stage was recorded on August 5 and 8. The lowest stage occured on August 14. The general trend was a gradual decrease in the stage from August 5 to 7 and from August 8 to 14, with some rise in between from the 7th to the 8th. The change in the stage is similar to that exhibited by the discharge in table 2.

MODEL RESULTS

The mathematical model discussed earlier was used to generate secondary velocities corresponding to seven different measurements in the Sangamon River. In this section the results of the model are discussed and compared to the measured values. The comparisons between the computed and measured values are shown in two sets of figures for each cross section. In the first figure, computed and measured secondary velocity vectors generated from the lateral and vertical components are compared. In the second figure, the primary velocity isovels approximated by the model are compared to the actual isovels developed using the measured values. A table is also presented that summarizes the magnitudes of the computed and measured velocities for all the cases considered. Finally, a section on shear stress distribution is presented.

Comparison of Computed and Measured Velocities

The comparisons for cross section 6 as measured on June 28, 1984, are shown in Figs. 23 and 24. Fig. 23a shows the secondary velocity vectors generated by the model, while Fig. 23b shows the vectors from the measured velocities. It should be mentioned here that the vectors for the measured velocities were obtained by interpolating the vertical components around the points where transverse velocities were measured. This was necessary because it was not possible to measure the vertical and transverse components at the same point simultaneously.

As shown in Fig. 23 the general pattern of secondary circulation generated by the model is similar to the measured one; however, the downward circulation indicated near both the right and left banks by the model (Fig. 23a) are not present in the measured velocities (Fig. 23b). Velocity measurements were not taken very close to the stream banks and the channel



Figure 22. Variation of stream stages during field data collection



Figure 23. Comparison between computed and measured secondary velocities, cross section 6, June 28, 1984



Figure 24. Comparison between isovels derived from computed and measured data, cross section 6, June 28, 1984 65

bottom because of the physical limitations of the equipment setup. The lack of these measurements makes it difficult to make conclusive remarks on the comparison of the model results with the measured values for areas close to the bank.

The measured secondary velocities are generally smaller than those generated by the model. This is true for all seven cross sections used for the model. The values of the primary, transverse, and vertical components of the velocity for all the cross sections are summarized in table 4. The percentages of the computed with respect to the measured transverse and vertical velocities are given in columns 9 and 11 in table 4. The percentages in both columns are all higher than 100 percent and in some cases extremely high. Because of the small magnitudes of the velocities, a very small error in velocity measurements or slight overestimation by the model results in very high percentages. However, further refinement of the model is needed and possible. Unfortunately this could not be accomplished within this project.

Table 4 also provides the relative magnitudes of the measured transverse and vertical velocity components with respect to the primary velocity component in columns 5 and 7, respectively. The transverse components range from 8 to 43.7 percent of the primary velocity, while the vertical components range from 4.1 to 40.4 percent. These values are in most cases higher than those reported in the literature, but might be because of the low flow conditions in the Sangamon River during data collection. Additional data during higher flows are required to more definitely define the relative magnitude of secondary velocity components as compared to the primary velocity.

Fig. 24 shows the primary velocity isovels generated by the model (Fig. 24a) and those developed from the measured velocities for cross section 6 discussed above. There is a clear similarity between these two plots; however, as can be seen in the figure, the isovels for the model are smooth curves, while the isovels from the measured velocities are irregular and show deformations at different places. This discrepancy was anticipated from the beginning since the model formulations assume smooth isovels which can be approximated by mathematical functions. The assumption of smooth isovels is the most limiting factor in secondary circulation models based on coordinate transformation.

Table 4. Mean Measured and Computed Velocities

col (10) col (6) 109 104 222 242 534 1227 311 ł ; i i ł (8) component Computed vertical (fps) .15 .12 .38 .16 .39 .27 ; .51 1 ł 1 ; col (4) col (8) 186296 509 160 146 330 317 (8) ; : ; ł transverse Component component (fps) .19 .16 .28 .22 .13 .40 .30 ; ł 1 ; ł <u>col (6)</u> col (3) (8) 23.0 29.5 40.4 9.0 31.9 4.1 32.7 16.233.6 19.6 27.5 8.7 component Measured vertical (fps) .138 .069 .022 .115 .088 .072 .137 .122 .067 .211 .101 .073 <u>col (</u>4) col (3) 8.0 41.8 12.9 15.3 18.2 6.3 18.2 43.7 24.4 22.1 23.1 21.5 (8) transverse Measured component (fps) .048 .118 .060 .054 .055 .064 .172 .274 .068 .105 .081 .091 longitudinal Measured component 1.0009 .599 .535 .465 .302 .744 .444 .352 .316 .373 (fps) .627 .251 6/27/84 6/28/84 6/29/84 7/10/84 7/12/84 7/13/84 8/05/85 8/06/85 8/07/85 8/08/85 7/09/84 8/09/85 <u>Date</u> Section Cross 9 ŝ \sim \mathbf{d} \sim ŝ 9 ~ ∞ δ 67

The comparisons of the model results with the measured values for the remaining six measurements are shown in Figs. 25 to 36. In general, the preceding discussion for cross section 6 (June 28, 1984) is applicable for all the comparisons. Some of the comparisons indicate a better correlation between measured and computed velocities and some are worse than the ones shown for cross section 6, as summarized in Table 4. A complete listing of the velocities generated by the model is included in Appendix C. Additional work both in data collection and in refinement of the model is definitely needed before the model adapted for this project can be applied to flow conditions in natural streams.

Shear Stress Distribution

As discussed in the section on approximation of shear stresses, the shear stresses were assumed to vary quadratically in the ξ direction according to Eq. 35. The coefficients for Eq. 35 are derived from boundary conditions, which depend on the flow conditions and the channel boundary. The shear stress distributions generated by the model for two sets of measurements at cross section 6 are shown in Figs. 37 and 38. τ/τ_0 is the ratio of the shear stress at any point (τ) to that at the channel bottom (τ_0). The figures clearly show higher shear stresses around the junctions of the stream bank with the stream bed, which might indicate areas of higher erosion rates as compared to the whole channel cross section.


Figure 25. Comparison between computed and measured secondary velocities, cross section 5, June 29, 1984



 $\leq A$

Figure 26. Comparison of isovels derived from computed and measured data, cross section 5, June 29, 1984



Figure 27. Comparison between computed and measured secondary velocities, cross section 3, July 9, 1984



Figure 28. Comparison of isovels derived from computed and measured data, cross section 3, July 9, 1984



Figure 29. Comparison between computed and measured secondary velocities, cross section 4 July 10, 1984



Figure 30. Comparison of isovels derived from computed and measured data, cross section 4, July 10, 1984 74

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Figure 31. Comparison between computed and measured secondary velocities, cross section 5, August 5, 1985



Figure 32. Comparison between isovels derived from computed and measured data, cross section 5, August 5, 1985



Figure 33. Comparison between computed and measured secondary velocities, cross section 6, August 6, 1985



Figure 34. Comparison of isovels derived from computed and measured data, cross section 6, August 6, 1985



Figure 35. Comparison between computed and measured secondary velocities, cross section 8, August 8, 1985



Figure 36. Comparison of isovels derived from computed and measured data, cross section 8, August 8, 1985



Figure 37. Shear stress distribution, cross section 6, June 28, 1985



Figure 38. Shear stress distribution, cross section 6, August 6, 1985

SUMMARY AND CONCLUSIONS

Secondary circulation is an important hydraulic phenomenon in natural streams. However, studies in the field have been limited, and in particular actual measurements of secondary circulation have been very few. This is because of the difficulty in measuring very small velocities under natural conditions. The development of mathematical models has also been limited, primarily because of the lack of reliable field data for calibration and verification.

This project is one of the few research projects in the United States that has attempted to measure secondary circulation in natural streams. As a result of this project, a reliable field data collection technique has been developed. The data collection system is computerized to reduce the time required in field data recording and reduction. The system has been used successfully several times to measure secondary circulation data in the Sangamon River near Mahomet, Illinois. All the data collected for the project are included in Appendix B. The data are also analyzed and presented graphically to simplify comparison of data with model results.

At the beginning of this project it was anticipated that an existing model would be available for calibration and verification. However, attempts to obtain existing models were not successful because of the early stages of model development by other researchers. Instead of delaying or dropping the task of calibration and verification of existing models, it was decided to develop a model based on information available in the literature. At that time, the model being developed by Chiu and his associates at the University of Pittsburgh was the best documented model available.

A mathematical model based on Chiu's methods has been developed and was used to compare measured secondary velocities with computed velocities. A complete listing of the computer program for the model is included in Appendix A. The mathematical derivations and the equations used in the model are thoroughly discussed in the report. The model has been used to simulate field conditions for all the cases where field data were collected, and the results are included in the report.

Comparisons of field data and model results show mixed results. In some cases the results are satisfactory while in others they are not. In general, the model is capable of generating similar secondary circulation patterns to those indicated by the field data. However, the magnitudes of the secondary

currents generated by the model are greater than those measured in the field. Comparison of model results with measured values was further hampered by the lack of data points very close to the banks, where the model indicates the existence of consistently downward velocity components. Because of the limitations of the physical setup for collecting field data, secondary circulation data very close to the stream banks could not be collected.

The overall results of the research can be summarized by the following statements:

- A reliable state-of-the-art secondary circulation data collection system has been developed and used successfully in the field. This will significantly increase the productivity of future research in the area.
- 2) A mathematical model for secondary circulation has been developed and compared with field measurements. The results of the comparison are not totally satisfactory, and further research is needed to improve the application of the model to flow conditions in natural streams.

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APPENDICES

(SECONDARY CIRCULATION IN NATURAL STREAMS)

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			Page
Appendix	A -	List of Computer Program	1
Appendix	в -	Measured Velocities	. 23
Appendix	с -	Computed Velocities	.85

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CONTENTS

Appendix A

List of Computer Program

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PROGRAM GIOLTAPE1, TAPE2, TAPE3, TAPE4, INPUT, OUTPUT, TAPE62. TAPES= INPUT TAPES=OUTPUT) ************************************ ****************** DEFINITION OF VARIABLES **************************************

 ALPHA1(N,KPART)
 : COEFFICIENTS 1 IN SHEAR STRESS APPROXIMATION

 ALPHA2
 : COEFFICIENT 2 IN SHEAR STRESS APPROXIMATION

 BETA(KPART)
 : GEOMETRY CORRECTION FACTORS

 BEI(KPART)
 : WIDTH OF THE RIVER SURFACE ON EACH COMPUTING SIDE

 CETA(K)
 : GEOMETRY CORRECTION FACTORS

 DELY
 : DEPTH ADJUSTING COEFFICIENT

 DUDX
 : INCREAMENT OF VELOCITY U IN X DIRECTION

 DUDT
 : U INCREAMENT OF TIME

 : DEPTH
 : DEPTH AT THE Y-AXIS

 DHETA(N,K,KPART)
 : FNC8
 K - SI INCREAMENTS, N - ETA INCREAMEMT

 : DEPTH
 : DECATION OF THE POINT OF MAXIMUM VELOCITY

 . DELTA(KPART)
 : WIDTH ADJUSTING FACTORS

 . DENSITY
 : DENSITY OF FLUID

 . DENSITY
 : DENSITY OF FLUID

 . DSIDZ(N,K,KPART):
 : DERIVATIVE OF ETA VALUE

 . DETADZ(N,K,KPART):
 : DERIVATIVE OF ETA W.R.T.Z

 . ETAMIN
 : LOWER BOUND OF ETA AT TIME OF COMPUTATION

 . ETAMIN
 : DEPER BOUND OF ETA AT TIME OF COMPUTATION

 . ETAMIN
 : UPPER BOUND OF ETA AT TIME OF COMPUTATION

 MANY FUNCTIONS WHICH ARE USED FREQUENTLY ARE DEFINED IN THE FNC CATAGORY RY FNC1 = DEPTH+DELTAY+EPSILON FNC2 = BI(K)+DELTA(K) FNC3 = COMPUTE THE SI VALUE FNC3DY = DERIVATIVE OF FNC3 W.R.T. YC FNC3DZ = DERIVATIVE OF FNC4 W.R.T. YC FNC4 = COMPUTE THE ETA VALUE FNC4DY = DERIVATIVE OF FNC4 W.R.T. YC FNC4DZ = DERIVATIVE OF FNC4 W.R.T. ZC FNC5 = SIMAX(ETA) - SI FNC5 = COMPUTE THE HSI VALUE ENC 1 FNC2 FNC3 FNC3DY : FNC3DZ FNC4 : FNC40Y . FNC4DZ FNC4DZ = DERIVATIVE OF FNC4 W.R.T. ZC FNC5 = SIMAX(ETA) - SI FNC5 = COMPUTE THE HSI VALUE FNC7 = COMPUTE THE HETA VALUE FNC6 = (1./HETA)*(D(HETA)/D(SI)) FNC9 = COMPUTE FO VALUE FNC10 = THE VALUE OF SI AT GIVEN ETA CURVE WHICH INTERSECTS WITH WATER SURFACE DAVITATIONAL ACCEL FDATION .FNC5 .FNC6 .FNC7 : FNC8 . FNC9 . FNC10 INTERSECTS WITH WATER SURFACE GRAVITATIONAL ACCELERATION HYDRAULIC RADIUS TRANSFORMATION FACTOR IN ETA DIRECTION TRANSFORMATION FACTOR IN SI DIRECTION TESTER. ISTOP LE. 1 ACCEPTABLE, REJECT OTHERWISE THE BEGINING NUMBER FOR ETA CURVE ON EACH XI LINE KPART = 1 OR 2 . G HR META(N,K,KPART) HSI (N,K,KPART) JBEG(2.40) .KPART NAME OF THE CROSS SECTION INTERVALS FOR VELOCITIES PLOTTER I = 1, SI INTERVALS I = 2, ETA INTERVALS KTOTAL = TOTAL ETA CURVES USED ON EACH SIDE NECHO = 0, PRINT OUT FOR MRF ERRORS OFTION = 1, NUMERICAL ACCURACY TEST ONLY OFTION = 2, PLOT ISOVELS AND GRADIENT LINES OFTION = 3, COMPUTE TOCOEFFICIENTS CHECK INTERV. OFTION = 4, CARRY OUT WHOLE COMPUTATIONS PPLOT = 1, PLOT ISOVELS AND TRAJECTORS PPLOT = 2, PLOT NO GRAPHS PPLOT = 2, PLOT NO GRAPHS NAME OF THE CROSS SECTION NAME NDIVIDE(2) KTOTAL NECHO OPTION 1 . PPLOT1 PPLOT = 1, PLOT SHEAR STRESSES DISTRIBUTION PPLOT = 2, PLOT NO GRAPHS PPLOT = 1, PLOT Y COMPONENTS OF VELOCITIES PPLOT = 1, PLOT X COMPONENTS OF VELOCITIES PPLOT = 1, PLOT X COMPONENTS OF VELOCITIES PPLOT = 1, PLOT SECONDARY VELOCITIES PPLOT = 2, PLOT NO GRAPHS CURVELINEAR COORDINATE EDGES ON EACH SIDE STREAM FUNCTION HYDROSTATIC PRESSURE THE ENERGE SLOPE SHEAR STRESSES . PPLOT2 . PPLOT3 . PPLOT4 PPLOT5 PETA(2) .PHI (N,K,KPART) PRESURE . SE SHEAR STRESSES DERIVATIVE OF SHEAR STRESS W.R.T. SI CORDINATE SHEAR VELOCITY OF THE CROSSSECTION = SQRT(G=R*SE) VELOCITY COMPONENT IN SI DIRECTION SHEAR(N.K.KPART) . SHEARD SHVELO VSI (N.K.KPART) SIDIFF(N,K,KPART): FNC5 FNC5 SI VALUE THAT APPROXIMATE THE CHANNEL GEOMETRY LOWER BOUND OF SI AT TIME OF COMPUTATION UPPER BOUND OF SI AT TIME OF COMPUTATION XI ZERO FOR VELOCITY DISTRIBUTION VELOCITY COMPONENT IN ETA DIRECTION COMPUTED VELOCITY COMPONENT IN Y DIRECTION MEASURED VELOCITY COMPONENT IN Y DIRECTION MEASURED VELOCITY COMPONENT IN Z DIRECTION MEASURED VELOCITY COMPONENT IN Z DIRECTION WETTED PERIMETER XI(N K KPAPT) = SI COMPDINATES SINUT SIMIN . SIMAX SI ZERO VETA(N,K,KPART) .VMY(N,K,KPART) .VZ(N,K,KPART) VM2(,K,KPART) . WP XI(N,K,KPART) XI(N,K,KPART) = SI COORDINATES MAX SI VALUE ALONG A ETA CURVE .XIMAX(N,KPART) ETA(N,K,KPART) MAX SI VALUE ALONG A ETA CURVE ETA COORDINATES Y COORDINATES AT COMPUTED POINT MODIFIED Y COORDINATE FULL DEPTH CONVERTED INTO Y DISTANCE FROM ORIGIN TO PT. OF MAX. VEL. CONVERTED INTO Y SMALLEST COMPUTATION VALUE UNIVERSAL CONSTANT K 2 COORDINATES AT COMPUTED MOINT Y(N,K,KPART) . ÝC . YMAX YMEAN YMIN . UNI VK Z(N,K,KPART) Z COORDINATES AT COMPUTED POINT

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: MODIFIED Z COORDINATE
: MAX. ZC POSSIBLE IN THE COMPUTATION
: MIN. ZC POSSIBLE IN THE COMPUTATION
с
с
                     ZC.
                   . ZMAX
Ċ
                     ZMIN
С
                      .....
 č • • •
 c
c
                                                           MAIN PROGRAM
 С
                 INTEGER NA, OPTION, PPLOT1, PPLOT2, PPLOT3, PPLOT4, PPLOT5
DIMENSION ALPHA1(50,2), VSI(50,40,2), VETA(50,40,2)
DIMENSION VY(50,40,2), VZ(50,40,2), ETAMAX(2), FNINE(40)
DIMENSION DSID2(50,40,2), ETAMIN(2), DSI(5)
DIMENSION DETADZ(50,40,2), PHI(50,40,2), SHEAR(50,40,2), DUDX(40)
DIMENSION XI(50,40,2), DETA(50,40,2), SHEAR(50,40,2), DUDX(40)
DIMENSION XI(50,40,2), DETA(2), DLTA(2), NDIVIDE(2)
DIMENSION JBEG(2,40), NB(2), BI(2), ZBOUND(20), YBOUND(20)
COMMON /R1/ FNC1
COMMON /R2/ WP
COMMON /R3/ SIDIFF(50,40,2)
COMMON /R4/ F2(50,2)
COMMON /R5/ ETA(2), CETA(2)
COMMON /R5/ ETA(50,40,2)
COMMON /R6/ ETA(50,40,2)
COMMON /R7/ DETTH, DELY, EPSILON
COMMON /R7/ DETTH, DELY, EPSILON
COMMON /MR7/ YMAX, ZMAX, YONE, YMIN
CHARACTER+28 NAME
č----
                                                                                 C1 -
С
 č
                                                           INPUT PHASE
с
с1--
                 --- CHANNEL GEOMETRY DATA

READ(1,10)DEPTH, YMEAN, HR, WP, BI(1), BI(2), AVGVEL, SE

--- GENERATED COORDINATES DATA

READ(1,11)EPSILON, DELY, SINUT, BETA(1), BETA(2),

1 DELTA(1), DELTA(2), UNIVK, SIZERO

--- MAXIMUM NUMBER OF DUDX, AND PLOTTING SELECTIONS

READ(1,13) NLINES, NDIVIDE(1), NDIVIDE(2)

--- VELOCITY GRADIENT IN X DIRECTION

READ(1,12)(DUDX(1),1=1,NLINES)

--- READ IN RANGE OF SPECIFICATION

READ(1,14)SCALE, SIMAX, ETAMIN(1), ETAMIN(2), ETAMAX(1), ETAMAX(2)

DSI VALUES FOR 5 ZONES

READ(1,16)OPTION, PPLOT1, PPLOT2, PPLOT3, PPLOT4, PPLOT3

--- CHANNEL BOUNDARIES

READ(1,13)(ZBOUND(1),1=1,MCBPTS)
                      С
С
c
¢
 c
 ¢
 ¢
 С
                   READ(1,15)(ZBOUND(1),I=1,MCBPTS)
                  READ(1,15)(YBOUND(1),1=1,MCBPTS)
--- CROSECTION NAME
¢
                   READ(1,303)NAME
 С
                 kPART = 2
DENSITY = 1.94
G = 32.174
DUDT = 0.
AVGSHEA = DENSITY=G=HR=SE
FNC1 = DEPTH + EPSILON + DELY
SHVELO = SQRT(G=HR=SE)
ISTOP = MONO = NECHO = 0
WRITE(4,203)NAME
WRITE(4,225)HR,SE,WP,OPTION
WRITE(4,225)HR,SE,WP,OPTION
                  WRITE(4,22)HR,SE,WF,OFILON
WRITE(4,23)
WRITE(4,232) (ZBOUND(1),I=1,MCBPTS)
WRITE(4,232) (YBOUND(1),I=1,MCBPTS)
WRITE(4,232) (YBOUND(1),I=1,MCBPTS)
WRITE(4,232) DSI(1),DELY,(BETA(1),I=1,2),(DELTA(1),I=1,2)
WRITE(4,233) DSI(1),DSI(2),DSI(3),DSI(4),DSI(5)
WRITE(4,205)AVGSHEA,AVOVEL
 C2-
                                                                                                     Ċ
                                                  GENERATING OVERALL COEFFICIENTS FOR EACH HALF
 ċ
                  YMAX = ((YMEAN-EPSILON)+DELY)/FNC1
YONE = (YMEAN + DELY)/FNC1
IF(EPSILON .GT. 0.) YONE = YMAX
YMIN = DELY/FNC1
DO 20 K = 1,KPART
FNC2 = B1(K) + DELTA(K)
ZMAX = B1(K)/FNC2
CETA(K) = BETA(K)=(FNC1/FNC2)==2
DEFINE A LOWER BOUND FOR ETA VALUE
TOL = .1E-07
YA = 0.9999999
IF(YA .GT. YMAX) YA = YMAX
                                                                                                                                                                          ----2
 C2-
 С
                   IF(YA.GT. YMAX) YA = YMAX
CALL MRF4(YA,ZA,TOL,BETA(K),SINUT,MONO,NECHO)
ETALOW = FNC4(YA,ZA,CETA(K))
 C
                     PETA(K) = ETALOW
 с
с
  С
                  ---- DEFINITION BLOCK; CHANNEL BOUNDARIES ----
                   [ = 3
FLG = 1.
MP = 1
                          MP1 = 5
```

```
3
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```
ZA = 0 5/FNC2

S11 = FNC3(YMAX, ZA, BETA(K))

ETA1 = FNC4(YMAX, ZA, CETA(K))

IF(FLG.EG.1.) ETAMIN(K) = ETA1

S12 = FNC3(YMIN, ZA, BETA(K))

ETA2 = FNC4(YMIN, ZA, CETA(K))

IF(FLG.EG.1.) ETAMAX(K) = ETA2

IF(OPTION.NE. 1) GOTO 122

WRITE(4, 229)

WRITE(4, 220)MP, S11, ETA1, MP1, S12, ETA2

IF(MP.EG.2) GOTO 120

WRITE(4, 226)I, PETA(K)

IF(MP.EG. 2) GOTO 121

FLG = 10.

MP = MP + 1

MP1 = 4.

GOTO 130
  130
  122
                      GOTO 130
                                                    .......................
с
с
                 ..........
                WRITE(4, 199)K
WRITE(4, 200)FNC1, FNC2, ETALOW
WRITE(4, 220)
  121
С
               IF(OPTION .EQ. 4) THEN

IF(K.EQ. 1) REWIND 3

READ(3)NA,NB(1),NB(2)

DG 100 I = 1, NA

READ(3) JBEG(K,I)

NC = JBEG(K,I)

DG 110 J = NC, NB(K)

READ(3) XI(J,I,K),ETA(J,I,K),Y(J,I,K),Z(J,I,K)

YURE = Y(J,I,K) = FNC2

WRITE(4,201) J,I,XI(J,I,K),ETA(J,I,K),Y(J,I,K),Z(J,I,K),

* YTURE, ZTURE

CONTINUE
                      CONTINUE
   110
                      WRITE(4,215)
CONTINUE
   100
                       GOTO 137
                 END IF
                     ASSIGN XI CURVE VALUES (NA = NUMBER OF ISOVELS)
С
                 NATIME = 1
TEST = SINUT
С
                      D0 149 1 = 1,40
IF(1 .EQ. 10) NATIME = 2
IF(1 .EQ. 18) NATIME = 3
IF(1 .EQ. 26) NATIME = 4
IF(1 .EQ. 32) NATIME = 5
С
                        IF(1 .NE. 1) TEST = TEST + DSI(NATIME)
IF(TEST .GT. SIMAX) GOTO 148
                       18(1
                       NA = 1
                       CONTINUE
   149
С
   148
                      NATIME = 1

IF(EPSILGN.LT.O.O) THEN

XI(NA,1,K) = SINUT

DO 150 I = 2, NA

IF(I .EQ. 10) NATIME = 2

IF(I .EQ. 26) NATIME = 3

IF(I .EQ. 26) NATIME = 4

IF(I .EQ. 32) NATIME = 5

I) = I - 1

I2 = NA - I1

I3 = 12 + 1

XI(I2,1,K) = XI(I3,I1,K) + DSI(NATIME)

CONTINUE
                       NATIME .
   150
                             CONTINUE
                 FLSE
                                X1(1,1,K) = SINUT
                            ACT, I, I = SINUT

DO 159 I = 2,NA

IF(I .EQ. 10) NATIME = 2

IF(I .EQ. 16) NATIME = 3

IF(I .EQ. 26) NATIME = 4

IF(I .EQ. 32) NATIME = 5

{1 = I - 1
                             IF(EPSILON.EO.O.O) XI(1,1,K) = XI(1,11,K) + DSI(NATIME)
IF(EPSILON.GT.O.O) XI(1,1,K) = XI(11,11,K) + DSI(NATIME)
                             CONTINUE
   159
                END IF
                          DETERMIN ETAMIN ON EACH SI CURVE IF EPSILON IS LE O
С
                 IF(EPSILON .GE. 0.0) GOTO 152
I = J = 0
                11 = NA - 1
J = J + 1
   155
                 Y(J,T1,K) = Z(J,T1,K) = 0.0
                J1 = J + 1
CALL MRF4(YMAX,ZA,TOL,BETA(K),X1(J,I1,K),MONO,NECHO)
ETAMINI = FNC4(YMAX,ZA,CETA(K))
IF(I1.EQ. NA) THEN
IF(ETAMINI .LT. ETAMIN(K)) STOP
ETA(J,I1,K) = ETAMINI
Y(J,I1,K) = YMAX
Z(J,I1,K) = YMAX
Z(J,I1,K) = ZA
DO 156 M = J1,NA
Y(M,I1,K) = Z(M,I1,K) = 0.0
XI(M,I1,K) = XI(J,I1,K)
FLSE
                  J1 🗴 J + 1
   158
                 ELSE
                      ETA(J,11,K) = ETAMINI
Y(J,11,K) = YMAX
```

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4
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```
Z(J,I1,K) = ZA

12 = 11 + 1

DO 157 M = J1,NA

Y(J,M,K) = Z(J,M,K) = 0.0

X1(M,I1,K) = X(J,11,K)

DO 158 M = 12,NA

ETA(J,M,K) = ETA(J,I1,K)

DO 15
   157
    160
    158
                    END 1F
                END IF

[ = [ + 1

IF(I.LT.NA) GOTO 155

---- DC A BOUNDARY LINE ----

J = J + 1

XI(J,1,K) = SINUT

ETA(J,1,K) = PETA(K)

NATIME = 1

DO 151 M = 2, NA

IF(M .EQ. 10) NATIME = 2

IF(M .EQ. 10) NATIME = 3

IF(M .EQ. 26) NATIME = 4

IF(M .EQ. 26) NATIME = 5

Y(J,M,K) = Z(J,M,K) = 0.0

ETA(J,M,K) = Z(J,M1,K) + DSI(NATIME)
С
    151
 с
                            JBASE = J + 1
                     GOTO 161
с
с
                      --- EPSILON .GT. 0.0 ---
C
152 J = 1
ETA(
                        TA(1,1,K) = PETA(K)
Y(1,1,K) = Z(1,1,K) = D.0
JBASE = 1
                    JBASE = 1

[ = 1

IF(EPSILON .GT. 0.0) THEN

JBEG(K,I) = J

J = J + 1

[ = I + 1

J1 = J - 1

If(I .GT. NA) GOTO 170

JBEG(K,I) = J

Y(J,I,K) = Z(J,I,K) = 0.0

-- DETERINE ETA ON WATER SURFACE ---

CALL MRF4(YMAX,ZA,TOL,BETA(K),XI(J,I,K),MONO,NECHD)

ETA(J,I,K) = FNC4(YMAX,ZA,CETA(K))

Y(J,I,K) = YMAX

Z(J,I,K) = ZA

DO 166 M = 1,J1

TUJ,D,N = 4(JJ,D,N = 4,VJ)
                      1 = 1
     169
 С
                - - -
                              T(J,M,K) = X((J1,M,K)

ETA(J,M,K) = ETA(J,1,K)
      168
                              GOTO 169
      170
                              J = 1
J = J1
   С
                       DERIVE DISTRIBUTION GRID SYSTEM,
Solve the y Z coordinate for the last point
   0000
                        Y1 = YMAX
                       Z1 # ZMAX
DUMY # 1.0
                    DUMY = 1.0
CALL CORDTF(CETA(K), BETA(K), XI(J,1,K), ETA(J,1,K),
+ BI(K), DELTA(K), Y1, Z1, YA, ZA, ISTOP, DUMY)
Y(J,1,K) = YA
Z(J,1,K) = ZA
  C
174
                              J1 = J + 1

IF(J1 .GT, 50) THEN

NB(K) = J

GOTO 172
                             GOTO 172
END IF
IF(YA.GT. 0.25) THEN
YA = YA = 0.1
CALL MRF4(YA,ZA,TOL,BETA(K),SINUT,MONG,NECHO)
ETA(J1,I,K) = FNC4(YA,ZA,CETA(K))
Y(J1,I,K) = YA
Z(J1,I,K) = YA
XI(J1,I,K) = XI(J,I,K)
ELSE
                            ELSE
ZA = ZA - 0.03
[F(ZA .LE. 0.0] THEN
NB(K) = J
GOTO 172
END IF
CALL MRF1(YA,ZA,TOL,E
                                   END IF

CALL MRF1(YA, ZA, TOL, BETA(K), SINUT, ETA(J, I, K), MONO, NECHO)

ETA(J1, I, K) = FNC4(YA, ZA, CETA(K))

IF(ETA(J1, I, K) .GT. ETAMAX(K)) THEN

NB(K) = J

GOTO 172

END IF

XI(J1, I, K) = X1(J, I, K)

Y(J1, I, K) = YA

Z(J1, I, K) = ZA

ND IF
                             END IF
DO 173 M = 2,NA
Y(J1,M,K) = Z(J1,M,K) = 0.0
X1(J1,M,K) = X1(J,M,K)
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5
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ETA(J), M, K) = ETA(J), I, K)
J = J + 1
GOTO 174
170
                END IF
С
              -- DEFINE DETA VALUES IN THE LOWER PORTION
DERIVE DISTRIBUTION ON ETA,
SET DYC = 0.1 MANUALLY FOR 10 INTERVALS (YA = .999999 INIT(ALLY)
С
С
с
с
   161
                 J1 & J + 1
JF(J1 .GT, 50) GOTO 154
IF(YA .GT. 0.25) THEN
YA = YA - 0.1
   153
                            YA = YA - 0.1

CALL MRF4(YA,ZA,TOL,BETA(K),SINUT,MONO,NECHO)

ETA(J1,1,K) = FNC4(YA,ZA,CETA(K))

X1(J1,1,K) = X1(J,1,K)

Y(J1,1,K) = YA

Z(J1,1,K) = ZA

PE
                     Z(J1,I,K) = ZA

ELSE

ZA = ZA - 0.03

IF(ZA .LE. 0.0) GOTO 154

CALL MRF1(YA,ZA,TOL,BETA(K),SINUT,ETA(J,I,K),MONO,NECHO)

ETA(J1,I,K) = FNC4(YA,ZA,CETA(K))

IF(ETA(J1,I,K) = FNC4(YA,ZA,CETA(K))

IF(ETA(J1,I,K) = FNC4(YA,ZA,CETA(K))

Y(J1,I,K) = YA

Z(J1,I,K) = YA

Z(J1,I,K) = YA

Z(J1,I,K) = ZA

END IF

J = J + 1

GOTO 133

B(K) = J
   154 NB(K) = J
J = JBASE
                J = JBASE

NATIME = 1

DO 165 M = 2, NA

IF(M .EQ. 10) NATIME = 2

IF(M .EQ. 16) NATIME = 3

IF(M .EQ. 26) NATIME = 3

IF(M .EQ. 32) NATIME = 5

ETA(J,M,K) = ETA(J,1,K)

M1 = M ~ 1

X1(J,M,K) = X1(J,H1,K) + OSI(NATIME)

Y(J,M,K) = Z(J,M,K) = 0.0

CONTINUE

J = J + 1
    167
   165
   165 CONTINUE

J = J + 1

[F(J, GT, NB(K)) GOTO 166

GOTO 167

166 DO 301 [ = 1,NA

IM = NA - I + 1

[F(EPSILON .EQ. 0.0) [M = 1
JOEU(K, I) = IM
301 CONTINUE
172 CONTINUE
C
                   IF (OPTION.EQ.2 . OR. OPTION.EQ.3) WRITE(2)NA, NB(1), NB(2)
  000
              -- DEFINE INITIAL GUESSES FOR SUBROUTINE CORDEF
                    YOPER + YMIN
                  YOPER = YMIN

ZOPER = 0.00001

DYY = 1.0

CO 132 ( = 1,NA

IF(OPTION.EQ.2.OR. OPTION.EQ.3) WRITE(2) JBEG(K,1)

NCHECK = JBEG(K,1)

DO 133 J = NCHECK,NB(K)

J1 = J - 1

ISTOP = 0

DEFENSE (IMPER) (IMITE FOR M. NO. 5
                    ...DEFINE UPPER (LOWER) LIMITS FOR Y AND Z
IF(J .EQ, NCHECK) THEN
WRITE(4,215)
 ¢
                        WRITELA,213J

IF(Y(J,1,K) .NE. 0.0 .AND. Z(J,1,K) .NE. 0.0) THEN

YC = Y(J,1,K)

ZC = Z(J,1,K) .

GOTO 136

END IF

IF(ETA(J,1,K) .GE. 0.) THEN
                             Z1 = ZMAX
ZC = Z1 - 0.1
Y1 = YONE
YC = Y1 - 0.1
                        ELSE

Z1 = 0.00001

ZC = Z1 + 0.1

Y1 = YMAX

YC = Y1 - 0.1
                        END IF
                   ELSE
                             E

IF(Y(J,I,K) .NE. 0.0 .AND. Z(J,I,K) .NE. 0.0) THEN

YC = Y(J,I,K)

ZC = Z(J,I,K)

GOTO 136
                       GOTO 136
END IF
IF(ETA(J,1,K) .GE. 0.) THEN
IF(ETA(J,1,K) .LT. 0.) THEN
Z1 = ZMAX
Y1 = YONE
END IF
YC = (Y1 + YOPER)/2.
ZC = (Z1 + ZOPER)/2.
                       EL SE
```

```
YC = (Y1 + YONE)/2.
ZC = (Z1 + ZMAX)/2.
         END IF
С
      --- - CORDTF FIND CORRESPONDING Y Z COORDINATES
- Y1 Z1 ARE THE GUESSED VALUES AND YC ZC ARE SOLUTIONS
С
č
c
         CALL CORDTF(CETA(K), BETA(K), XI(J, 1, K), ETA(J, 1, K),
        000
c
č
C
             GOTO 999
          END IF
          HF(ISTOP .EQ. 2) THEN
WRITE(4,227)ISTOP,XI(J,I,K),ETA(J,I,K),YI,ZI,YC,ZC
        GOTO 999

END IF

Y(J,I,K) = YC

Z(J,I,K) = ZC

YTRUE = YC4FNC1-DELY

ZTRUE = ZC4FNC2

WRITE(4,201)J,I,XI(J,I,K),ETA(J,I,K),YC,ZC,YTRUE,ZTRUE

IF(OPTION.EQ.2.OR. OPTION.EQ.3) WRITE(2)XI(J,I,K),ETA(J,I,K),

Y(J,I,K),Z(J,I,K)

Y(J,I,K),Z(J,I,K)
             GOTO 999
 28
 136
 134
 133
0000
               FOR ALL ETA .GE. D.O. DETERMINE THE MIN Y FOR
The NEXT SI CURVE BY USING CURRENT Z AS Z UPPER
            JNOW = NB(K)
            IF(ETA(JNGW, I, K) .GE. 0.0) THEN (PLUS = 1 + 1
               IF(IPLUS. GT, NA) GOTO 132
IF(Y(JNOW,IPLUS,K).NE.O., AND.Z(JNOW,IPLUS,K).NE.O.) GOTO 132
              - 22 = 2(JNOW, I,K)
DUMY = 0.00001
               CALL CORDIF(CETA(K), BETA(K), XI(JNOW, IPLUS, K), ETA(JNOW, I, K),
                 BI (K) , DELTA(K) , Y1 , Z2 , YOPER , ZOPER , MONO , DUMY )
            END IF
  132 CONTINUE
С
          IF(OPTION .EQ. 2) GOTO 2D
c
  137 CONTINUE
               COMPUTE OVERALL COEFFICIENTS
C
С
           XITERM . SHVELO / UNIVK
          DO 26 1 = 1,NA
WRITE(4,215)
          WRITE(4,213)

NC = JBEG(K,1)

XIVELO = XITERM = ALOG(XI(NC,1,K)/SIZERG)

FNINE(1) = FNC9(DENSITY,DUDT,DUDX(1),XIVELO,HR,AVGSHEA)

DO 21 J = NC,NB(K)

HS1(J,1,K) = FNC6(K,FNC2,XI(J,1,K),Y(J,1,K),Z(J,1,K))

DFTA(1 + K) = FNC6(K,FNC2,XI(J,1,K),Y(J,1,K),Z(J,1,K))
  18
          HETA(J, I, K) = FNC7(K, FNC2, XI(J, I, K),
ETA(J, I, K), Y(J, I, K), Z(J, I, K), HSI(J, I, K))
с
с
          DETERMINE SIMAX - SI VALUES
ċ
          IF(EPSILON .LE. 0.0) THEN
XIMAX(J,K) = 1.0
SIDIFF(J,1,K) = FNC5(XIMAX(J,K),X1(J,1,K))
           ELSE
                TF(1 .EQ. 1) THEN
                   XIMAX(J,K) = FNC10(ETA(J,I,K), CETA(K), BETA(K))
                END IF
SIDIFF(J,1,K) = FNC5(XIMAX(J,K),XI(J,1,K))
          END IF

DHETA(J,1,K) = FNC8(K,X1(J,1,K),Y(J,1,K),Z(J,1,K))

DSIDZ(J,1,K) = FNC8DZ(Y(J,1,K),Z(J,1,K),BETA(K))

DETADZ(J,1,K) = FNC4DZ(Y(J,1,K),Z(J,1,K),CETA(K))
           CONTINUE
  21
  26
           CONTINUE
с
           WRITE(4,207)
WRITE(4,203) (FNINE(10),10=1,NA)
WRITE(4,202)
DO 22 I = 1,NA
WRITE(4,215)
WRITE(4,215)
           NC = JBEO(K, I)
           DO 23 J = NC,NB(K)
WRITE(4,201)J,1,HS1(J,1,K),HETA(J,1,K),S1D1FF(J,1,K),
DHETA(J,1,K),DS1D2(J,1,K),DETADZ(J,1,K)
CONTINUE
  23
  22
           CONTINUE
           CONTINUE
 С
           IF (OPTION . EQ. 1) GOTO 999
IF (OPTION . EQ. 3) GOTO 998
```

```
7
```

```
IF (OPTION . EQ. 2) THEN
CALL DIBOISO (NA MOBPTS DEPTH EPSILON, Y. Z. JEEG.
      +
                              NB, B1, ZBOUND, YBOUND, FNC1, DELTA, DELY)
       END IF
C
-
C3------
   COMPUTATION ALONG BOUNDARY LINE. DETERMINE SHEAR COEFFICIENTS
C
č
.
C3-----3
с
с
            ALPHA2 IS COMPUTED FOR EAT CURVES ON BANKS AND BED ONLY
       ALPHA2 = 0.
       DO 30 K = 1,KPART
       IF(EPSILON ,LT. 0.0) J1 = NA
DO 31 J = J1,N8(K)
F2(J,K) =
           SIDIFF(J,I,K)/(1.-DHETA(J,I,K)=SIDIFF(J,I,K))
      1
       CONTINUE
 31
       CALL SHEAR2(K, J1, NB(K), AVGSHEA, FNINE(1), ALPHA2, T1, T2) WRITE(4, 206) K, J1
       WRITE(4, 203) (F2(J,K), J=J1, NB(K))
 30
       CONTINUE
       WRITE(4,208)ALPHA2
C4
C
       COMPUTE SHEAR STRESS AND SECONDARY FLOW COMPONENTS VSI, AND THE
с
с
       STREAM FUNCTION FOR EACH HALF OF THE CROSS-SECTION
с
с4
                                                                     -----
       DG 40 K = 1,KPART
COMPUTE ALPHA1 COEFFICIENT FOR EACH ETA CURVE
C
        1 . . . . 1
        IF (EPSILON . LT. 0.0) THEN
         1 = NA - 1
J1 = NA - 1
DG 44 J = 1,J1
ALPHA1(J,K) = -ALPHA2=SIDIFF(J,I,K)
         1 = 1 - 1
 44
       CONTINUE
       I = NA
END IF
       D5 45 J = 1,NB(K)
        CALL SHEARI(K, J, FNINE(1), ALPHA2, ALPHA1(J,K))
 45
       CONTINUE
C
       DO 41 I = 1, NA
NC = JBEG(K, 1)
        06 42 J = NC, NB(K)
     COMPUTE SHEAR STRESS AND VSI ALONG EACH ETA LINE
С
С
C..... COMPUTE VSI (AS U FOLLOWS LOGRITHMATIC DISTRIBUTION)
        RODUDSI = (DENSITY=SHVELO)/(XI(J,1,K)=UNIVK)
        TERM1 = DHETA(J,I,K)=SHEAR(J,I,K)
TERM2 = SHEARD - HETA(J,I,K)=FNINE(I)
VSI(J,I,K) = (TERM1 + TERM2)/RODUDSI
с...
с...
       COMPUTE STREAM FUNCTION ALONG EACH SI LINE, NAMELY SUM(VSI=HETA)
ALSO COMPUTE SUM(HETA=HSI) AND TEMPORARY STORE INTO VETA
        J1 = J = 1
IF(J .EQ. NC) THEN
PH1(J,I,K) = 0.
VETA(J,I,K) = 0.
        ELSE
          TVX11 = VSI(J1,1,K)
          TVX11 = V31(J1,I,K)

TVX12 = V31(J1,I,K)

THETA1 = HETA(J1,I,K)

THETA2 = HETA(J,I,K)

DETA = ETA(J,I,K) - ETA(J1,I,K)

CALL INTPHI(TVX11,TVX12,THETA1,THETA2,PQ,DETA)

PHI(J,I,K) = PHI(J1,I,K) + PQ

TVX11 = HSI(J1,I,K)

TVX12 = HSI(J,I,K)

CALL INTPHI(TVX11,TVX12,THETA1,THETA2,PD,DETA)
          CALL INTPHI(TVXI1, TVXI2, THETA1, THETA2, PQ, DETA)
VETA(J, [,K) = VETA(J1, [,K) + PQ
        END IF
 42
        CONTINUE
  41
  40
        CONTINUE
                   C5-
c
c
                        COMPUTE VETA'S
Ċ
C5.
                                         DC 50 K = 1,KPART
WR(TE(4,199)K
NA1 = NA - 1
DC 51 [ = 2,NA1
NC = JBEG(K,1)
000
                               COMPUTE VETA AT WATER SURFACE
                         . . .
        TX = HSI(NC,I,K)=DSIDZ(NC,I,K)
VZ(NC,I,K) = VSI(NC,I,K)/TX
```

```
VETA(NC, I, K) = HETA(NC, I, K) + DETADZ(NC, I, K) + VZ(NC, I, K)
 С
               NCCA + NB(K)
               11 = 1 = 1
12 = 1 + 1
               DS11 = XI (NCCA, 12, K) - XI (NCCA, 11, K)
               DS(1 * X((NCCA, (2,K) * X((NCCA, (1,K)
ND * NC + 1
DO 52 J = ND, NB(K)
TERM * (-PH((J, (1,K) + PH((J, (2,K))/DS))
TERM * TERM/HSI(J, (,K) + PH((J, (2,K))/DS))
VETA(J, (,K) * VETA(NC, (,K) - DUDX(())*VETA(J, (,K)/HSI(J, (,K) +
TERM -
               CONTINUE
52
51
C
   52
               CONTINUE
               WRITE(4,223)
WRITE(4,203)
WRITE(4,203)(ALPHA1(L1,K),L1=1,NB(K))
WRITE(4,215)
               WRITE(4,213)
WRITE(4,211)
DO 53 IP = 1,NA
MNC = JBEG(K,IP)
DO 54 IQ = MNC,NB(K)
IF(IP.EQ.1.0R. IP.EQ.NA) THEN
WRITE(4,201)IQ, IP, SHEAR(IQ,IP,K)
                ELSE
                    -SE
WRITE(4,201)10,1P,SHEAR(10,1P,K),VSI(10,1P,K),VETA(10,1P,K)
               END IF
CONTINUE
WRITE(4,215)
CONTINUE
    54
    53
                CONTINUE
    50
                 C6--
  č
  č
                                                     COMPUTE VY AND VZ AT COORDINATE (SI, ETA)
  С
               DO 60 K = 1, KPART

FNC2 = B1(K) + DELTA(K)

WRITE(4,199) K

WRITE(4,222)

WRITE(4,217)

DO 62 [ = 2,NA1

WRITE(4,215)

NC = JBEG(K, []

DO 61 J = NC, NB(K)

YTEMP = Y(J,1,K)

ZTEMP = Z(J,1,K)

YTR = YTEMP=FNC1 - DELY

ZTR = ZTEMP=FNC2
  Č6--
               IF(K .EQ. 1) ZTR = B1(1) - ZTR

IF(K .EQ. 2) ZTR = ZTR + B1(1)

OXIDY = FNC3DY(YTEMP,ZTEMP,BETA(K))

DXIDZ = DSIDZ(J,1,K)

PETAPY = FNC4DY(YTEMP,ZTEMP,CETA(K))

PETAPZ = DETADZ(J,1,K)

IF(J .EQ. NC) THEN

VY(J,1,K) = 0.

ELSE
              VY(J,[,K) = 0.
ELSE
D0 = HSI(J,[,K)=HETA(J,],K)=(DX1DY=PETAPZ-DX1DZ=PETAPY)
D1 = VSI(J,[,K)=HETA(J,],K)=PETAPZ -
+ VETA(J,],K)=HETA(J,],K)=DX1DZ
D2 = -VSI(J,],K)=HETA(J,],K)=PETAPY +
+ VETA(J,],K)=HETA(J,],K)=PETAPY +
+ VETA(J,],K)=HETA(J,],K)=PETAPY +
+ VETA(J,],K)=HETA(J,],K)=PETAPY +
+ VETA(J,],K)=HETA(J,],K)=PETAPY +
+ VETA(J,],K)=D1 / D0
VZ(J,],K) = D2 / D0
END IF
              ٠
               WRITE(4,218)1,J,YTR,ZTR,VY(J,1,K),VZ(J,1,K)
    61
                CONTINUE
   62
60
                CONTINUE
                CONTINUE
  c
                     DRAW SHEAR STRESSES AND VELOCITY COMPONENTS
  č
               CALL DI303V(NA, MCBPTS, DEPTH, EPSILON, Y, Z, JBEG,

NB, BI, ZGOUND, YBOUND, FNCI, DELTA, DELY, SHEAR, VY, VZ,

PPLOT2, PPLOT3, PPLOT4, PPLOT5, AVGSHEA, NDIVIDE,
                     AVOVEL, SCALE, NAME)
  С
                GOTO 999
  С
               DO 63 K = 1,2
WRITE(4,199)K
WRITE(4,300)
    998
  C
                        = NA -
               NA1 = NA - 1

DO 64 | = 2,NA1

ADD 1 TO NCI SO THAT IT START AT CORRECT POINT ([+1) WILL

HAVE CORRESPONDING NUMBERS

NC1 = JBEG(K, 1) + 1

DO 65 J = NC1,NB(K)

DSITEMP = XI(J,I+1,K) - XI(J,I-1,K)

A1 = ALOG(HETA(J,I+1,K))

A2 = ALOG(HETA(J,I+1,K))

A3 = (A2-A1)/DSITEMP

WRITE(4 3021 J. HETA(J,I,K) DSITEMP DHETA(1,I,K),A3
                NA1
  C
C
                              WRITE(4, 302) J, I, HETA(J, I, K), DSITEMP, DHETA(J, I, K), A3
    65
                          CONTINUE
  64
63
C
                     CONTINUE
                CONTINUE
```

```
9
```

```
FORMAT(/7F8.2,F9.6)
FORMAT(/2F6.2,7F8.4)
FORMAT(/10F6.3)
 10
 12
  13
           FORMAT(/315)
           FORMAT(76F6.3)
 14
           FORMAT(10F6.2)
  15
        FORMAT(10F6.2)

FORMAT(7613)

FORMAT(7613)

FORMAT(715X, 'CROSS-SECTION PART',13)

FORMAT(715X, FNC1',E10.4,3X, 'FNC2',E10.4//,

* 15X, 'ETALOW',E10.4/)

FORMAT(715X, FNC1',10.4)

FORMAT(74X, J',4X, '1',

* 7X, 'HSI',14X, 'HETA',13X, 'SIDIFF',11X, 'OHETA',12X,

* 'DSID2',12X, 'DETAD2'/113(1H-))

FORMAT(74)2X, 'DETAD2'/113(1H-))
 16
 199
 200
 201
 202
         + 'DSIDZ',12X,'DETADZ'/113(1H-))
FORMAT(8(2X,E10.4))
FORMAT(//10X,'CHANNEL INFORMATION',5X,'AVERAGE SHEAR STRESS',
+ 'DSX,FI0.6,10X,F10.4/)
FORMAT(//9X,'F2',3X,'PART',13,'**** ETA SEG.',13/)
FORMAT(//9X,'F2',3X,'PART',13,'**** ETA SEG.',13/)
FORMAT(//9X,'INTERMEDIATE COEFFICIENTS'//10X,'FNINE'//)
FORMAT(//9X,'ALPHA1'/)
FORMAT(//4X,'J',4X,'1',7X,'SHEAR',12X,'VSI',14X,'VETA',
+ 'G6(1H-1/)
 203
 205
 206
207
 208
 209
 211
          + /96(1H-)/)
FORMAT(/)
215
  231
           FORMAT(/10X, 'MEASURED CHANNEL BOUNDRIES'/)
  232
           FORMAT(/10X.10F6.2)
           FORMAT(//IOX, JOFD.2)
FORMAT(//IOX, 'ZONE INTERVALS (DSI VALUES)',/IOX,5F6.4/)
FORMAT(/,15X, 'PERFORM NUMERICAL INTERVAL CHECK',//)
  233
  302
           FORMAT(20X,215,3X,4E10.4)
           FORMAT (A28)
  303
  999
           STOP
           END
C - - - - -
                                                   FNC3
           FUNCTION FNC3(Y,Z,BA)
AO = 1. - Z
A1 = AO**BA
A2 = EXP(BA*Z-Y+1.)
           FNC3 = Y#A1#A2
 8
            RETURN
            END
              ....
                                  -----FUNCTION FNC4
C----
           FUNCTION FNC4(Y, Z, CETA)
           A2 = EXP(Z+CETA+Y)
A0 = 1, - Y
IF(A0.LT.D.)GOTO7
A1 = A0==CETA
FNC4 = (1./Z)+A3=A2
RETURN
           A1 = (-A0)==CETA
FNC4 = - A1=A2/Z
  7
            RETURN
           END
c----
                                                            -----FUNCTION FNC3DY
           FUNCTION ENCODY(Y, Z, BA)
           A1 = (1.-Z)==BA
A2 = EXP(BA=Z-Y+1.)
FNC3DY = A1=A2=(1.-Y)
RETURN
           END
                                                       -----FUNCTION FNC3DZ
                 . . . . . . . . . .
C----
           FUNCTION FNC3DZ(Y, Z, BA)
           A1 = 1.-Z
A2 = A1==BA
A3 = EXP(BA=Z-Y+1.)
           FNC3DZ = BA=A2=Y=A3=(-Z)/A1
RETURN
           END
C----
                                                              -----FUNCTION FNC4DZ
          FUNCTION FNC4DZ(Y,Z,CETA)
A2 = EXP(Z+CETA=Y)
IF(Y.GE.1.) THEN
A0 = Y - 1.
A3 = (1. - Z) / Z
           FI SE
                  AO = 1. - Y
```

```
A3 = (2 - 1.) / Z
END (F
           A4 = (A0==CETA) / Z
FNC4DZ = A0 = A2 = A3
           RETURN
           END
               -----FUNCTION FNC40Y
C - - - -
          FUNCTION FNC4DY(Y, Z, CETA)
с
              F Y = 1. ASYMTOTICALLY VALUES ARE CONVERGED
THEREFORE IT SET Y = 0.99999
0 C C C
           IF Y
           IF(Y .EQ. 1.) Y = 0.99999
          IF(Y .EWF(Z+CETA=Y)
A0 = 1. - Y
IF(Y .GT. 1.) A0 = Y - 1.
A1 = A0 == (CETA - 1.)
FNC4DY = (CETA/Z) = A1 = A2 = (-Y)
           RETURN
           END
            FUNCTION FNC5
C----
           FUNCTION FNC5(V51,V52)
FNC5 = V51 - V52
  30
           RETURN
           END
c----
                        -----FUNCTION FNC6
           FUNCTION FNC6(K, FNC2, ASI5, AY5, AZ5)
           FUNCTION FROSIK, FRC2, ASIS, A
COMMON /R1/ FNC1
COMMON /R5/ BETA(2), CETA(2)
AA53 = FRC1*AY5=(1.-AZ5)
AA53 = (1.-AY5)=(1.-AZ5)
AA54 = AA53*AA53
           AA55 = BETA(K) = (FNC1/FNC2) = AY5=AZ5
AA56 = AA55=AA55
            AA52 = AS15=SORT(AA54+AA56)
           FNC6 = AA51/AA52
            RETURN
           END
c - - - -
          FUNCTION FNC7(K,FNC2,ASI7,AETA7,AY7,AZ7,HFNC6)
COMMON /R1/ FNC1
COMMON /R5/ BETA(2),CETA(2)
AA71 = FNC2/FNC1
AA72 = ASI7/AETA7
AA73 = AZ7=(1.-AY7)
AA74 = AY7=(1.-AZ7)
FNC7 = A471=A473(AA77/AA74)-UFNC4
                                                                       ----FUNCTION FNC7
            FNC7 = AA71+AA72+(AA73/AA74)=HFNC6
            RETURN
           END
                                   FUNCTION FNC8
C----
            FUNCTION FNCS (K, ASI8, AY8, AZ8)
           COMMON /R5/ BETA(2), CETA(2)
AA81 = (1.-AY8)=(1.-AY8)=(1.-AZ8)=(1.-AZ8)
AA82 = AY8=AY8=AZ8=AZ8
С
           AA83 = AA81=CETA(K)=AY8=AY8
AA84 = CETA(K)=BETA(K)=AA82=(1.-AZ8)=(1.-AZ8)
AA85 = AS18=((AA81+CETA(K)=BETA(K)=AA82)=
                                           (AA81+CETA(K)#8ETA(K)#AA82))
С
            FNC8 = - (AA83+AA84)/AA85
            RETURN
            END
                        FINCE FUNCTION FINCE
0000
             . . . .
           FUNCTION FNC9 COMPUTES FO OR F VALUES ACCORDING TO THE
MOMENTUM EQUATION. ALL INPUT VALUES ARE PRECALCULATED
FROM FIELD MEASURED DATA.
FUNCTION FNC9(DENSITY, DELUT, DELUX, XITERM, HR, SHEAR)
AMT = DELUT + XITERM = DELUX
AMS = DENSITY=AMT
FNC9 = AMS-SHEAR/HR
            RETURN
            END
 C----
                  FINCTION FICTO
            FUNCTION FNC10(V12,V15,V16)
COMMON /R7/ DEPTH, DELY, EPSILON
COMMON /MRF/ YMAX, ZMAX, YONE, YMIN
            EXTERNAL FNC3
            TOL * 1E-OT
on the water surface, y_1 = (\text{Depth+Dely})/(\text{dep+Dely+EP1}). Therefore
we have eta and y_1 specified.
z = 0.
C...
            NECHO . O
            CALL MRF2(YMAX, Z, TOL, V15, V12, 1STOP, NECHO)
FNC10 = FNC3(YMAX, Z, V16)
   99
            RETURN
            END
           -SUBROUTINE CORDIFICETA, BEDA, SI, ETA, BI, DELI, YU, ZU, YI, ZI, MONO, DYY)

THIS SUBROUTINE IS THE COMBINATION OF NONLINI AND

NONLIN2, I.E. IT USES MODIFIED REGULA FALSI METHOD

TO FINO Y OR Z, THEN USE BISECTION METHOD ON THE

OTHER DIRECTION.

MONO = 0 COMPLETES THE LOOP

MONO = 1 NO SOLUTION FROM MRF4 - MRF1

MONO = 2 THIS ETA VALUE DOES NOT EXIST (FAIL IN 'CIRCL')

COMMON /R7/ DEPTH, DELY, EPSILON

COMMON /CIR/ PASI, PAETA

COMMON /LL/ TOL, TOLETA
 C----
                                  -----SUBROUTINE CORDTF
 00000000
```

PASI = SI PAETA = ETA TOL = IE-07 TOLETA = IE-3 MONO = ITER = II4 = NECHO = 0 ... - START SEEKING SEQUENCE; DETERMINE UPPER AND LOWER BOUNDS BEFORE CALL SUBROUTINE CIRCLES DEFINE UPPER OR LOWER REGION DEFINE UPPER OR LOWER REGION PAS! = S! С c IF(PAETA.LT.O.)GOTO70 0000 LOWER REGION Y = YU Z = ZU IF(DYY .LT. 0.01 .AND, Y .LT. 0.9) GOTO 24 CALL MRF4(Y,Z,TOL,BEDA, PASI, MONG, NECHO) IF(MONO.NE.1)GOTO 23 YU = Y ZU = Z 23 BANK AND UPPER BOTTOM REGION COMPUTATION ¢ 114 # 4 CHECK A LOWER BOUND BEFORE CALL CIRCLL c NECHO . NECHD = 1 CALL MRF4(Y1,Z,TOL,BEDA,PASI,MONO,NECHO) ITER = ITER + 1 IF(ITER.EG. 1) DY = (YU - Y1)/8. IF(ITER.GT. 6) THEN WR(TE(4,52) 26 GOTO 50 END IF Y1 # Y1 + DY GOTO 26 25 FIE: = FNC4(Y1,Z,CETA) - PAETA (F(ABS(FIE)).LT.TOLETA) THEN Y1 = Y1 Z1 = Z RETURN IF (FIEL)11,12,13 IF (FIEL)11,12,13 Lower Bottom Region Check Lower Bound Before Call Circle C C 24 CHECK LOWER BOUND BEFORE CALL CIRCLL YU = Y I14 = 1 NECHO = 0 CALL MRF1(Y, Z1, TOL, BEDA, PASI, PAETA, MONG, NECHO) IF(MONO.NE. 1) GOTO 27 WRITE(4,55) GOTO 50 CHECK IF THIS IS LOWER BOUND С FIE1 = FNC4(Y,Z1,CETA) - PAETA IF(ABS(F)E1) LT.TOLETA) THEN Y1 = Y Z1 = Z1 27 RETURN END IF IF(F1E1)20,12,21 ¢ UPPER REGION С FIE1 = FNC4(YU,ZU,CETA) - PAETA [F(ABS(FIE1) .LT. TOLETA) GOTO 12 CALL MRF4(Y1,Z,TOL,BEDA,PASI,MONO,NECHO) [F(MONO,NE.1) GOTO 40 FIE2 = FNC4(Y1,Z,CETA) - PAETA IF(ABS(FIE2) .LT. TOLETA) GOTO 12 [F(F(E1=FIE2) 18,12,18 c 70 40 YU = Y1 ZL = Z ZU = ZMAX YL = YONE GOTO 77 18 GOTO 77 YL = Y1 ZU = Z GOTO 77 YU = Y ZU = Z1 YL = YMIN ZL = 0.00001 18 20 GOTO 60 YL = YZL = Z1 21 GOTO 60 ZU = Z YU = Y1 11 YL = YMIN ZL = 0.00001 GOTO 60 ZL = Z YL = Y1 CALL CIRCLL(BEDA,CETA,YU,YL,ZU,ZL,Y,Z,114) IF(114 .EQ. 100) THEN MONO = 2 WRITE(4,56)BEDA,CETA,YU,YL,ZU,ZL,Y,Z 13 60 С GOTO 12 ELSE GOTO 12 END IF CALL CIRCLU(BEDA, CETA, YU, YL, ZU, ZL, Y, Z, 114) IF(114 .EQ. 100) THEN 77

```
MONO # 2
                    WRITE(4,56)BEDA, CETA, YU, YL, ZU, ZL, Y, Z
GOTO 12
С
                ELSE
                    BOTO 12
                END IF
                Y1 + Y
Z1 + Z
  12
             21 = Z

RETURN

WRITE(4,201)]TER,YU,YL,ZU,ZL

WRITE(4,230)Y,Z

FORMAT(5X,'ITER',IOX,'YU',IOX,'YL',IOX,'ZU',IOX,'ZL'/

1 3X,I4,12X,4(3X,F8.5))

FORMAT(5X,'UPPER REGION FAILED - FROM MRF4'/,5X,

* 'CHECK INPUT DATA'/)

FORMAT(5X,'UPPER REGION FAILED - FROM MRF2'/5X,

* 'NO SOLUTION FOR THIS ETA VALUE....'/)

FORMAT(5X,'LOWER BOTTOM REGION FAILED - FROM MRF1 '/5X,

* 'SHOULD NOT HAPPEN'/)
  50
   59
  201
  51
  56
  53
  55
               FORMAT(SX, LOWER BOTTOM REGION FAILED - FROM M

'SHOULD NOT HAPPEN'/)

FORMAT(SX, UPPER REGION FAILED - FROM MRF2 '/,

5X, 'THIS IS (MPOSSIBLE, Z1 .LE. ZL')

FORMAT(5X, LOWER REGION FAILED - FROM MRF1 '/)

DETUDION
  54
  52
                 RETURN
                FORMAT(2(2X,E11.5), 'DIVERGE! -- BAD ETA VALUE')
  238
                END
                                              . . . . . . . . . . . . . . . . . . . .
C ....
               SUBROUTINE CIRCLL(BEDA,CETA,YU,YL,ZU,ZL,Y,Z,I14)

COMMON /CIR/ PASI,PAETA

COMMON /LL/ TOL,TOLETA

COMMON /LL/ TOL,TOLETA

COMMON /MRF/ YMAX,ZMAX,YONE,YMIN

10 = 100 = 0

NECHO = 1

A = A1 = 0.

10 = 10 + 1

WRITE(4,140)1D,YU,YL,ZU,ZL,Y,Z,114

IF(10 GT. 10) THEN

10 = 0

INO = 1NO + 1

IF(114 :EQ. 4) THEN
                                                                                               -----CIRCLL
   100
С
                           NO = [NO + 1

IF(114 :EO. 4) THEN

IF(ABS(YU-YL).LT..1E-01) THEN

IF(ABS(A) .LE. ABS(A1)) THEN

IF(ABS(A) .LT. .1E-01) GGTO 110
                                                           IF(ING LT. 6) GOTO 150
IF(ING LT. 6) GOTO 150
IF(ABS(A) LT. 1) GOTO 110
                                                           GOTO 101
                                                      ELSE
                                                      MAN = 0
Y = (YU + YL)/2.
   15
                                                      Y = (YU + YL)/2.

MAN = MAN + 1

IF(MAN, GT. 10) GOTO 110

A = FNC4(Y,Z,CETA) - PAETA

IF(ABS(A).LT. 1)GGTO110

IF(A) 11,110,13

YU = Y

GOTO 15

VI = Y
    11
                                                       YL + Y
GG TG 15
    13
                                            END IF
                               END I
  С
                           1F(114 EQ. 1) THEN
[F(ABS(ZU-ZL),LT. 1E-01) THEN
[F(ABS(A) LE. ABS(A1)) THEN
[F(ABS(A) LT. 1E-01) GOTO 110
                                                             A1 = A
                                                             IF(1NO .LT. 6) GOTO 150
IF(ABS(A) .LT. .1) GOTO 110
                                                          END IF
                                          END IF
                            END IP
               END (F
END (F
IF(1)4.EQ.1) Z = (ZU + ZL)/2.
IF(1)4.EQ.4) Y = (YU + YL)/2.
IF(1)4.EQ.1) THEN
     150
                               CALL MRF1(Y, Z, TOL, BEDA, PASI, PAETA, MONG, NECHO)
IF(MONG.EQ.1) THEN
                              ZU * Z
GOTO 100
END 1F
                  ELSE
                               CALL MRF4(Y,Z,TOL,BEDÁ,PASI,MONO,NEGHO)
IF(MONO.EQ.1) THEN
YL_= Y_
                               GOTO 100
END IF
    END IF
END IF
114 A = FNC4(Y,Z,CETA) - PAETA
IF(A1 .EO. 0.0) A1 = A
IF(ABS(A).LT.TOLETA) GOTO 110
IF(A) 120,110,130
                 RETURN
YU = Y
ZU = Z
GOTG 100
YL = Y
ZL = Z
     110
     120
     130
                  GOTO 100
```
```
140
          FORMAT(5X, 13, 6(1X, E10.4), 14)
  101
          114 = 100
           RETURN
          END
c - - - -
                                                         -----CIRCLU
          SUBROUTINE CIRCLU(BEDA, CETA, YU, YL, ZU, ZL, Y, Z, 114)
          COMMON /CIR/ PASI, PAETA
COMMON /CIR/ PASI, PAETA
COMMON /LL/ TOL, TOLETA
COMMON /MRF/ YMAX, ZMAX, YONE, YMIN
          LORMON /MRF/ YMAX,Zr

1D = 114 = 1NO = 0

NECHO = 1

A = A1 = 0.

10 = 10 + 1

1F(1D.GT. 10) THEN

1D = 0

1NO = 1NO = 1
  100
               IND = IND + 1
                  IF (ABS (YU-YL) . LE . . 1E-2) THEN
                                    IF(ABS(A) .LE. ABS(A1)) THEN
IF(ABS(A) .LT. .1E-01) GOTO 110
                                       IF(ABS(A) .LT. 6) GOTD 150
IF(ING .LT. 6) GOTD 150
IF(ABS(A) .LT. 1.0) GOTO 110
GOTO 101
                                    ELSE
                                               DIVERGED COMPUTATION
С
                                        LF(INO .GT. 6) GOTO 101
                                    END 1F
                    END IF
 END 1F

END 1F

150 Y = (YU + YL)/2.

CALL MRF4(Y,Z,TOL,BEDA,PASI,MONO,NECHO)

1F(MONO.EQ.1) GOTO 101

A = FNC4(Y,Z,CETA) - PAETA

1F(A1 .EQ. 0.0) A1 = A:

1F(ABS(A).LT.TOLETA) GOTO 110
С
           NEED TO CKECK WHEN Y IS VERY CLOSE TO 1
c
           TESTA = ABS(A - A1)
[F((Y-1.) .LT. .1E-05 .AND. TESTA .LE. .1E-02) GOTO 110
C
            IF(A) 120.110.130
   110
           RETURN
           YL = Y
ZU = Z
  130
            GOTO 100
           YU = Y
ZL = Z
   120
            GOTO 100
C101
           WRITE(4, 140) 10, Y, Z, A, YL, ZL
           FORMAT(5X, 'SUBROUTINE CIRCLU FAIL',3X, 'ID * ',44,3X,

'Y = ',E12.5,'Z * ',E12.5,/12X,'FIE * ',E12.5,3X,

'YLOWER * ',E12.5,3X, 'ZLOWER * ',E12.5)
C140
С
          1
č
          1
          114 = 100
RETURN
  101
           END
C-
                                                              -----SUBROUTINE SHEAR2
          SUBROUTINE SHEAR2(KN,KJ,KF,S1,S9,COEFSHE,SUM1,SUM2)
THIS SUBROUTINE CALCULATES THE ALPHA2 COEFFICIENT IN
THE SHEAR STRESS DISTRIBUTION APPROXIMATION ACCORDING
TO CHIU'S ASSUMPTION. THIS COEFFICIENT IS A CONSTANT
С
C
C
C
C
                    AND THE COMPUTATION IS CARRIED ALONG THE CHANNEL BED & BANKS.
           AND THE COMPOTATION IS CARRIED ALDING THE CP
COMMON /R2/ WP
COMMON /R3/ SID(FF(50,40,2)
COMMON /R4/ F2(50,2)
COMMON /R6/ ETA(50,40,2)
COMMON /R10/ HS((50,40,2), HETA(50,40,2)
COMMON /R10/ HS((50,40,2), HETA(50,40,2)
          c - -
  30
   100
            RETURN
            END
 C - - -
                                                                      -----SUBROUTINE MRF1
           SUBROUTINE MRF1(Y1,Z1,TOL,BETA,PASI,PAETA,ISTOP,NECHO)
THE SUBROUTINE IS THE MODIFIED REGULA FALSI ALGORITHM
FOR FUNCTION 3
00000
                    Z IS FIXED VALUE AND Y IS THE INTERATIVE VALUE.
           DIMENSION A(300), 8(300), P(300)
```

an air a chun

```
14
```

```
COMMON /MRF/ YMAX, ZMAX, YONE, YMIN
             EXTERNAL FNC3
С
                         INITIAL VLAUE TO ZEROS
С
ċ
             00.10 [M = 1,300
             A(1M) = D.
B(1M) = D.
  10
              P(IM) = 0
              ISTOP . 0
С
                           GIVE EITHER 0.01 OR 0.99 FOR + OR - TEST THE A IS ALWAYS SEATS AT THE LEFT OF B
C
C
              \begin{array}{l} \text{In } x = 1 \\ \text{In } x = 1 \\ \text{IF}(\text{PAETA } \cdot \text{GE}, \text{ O}, \text{)} \text{ THEN} \\ \text{B}(\text{IN}) = \text{YONE} \\ \text{A}(\text{IN}) = \text{YMIN} \end{array} 
              ELSE
B([N] T YMAX
A([N] T YONE
              END IF
             END IF

TL = FNC3(A(IN),Z1,BETA) - PASI

TR = FNC3(B(IN),Z1,BETA) - PASI

IF(TL=TR)11,12,13

P(IN) = A(IN)

DO 20 IN = 1,999

P(IN+1) = (A(IN)=TR-B(IN)=TL)/(TR-TL)

CHECK = FNC3(P(IN+1),Z1,BETA) - PASI

IF(ABS(CHECK).GT.TOL)GOTO98

Y1 = P(IN)

GOTO99
  11
              GOTO99
             GOTO99

!F(TL=CHECK)14,99,16

A(IN+1) = A(IN)

B(IN+1) = P(IN+1)

TR = CHECK

CHK = FNC3(P(IN),Z1,BETA) - PASI

!F(CHK=TR.GT.O.)TL = TL/2.
  98
  14
              GOTO20
A(1N+1) = P(1N+1)
  16
              B(IN+1) = B(IN)
              TL = CHECK
CHK = FNC3(P(IN), Z1, BETA) - PASI
IF(TL=CHK.GT.O.)TR = TR/2,
              GOTO20
              IF(TL.EQ.0.)GOTO15
Y1 = B(IN)
  12
              GOTO99
Y1 + A(IN)
  15
              GOTO99
              CONTINUE
  20
             ISTOP = 1
IF(NECHO .EO. 0 ) WRITE(4,77)A(IN),B(IN),TL,TR
FORMAT(/5X,'MRF1 ','YL ',E10.4,'YU ',E10.4,'TL ',E10.4,
TR ',E10.4/)
= ___
   13
   77
   99
             RL
END
c----
                                                          -----SUBROUTINE MRF2
              SUBROUTINE MRF2(Y1, Z1, TOL, CETA, PAETA, ISTOP, NECHO)
THE SUBROUTINE IS THE MODIFIED REGULA FALSE ALGORITHM
FOR FUNCTION 4
С
0000
                         Y IS FIXED VALUE AND Z IS THE INTERATIVE VALUE.
              DIMENSION A(300), 8(300), P(300),
              COMMON /MRF/ YMAX, ZMAX, YONE, YMIN
EXTERNAL FNC4
C
                         INITIAL VLAUE TO ZEROS
c
             DG 10 IM = 1,300
A(IM) = 0.
B(IM) = 0.
P(IM) = 0.
ISTOP = 0
   10
С
C
C
C
C
                           GIVE EITHER 0.0001 OR 0.9999 FOR + OR - TEST
THE A IS ALWAYS SEATS AT THE LEFT OF B
              IN = 1
             [N = 1
B(IN) = ZMAX
A(IN) = 0.00001
TL = FNC4(Y1,A(IN),CETA) - PAETA
TR = FNC4(Y1,B(IN),CETA) - PAETA
IF(TL=TR)11,12,13
P(IN) = A(IN)
DO 20 IN = 1,999
P(IN+1) = (A(IN)=TR-B(IN)=TL)/(TR-TL)
CHECK = FNC4(Y1,P(IN+1),CETA) - PAETA
IF(ARS(CHECK),GT.TGL.0GT096
   11
               IF (ABS(CHECK) . GT. TOL) GOTO98
              21 = P(IN)
              Z1 * P(IN)

GOTO99

IF(TL=CHECK)14,99,16

A(IN+1) = A(IN)

B(IN+1) = P(IN+1)

TR = CHECK
  98
14
              CHK = FNC4(Y1,P(IN),CETA) - PAETA
IF(CHK=TR.GT.O.)TL = TL/2.
              GOTO20
A(IN+1) = P(IN+1)
B(IN+1) = B(IN)
  16
```

```
TL + SHECK
CHK = FNC4(Y1,P(IN),CETA) - PAETA
IF(TL=CHK.GT.O.)TR = TR/2.
            GOTO20
            IF(TL.EQ.0.)G0T015
21 * B(IN)
  12
            GOTO99
            Z1 = A(1N)
G0T099
 15
            CONTINUE
 20
            ISTOP = 1
IF (NECHO
            IF(NECHO'.EQ. 0) WRITE(4,77)A([N),B([N),TL,TR
FORMAT(/5X,'MRF2','ZL',E10.4,'ZU',E10.4,'TL',E10.4,
TR',E10.4/)
  13
  77
            RETURN
 99
            END
                                                -----SUBROUTINE MRF3
C - - - - -
            SUBROUTINE MRF3(Y1, Z1, TOL, CETA, PAETA, ISTOP, NECHO)
THE SUBROUTINE IS THE MODIFIED REGULA FALSI ALGORITHM
FOR FUNCTION 4
С
0000
                      Z IS FIXED VALUE AND Y IS THE INTERATIVE VALUE.
            DIMENSION A(300), B(300), P(300)
Common /MRF/ YMAX, ZMAX, YONE, YMIN
EXTERNAL FNC4
С
                      INITIAL VLAUE TO ZEROS
С
Ċ
            DO 10 (M # 1,300
            A(|M) = 0.
B(|M) = 0.
P(|M) = 0.
  10
            ISTOP = 0
000
                       GIVE EITHER 0.01 OR 0.99 FOR + OR - TEST
THE A IS ALWAYS SEATS AT THE LEFT OF B
            THE A IS ALWAYS SEATS AT TH

IN = 1

IF (PAETA .GE. O.) THEN

B(1N) = YONE

A(1N) = YMIN

ELSE

B(1N) = YMAX

A(1N) = YMAX

A(1N) = YMAX

A(1N) = YMAX

TL = FNC4(A(1N), Z1, CETA) - PAETA

TR = FNC4(B(1N), Z1, CETA) - PAETA

IF (TL=TR)11,12,13

P(1N) = A(1N)

DO 20 IN = 1,999
  11
             PLIN+1) = LALINJ#IR-BLINJ#ILJ/LIR-1LJ
CHECK = FNC4(P(IN+1),Z),CETA) - PAETA
             IF (ABS (CHECK) . GT . TOL) GOTOSS
             Y1 = P((N)
            Y1 = P(1N)

GOTO99

(F(TL=CHECK)14,99,16

A(IN+1) = A(IN)

B(IN+1) = P(IN+1)

TR = CHECK

OUT = FUNC(B(IN) 71 (
  98
14
             CHK = FNC4(P(IN),Z1,CETA) - PAETA
{F(CHK+TR.GT.0.)TL = TL/2.
            IF(CHK*TR.GI.U.JIL - 12.2.
GOTO20
A(IN+1) = P(IN+1)
B(IN+1) = B(IN)
TL = CHECK
CHK = FNC4(P(IN),Z1,CETA) - PAETA
IF(TL*CHK.GT.O.)TR = TR/2.
COTO20
   16
             IF(TL.EQ.0.)GOT015
   12
             YI = B(IN)
             001099
             Y1 = A(IN)
G0T099
   15
             CONTINUE
ISTOP = 1
   20
   13
             ISIOF # :
IF(NECHO .EQ. 0) WRITE(4,77)A([N),B(IN),TL,TR
FORMAT(/5X,'MRF3 ','YL ',E10.4,' YU ',E10.4,' TL ',E10.4,
FORMAT(/5X,'MRF3 ','YL ',E10.4,' YU ',E10.4,' TL ',E10.4,
FRETURN
   77
   99
            END
                                                -----SUBROUTINE MRF4
 c - - - -
             SUBROUTINE MRF4(Y1, Z1, TOL, BETA, PASI, ISTOP, NECHO)
THE SUBROUTINE IS THE MODIFIED REGULA FALSI ALGORITHM
FOR FUNCTION 3
 C
C
 C
C
C
                       Y IS FIXED VALUE AND Z IS THE INTERATIVE VALUE.
             DIMENSION A(300),B(300),P(300)
Common /MRF/ YMAX,ZMAX,YONE,YMIN
EXTERNAL FNC3
 c
                       INITIAL VLAUE TO ZEROS
 c
c
             DO 10 (M = 1,300
A(1M) = 0.
B(1M) = 0.
F(1M) = 0.
ISTOP = 0
   10
 C
C
```

```
16
```

```
Give Either 0.0001 or 0.9399 for + or - Test the a is always seats at the left of {\bf B}
с
с
              IN = 1
             IN = 1

B(IN) = ZMAX

A(IN) = 0.00001

TL = FNC3(Y1,A(IN),BETA) - PASI

IF = FNC3(Y1,B(IN),BETA) - PASI

IF(TL=TR)11,12,13
             |F(IL=IR)|,(2,13
P(IN) = A(IN) =
D0 20 IN = 1,999
P(IN+1) = (A(IN)=TR-B(IN)=TL)/(TR-TL)
CHECK = FNC3(YI,P(IN+1),BETA) - PASI
IF(ABS(CHECK).GT.TOL)GOTO98
  11
              Z1 = P(IN)
              GOTO99
             G01099
IF(TL*CHECK)14,99,16
A(IN+1) = A(IN)
B(IN+1) = P(IN+1)
TR = CHECK
  98
   14
              TH = UNEUN
CHK = FNC3(Y1,P(IN),BETA) - PASI
IF(CHK=TR.GT.O.)TL = TL/2.
              GOTO20
A([N+1) = P([N+1)
   16
              B([N+1) = B([N)
              TL = CHECK
CHK = FNC3(Y1,P(IN),BETA) - PASI
               IF(TL=CHK.GT.O.)TR = TR/2.
              601020
              1F(TL.EQ.0.)GOTO15
Z1 = B(IN)
   12
               GOTO99
              71 = A(1N)
   15
               GOTO99
               CONTINUE
   20
              CONTINUE
ISTOP = 1
IF(NECHO .EQ. 0) WRITE(4,77)A(IN),B(IN),TL,TR
FORMAT(/5X,'MRF4 ','ZL ',E10.4,' ZU ',E10.4,' TL ',E10.4,
FORMAT(/5X,'MRF4 ','ZL ',E10.4,' ZU ',E10.4,' TL ',E10.4,
FORMAT(/5X,'MRF4 ','ZL ',E10.4,' ZU ',E10.4,' TL ',E10.4,
FORMAT(/5X,'MRF4 ','ZL ',E10.4,' ZU ',E10.4,' TL ',E10.4,'
METURN'
    13
   77
   99
               END
                ____
                                                                                                -----SUBROUTINE SHEAR1
c----
              SUBROUTINE SHEAR1(K1,K2,FNINE,ALPHA2,ALPHA)
COMMON /R3/ SIDIFF(50,40,2)
COMMON /R4/ F2(50,2)
COMMON /R10/ HSI(50,40,2),HETA(50,40,2)
                1 = 1
               I = I
XBON = SIDIFF(K2,1,K1)/F2(K2,K1)
T1 = -HETA(K2,1,K1)=FNINE
T2 = -ALPHA2=(1.+XBON)=SIDIFF(K2,1,K1)
ALPHA = (T1 + T2)/XBON
               RETURN
               END
                  SUBROUTINE INTPHI
  c----
              SUBROUTINE INTPHI(A1, A2, B1, B2, C, D)
---GAUSSINA INTEGRATION FOR PHI
DIMENSION SHAPE(2), W(3), PT(3)
DATA PT/-0.77459667, 0.000000,0.77459667/
DATA W/0.5555555,0.88888888,0.55555555
  С
              DATA W/0.5555555,0.86868686,0.5555555/

C = 0.

D1 = D/2.

D0 100 ! = 1,3

SHAPE(1) = (1. - PT(1))/2.

SHAPE(2) = (1. + PT(1))/2.

C=C+W(1)=(SHAPE(1)=A1+SHAPE(2)=A2)=(SHAPE(1)=B1+SHAPE(2)=B2)=D1

CONTINUE

DETUDN
    100
                RETURN
                END
  c ---
                       -----DI3000 ROUTINE
             SUBROUTINE DI30SV(NA, MCBPTS, DEPTH, EPSILON, Y, Z, JBEG,
NB, BI, ZBOUND, YBOUND, FNC1, DELTA, DELY, SHEAR, VY, VZ,
PPLOT2, PPLOT3, PPLOT4, PPLOT5, AVGSHEA, NDIVIDE,
AVGVEL, SCALE, STATION)
INTEGER PPLOT2, PPLOT3, PPLOT4, PPLOT5
DIMENSION Y(50, 40, 2), Z(50, 40, 2), DELTA(2), NDIVIDE(2)
DIMENSION JBEG(2, 40), NB(2), BI(2), ZBGUND(20), YBGUND(20)
DIMENSION SHEAR(30, 40, 2), VY(50, 40, 2), VZ(50, 40, 2), XYL(4)
  с
с
с
                INITIAL DIGOOD PLOT ROUTINE
                CHARACTER=28 STATION
CHARACTER=21 NOTATU(3)
  с
                DATA NOTATU/ VERTICAL COMPONENTS ', 'TRANSVERSE COMPONENTS'
               SECONDARY VELOCITIES '/
DATA PIE/3 14159265/
  С
                CALL UUINI(1)
  c
                XYL(1) = 0.0
                XYL(1) = 0.0
XYL(2) = 100.
XYL(3) = 0.0
XYL(4) = 5.00
VMAX = 0.0
CALL JWINDO(XYL(1),XYL(2),XYL(3),XYL(4))
CALL JWINDO(XYL(1),XYL(2),XYL(3),XYL(4))
                 CALL JVPORT(-1.0,1.0,-0.7,0.7)
  С
                 RATIO = XYL(2)/XYL(4)
  с с с
с
                     AXES ATTRIBUTES
```

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17
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CALL UUSTI('ALL.EXIST',1) CALL UUSTI('TOP.LABEL.TYPE',0) CALL UUSTI('TOP.TITLE.FONT',5) CALL UUSTI('TOP.TITLE.HORIZONTAL',1) CALL UUSTI('ALL.TITLE.QUALITY',3) CALL UUSTR('TOP MAJOR.INWARD',0.0) CALL UUSTR('TOP TITLE.POSITION',0.1) CALL UUSTR('TOP.TITLE.POSITION',0.1) CALL UUSTR('TOP.TITLE.STRING',STATION) CALL UUSTR('LEFT.MAJOR.INWARD',0.01) CALL UUSTR('LEFT.SCALE.INTERVAL',1.0) DRAW LEFT AXIS FOR DEPTH CALL UUSTR('LEFT.SCALE.INTERVAL',1.0) DRAW LEFT AXIS FOR DEPTH CALL UUSTR('LEFT.SCALE.INTERVAL',1.0) CALL UUSTR('LEFT.SCALE.HSIZE',0.015) CALL UUSTR('LEFT.LABEL.HSIZE',0.015) CALL UUSTR('LEFT.TITLE.FONT',1)) CALL UUSTR('LEFT.TITLE.FONT',1) CALL UUSTR('LEFT.LABEL.GONT',9) DRAW BOTTOM AXIS FOR DEPS SECTION VIDTH с 000 DRAW BOTTOM AXIS FOR CROSS SECTION WIDTH CALL UUSTR('BOTTOM.SCALE.INTERVAL',10.) CALL UUSTR('BOTTOM.TITLE.STRING',' DISTANCE FROM LEFT BANK OF *RIVER, FT.') CALL UUSTR('BOTTOM.LABEL.MSIZE',0.015) CALL UUSTR('BOTTOM.LABEL.VSIZE',0.020) CALL UUSTR('BOTTOM.TITLE.VSIZE',0.020) CALL UUSTR('BOTTOM.TITLE.VSIZE',0.020) CALL UUSTR('BOTTOM.MAJOR.UTVARD',0.02) CALL UUSTR('BOTTOM.MAJOR.INWARD',0.02) CALL UUSTR('BOTTOM.TITLE.DISTANCE',0.09) CALL UUSTR('BOTTOM.TITLE.FONT',11) CALL UUSTI('BOTTOM.TITLE.FONT',11) CALL UUSTI('BOTTOM.TITLE.FONT',1) CALL UUSTI('BOTTOM.TITLE.FONT',3) c c TRANSFER COORDINATES ċ DEFINE UNIT IN PLOTS UNIT = (XYL(2) - XYL(1)) / 15, RATIO = (XYL(4)-XYL(3))/(XYL(2)-XYL(1)) ċ c NA1 = NA - 1 C ADJUST COORDINATES DO 100 K = 1,2 FNC2 = BI(K) + DELTA(K) DO 110 [= 1.NA DO 110 [= 1,NA NC = JBEG(K,I) DO 120 J = NC,NB(K) Y(J,I,K) = - DEPTH+(Y(J,I,K) = FNC1 - DELY) + 4.0 IF(K.EG.1) Z(J,I,K) = BI(1) - Z(J,I,K)=FNC2 F(K.EG.2) Z(J,I,K) = Z(J,I,K)=FNC2 + BI(1) IF(I .GE. 2 .AND. I .LE. NA1) THEN IF(K.EG.1) VZ(J,I,K) = -VZ(J,I,K) IF(VMAX .LT. ABS(VZ(J,I,K))) VMAX = ABS(VZ(J,I,K)) IF(VMAX .LT. ABS(VY(J,I,K))) VMAX = ABS(VY(J,I,K)) END IF 120 110 CONTINUE 100 CONTINUE с DO 111 1 = 1,MCBPTS YBOUND(1) = -DEPTH+YBOUND(1) + 4.0 111 CONTINUE RAT = UNIT / VMAX с с DRAW SHEAR STRESSES ALONG BOUNDARY ċ IF (PPLOT2 .EQ. 1) THEN CALL UUAXIS(XYL(1), XYL(2), XYL(3), XYL(4)) С C C C C DRAW BOUNDARY CALL JOPEN CALL JDFEN CALL JLWIDE(16363) CALL JMOVE(286UND(1),YB6UND(1)) CALL JPOLY(286UND,Y86UND,MC8PTS) CALL JDRAW(286UND(1),Y86UND(1)) CALL JCLOSE С D0 150 K = 1,2 D0 160 J = 1,NB(K) SHEAR(J,1,K) = SHEAR(J,1,K) / AVGSHEA D0 160 I = MK,NA CONTINUE 170 180 160 CONTINUE 150 CONTINUE С С С С LABLES WRITE(6,77) CALL JPAUSE(1) END IF Ċ

18

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```
č
             DRAW VELOCITIES COMPONENTS
č
             1F(PPL0T3.EQ.1..OR. PPL0T4.EQ.1..OR. PPL0T5.EQ.1) THEN
ANG = 30./180.=3.14159265
1F(PPL0T3..EO. 1) THEN
С
                             1FRAME
                    CALL
               CALL UUAXIS(XYL(1), XYL(2), XYL(3), XYL(4))
С
              DRAW BOUNDARY
c
             CALL JOPEN
                 ALL JOPEN
CALL JLWIDE(16383)
CALL JMOVE(ZBOUND(1),YBOUND(1))
CALL JPOLY(ZBOUND,YBOUND,MCBPTS)
CALL JDRAW(ZBOUND(1),YBOUND(1))
              CALL JCLOSE
С
                    THETA = 3.14159265 / 2.
С
                  CALL JOPEN
              CALL JLWIDE (8000)
С
                    D0 200 K = 1,2
D0 210 I = 2,NA1,NDIVIDE(1)
NC = JBEG(K,I)
                    RU = JBEG(K,1)
D0 220 J = NC,NB(K),NDIVIDE(2)
CALL JMOVE(Z(J,1,K),Y(J,1,K))
PVY = VY(J,1,K) = RAT = RATIO + Y(J,1,K)
PVZ = Z(J,1,K)
                    CALL JDRAW(PVZ, PVY)
С
                    RT = SQRT(VY(J,:,K)=VY(J,:,K)+0.0=0.0)/8.0
С
                    DX = RT = COS(THETA) = RAT
                   DX = RT = COS(THETA) = RAT

DY = RT = SIN(THETA) = RAT = RATIO

IF(VY(J,I,K) LT. 0.0) DY = - DY

TY = PVY - DY

TZ = PVZ - DX

DX = RT = SIN(ANG) = COS((PIE/2.-ANG)+THETA) = RAT

DY = RT = SIN(ANG) = SIN((PIE/2.-ANG)+THETA) = RAT=RATIO

DX1 = RT = SIN(ANG) = COS(THETA-(PIE/2.-ANG)) = RAT

DY1 = RT = SIN(ANG) = SIN(THETA-(PIE/2.-ANG)) = RAT=RATIO

VA = TY + DY
                        YA = TY + DY
ZA = TZ + DX
YB = TY + DY1
ZB = TZ + DX1
C
              CALL JHOVE (PVZ, PVY)
              CALL JURAWIZA, YA)
CALL JMOVE(PVZ, PVY)
CALL JDRAW(ZB, YB)
C
220
210
                    CONTINUE
200
C
                    CONTINUE
              CALL JCLOSE
C
C
C
                WRITE IDENTIFICATION
                XPOST = XYL(1) + UNIT
YPOST = XYL(4) - 0.5
CALL JOPEN
                      CALL JUSTEN

CALL JJUST(1,2)

CALL JFONT(3)

CALL JMOVE(XPOST,YPOST)

CALL JSIZE(3.3,3.3=RATIO)

CALL JSIZE(3.3,3.3=RATIO)
                      CALL JCLOSE
C
C
C
             DRAW AN UNIT LENGTH TO EXPRESS SCALE
             XPOST = XYL(1) + 2. = UNIT
YPOST = XYL(3) + .60
                POST = XYL(3) + .60

CALL JOPEN

CALL JMOVE(XPOST,YPOST)

CALL JJUST(1,2)

CALL JSIZE(2.8,3.0=RAT10)

CALL JHSTRG('SCALE')

CALL JMOVE(XPOST,YPOST-.2)

CALL JMOVE(XPOST+1.0=RAT,YPOST-.2)

CALL JMOVE(XPOST+1.2=RAT,YPOST-.2)

CALL JHSTRG('=1FPS')

CALL JCLOSE
¢
             CALL JPAUSE(1) - END (F
Ċ
             IF (PPLOTA .EQ. 1) THEN
C
                    CALL JERAME
               CALL UUAXIS(XYL(1), XYL(2), XYL(3), XYL(4))
000
             DRAW BOUNDARY
             CALL JOPEN
                  CALL JLWIDE(16383)
```

```
CALL JMOVE(2BOUND(1),YBCUND(1))
CALL JPOLY(ZBOUND,YBOUND,MCBPTS)
CALL JDRAW(ZBOUND(1),YBOUND(1))
               CALL JCLOSE
 С
                       THETA = 0.0
 с
                    CALL JOPEN
               CALL JLWIDE (8000)
 с
                     D0 300 K = 1,2

D0 310 I = 2,NA1,NDIVIDE(1)

NC = JBEG(K,1)

D0 320 J = NC,NB(K),NDIVIDE(2)

CALL JMOVE(Z(J,1,K),Y(J,1,K))

PVY = Y(J,1,K)

PVZ = Z(J,1,K) + VZ(J,1,K) = RAT

CALL IORAL(PVZ PVX)
                      CALL JDRAW(PVZ, PVY)
 С
                       RT = SQRT(VZ(J, [,K)=VZ(J, ],K)+0.0=0.0)/8.0
 С
                      DX = RT = COS(THETA) = RAT

IF(VZ(J,1,K) .LT. 0.0) OX = -DX

DY = RT = SIN(THETA) = RAT = RATIO

TY = PVY - DY

TZ = PVZ - DX
                          Z = PVZ - DX

DX = RT = SIN(ANG) = COS((PIE/2.-ANG)+THETA) = RAT

DY = RT = SIN(ANG) = SIN((PIE/2.-ANG)+THETA) = RAT=RATIO

<math>DX1 = RT = SIN(ANG) = COS(THETA-(PIE/2.-ANG)) = RAT=RATIO

YA = TY = DY

ZA = TZ + DY

YB = TY + DY1

ZB = TZ + DX1
 С
               CALL JMOVE (PVZ, PVY)
               CALL JDRAW(ZA,YA)
CALL JMOVE(PVZ,PVY)
CALL JDRAW(2B,YB)
 С
   320
                       CONTINUE
   310
                       CONTINUE
  300
                       CONTINUE
 c
               CALL JCLOSE
 С
                 WRITE IDENTIFICATION
 C
 ē
                 XPOST = XYL(1) + UNIT
                  YPOST = XYL(4) - 0.5
                        CALL JOPEN
                        CALL JOPEN
CALL JJUST(1,2)
CALL JFONT(3)
CALL JMOVE(XPOST,YPOST)
CALL JSIZE(3.3,3.3¤RATIO)
CALL JSIZE(3.3,3.9¤RATIO)
                         CALL JCLOSE
С
ċ
              DRAW AN UNIT LENGTH TO EXPRESS SCALE
ř
              XPOST = XYL(1) + 2. = UNIT
YPOST = XYL(3) + .60
CALL JOPEN
                  CALL JOPEN

CALL JOVE(XPOST,YPOST)

CALL JJUST(1,2)

CALL JFONT(3)

CALL JSTRG(SCALE')

CALL JHSTRG(SCALE')

CALL JMOVE(XPOST,YPOST-2)

CALL JMOVE(XPOST+1.0=RAT,YPOST-2)

CALL JMOVE(XPOST+1.2=RAT,YPOST-2)

CALL JHSTRG('=1FPS')

CALL JCLOSE
С
              CALL JPAUSE(1)
END IF
С
              IF (PPLOTS .EQ. 1) THEN
C
                      CALL JERAME
                CALL UUAXIS(XYL(1),XYL(2),XYL(3),XYL(4))
C
C
C
C
              DRAW BOUNDARY
              CALL JOPEN
                  CALL JUHIDE(16383)
CALL JMOVE(2BOUND(1),YBOUND(1))
CALL JPOLY(2BOUND,YBOUND,MCBPTS)
CALL JDRAW(2BOUND(1),YBOUND(1))
              CALL JCLOSE
С
              CALL JOPEN
CALL JLWIDE(8000)
С
                     D0 400 K = 1,2

D0 410 I = 2,NA1,NDIVIDE(1)

NC = JBEG(K,I)

D0 420 J = NC,NB(K),NDIVIDE(2)

CALL JMOVE(2(J,I,K),Y(J,I,K))

PVY = VY(J,I,K) = RAT = RATIO + Y(J,I,K)
```

```
PVZ = Z(J,T,K) + VZ(J,T,K) = RAT
                     CALL JDRAW(PVZ, PVY)
С
                   RT = SQRT(VY(J,I,K)=VY(J,I,K)+VZ(J,I,K)=VZ(J,I,K))/8.0
THETA = ATAN(VY(J,I,K)/VZ(J,I,K))
                     DX = RT = COS(THETA) = RAT

DY = RT = SIN(THETA) = RAT

DY = RT = SIN(THETA) = RAT = RATIO

(F(VY(J,1,K) .LT. 0.0) DY = - DY

(F(VZ(J,1,K) .LT. 0.0) DX = -DX

TY = PVY - DY

TZ = PVZ - DX

DX = RT = SIN(ANG) = COS((PIE/2.-ANG)+THETA) = RAT

DY = RT = SIN(ANG) = SIN((PIE/2.-ANG)+THETA) = RAT

DX1 = RT = SIN(ANG) = COS(THETA-(PIE/2.-ANG)) = RAT

DY1 = RT = SIN(ANG) = SIN(THETA-(PIE/2.-ANG)) = RAT

DY1 = RT = SIN(ANG) = SIN(THETA-(PIE/2.-ANG)) = RAT

DY1 = RT = SIN(ANG) = SIN(THETA-(PIE/2.-ANG)) = RAT

TZ = TZ + DY

ZA = TZ + DX

YB = TZ + DX1
С
с
              CALL JMOVE(PVZ, PVY)
CALL JDRAW(ZA, YA)
CALL JMOVE(PVZ, PVY)
                    CALL JDRAW(ZB, YB)
С
  420
                      CONTINUE
  410
                      CONTINUE
                      CONTINUE
С
               CALL JOLOSE
C
C
C
C
                  WRITE IDENTIFICATION
                  XPOST = XYL(1) + UNIT
                  YPOST = XYL(4) - 0.5
CALL JOPEN
                        CALL JOPEN

CALL JJUST(1,2)

CALL JFONT(3)

CALL JMOVE(XPOST,YPOST)

CALL JSIZE(3,3,3:3=RATIO)

CALL JSIZE(3,3,3:3=RATIO)

CALL JSIZE(3,3,3:3=RATIO)
                         CALL JCLOSE
000
                DRAW AN UNIT LENGTH TO EXPRESS SCALE
               XPOST = XYL(1) + 2. = UNIT
YPOST = XYL(3) + .60
CALL JOPEN
                        CALL JMOVE (XPOST, YPOST)
                        CALL JHOVE(XPOST, FPOST)

CALL JFONT(3)

CALL JFSTRG('SCALE')

CALL JHSTRG('SCALE')

CALL JHSTRG('SCALE')

CALL JMOVE(XPOST, YPOST-.2)

CALL JMOVE(XPOST+1.2=RAT, YPOST-.2)

CALL JHSTRG('=1FPS')

ALL (OSE
                    CALL JCLOSE
С
               CALL JPAUSE(1)
¢
               END IF
С
               CALL JEND
c
   77
               FORMAT(12X, 'CHANGE PAPER ON PLOTTER. ENTER STOP TO STOP RUNS')
               STOP
               END
C .....
                                                        -----DI3000 ROUTINE
               SUBROUTINE DIGOISO (NA, MCBPTS, OEPTH, EPSILON, Y, Z, JBEG,

        NB.BI.ZBOUND.YBOUND.FNC1.DELTA.DELY)

        DIMENSION ZP(3D).YP(3D)

        DIMENSION Y(50,40,2),Z(50,40,2),DELTA.DELY)

        DIMENSION Y(50,40,2),Z(50,40,2),DELTA(2)

        DIMENSION JBEG(2,40),NB(2),BI(2),ZBOUND(20),YBOUND(20)

с
с
с
                INITIAL DISDOG PLOT ROUTINE
               CALL UUINI(1)
YMIN = XMIN = 0.0
YMAX = DEPTH = 1.2
XMAX = (B1(1) + B1(2)) = 1.1
С
               CALL JWINDO(XHIN, XMAX, YMIN, YMAX)
CALL JVPGRT(-1.0,1.0, -1.0,1.0)
CALL UUSTI('ALL EXIST',0)
CALL UUAXIS(XMIN, XMAX, YMIN, YMAX)
С
с
с
                DRAW BOUNDARY
                CALL UUCURV(280UND, YBOUND, MCBPTS)
0000
                DRAW XI CURVES
               00 100 K = 1,2
FNC2 = B1(K) + DELTA(K)
D0 110 I = 1,NA
NC = JBEG(K,1)
```

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```

```
ND = 1

DO 120 J = NC,NB(K)

YP(ND) = Y(J,1,K) = FNC1 - DELY.

IF(K.EQ.1) ZP(ND) = BI(1) - Z(J,1,K) = FNC2

IF(K.EQ.2) ZP(ND) = Z(J,1,K) = FNC2 + BI(1)

ND = ND + 1

120 CONTINUE
¢
  ND = ND - 1
CALL UUCURV(ZP,YP,ND)
110 CONTINUE
100 CONTINUE
c
                    CALL JPAUSE(1)
000
                    DRAW ETA CURVES
 C DRAW ETA CONVES

C

DO 150 K = 1,2

FNC2 = B1(K) + DELTA(K)

MK = NA

MP = 1

DO 160 J = 2,N8(K)

IF(J_LE. NA) THEN

IF(EPSILON .GT. 0.0) MK = J

IF(EPSILON .LT. 0.0) MP = NA - J + 1

END IF

ND = 1

170 DO 180 L = MP,MK

YP(ND) = Y(J,I,K) = FNC1 - DELY

IF(K.E0.2) ZP(ND) = B1(1) - Z(J,I,K) = FNC2

IF(K.E0.2) ZP(ND) = Z(J,I,K) = FNC2 + BI(11)

ND = ND + 1

180 CONTINUE

C
 С
                    ND = ND - 1
CALL UUCURV(ZP,YP,ND)
CONTINUE
   160
150
                     CONTINUE
000
                     LABELS
                    CALL JPAUSE(1)
CALL JEND
с
                     STOP
                     END
```

Appendix B

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Measured Velocities





Figure B-1.

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Distribution of measured longitudinal, transverse and vertical components of velocity at cross section 7, June 27, 1984



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Figure B-2. Isovels of longitudinal velocity at cross section 7, June 27, 1984



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Figure B-3. Measured secondary current vectors at cross section 7, June 27, 1984



Figure B-4. Distribution of measured longitudinal, transverse and vertical components of velocity at cross section 6, June 28, 1984



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Figure B-5. Isovels of longitudinal velocity at cross section 6, June 28, 1984



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Figure B-6. Measured secondary current vectors at cross section 6, June 28, 1984









B-7. Distribution of measured longitudinal, transverse and vertical components of velocity at cross section 5, June 29, 1984



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Figure B-8. Isovels of longitudinal velocity at cross section 5, June 29, 1984



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Figure B-9. Measured secondary current vectors at cross section 5, June 19, 1984



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Figure B-11. Isovels of longitudinal velocity at cross section 3, July 9, 1984



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Figure B-12. Measured secondary current vectors at cross section 3, July 9, 1984







Figure B-13. Distribution of measured longitudinal, transverse and vertical components of velocity at cross section 4, July 10, 1984



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DISTANCE, ft

Figure B-14. Isovels of longitudinal velocity at cross section 4, July 10, 1984



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Figure B-15. Measured secondary current vectors at cross section 4, July 10, 1984





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Distribution of measured longitudinal, transverse and vertical components of velocity at cross section 2, July 12, 1984



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Figure B-17. Isovels of longitudinal velocity at cross section 2, July 12, 1984



Figure B-18. Measured secondary current vectors at cross section 2, July 12, 1984







Figure B-20. Isovels of longitudinal velocity at cross section 1, July 13, 1984

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Figure B-21. Measured secondary current vectors at cross section 1, July 13, 1984





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Figure B-22. Distribution of measured longitudinal, transverse and vertical components of velocity at cross section 5, August 5, 1985



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Figure B-23. Isovels of longitudinal velocity at cross section 5, August 5, 1985



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Figure B-24. Measured secondary current vectors at cross section 5, August 5, 1985



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Figure B-25. Distribution of measured longitudinal, transverse and vertical components of velocity at cross section 6, August 6, 1985

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DISTANCE, ft

Figure B-26. Isovels of longitudinal velocity at cross section 6, August 6, 1985


Figure B-27. Measured secondary current vectors at cross section 6, August 6, 1985



Figure B-28. Distribution of measured longitudinal, transverse and vertical components of velocity at cross section 7, August 7,

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DISTANCE, ft

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Figure B-29. Isovels of longitudinal velocity at cross section 7, August 7, 1985



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Figure B-30. Measured secondary current vectors at cross section 7, August 7, 1985





Figure B-31. Distribution of measured longitudinal, transverse and vertical components of velocity at cross section 8, August 8, 1985



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Figure B-32. Isovels of longitudinal velocity at cross section 8, August 8, 1985



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Figure B-33. Measured secondary current vectors at cross section 8, August 8, 1985









Figure B-34. Distribution of measured longitudinal, transverse and vertical components of velocity at cross section 9, August 9, 1985



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Figure B-35. Isovels of longitudinal velocity at cross section 9, August 9, 1985



Figure B-36. Measured secondary current vectors at cross section 9, August 9, 1985

TABLE B-1 SANGAMON RIVER CROSS SECTION 7 JUNE 27 1984

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MEASURED DATA

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U (FPS)	. 71	. 68	. 77	. 71	. 70	. 74	98.	. 97	96 .	. 85	. 96	, 98	. 97	. 85	. 90	. 82	. 78	. 62	. 77	. 76	. 67		
(FPS)	.21	. 19	.10	. 13	.12	. 17	. 22	. 27	. 13	. 19	. 22	. 24	. 20	. 25	. 24	. 24	.21	. 15	. 27	. 23	.21		
DEPTH (FT)	. 25	. 75	. 25	. 75	1.25	1.75	50	1,00	1.50	2.00	. 25	. 75	1.25	1.75	.58	1.08	1.58	2.08	. 58	1.08	1.58		
DISTANCE (FT)	8.25	8.25	10.95	10.95	10.95	10,95	25, 75	25, 75	25.75	25.75	35.75	35.75	35.75	35,75	50.75	50,75	50.75	50, 75	65, 75	65.75	65, 75		
U (FPS)	61	. 72	68	. 64	. 70		54	06	50	94	. 96	06	. 97	.92	. 81	. 83	. 82	. 79	74	76	75	. 74	. 70
V U (FPS) (FPS)	. 05	03 72	. 13	. 01 . 64	06 .70	01 59	07 . 54	. 21 . 90	. 04	- 01 . 94	.08 .96	06 60.	.12	, 00 , 92	. 61	. 02 . 83	. 02 . 82	- 04 . 79	- 06 . 74	- 19 76	- 02 . 75	- 03 74	. 09
DEPTH V U (FT) (FPS) (FPS)	.50 .05 .61	1.00 .03 .72	1.50 .13 .68	2.00 .01 .64	1.0006 .70	1.50 01 59	2.0007 .54	. 25 . 21 . 90	. 75 04	1,2501 .94	1, 75 , 08 , 96	2.25 .09 .90	.50 .12 .97	1,00 ,00 ,92	2.00 .06 .81	. 33 . 02 . 83	. 83 . 02 . 82	1.3304 .79	1.83 - 06 .74	33 - 19 76	. 83 02 . 75	1.3303 .74	1.83 .09 .70

TABLE B-2 SANGAMON RIVER CROSS SECTION 6 JUNE 28 1984

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MEASURED DATA

5	(FPS)		. 36	. 24	.40	.37	.27	89	. 85	. 82	. 71	1.07	66 '	90	. 79	. 77	. 78	. 70	. 59	. 81	. 78	.67	۲۲.	. 65	54	. 57	. 59	. 57	. 57	48	. 42
3	(FPS)		<u>.</u> 05	.03	.03	. 05	. 02	.03	.04	.01	.04	00.	.01	. 07	.03	.04	.04	. 06	.08	. 05	. 05	.08	.06	. 05	.08	. 06	04	.05	.05	.07	. 07
DEPTH	(FT)		. 20	. 70	1.20	1.70	2.20	. 50	1.00	1.50	2,00	35	. 85	1.35	. 35	. 85	1.35	1.85	2.35	. 20	. 70	1.20	1.70	2.20	2.70	. 35	. 85	1.35	1.85	2.35	2.85
DISTANCE	(FT)		8.75	8.75	8,75	8.75	8.75	18.75	18.75	18.75	18.75	33.75	33.75	33.75	48.75	48.75	48.75	48.75	48.75	58.75	58.75	58, 75	58,75	58.75	58.75	66.25	66.25	66,25	66.25	66.25	66.25
5	(FPS)		. 53	. 25	.39	42	28	. 78	. 83	. 79	. 76	۲۲.	06.	1.00	. 96	.97	. 71	. 79	71	.68	.57	. 85	. 74	. 70	. 76	. 63	.51	. 56	. 50	.46	. 49
>	(FPS)	•	- , 08	08	- 06	07	80.1	60	- 12	12	10	16	51	. 05	16	- , 06	-, 19	- 12	- 10	- , 16	16	24	- , 23	- , 22	- , 18	- , 17	17	14	15	- , 18	12
DEPTH	(FT)		. 45	. 95	1.45	1.95	2.45	. 25	. 75	1,25	1.75	2.25	. 60	1.10	1.60	2.10	. 60	1.10	1.60	2.10	2.60	. 45	. 95	1.45	1,95	2.45	. 60	1.10	1.60	2.10	2.60
DISTANCE	(FT)		8,00	8.00	8.00	8.00	8,00	18.00	18.00	18.00	18.00	18.00	33.00	33, 50	33,50	33.50	48.00	48.00	48.00	48.00	48.00	58.00	58.00	58.00	58,00	58.00	65.50	65.50	65, 50	65.50	65.50

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TABLE B-3 SANGAMON RIVER CROSS SECTION 5 JUNE 29 1984

Sec. 2

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MEASURED DATA

U (FPS)	. 43	. 45	.41	34	. 61	. 62	. 56	.51	. 84	. 76	. 73	. 64	. 65	. 77	02.	. 66	60	. 68	69	. 61	. 63	. 60	. 60		40. 04	5 C		64.	4	. 45	. 35	. 37	. 30
K (FPS)	60 '	60 '	.07	. 07	. 13	Ξ.	=	60	. 18	12	.11	=		61.	.12	10	- 11	- 16 - 16	4		. 15	.07	. 15		- 5	2	-	. 12		.12	. 06	05	.07
DEPTH (FT)	.60	1.10	1.60	2.10	. 70	1.20	1.70	2.20	. 25	. 75	1.25	1.75	2.25	.45	2 6 .	1.45	1.95	1 1 1 1	1.00	1.50	2.00	2.50	20	00.1	00.1		00.2	. 58	1.08	1.58	2.08	2, 58	3.08
DISTANCE (FT)	8.55	8,55	8.55	8.55	18.55	18.55	18.55	18.55	28.55	28.55	28.55	28.55	28.55	38.55	38.55	38.55	38,55		48.55	48.55	48.55	48.55	58.55	58.55	58. 55 55	58.35	58, 55	66.55	66.55	66.55	66.55	66.55	66, 55
	,																																
U (FPS)	2	. 48	.47	. 62	.58	. 59	. 55	.51	. 86	. 78	. 75	. 77	. 65	.67	. 68	. 65	.60			74	. 62	. 64	.61	20	. 53	53	. 52	.51	44.	. 43	.39	. 44	40 20
(FPS) (FPS)	- 07 32		02 .47	- 04 .62	07 .58	- 06 .59	03 .55	01 .51	-,03	02 .78	03 .75	.02 .77	- 02 , 65	10 .67	11 .68	-,14 ,65	12 .60	- 12 - 62	- 06 - 66	02 74	14 .62	10 .64	13 .61	- 18	17 .53	- 12 - 53	- 11 , 52	11 .51	10.44	12 .43	07 .39	08 .44	- 04 . 40
DEPTH V U (FT) (FPS) (FPS)	ак - 07 32	- 20 · 20 · 48	1 35 . 02 . 47	. 45 04 62	. 95 07 . 58	1,4506 .59	1.9503 .55	2.4501 .51	50 - 03 . 86	1.0002 .78	1.5003 .75	2.00 02 77	2.5002 .65	. 20 10 , 67	.7011 .68	1.2014 .65	1.7012 .60	2.2012 .62		75 .02 .74	1.25 - 14 .62	1.7510 .64	2.2513 .61	2.75 - 18 .50	. 75 17 . 53	1.25 - 12 - 53	1, 75 -, 11 , 52	2.2511 .51	2.7510 .44	.3312 .43	.8307 .39	1.3308 .44	1.8304 .40 2.83 - 01 .32

TABLE B-4 SANGAMON RIVER CROSS SECTION 3 JULY 9 1984

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U (FPS)	 .37	.40	. 53	. 37	.36	. 49	.51	00	10	/01	0.	00.	20	43	45	53	. 51	- 20	. 50	. 46	. 42	. 50	. 52	. 50	. 46	. 47	43	44	.40	. 46	. 45	. 45	. 45	.39	.43	.41	. 42	.40	.40	. 28
W (FPS)	. 15	. 15	.27	.17	. 16	. 13	. 15	. 15	. 16	. 15	. 16	- 17	. 17	12	. 16	14	. 13	14	. 15	. 15	.14	. 16	. 18	. 18	. 16	. 17	10	. 18	. 16	. 18	. 18	. 18	. 17	. 17	. 17	. 17	. 17	. 18	. 15	.13
DEPTH (FT)	 . 25	. 50	. 75	1.00	1.25	<u>.</u> 05	. 25	. 50	. 75	.08	. 58	. 83	1,08	1.28	1.53	. 25	. 50	. 75	1.00	1.25	1.50	. 15	.40	. 65	. 90	. 25	. 50	. 75	1.00	. 25	. 50	. 75	1.00	1.25	. 33	. 58	. 83	1.08	1.33	1.58
DISTANCE (FT)	 6.25	6,25	6,25	6,25	6.25	16.25	16.25	16.25	16.25	26.25	26.25	26,25	26.25	26.25	26.25	36,25	36,25	36.25	36.25	36.25	36.25	46.25	46.25	46.25	46.25	56.25	56.25	56.25	56.25	66, 25	66.25	66.25	66.25	66.25	72,75	72.75	72.75	72.75	72.75	72.75
U FPS)	41	35	35	34	46	44	45	51	47	47	47	52	52	47	50	46	47	52	47	46	48	45	44	35	43	42	43	40	35	40	39	37	39	35						
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DEPTH (FT)	10		1.00	22	40	65	06	33	58	. 83	1,33	1.58	. 25	50	. 75	1.00	1.25	40	.65	06	.25	50	. 75	. 25	. 50	. 75	1.00	1.25	1.50	. 33	. 58	. 83	1.08	1.33						
DISTANCE					15.50	15.50	15.50	25,50	25.50	25, 50	25.50	25, 50	35,50	35,50	35.50	35.50	35,50	45.50	45.50	45.50	55.50	55.50	55.50	65, 50	65,50	65.50	65,50	65.50	65, 50	72.00	72.00	72.00	72.00	72.00						

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TABLE B-5 SANGAMON RIVER CROSS SECTION 4 JULY 10 1984

MEASURED DATA

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U (FPS)	50	. 26	2 4	, , , ,	10	38	36	200	166	36	20	22	26	44	643	57	10		200			05.	4	. 40	4	. 39	. 38	40	38	35	87.	. 4 .	. 42	66.	. 40	.4	.40	. 37	. 35	53	. 35	. 39	. 37	. 33	38	35	34	
K (FPS)	. 15	4	4	 4 u		21		2 -	- 10	- u	2	2 0			<u>-</u>	2 -						4	4	4 1	61	16	2	16	16	16	.17	20	- 17	. 17	. 16	.18	. 16	2	.17	15	. 17	. 19	. 17	. 17	16	9	9	>
DEPTH (FT)	. 25	20	c/ .					1 (1													1. /3	2.00	07.	20	. 75	1.00	1.25	1.50	1.75	2.00	2.25	22	. 50	. 75	1.00	1.25	1.50	1.75	2.00	2.25	. 25	. 50	. 75	1.00	1.25			2
DISTANCE (FT)	7.25	7.25	7.25	62.7					10, 20		20.20		10.20		20, 20 RC 40				28.23	29.23	28,25	28,25	38, 25	38, 25	38,25	38.25	38,25	38,25	38.25	38,25	38,25	48.25	48,25	48.25	48.25	48.25	48.25	48.25	48.25	48.25	58,25	58,25	58,25	58.25	58.25			20.60
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	:											•							7.4																													
U (FPS)	24	22	. 24	. 23	. 24	202.	16.	. 34	.36	34	. 36	36	. 35	4.	4	4	4	.40	36	39	.37	. 35	39	.37	. 38	38.	.37	. 37	. 36	. 35	. 36	38	40	. 38	. 38	. 41	. 41	. 36	. 41	. 36	. 32	. 36	38	36		-	000	. 34
V (FPS)	95	90	.07	.06	=	0.0	.03	.03	.03	60	66	= :	.13		9	00	8	12	8	08	8	9	<u>.</u>	02	.02	.07	90.	60'	.06	=	6 0 '	02	03	66.	.07	03	02	02	07	04	90	04	10		2	2.9	2.0	90.
DEPTH (FT)	R.C.	50	. 75	1.00	1.25	1.50	.40	. 65	06 '	1.15	1.40	1,65	1,90	52	20	. 75	1.00	1.25	1.50	1.75	2.00	. 25	. 50	. 75	1.00	1.25	1.50	1.75	2.00	2.25	. 20	45	. 70	. 95	1.20	1.45	1.70	1.95	2.20	22	02	25					1. /3	Z, 00
TANCE			_	_	_	_	_	_	_	~	~	~	~	~	0	0	0	0	0	0	0	0	0	0	0	0			0	0	~	0	0	0												~ ^		•

TABLE B-5 SANGAMON RIVER CROSS SECTION 4 JULY 10 1984 - (CONCLUDED)

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MEASURED DATA

U (FPS)	.32	.27	. 29	. 25	. 27	.27	. 30	. 22	. 23	. 24	. 22
(FPS)	. 14	. 16	. 16	. 14	.17	. 17	. 16	. 15	. 17	. 18	.17
DEPTH (FT)	2.00	2.25	. 33	. 58	. 83	1.08	1.33	1.58	1.83	2.08	2.33
DISTANCE (FT)	58.25	58.25	65.25	65.25	65,25	65.25	65.25	65, 25	65.25	65.25	65.25
U (FPS)	32	31	29	32	26	. 27	. 28	.28	29		
V (FPS)	01	20	03	8	06	60	07	90	0.8		
DEPTH (FT)	50.0	33		83	1.08	1.33	1.58	1.83	2.08		
DISTANCE (FT)	57 BO	64 50	64.50	64.50	64.50	64.50	64.50	64.50	64.50		

TABLE B-6 SANGAMON RIVER CROSS SECTION 2 JULY 12 1984

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MEASURED DATA

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ANGLE (DEGREE)	0.000	87.529 74.824 77.647	76.235 0.000	77.647 81.882 80.471 80.471 0.000	84.706 80.471 86.118 0.000	81,882 81,882 0,000	86.118 0.000 0.000	76, 235 76, 235 74, 824 74, 885 74, 885 0, 000
s.D.	0.000	0.000	0.000 0.000	0.0000000000000000000000000000000000000	0.000 0.000 0.000 0.000	0.000	0.000 0.000 0.000 0.000	. 109 . 026 . 016 . 014 0. 000
(FPS)	(-)						() 000.0	
s.D.	0, 000	.047 .033 .051	.026 0.000	. 124 . 087 . 064 . 034 0. 000	.121 .063 .148 0.000	104 036 0.000	0.000 0.000	
V (FPS)	0.000	374 044 141	- , 093 0 , 000	208 130 194 113	078 . 071 192 0. 000	037 .119 0.000	- 003 0.000	
S .D.	0.00	.040 .025 .068	.054	. 105 . 081 . 055 046	. 063 . 061 . 114 0. 000	059 059 000	0.000	150 047 043 043 000
U (FPS)	0,000	2.176 2.218 2.139	2.054	2.598 2.316 2.296 2.004 0.000	1.979 2.046 1.839 0.000	1.548 1.677 0.000	1.350 0.000 0.000	2.296 2.163 2.163 2.050 0.000
DEPTH (FT)	0.00	22	1.00	23 20 1.00 1.00	- 50 - 70 - 20	98 98	. 70	
DISTANCE (FT)	2.00	7.00	7.00	12.00 12.00 12.00	17.00 17.00 17.00	22.00 22.00 22.00	27.00 27.00	44444 9.82.444 8.82.92.44 8.82.82.83 8.82.82 8.82.82 8.82.83 8.82.83 8.83 8

TABLE B-6 SANGAMON RIVER CROSS SECTION 2 JULY 12 1984 - (CONCLUDED)

MEASURED DATA

DISTANCE (FT)	DEPTH (FT)	U (FPS)	s.D.	V (FPS)	s. D.	(FPS)	s.D.	ANGLE (DEGREE)	S.[
0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	23 23 23 23 23 23 23 23 23 23 23 23 23 2	2.676 2.615 2.364 2.004 0.000	036 036 084 000 000		00000000000000000000000000000000000000	.371 .180 .098 .124 0.000	055 044 131 0.000	79.059 79.059 79.059 79.059	
14.25 14.25 14.25	20 45	2.483 2.189 1.667 0.000	055		0.000.0	372 187 138 0.000	. 063 . 030 . 039 0. 000	83.294 84.706 84.706 0.000	000.0
19.25 19.25 19.25	20 95	2.153 1.644 0.000	. 067 . 043 0. 000		0.000 0.000 0.000	. 169 . 055 0. 000	. 042 . 034 0. 000	74.824 74.824 0.000	0.000 0000.000
24.25 24.25	50	1.860	. 07 9 0. 000		0.000	. 19 0 0.000	.07 3 0.000	87.529 0.000	. 000 000 .
29.25	50	1.589	085	(-)	0.000	.076	. 050	69.176	000

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TABLE B-7 SANGAMON RIVER CROSS SECTION 1 JULY 13 1984

MEASURED DATA

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(-) 0.000 0.000 (-) 0.000 87.52 (-) 0.000 0.00	(-) 0.000 0.000 (-) 0.000 87.52 (-) 0.000 84.70	(-) 0.000 0.000 (-) 0.000 87.52 (-) 0.000 87.52 (-) 0.000 84.70 (-) 0.000 81.88	() 0.000 0.000 () 0.000 87.52 () 0.000 84.70 () 0.000 84.70 () 0.000 84.70 () 0.000 91.86	- 0.000 0.000 - 0.000 0.000 - 0.000 0.000 - 0.000 0.000 - 0.000 0.000 - 0.000 0.000 - 0.000 0.000 - 0.000 0.000 - 0.000 0.000 - 0.000 0.000 - 0.000 0.000 - 0.000 0.000 - 0.000 0.000 - 0.000 0.000
(-) 0.000 87.52 (-) 0.000 0.00	(-) 0.000 87.52 (-) 0.000 0.000 (-) 0.000 84.70	() 0.000 87.52 (-) 0.000 87.52 (-) 0.000 84.70 (-) 0.000 81.88	() 0.000 87.52 () 0.000 84.70 () 0.000 84.70 () 0.000 81.80 () 0.000 91.80	- 0.000 0.752 - 0.000 0.752 - 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000
(-) 0.000 0.00	(-) 0.000 0.000 (-) 0.000 84.70	(-) 0.000 0.000 (-) 0.000 84.70 (-) 0.000 81.88	(-) 0.000 0.000 (-) 0.000 84.70 (-) 0.000 81.86 (-) 0.000 0.000	0.000 0.000 0.000 1 0.000 84.706 1 0.000 81.882 1 0.000 91.882 1 0.000 91.382 1 0.000 91.706
	(-) 0.000 84.70	(-) 0.000 84.70 (-) 0.000 81.88	(-) 0.000 84.70 (-) 0.000 81.86 (-) 0.000 81.86	- 0.000 84.706 - 0.000 81.885 - 0.000 0.1.885 - 0.000 0.000 - 0.000 0.000 - 0.000 84.706

TABLE B-7 SANGAMON RIVER CROSS SECTION 1 JULY 13 1984 - (CONTINUED)

1.0

No. of Concession, Name

S.D	000000000000000000000000000000000000000	0.000	0.000	. 000 . 000 0. 000	0000.0	0.000	. 654 . 000 . 000 0. 000	000.0	000.0
ANGLE (DEGREE)	81.882 81.882 81.882 0.000	76.235 77.714 80.471 76.235 0.000	0.000	73.412 73.341 73.412 0.000	74.824 74.824 74.824 0.000	81.882 81.882 81.882 0.000	81.479 81.882 81.882 0.000	81.882 81.882 0.000	80.471 81.882 0.000
s.D.	000000000000000000000000000000000000000		0.000	0.050	123 054 0.000	. 070 . 044 . 098 0. 000	103 115 0.000	. 051 . 039 0. 000	. 053 . 048 0. 000
(FPS)			0.000	188 035 0.000	. 303 . 172 . 266 0. 000	151 113 0.000	.139 .055 .133 0.000	107 074 0.000	203
s.D.	.075 .057 .086 0.000	113 068 063 0000	0,000	0.0000000000000000000000000000000000000	000 000 000 000 000 000 000 000 000 000 000 0000	0.000 0.000 0.000 0.000	0.000	0.000	0,000
V (FPS)	. 231 - 024 . 013 0. 000	-,008 051 148 148 0,000	(-)						
S.D.	0.000	.073 .029 .064 .049 0.000	000 .0	. 045 . 035 . 027 0. 000	.072 .066 .032 0.000	049 048 048 048 048	076 048 049 0.000	. 053 . 032 0. 000	.047 .052 0.000
U (FPS)	715 488 477 0.000	.524 .472 .572 .570 .000	0.000	2.164 2.022 1.847 0.000	2.118 1.798 1.586 0.000	2.049 1.765 1.463 0.000	2.056 1.737 1.602 0.000	1.516 1.394 0.000	1.535 1.147 0.000
DEPTH (FT)	. 17 . 42 . 67 1. 17	1. 67 92 22	0.00	. 25 . 75 1. 25	. 25 . 75 . 75	- 25 - 75 - 25	- 17 08 83	. 25 1.00	17 42 92
DI STANCE (FT)	75.00 75.00 75.00 75.00	000000000000000000000000000000000000000	0.00	4 4 4 4 26 26 26 26 26 26 26	9.25 9.25 9.25	14.25 14.25 14.25	19.25 19.25 19.25	24.25 24.25 24.25	29.25 29.25 29.25

TABLE B-7 SANGAMON RIVER CROSS SECTION 1 JULY 13 1984 - (CONCLUDED)

MEASURED DATA

.

DISTANCE (FT)	DEPTH (FT)	U (FPS)	s.D.	V (FPS)	s.D.	K (FPS)	s.D.	ANGLE (DEGREE)	s.D.
34.25 34.25	.17 .67	888 0.000	. 112 0. 000		0, 000 0, 000 0, 000	.027 0.000	374	74.824 0.000	000.0
69. 25 69. 25	. 25	, 697 0.000	. 028 0. 000		000 000	. 068 0. 000	0.000	79,059 0.000	.000
74.25 74.25 74.25	.17 .67 1.17	.678 .530 0.000	. 052 . 088 0. 000		0.000 0.000 0.000	023 005 0.000	.067 .065 0.000	83.697 79.059 0.000	21.257 .000 0.000
79.25 79.25 79.25 79.25		. 665 . 639 . 506 . 290	050.065.067.0077		0,000 0,000 0,000 0,000 0,000	181 187 078 019	.062 .068 .052 .052	76.235 76.165 74.824 76.235	0.000 316 0.000

TABLE B-8 SANGAMON RIVER CROSS SECTION 5 AUGUST 5 1985

Surger States

MEASURED DATA

DISTANCE (FT)	DEPTH (FT)	U (FPS)	S .D.	(FPS)	S. D.	K (FPS)	s.D.	ANGLE (DEGREE)	S. [
0.00	0.00	0.000	000.0	0.000	0.000	<pre></pre>	0.000	0.000	0.00
							000 0	77 EA7	č
9.00	. 25	. 580	353	- , 063	902				
.9.00	. 75	. 654	, 329	037			0.000	100.00	6 6
9,00	1.25	. 648	. 394	067	. 292	-	0.000	83.294	
6 , 00	1.75	0,000	0.000	0.000	0,000	- -	0.000	0.000	0.00
									Ŏ
33.50	. 35	.678	. 357	142	. 222		0.000	76.235	
33, 50	. 85	.570	. 362	190	. 203	, ,	0.000	76.235	
33.50	1.35	. 586	. 287	125	.201		0.000	76,235	00,00
33.50	1.85	0.000	0.000	0,000	0,000	- -	0.000	0.000	0.00
					-				č
50.00	.20	.901	.442	218	359		0.000	140 11	
50.00	. 70	. 575	. 368	106	. 274		0.000	84.706	B F
50.00	. 9 5	. 674	. 390	-,025	. 252	-	0,000	79.710	
50.00	1.45	0.000	0.000	0,000	0.000		0.000	0.000	0,00
63.50	30	.814	. 543	. 080	.404	(-)	0.000	72.000	00.
63.50	90	. 764	. 458	- , 050	.271	<u>.</u>	0.000	83,294	õ.
63.50	1.05	. 587	. 260	- , 089	. 181		0.000	79.007	. 27
63.50	1.55	0.000	0.000	0.000	0,000	- -	0.000	0.000	0.00
	ł							78 053	R 7.
74.00	22	.619	. 429	- 125	855.			10.900 ACC 37	
74.00	222	.447	200	- 305	103			76 074	450
. 4 . 00. 4 .		7070.			1 60			76.191	. 250
74.00	2.25	0.000	0.000	0.000	0.000		0.000	0.000	0.000
						•			
0.00	0.00	0,000	0.000	(-)	0,000	0.000	0.000	0.000	0.00
8,25	50	. 746	. 299	(-)	0.000	. 160	. 324	81,882	000.
		805	294	:	0.000	. 182	. 305	81,882	ŏo.
	100	801	391		0.000	. 197	.287	81.882	00.
8.25	1.75	0.000	0.000	- · ·	0.000	0.000	0.000	0.000	0.00
32.75	.60	. 685	. 307	(-)	0,000	. 207	. 256	70.588	000
32.75	1.10	.518	. 301		0,000	002	. 234	69.25/	5.0
32.75	1.35	. 809	. 397		0.000	181.0	EZE.	0/1/60	
32.75	1.85	0.000	0,000		0.000	0.000	0.000	0.000	0.00

TABLE B-8 SANGAMON RIVER CROSS SECTION 5 AUGUST 5 1985 - (CONCLUDED)

Service Service

記録を書

C. Allen

S.D.	0000.0	0000.000	0.000 .258 .638
ANGLE (DEGREE)	76.235 74.824 74.824 0.000	72.000 72.000 72.000 0.000	76.2 35 76.188 70.203
s.D.	. 395 . 323 . 209 0. 000	.301 .344 .227 0.000	. 292 . 194 . 240
(FPS)	014 117 029 0.000	076 111 169 0.000	049 .071 .055
s.D.			000 0
V (FPS)			
s.D.	. 336 . 315 . 220 0.000	.322 .329 .281 0.000	. 323 . 239 . 324
U (FPS)	.633 .677 .718 0.000	. 557 . 708 . 847 0.000	. 539 . 562 . 562
DEPTH (FT)	. 20 . 70 . 95	. 30 1.05 1.55	.50 1.50
DISTANCE (FT)	49,25 49,25 49,25 49,25	62.75 62.75 62.75 62.75	73.25 73.25 73.25

TABLE B-9 SANGAMON RIVER CROSS SECTION 6 AUGUST 6 1985

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S.D. ANGLE S.D. (DEGREE)	.000 0.000 0.000	.000 84.711 0.000	.000 81.888 .000	000 83.299 000	27 EK2 000		000 75.339 680	.000 87.535 .000	.000 0.000 0.000	000 83 299 000					.000 77.652 .000	.000 76.241 .000	000 77.652 ,000	.000 0.000 0.000
K (FPS)	(-)	(')	~ - _					(-)	(-)						-	- -	(·)	- ·)
с, D.	0.000	1.365	677	140		503	. 492	. 532	0.000	102		194.	COE	.447	. 227	. 311	. 249	000 0
V (FPS)	0.000	. 732	179		120.	.043	. 358	019	0.000			160	463	278	627	. 189	340	0.000
s.D.	0.000	.829	216			. 184	. 292	459	0.000		199.	. 666	. 449	. 550	. 346	537	320	000 0
U (FPS)	0.000	934	108		. 604	.514	. 719	543	0.000		1.644	, 679	1.154	. 991	1.574	458	262	
DEPTH (FT)	0.00	00		1	. 70	36.	1 20		1.95		.10	.35	. 85	1.10	1.35			- C
DISTANCE (FT)	0.00	00 9		9.00	6.00	6.00	00 9		6.00 6		12.25	12,25	12.25	12.25	12.25	10.0		22.21

TABLE B-9 SANGAMON RIVER CROSS SECTION 6 AUGUST 6 1985 - (CONTINUED)

" And Bridger

ALC: NO

19 - S.

n ar an

DISTANCE (FT)	DEPTH (FT)	U (FPS)	s.D.	V (FPS)	s.D.	H (FPS)	s.D.	ANGLE (DEGREE)	S I
63.00	.20	.617	212	013	.241		0.000	87.535	000.
63.00	. 70	. 672	.179	077	. 159		0.000	008 80	
63,00 63,00	1.20	926.0	0.000	0.000	0.000		0,000	000 .00	0.000
73.00	.30	. 568	. 307	- , 302	. 234	(-))	0.000	86.123	00.
73.00	. 55	.571	325	- , 289	356		0.000	87.535	000
73.00	1.05	. 748	. 366	073	. 363	-	0.000	87.535	
73.00	1.30	. 688	.311	.032	. 397			72 652 77 653	
73.00 73.00	2.05	0.000	000.0	0.000	0.000		0.000	0.000	0.000
			•						
77.00	02	.950	. 839	.574	.814	(-)	0 .000	74.829	. 000
77.00	30	519	. 335	025	. 502	(-)	0 0 0 0 0 0 0	79.064	8
77.00	*. 55	. 533	.132	013	. 185	- - -	0.000	79.064	
77.00	80	511	.268	015	. 393		0.000	79.064	
77.00	1.05	. 542	.240	600 [.] -	162.		0.000	81.888 70 064	
00.12		. 628	.460	977	. / 10			BD 476	
00.77		- 44 		101	419			80.476 80.476	
77.00	2.30	0.000	0.000	0.000	0.000		0.000	0.000	0.000
0.00	0.00	0.000	0,000	(-)	0.000	0.000	0.000	0.000	0.000
5,25	20	545	. 163		0.000	- 212	. 283	77.652	
2.52 2.52	40	962.	. 330		0.000	100.	325	200.//	
		727	445			130	437	84.437	. 567
5.25	1.20	568	311		0.000	053	. 293	80.476	000
5.25	1.45	. 553	. 267	(-)	0.000	. 001	. 242	80.476	000.
5.25	1.70	. 590	.373	-	0,000	. 107	. 332	80.476	000.
5,25	1.95	. 497	. 489		0.000	0.001	. 403	80.476	
5.25	2.45	0.000	0.000		0.000	0.000	0.000	0.000	0.000
11.50	. 35	. 420	.401	(-)	0.000	. 285	. 260	81.928	. 239
11.50	. 60	. 944	. 752	- -	0.000	- , 313	. 765	77.652	000.
11.50	. 85	. 819	.459		0.000	037	478	77.652	000
11.50	1.10	. 641	.429	- - 	0.000	.026	401	77.652	
00.11	1.60	1. UZU	250		000.0	- 409 082	162.	200.11	
11.50	1.65 2.35	0.000	0.000		0.000	0.000	0.000	0.000	0.000

TABLE B-9 SANGAMON RIVER CROSS SECTION 6 AUGUST 6 1985 - (CONCLUDED)

DEPTH (FT)	U (FPS)	G .D	(FPS)	S.D.	(FPS)	S. D.	ANGLE (DEGREE)	S.D.
	1.390	. 337	•	0.000	. 727	336	80.476	000.
		404 104				200.	AD 476	
	1.003	585			1	- - 	80.637	456
	525	532		0.000	100	479	81.846	. 242
Ū	0.000	0.000	(-)	000 0	0,000	0.000	0.000	0.000
		000			001			000
	280	362		0.000	- 439	1957	19, U54	. 000
			 			1077.	76 454	512
	426.	- 200 1995			- 159	472	77.652	000
Ū	0.000	0.000		0.000	0.000	0.000	0.000	0.000
	. 263	. 311	(-)	0,000	- , 582	. 298	79.064	. 000
•	1.183	. 414		0.000	. 530	. 369	62.081	. 242
•	1.254	. 353	(-))	0.000	. 619	.316	62.123	. 000
-	. 368	. 548	- - -	0.000	. 715	. 401	62.123	. 000
0	0000.	0,000	 -	0.000	0.000	0.000	0.000	0.000
	. 763	. 160	(-)	0,000	227	. 241	81.888	000
	. 717	258	- - -	0,000	- , 038	. 246.	81,888	. 000
	. 581	. 364	(· _)	0.000	076	. 384	94.549	. 250
-	1.370	.460	$\left(\begin{array}{c} \cdot \\ \cdot \end{array}\right)$	0.000	.679	.416	93.229	. 258
	. 734	. 552	~ · _	0,000	. 313	. 442	94.543	. 267
0	000.0	0,000		0.000	0.000	0.000	0.000	0.000
	.614	. 277	(~ -)	0,000	- , 075	. 258	84.711	0.000
	. 752	.487	(- -	0.000	. 049	. 452	84.711	0.000
	. 694	, 903	(- <u>)</u>	0.000	.049	. 855	75,698	200
	.213	. 644	(- -	0,000	489	. 569	74,829	000 '
	. 874	.677	(·)	0, 000	. 226	. 603	74.829	0.000
	.874	, 436		0.000	. 245	. 363	74.829	. 000
0	0.000	000.0	- ·	0.000	0.000	0.000	0.000	0,000
	. 578	, 200	(-)	0.000	. 208	. 232	79.064	. 000
	. 539	390	$\hat{\cdot}$	0,000	. 139	.417	77.652	.000
	505	292	-	0,000	660	. 283	77.652	000
	. 464	222		0.000	. 046	203	78,098	. 665
	. 600	.416		0.000	241	. 365	75,906	. 608
	.440	. 192	(-)	0.000	. 068	. 142	74.829	. 000

TABLE B-10 SANGAMON RIVER CROSS SECTION 7 AUGUST 7 1985

MEASURED DATA

DISTANCE (FT)	DEPTH (FT)	U (FPS)	ຕ ຸ ອ	V (FPS)	s.D.	W (FPS)	s. D.	ANGLE (DEGREE)	s. S
1.00	0.00	0.000	0.000	0.000	0.000	(-)	0.000	0.000	0.000
8.50 8.50 8.50	. 40 1.90	.220 .427 0.000	358 277 0.000	215 .103 0.000	.411 .274 0.000		000 000 0	86,919 87,529 0,000	600 · 0
441 441 000 000 000 000 000 000 000 000	2	239 215 286 352 000 000	2 4 7 7 7 7 7 7 7 7 7 7 7 0 0 0 0 0 0 0 0	- 090 - 070 - 035 - 119 039 0.000	. 333 . 124 . 3392 . 455 0 000			79.059 80.426 83.294 84.706 84.578 0.000	0.000 0000 0000 0000 0000 0000 0000
28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 28.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.000 20.000 20.000 20.00000000	2303003	846 846 9325 930 930 930 930 930 930 930 930 930 930	598 343 372 312 298	. 148 . 221 . 335 . 025	668 391 391 202 80 204 200 204 200 200 200 200 200 200 20			83.294 84.706 86.125 88.941 88.941 84.706	000000000000000000000000000000000000000
00000000000000000000000000000000000000	02202220 02202220 02202220 02202220 022022	000 000 000 000 000 000 000 000	0.000 332 0.000 336 0.000	0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000	277 277 277 277 277 277 277 277 277 277			81.882 89.590 89.590 89.471 88.941 0.000	0000
50,00 50,00 50,00	20	. 507 . 497 . 482 0.000	. 480 . 415 0. 000	070 058 049 0.000	508 433 0.000		000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 000 0000	84.706 84.706 84.706 0.000	000 000 000 000 000 000 000 000 000 00
63 00 63 00 60 63 00 60 60 60 60 60 60 60 60 60 60 60 60 6	1 1 1 0 3 3 3 3 1 1 0 3 3 1 1 0 3 3 1 1 0 3 3 1 1 0 3 3 1 0 3 3 1 0 1 0 1 0 1 0 1 0 1 0 1 0 1 0 1 0 1 0	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	214 2388 257 206 385 000	- 043 - 037 - 063 - 063 000	.176 .440 .331 .345 .334		0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	84.706 84.749 84.706 84.706 83.294 0.000	000 000 000 000 000 000 000 000

TABLE B-10 SANGAMON RIVER CROSS SECTION 7 AUGUST 7 1985 - (CONTINUED)

MEASURED DATA

S.D	000	000.		0.000	0.000		0.000	000	0.000	0.00		0.000		000.	000.0	0,000	000	000	. 626	000.	000	000 .	.687	0.000		000	000.	000 .	000.	000	000.	0.000	000 .	0.000	0.000	0.000	0,000
ANGLE (DEGREE)	87,529 86,118	87.529	87.529	01.002	81.882	00.341	0.000		0.000	 700.10	81.882 	19.005	77.647	77.647	77.647	0,000	83.294	83.294	84.339	83.294	80.471	80.471	80.582	0.000		87,529	87.529	86.118	86,118	86,118	86.118	0.000	81.882	79.059	79.059	79.059	0.000
s.D.	0.000 0.000	0.000	0.000	0,000	0.000	0,000	0,000		0.000	40	212	.411	, 603	481	1,050	0.000	.411	. 243	. 359	. 376	. 549	. 248	. 306	0.000		. 335	. 498	. 433	.489	. 493	. 502	0.000	. 569	. 379	. 357	. 225	0.000
W (FPS)		- - -				-	-		0.000	020	.070	-, 066	.471	. 053	-,151	0.000	. 139	.017	- , 556	- , 173	.154	- 001	- 102	0.000		146	019	. 098	410	. 006	. 151	0.000	.158	.012	140	. 030	0.000
s.D.	192	. 455	385	.371	. 204	174	0,000		0.000	0.000	0,000	0.000	0.000	0,000	0.000	0, 000	0.000	0.000	0.000	0,000	0.000	0.000	0.000	0.000		0,000	0,000	0.000	0.000	0.000	0,000	0.000	0.000	0.000	0.000	0.000	0.000
V (FPS)	-, 132 -, 068	.041	-, 029	- , 050	- , 048	- ,049	0,000		(-)	-	- - -		$\hat{\cdot}$	- - -	<u>.</u>	(-)	(-)	<u>-</u>		•						(-)	^ · _	 	 	(-)	- - -	- - -	(-)	- - -	- - -	- <u>)</u>	
S.D.	234	430	.468	. 338	. 215	. 229	0,000	•	0,000	, 193	. 209	.460	. 658	.517	1.289	0.000	. 324	. 231	328	429	585	212	346	0.000	-	.291	516	479	492	. 592	. 701	0,000	.412	. 434	. 436	337	0.000
U (FPS)	261 360	449	. 355	.374	. 284	. 299	0.000		0.000	. 499	.461	. 325	. 900	. 423	. 129	0.000	. 648	563	- 038	300	629	382	0.00	0.000	-	. 462	547	622	. 860	498	577	0.000	. 688	. 458	319	534	0.000
DEPTH (FT)	20	22.	90	1.05	1.30	1.55	2.05		0.00	. 35	. 60	1.10	1.35	1,60	2.10	2.60	50		2 2 2					2.25	Ì	30		BO BO	1.05	1.30	1.55	2.05	. 20	. 45	20	. 0	1.45
DISTANCE (FT)	75.25 78.25	75.25	75.25	75,25	75.25	75,25	75.25		1.00	12.25	12.25	12.25	12.25	12.25	12.25	12.25	24 25	24 25	24.04	21010	0 1 7 0 0 1 7 0 0 1 7 0	20.00	24. FC	24.25		33 25	33 25	33.25	33.25	33.25	33.25	33.25	48.25	48.25	48.25	48.25	40.25

TABLE B-10 SANGAMON RIVER CROSS SECTION 7 AUGUST 7 1985 - (CONCLUDED)

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MEASURED DATA

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S.D	000000000000000000000000000000000000000	000 242 000
ANGLE (DEGREE)	80.511 90.353 88.941 88.941 0.000	80.471 80.471 87.571 87.529
s.D.	. 329 . 545 . 373 . 373 0 000	323 183 240
(FPS)	. 079 208 025 032 0. 000	085 016 150
s.D.		0.000 0.000 0.000
V (FPS)		
s, D.	344 592 506 428 000	300 235 282 371
U (FPS)	483 260 269 000	343 - 202 - 275
DEPTH (FT)	800 800 800 800 800	333
DISTANCE (FT)	63.25 63.25 63.25 63.25 63.25 63.25	73.50 73.50 73.50 73.50

TABLE B-11 SANGAMON RIVER CROSS SECTION 8 AUGUST 8 1985

· TABLE B-11 SANGAMON RIVER CROSS SECTION 8 AUGUST 8 1985 - (CONTINUED)

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s.D.	000	. 000 000 000 000	0000.0000.000	703	. 000	000000000000000000000000000000000000000	0.000
ANGLE (DEGREE)	104.476 103.064 103.064	100.241 100.241 111.535	112,946 112,946 0.000	103.904 103.064 104.476 105.888 114.358	113.079 112.946 0.000	117.182 117.182 114.358 114.358 114.358 114.358 118.594 118.594 118.594	88.946 90.358 84.711 84.711 96.005 96.005 98.102
s.D.	431 381 347	. 481 . 268 . 361	.325 .301 0.000	295 383 358 431	451	296 201 201 201 201 201 201 201 201 201 201	243 297 290 282 282 310
W (FPS)	.115 .025 .128	. 098 . 117 . 063	001 .124 0.000	- 123 - 031 122 074	- 000 - 000 - 000	.080 .027 .033 .058 .058 .050 .039	- 065 - 062 128 157 049
S.D.	0.00 0.00 0.00 0.00	0.000 0.000 0.000	0.000		00000		
V (FPS)							
s.D.	1, 299 391 389	. 325	. 323 . 408 0. 000	273 315 395 481	. 524 . 524 0. 000	256 255 255 255 255 255 255 295 295 295 295	255 285 287 287 287 344 313
U (FPS)	1.469 .610 .662	658 686 554	. 442 . 636 0. 000	577 495 533 501	398	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	273 319 336 338 338 338 383 383 383
DEPTH (FT)	15 40	1.15 1.65 2.15	2.65 3.15	65 1.15 1.65	2.65 3.15	.50 .75 3.25 3.25 3.25	22.20 22.20 22.20 23.20
DI STANCE (FT)	41.75 41.75 41.75	41.75	41.75 41.75 41.75	57,25 57,25 57,25 57,25	57.25 57.25 57.25	85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25 85.25	66.75 66.75 66.75 66.75 66.75 66.75 66.75

TABLE B-11 SANGAMON RIVER CROSS SECTION 8 AUGUST 8 1985 - (CONCLUDED)

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MEASURED DATA

s.D		347 00000 000000000000000000000000000000	o 000000000000000000000000000000000000	0.0000000000000000000000000000000000000
ANGLE (DEGREE)	115.770 115.770 115.770 115.770 115.770 117.182 114.358 114.358 114.358 117.182 0.000	86.035 86.123 87.535 88.946 93.182 88.946 96.005 0.000	0.000 105.888 105.888 105.888 103.064 103.064 103.064 103.064 0.000	105.888 105.888 105.888 108.711 108.711 98.787 98.829 105.888
s.D.			0.000 385 373 285 285 285 285 285 285 285 285 285 285	. 484 . 416 . 425 . 373 . 373 . 373 . 353 . 363 . 000
H (FPS)			0.000 0.060 0.046 0.046 0.095 - 1139 - 015 0.000	.156 .027 .157 .157 .135 .006
s.D.	293 201 201 201 197 168 168 168 142 228 0.000	613 3357 3357 297 297 291 291 291 0.000		
V (FPS)	153 173 173 177 177 177 177 177 177 177 17	- 084 083 - 111 - 141 - 067 - 067 - 022 0000		
s.D.	0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.000000	477 269 273 273 241 268 268 0.000	0.000 322 322 323 323 323 323 323 323 323	507 338 475 475 413 413 0 000
U (FPS)	0.00912 0.00912 0.00912 0.00912 0.00912 0.00912	0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,008 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000 0,000000	0.000	715 601 661 671 786 786 727 727 727 727 700
DEPTH (FT)		20 20 20 20 20 20 20 20 20 20 20 20 20 2	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	3.50 3.50 3.50 3.50 3.50 3.50 3.50 3.50
DISTANCE (FT)		70.00 70.00 70.00 70.00 70.00 70.00 70.00 70.00 70.00 70.00 70.00 70.00 70.00 70.00 70.00 70.00 70.00 70.00 70.00 70.00	5333333333333 7 7 7 7 7 7 7 7 7 7 7 7 7 7	26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50 26.50

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TABLE B-12 SANGAMON RIVER CROSS SECTION 9 AUGUST 9 1985

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MEASURED DATA

DISTANCE (FT)	DEPTH (FT)	U (FPS)	S.D.	V (FPS)	s.D.	(FPS)	S.D.	ANGLE (DEGREE)	°.
0.00	0.00	000 0	0.000	0.000	0.000	(-)	0,000	0,000	0.000
	00	178	262	- 253	168	(-)	0.000	124.235	000.
			152	- 299	115		0.000	125.862	.513
	n o	192	122	285	060	- - -	0.000	127.059	000.
	00.1	136	.132	202	.123	- J	0.000	125.647	000.
11.50	1.70	0.000	0.000	0.000	0.000	(-)	0.000	0.000	0.000
							000 0	117 176	100
26.50	10	. 182	.319	- , 420	523.	-	0.000		, ,
26.50	.35	. 326	.108	- , 355	. 104	-	0.000	121.412	00.
26.50	.35	.317	129	388	, 086		000 0	121.412	00.
26.50	1,35	.291	. 092	321	.087	- -	000.0	117.176	0.00
26.50	1,85	. 245	.126	- , 252	. 1:08	- - -	0.000	120.000	0.00
26.50	2.10	.057	.177	- , 199	.134	-	0,000	122.012	02.
26.50	2.60	0.000	0.000	0,000	0,000	- -	0.000	0.000	0.000
41.50	.25	. 282	. 187	- , 233	. 139	(0,000	120.409	1.464
41.50	20	.247	.108	- , 269	.075		000.0	118.588	000.
41.50	25	. 275	. 115	225	, 096		0.000	115.765	00.
41.50	1.25	. 258	. 107	-,189	.077	- - -	0.000	118.588	000.
41.50	1.75	0.00	0.000	0,000	0,000	- - -	0.000	0,000	0.000
51,00	15	390	. 142	- 051	660 '	(-)	0.000	120.000	0.000
51.00	40	.437	.120	- ,071	. 100	- - -	0, 000	116.887	. 662
51.00	1.40	. 453	. 023	- , 009	, 032	- -	0.000	114.353	0.000
51.00	1.90	450	030	010	. 024	- - -	000 000	118.657	308
51.00	2.40	.447	.053	.052	.040	- - -	0.000	115.765	0.000
51.00	2.90	402	.063	. 030	. 048	- - -	0.000	118.588	000.
51.00	3.40	423	. 060	1058	.037	- - -	0.000	122.824	000.
51.00	3.90	0.000	0.000	000-*0	0,000	- J	0.000	0,000	0.000
	ac S	6 L Y	670	101 -	850	1		116.016	546
00.70		11	410.		-				
57.00	20	. 515	.041	092	620		0.000		
57.00	1.00	.471	.039	070	. 044	-	0,000	115.765	000.
57.00	1.50	.474	. 031	- , 003	.031	-	0.000	111.529	
57.00	1.75	430	.026	.015	. 051	-	0.000	116.336	102.
57.00	2.25	.370	. 037	.067	.041	- - -	0.000	111.529	000.
57.00	2.75	. 421	. 086	. 061	. 060	- - -	0.000	115.154	1.080
57,00	3.25	0.000	0.000	0.000	0.000	- - -	0,000	0.000	0.000

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TABLE B-12 SANGAMON RIVER CROSS SECTION 9 AUGUST 9 1985 - (CONTINUED)

			an						
DI STANCE (FT)	DEPTH (FT)	U (FPS)	8 D.	(FPS)	s.D.	W (FPS)	s.D.	ANGLE (DEGREE)	s.D
66. 75 66. 75 66. 75	. 10 1.10	. 187 . 180 0.000	. 185 . 069 0. 000	- 008 111 0.000	. 136 . 061 0.000		0.000	113.682 114.353 0.000	. 714 0.000 0.000
0.00	0.00	0.000	000 0	(-)	0,000	0,000	0,000	000.0	0,000
10.75 10.75 10.75 10.75	20 45 1.20	319 343 250 0.000	.060 .092 .124 0.000		000000000000000000000000000000000000000	-, 057 -, 036 -, 055 0, 000	.085 .112 .083	124.235 124.235 136.412 0.000	0.000
25.75 25.75 25.75	35 35	589 353	749		000.0	- 019 - 045 - 025	. 120 . 089 . 109	118.588 117.341 117.176	.000 .458 .000
25.75 25.75 25.75 25.75	1.10 1.60 2.35 2.35	349 242 218 0.000	118 094 137 0.000			- 024 - 046 - 082 0.000	079 237 0.000	119.872	0.000
40.75 40.75 40.75 40.75 40.75 75 75 75 75	2200 200 200 200 200 200 200 200 200 20	356 266 2499 0033 000 000	079 105 228 066			044 071 .101 .102 .037 0.000	088 105 105 189 0.056	114 353 112 941 117 169 117 176 118 588 0.000	000000000000000000000000000000000000000
50 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5	9 9 3 3 5 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	075 064 064 035 035 037 039 000 000			229 107 118 1131 1131 083 083 077	022 022 022 022 022 022 022 022 022 022	114.353 114.353 114.284 111.529 111.529 115.7529 115.765 115.765 0.000	000000000000000000000000000000000000000
56.25 56.25 56.25 56.25	22 20 1 00	534 496 498	039 041 042		000000000000000000000000000000000000000	121 129 127	025	110.118 110.118 110.118 110.118	0000.0000

TABLE B-12 SANGAMON RIVER CROSS SECTION 9 AUGUST 9 1985 - (CONCLUDED)

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DISTANCE (FT)	DEPTH (FT)	U (FPS)	s.D.	V (FPS)	S.D.	H (FPS)	s.D.	ANGLE (DEGREE)	S.D
56.25 56.25 56.25 56.25 56.25	- 50 2.50 3.25		048 054 056 056 154		000000000000000000000000000000000000000	113 065 093 063	039 043 053 072	197.647 111.529 111.529 111.529 112.543	. 000 . 000 . 000 . 000

Appendix C

Computed Velocities

No. of Street, or Stre

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Figure C-2. Computed secondary current vectors at cross section 5, June 29, 1984



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Figure C-3. Computed secondary current vectors at cross section 3, July 9, 1984



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Section of





August 6, 1985

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Figure C-7. Computed secondary current vectors at cross section 8, August 8, 1985

TABLE C-1 SANGAMON RIVER CROSS SECTION 6 JUNE 28 1984

TABLE C-1 SANGAMON RIVER CROSS SECTION 6 JUNE 28 1984 - (CONTINUED)

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COMPUTED SECONDARY VELOCITY

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K (FPS)	51		.26	44.	- 17	- , 35	- 44	- 05	. 49	.08	. 28	16	- , 35	32	- , 39	- , 06	. 18	. 67	17	60 -		48		38			10	22	22	16	00.00	- 08	- 30	.09	. 23	. 50	. 43	- 08	0.8	20	a c	000				 -	33	-
V (FPS)	2	5 -	01	.08	30	28	21	- 02	1.	. 12	60.	- , 36	45	- 34	27	04	. 07	Ξ	- 40		- 0	14		- 6			20			80	17	16	- 47	08	14	.18	07	- 18	0			2	ח ת - כ		<u>.</u>	- 20	24	5
DEPTH (FT)	-	0 0 -	22	1.17	2.84	2,52	2.23	2.12	1.63	. 56	. 82	2.88	2.68	2.61	2.37	2.27	1.91	33	2 86	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0		2.03 1				- 0 - 0	1000	200	1, 30			0 21	2,60	2.32	2.05	1 69	1 03	69 6	0 40					201	001	2.0	1 00	
DISTANCE (FT)		20.11	11.04	11.35	12.07	12.10	12.13	12.14	12.26	12.84	12.90	13.26	13.26	13.27	13.29	13.30	12 25	20.00		11. 11. 11.		14,00		14. /4	500 T	29.01	10,03	00.0	10.70	10.01	16.70	16 91	16.01	16 82	16.85	16 91	17 14	17 00				18,05	16.13	19,08	19.08	19.17	19.10	19.60
FPS)	5 6 7 1 6	.07	80.		2		2				- 6	5	20			n u		200	12.	29	24	. 15	. 06	10	02	12	4	24	42	. 44		2.0		5.0	56			04	. 4.		01	40	.48	19	43	13	.48	. 02
V FPS) (0	, 8	Be	88	3.5	5.5	5.6	36			8			30		5	5	- 60	- 02	.02	-	.04	04	40	.06	.07	. 16	-12		.07	5	6	60.0	80.0			2		.03	.02	. 06	.09	. 25	.30	50	- 15	60.
EPTH (FT) (1	L 4 4 4 4 4 4 4 4 4 4 7 1 4	.75 0	00	66 66	26 0	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	50.	44			0.0	44		2		- 8/	68	47	35	4	. 84	56	84	25	12	.97	78	- 19	- 33	0	67	6/	55	21	61	99		37	24	0	- 63 ·	- 74	52	- 20	- 21	- 18	64	56
ANCE DE		17	23 1.	58	36	48		299		02		24	6	980	66 I.	1		92	60 2 [.]	65 2.	76 1.	91 1.	94	1.9	20	50	10	53 2.	57 2.	65 2.	65		82			34	1	75 2.	77	79 2.	84 1.	90	99 1.	69 2.	92 2.	96 2	98 2.	1 0
DIST/ (F)		4	4	4	4	4	4	4	4 1	0	n i	0	0	0	0	'n	ю.	5	9	9.	9	9	9	7.	7	~	e.	80	60	8	8	80	80	0	0		5	0	o,	Ō	6	0	6	10.	10.	10.	10.	11.

TABLE C-1 SANGAMON RIVER CROSS SECTION 6 JUNE 28 1984 - (CONTINUED)

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K (FPS)	. 29	. 15	83	77	39	40	52.		32	EN C	20.	. 27	02.	-1.13	97	76	4 4 4			60	.15	18	34	. 49	. 13	.11		60 -		19	. 16	14	.13	.13	. 15	. 13	. 13	13	. 13	.12	37		57
V (FPS)	.12	. 22	- 39	51	21	14	. 19	. 13	. 13	21.0	.06	.02	- ' 08	17	17	17	6L -	- 28		02	90.	03		60. -	02	.0	0.00	60	- 12	02	.02	.01	0.00	04	.02	0.	0.00	.01	.02	.01	01	0,00	0. 0 10
DEPTH (FT)	1.12	. 57	2.41	2.59	. 61	. 56	. 75	59	1,06	. 66	29	1.47	1.81	1,90	1.98	2.13	2,20	2,29	. 85	61	1.23	1.37	1.50	1.63	. 71	. 65	. 57	C/ . [1.96	96	. 77	. 68	. 63	1.08	. 84	. 73	. 65	1.36	. 94	. 78	1.50	.69	. 61 1,64
DISTANCE (FT)	69 . 94	70.24	70.37	70, 39	71.04	71.04	71.82	71.84	72.46	72.65	72.65	72.79	72.94	72,97	72,99	73.03	73.04	73.06	73.42	73.47	74.00	74.12	74.21	74.29	74.29	74.30	74.31	74.34	74.42	74.48	74.55	74.57	74.58	74.69	74.81	74.84	74.85	74.99	75.04	75.10	75.11	75.12	75.19
;																																											
:																																							×.				
(FPS)	17	. 13	.10	.51	. 47	. 43	. 94	. 15	. 56	. 33	60'-		21	. 63	6 0 '-	. 12	43	. 38	E/ .	4 02	- 10	36	.17	11 -	60 -	. 56	.37	53	57	. 48	42	.37	- , 13	. 57	47	54	- , 63	- ,09	. 71	71	71	- 53	. 69 . 02
V W (FPS) (FPS)	0817	. 11 . 13	. 18 . 10	.60 .51	.21 .47	. 26 . 43	. 30 94	. 09	. 52 56	. 65	- 15 - 09	11 11	.09	. 44 63	-,13 -,09	. 14 12	. 24 . 43	.28 .38	. 37 . 73	40 30	- 11 - 10	-,85 -,36	.10 .17	11 - 60 - 43	- 73 - 39	. 27 56	. 30	. 10 . 23		6748	6242	.30 .37	-,05 -,13	. 29 . 57	- 59 - 47	- 56 - 54	6163	- 0109	. 23 . 71	5071	6871	- 42 - 53	. 23 . 69
DEPTH V W (FPS) (FPS)	1.7908	2.12 .11 .13	2.57 .18 .10	2.37 .60 .51	.64 .21 .47	.55 .26 .43	1.59 .30 .94	2.00 .09 .15	2.29 .52 .56	2.69 .65 .33	2.64 -,15 -,09	2.51 .16 .11	1.84 .09 .21	2.19 .44 .63	2.58 -,1309	2.45 .14 .12	. 63 24 . 43	.55 .28 .38	2.05 .37 .73	2.30 .12 .14 2 RG 40 30	2.521110	2.858536	2.25 .10 .17	2.43 - 09 - 11 2.50 42 43	2.81 - 73 - 39	. 59 . 27 . 56	.55 .30 .37	2.09 .10 .23		2.746748	2.776242	.55 .30 .37	2.16 - 05 - 13	2.29 .29 .57	2.765947	2.68 -,5654	2.656163	1.910109	2.09 .23 .71	2.555071	2.706871	2.624253	. 76 23 . 69 . 55 . 17 . 02

TABLE C-1 SANGAMON RIVER CROSS SECTION 6 JUNE 28 1984 - (CONCLUDED)

COMPUTED SECONDARY VELOCITY

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. 06	.07	.07	.07	11	- , 58	18	14	52	12	62	30	- 41	- 59	43	52
 6	0	0.00	0,00	0.00	03	0.00	0.00	- 02	0.00	02	0.00	01	01	0.00	0.00
1.06	. 86	. 73	. 63	1,19	1.37	1.08	. 87	1.15	. 69	1.18	. 89	. 92	96	. 72	. 73
/5.26	75.36	75,39	75.40	75.45	75.65	75.82	75.92	75.93	75.96	75.98	75,99	76.05	76.18	76.18	76.25
	75.26 1.06 .01 .06	75.26 1.06 .01 .06 75.36 .86 .01 .07	75.26 1.06 .01 .06 75.36 .86 .01 .07 75.39 .73 0.00 .07	75.26 1.06 .01 .06 75.36 1.06 .01 .06 75.39 .73 .01 .07 75.39 .73 0.00 .07 75.40 .63 0.00 .07	75.26 1.06 .01 .06 75.36 1.06 .01 .06 75.36 .86 .01 .07 75.39 .73 0.00 .07 75.40 .63 0.00 .07 75.45 1.19 0.00 .07	75.26 1.06 .01 .06 75.36 .86 .01 .07 75.39 .73 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.65 1.37 03 11	75.26 1.06 .01 .06 75.36 .86 .01 .07 75.39 .73 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.65 1.37 03 11 75.65 1.08 0.00 11	75.26 1.06 .01 .06 75.36 .86 .01 .07 75.39 .73 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.65 1.37 0.00 .07 75.65 1.37 03 11 75.65 1.08 0.000 11 75.65 1.08 0.000 14 75.92 .87 0.000 14	75.26 1.06 .01 .06 75.36 .86 .01 .06 75.39 .73 0.01 .07 75.40 .63 0.00 .07 75.45 1.19 0.00 .07 75.55 1.37 0.00 .07 75.65 1.19 0.00 .01 75.65 1.37 03 11 75.62 1.37 03 11 75.62 1.37 03 14 75.92 .87 0.000 14 75.93 1.15 02 14	75.26 1.06 .01 .06 75.36 1.06 .01 .06 75.36 .86 .01 .07 75.39 .73 0.00 .07 75.40 .63 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .11 75.45 1.37 03 .11 75.92 1.08 0.00 .11 75.93 1.18 .75.93 .168 75.93 1.15 .000 .14 75.93 1.15 .000 .14 75.93 1.15 .000 .14	75.26 1.06 .01 .06 75.36 1.06 .01 .06 75.36 .86 .01 .07 75.39 .73 0.00 .07 75.40 .63 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.45 1.37 03 .11 75.65 1.08 0.00 .11 75.92 1.08 0.00 .14 75.93 1.15 02 .16 75.93 1.15 02 .14 75.93 1.16 .020 .14 75.93 1.16 .02 .14 75.93 1.16 .02 .12 75.96 .69 .020 .12 75.93 .116 .02 .12 75.96 .66 .02 .12	75.26 1.06 .01 .06 75.36 1.06 .01 .06 75.39 .73 0.00 .07 75.39 .73 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .11 75.45 1.08 0.00 .11 75.92 1.08 0.000 18 75.93 1.15 02 .14 75.93 1.15 02 .14 75.99 1.16 02 .12 75.99 1.16 02 .12 75.99 1.16 02 .12 75.99 1.16 02 .12 75.99 .000 .02 .12 75.99 .000 .02 .12 75.99 .000 .02 .12 .00 <td>75.26 1.06 .01 .06 75.36 1.06 .01 .06 75.39 .73 0.00 .07 75.40 .63 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.65 1.19 0.00 .11 75.65 1.08 0.00 .11 75.92 1.137 03 .14 75.93 1.15 02 .14 75.93 1.15 .020 .14 75.93 1.15 .022 .52 75.93 1.15 .022 .12 75.93 .93 .118 .12 75.93 .93 .900 .92 .12 75.93 .93 .900 .91 .92 76.05 .92 .900 .91 .91 75.93 .92 <</td> <td>75.26 1.06 .01 .06 75.36 .86 .01 .07 75.39 .73 0.00 .07 75.40 .63 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .11 75.92 1.13 03 .14 75.93 1.15 02 .14 75.93 1.15 02 .14 75.93 1.15 .020 .14 75.93 1.15 .020 .14 75.93 1.18 .020 .12 75.93 .118 .020 .12 75.93 .19 .020 .12 76.19 .96 .010 .12 76.19 .96 .010 .12</td> <td>75.26 1.06 .01 .06 75.36 .86 .01 .07 75.39 .73 0.00 .07 75.40 .63 0.00 .07 75.45 1.19 0.00 .07 75.65 1.19 0.00 .07 75.65 1.19 0.00 .07 75.65 1.19 0.00 .07 75.65 1.19 0.00 .014 75.92 1.19 0.000 .114 75.93 1.15 02 .14 75.96 1.18 02 .14 75.99 1.18 02 .14 75.99 1.18 02 .12 75.99 .118 02 .12 76.18 .72 .001 .12 76.18 .72 .010 .43 76.18 .72 .010 .43</td>	75.26 1.06 .01 .06 75.36 1.06 .01 .06 75.39 .73 0.00 .07 75.40 .63 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.65 1.19 0.00 .11 75.65 1.08 0.00 .11 75.92 1.137 03 .14 75.93 1.15 02 .14 75.93 1.15 .020 .14 75.93 1.15 .022 .52 75.93 1.15 .022 .12 75.93 .93 .118 .12 75.93 .93 .900 .92 .12 75.93 .93 .900 .91 .92 76.05 .92 .900 .91 .91 75.93 .92 <	75.26 1.06 .01 .06 75.36 .86 .01 .07 75.39 .73 0.00 .07 75.40 .63 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .07 75.45 1.19 0.00 .11 75.92 1.13 03 .14 75.93 1.15 02 .14 75.93 1.15 02 .14 75.93 1.15 .020 .14 75.93 1.15 .020 .14 75.93 1.18 .020 .12 75.93 .118 .020 .12 75.93 .19 .020 .12 76.19 .96 .010 .12 76.19 .96 .010 .12	75.26 1.06 .01 .06 75.36 .86 .01 .07 75.39 .73 0.00 .07 75.40 .63 0.00 .07 75.45 1.19 0.00 .07 75.65 1.19 0.00 .07 75.65 1.19 0.00 .07 75.65 1.19 0.00 .07 75.65 1.19 0.00 .014 75.92 1.19 0.000 .114 75.93 1.15 02 .14 75.96 1.18 02 .14 75.99 1.18 02 .14 75.99 1.18 02 .12 75.99 .118 02 .12 76.18 .72 .001 .12 76.18 .72 .010 .43 76.18 .72 .010 .43

TABLE C-2 SANGAMON RIVER CROSS SECTION 5 JUNE 29 1984

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(FPS)	0	- C	2 4	- -	9 - 1 - 1	.37	. 38	. 33	22	. 27	. 56	. 58	. 50	.49	.87	.70	.33	. 58	.46	.68	43	.34	- , 69	34	27	0	. 27		45		23	12	- , 38	. 13	. 81	. 37	. 15	. 19	- 10	Ξ.	. 13	. 16	.27	11.	.23	14	. 52	- , 25
DEPTH (FT)	6				c/ .2	2,59	2.40	2.17	2.90	2.76	2.59	2.41	2.18	2.90	2.41	2.18	2.90	2.40	2.17	2.58	2,38	2,15	3,03	2.37	2.12	68	2.34	000	2.55	02	2,31	3.09	3.00	2.86	2.71	2.52	2.00	2.28	3.08	2.85	1.93	2.23	2.46	1,84	2.42	2.17	2.63	2.96
DI STANCE (FT)	22 27	20.02	24.04	1 - 1 - 1	25.11	25,11	25.11	25,11	25,98	25.98	25.98	25.99	25,99	31.28	31.28	31.28	32.96	32.96	32.96	34,64	34.64	34,64	34,65	36.32	36.32	36.33	38.00	28.00	38.01	39,67	39,68	39,69	39,69	39,69	39,69	39,69	41.35	41.36	41.37	41.37	43.03	43.03	43.05	44.70	44.72	44.72	44.73	44.74
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DEPTH (FT)			00	00.	1,26	1,63.	68.	1.28	1.61	. 84	1.84	1.23	2.01	1.53	2.14	1.74	1.08	2,23	1.90	1,39	2.31	2.01	1.59	2.37	2.10	1 74	2.42	10	1.85	2,83	2.46	2.23	1.93	2.69	2.28	2.00	2.86	2.71	2.53	2.31	2.05	2.88	2.55	2.09	2.89	2.74	2.56	2.37
DISTANCE (FT)					10.75	11.34	11.99	12.38	13.05	13,63	13,84	14.07	14.66	14.78	15.51	15.58	15.82	16.37	16.42	16.53	17.24	17.27	17.33	18.11	18.13	18.17	18.98		19.02	19.84	19.85	19,86	19.88	20.72	20,74	20.75	21.59	21.60	21.60	21.61	21.62	22.47	22.48	22.49	23.35	23, 35	23.35	23.36

TABLE C-2 SANGAMON RIVER CROSS SECTION 5 JUNE 29 1984 - (CONTINUED)

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Ţ	DEPTH (FT)	2.58 2.52 2.52 2.53 2.53 2.53 2.53 2.53 2.53	2,70 2,63 1,75 2,98 2,98 2,98 2,98 2,98 2,98 2,98 2,98	22822232827 2328225328 2532827 2532827 2532827 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 253287 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 25327 2		
	STANCE FT)	9.96 0.10 17 18 18 23 23	. 24 . 94 . 05 . 05	112 112 112 112 112 112 112 112 112 112	51 51 51 51 51 51 51 51 51 51 51 51 51 5	
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JULY 9 1984 CROSS SECTION 3 TABLE C-3 SANGAMON RIVER

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TABLE C-3 SANGAMON RIVER CROSS SECTION 3 JULY 9 1984 - (CONTINUED)

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V (FPS)	00	. 16	.10	.12		5.6		90.				82	- C	0.4	.03	.03	.04	.02	.03	- , 06	. 02	.03	- 13	07	.02	.03	03	- 07	- 12	10	.02	.03	.03	.03	- , 05	60'-	02	.02	03	.03	.02	03	08	.01	.02	. 02	.02
DEPTH (FT)	đ	. 76	, 96	49	62.	5		20.1					4	. 77	. 95	29	. 63	. 28	. 63	1.48	. 26	64	1.56	1,46	. 24	64	1.05	1.40	1.54	1.48	. 22	. 65	. 83	.91	1.26	1.41	. 18	. 33	.47	. 59	. 70	1,09	1.34	12	03	. 08	- 04
DISTANCE (FT)		51.33	51.35	54,63	54,65	57,39	24.70	57.44 00.00	29,0C		52.04 00.00	60' 79	64.11	64.16	64.19	65,99	66.11	67.09	67.23	67.35	68.17	68.34	68.47	68.47	69.24	69.44	69 53	69.57	69.58	69, 59	70,29	70.54	70,60	70.62	70,68	70.69	71.34	71.48	71.56	71.63	71.67	71.78	71.81	71.82	72.08	72.31	72.37
K (FPS)		- 04	02	04	.17	02	- 04	- 04	. 20 -	. 03	- 02	02	04	-, 05	- , 03	01	- 02	100		90	90		6 -	 	90			3	5.6	200	50	- 05	50	01	2 0	05	05	.31	.01	11	50.1		- 04	50	45	02	17
V V V (FPS) (FPS)		- 23 - 04	- 02	. 44	. 78 17	29 02	35 04	1204	13 - 03	.51 .03	4002	1102	-,5804	2605	03	62 .01	- 30 - 02		- 47 - 05										10. 10. AO			- 42 - 05	- 22 - 05	- 01	- 30 - 05	- 22 - 05	- 1305	. 55 31	. 01	34 11	- 10	31	- 10 - 04	16	41 45	16 .05	.29
DEPTH V W (FT) (FPS) (FPS)		1.39 - 23 - 04 • 26 - 17 - 04		. 87 44 04	.67 .78 .17	1.632902	1.393504	1.171204	1,04 ,13 -,03	. 69 . 51 . 03	1.674002	1.541102	1.395804	1.182605	1.050903	an 62 01	1 54 - 30 - 02		1.10											-0/ -10 -0 -0 -0 -0 -0 -0 -0 -0 -0 -0 -0 -0 -0		1 48 - 42 - 05	1 28 - 22 - 05	1 03 14 - 01	1.473005	1 38 - 22 - 05	1.27 - 1305	60 55 .31	16 . 01	A2 34 11		70 31 13				93 16 05	74 . 29 . 17

TABLE C-3 SANGAMON RIVER CROSS SECTION 3 JULY 9 1984 - (CONCLUDED)

K (FPS)	17	4	- 17	21	21	- , 23	. 13	12	10	21	- 22	- , 09	.08	19	24	25	12	- 23	16	- 11
V (FPS)	02	02	- 02	. 04	- , 06	- , 06	. 02	01	. 02	- , 03	- 05	0.00	.02	02	05	04	0,00	02	- 01	01
DEPTH (FT)	22	37	. 92	1.05	1.21	1.24	. 20	. 76	. 06	.97	1.15	. 64	. 05	. 82	1.11	1.03	. 93	. 82	.64	.51
DISTANCE (FT)	72.49	72.62	72.88	72.92	72.95	72.96	73.10	73.16	73.20	73.23	73.27	73.45	73.50	73.52	73.60	73.93	74.25	74.55	74.83	75.08

TABLE C-4 SANGAMON RIVER CROSS SECTION 4 JULY 10 1984

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COMPUTED SECONDARY VELOCITY

(FPS)	.17	.21	. 12	. 13	32	35	38.	18	<u>61</u> .	42	. 16	. 33	. 53	. 35	. 18	. 29	. 20	. 38	. 30	. 23	. 42	.32	. 30	.45	.01	54	. 46	- , 32	. 35	. 53	. 38	- 34	. 68	37	.4	0.00	1.72	. 46	- , 40	1.48	. 52	43	0.00	.61	48	0.00	- , 54
V (FPS)	88.	. 76	. 65	.51	. 67	56	. 46	. 82	. 29	. 39	. 25	. 70	.37	. 61	.21	. 92	. 18	. 53	. 80	16	46	. 70	. 88	08	. 30	.26	. 39	77	. 62	. 34	24	- , 68	.31	- , 59	.47	.37	. 48	.41	51	.20	. 35	44	.31	. 30	37	. 25	30
DEPTH (FT)	1.78	1.53	1.77	1.77	1.48	1.44	1.37	1.90	1,67	1.30	1.62	1,86	1.19	1.82	1.56	2.21	1.47	1.76	2.18	1.37	1.69	2.14	2.31	36	13	1.22	1,61	2.28	2.10	1.49	2,04	2.24	1.33	2.20	1,98	. 13	1.09	1.89	2.14	.67	1.78	2.07	. 13	1.63	1.97	. 13	1.84
DISTANCE (FT)	23, 14	23.15	34.03	35.73	37,44	39.14	40.84	40,85	40.85	42.54	42.55	42.56	44.23	44.26	44.26	44.27	45,95	45.97	45.98	47.65	47 67	47.68	47.69	49.11	49.26	49.34	49.37	49,39	49,39	51.06	51,09	51,10	52,75	52,80	52,80	53, 18	54.41	54.50	54.51	55,98	56,19	56.21	56.55	57,88	57.91	59, 59	59,61
K (FPS)	- 03	80.	22	14	. 06	.10	10	80	. 22	.21	.03	. 25	.40	13	06	27	22	53	61	0			41		30	64	28	20	0	. 15	. 25	. 34	. 45	. 23	. 29	. 56	. 48	. 24	. 33	23	. 29	.27	. 20	. 25	. 23	. 18	. 22
V (FPS)	ġ		01	.02	.07	.08	- , 02	.11	60'	.07	. 13	11.	10	14	16		0	2 =				2		- C	00		53	24	12	.07	.27	. 22	. 12	.31	. 25	.24	. 30	. 34	. 28	39	32	. 37	. 52	. 44	. 52	. 72	. 62
DEPTH (FT)			1.38	1.17	. 74	. 56	1.61	. 24	. 79	1.16	15	44	. 81	91	4	08	95.1	22.1	- F			90.1	201				1 56	1.16	5 F -	34	1.65	1.31	. 71	1.72	1.42	. 94	1.10	1.57	1.22	1.63	1.30	1.37	1.71	1.43	1.47	1.76	1.50
DISTANCE (FT)		40.0	0. 30 7.2	5.88	6, 35	6,68	7.80	7,98	8.27	8,81	9.45	9.52	9.84	11.06	11 06	1 10	1 0 0	11 10	at c -	10 01		12 00		10 01	13.25	13,80	14 60	14.66	15.02	15.09	15.44	15.48	15,64	16.29	16,32	16.40	17.21	18.01	18.04	18.87	18.89	19.73	20.57	20.59	21.44	22.28	22,29

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(FPS)		
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DEPTH (FT)	2 - 2 2	
DISTANCE (FT)	8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8	> · · · ·

FABLE C-5 SANGAMON RIVER CROSS SECTION 5 AUGUST 5 1985

TABLE C-5 SANGAMON RIVER CROSS SECTION 5 AUGUST 5 1985 - (CONCLUDED)

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W (FPS)	00	15	22	00				- 26	000		10						60 · -		20. -		2																											
V (FPS)			50.									8 C	- C	0.0	5.0	5	6.0	5.0		0.00	- · ·																											
DEPTH (FT)		- 60	1 6	20					04.			a/ .			55	. 81	.91	.63	5.		1.34																											
DISTANCE (FT)		66.62 CC 74		00 AU	61.13	67.33	67.85	67.88	67.91	67,93	67,96	67,98	69, 29	69.37	69.48	69.49	69.59	69.61	69.75	69.78	69,94																											
W (FPS)		Ξ	02	30	- 03	02	0.00	. 55	. 25	10	0.00	.03	.45	. 22	.04	40	. 20	60.	. 36	19	.10	=	. 33	. 19	.17	15	. 32	.31	. 23	.43	. 15	32	41		- 6	62.	5	07.	. 13	. 35	. 28	12	18	17	- 20	. 24	30	.27
V (FPS)		.27	. 16	. 39	8	60.	.02	. 28	. 24	. 18	60.	. 13	. 19	. 17	Ξ.	. 14	. 13	=	=	Ξ.	-10	14	60.	.10	. 13	. 13	.08	. 13	. 12	. 12	.07	.06	80.	6.6	6.	60.	20.	.07	02	.07	.07	- 02	06	07	13	2 0.	. 06	. 06
DEPTH (FT)		1.07	1.23	.87	1.49	1.37	0.00	. 60	60.	1.11	1.41	1.27	. 62	. 92	1.32	.63	. 95	1.18	. 63	. 98	1.23	1.42	.64	1.02	1.28	1,39	.64	1.05	1.22	. 92	1.13	.32	. 69	66.	1.22	. 28	. 58	1.03	1.31	.67	1.06	1.41	1.55	1.62	1.74	.24	. 65	. 78
DI STANCE (FT)		47.96	47.97	47.97	47.98	47.98	50.50	50.70	50.73	50.75	50.76	50.76	52.93	52.98	53.01	54.87	54.93	54.95	56.63	56.70	56.74	56.75	58,28	58.37	58.40	58.41	59,85	59.96	59,99	61.48	61.53	62,66	62.93	63,05	63.11	64.24	64.63	64.81	64,89	65.51	65.73	65,83	65.86	65.87	65.88	65.88	66.43	66.54

TABLE C-6 SANGAMON RIVER CROSS SECTION 6 AUGUST 6 1985

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W (FPS)	 23	. 98	. 50	44	66	0.4	.20	. 40	80.	76.	55.		1 20	2.5		200	20	00.	0.0	29.	5.6		202	88	.02	. 76	. 93	. 88	1.01	.64	5 I		<u>5</u>	5	<u>.</u>	.33	.07	.51	0.00	.22	44	1.22	04	80.	.67
V (FPS)	38	. 87	. 28	. 34	1.04	. 79	.51	. 40	. 66	1.10	. 66	55. 10		0		5	2.	- 6	32	. 83	20.		26	31	. 33	.45	. 84	.37	. 69	. 19	.41	30	. 49	22.	. 32	. 36	. 34	41	60.	.31	. 29	32	.16	. 15	. 28
DEPTH (FT)	 1.94	1.70	1.46	1.56	1.76	.10	2.00	1,64	.10	1.58	1,78	1.30		1.20	. 30		- 40	40	1.34	1.53		87 I	07.1	76	0	1,31	1.53	1.18	1.45	22	.10	. 16	0.	0.	.10	. 16	. 12	Ξ.	.10	. 14	. 19	44	13	.16	.31
DISTANCE (FT)	22,40 22 48	22.49	22,51	24,09	24.08	25.07	25.66	25.67	28.05	28,86	30,44	32.05	32, 30	30.22		0/ · /]	0/./0	58.81	58,94	58,95	59.25	50.19 50.20	61 44	62.40	62.53	62,68	62.70	63,90	63.93	64.70	64.74	66.75	66.75	68,66	70.51	74.33	74.33	72.39	74.33	74.97	75.62	75.55	75.63	76.31	76.29
W (FPS)	98		- 28	17	- 17	- , 22	- , 86	79	-1.11	71	33	=:	. 14	-12	<u>.</u>	-1,24	24	14	- 94	- 33	. 45		111	2	80	60	- , 68	- , 76	. 32	- 68	- , 42	-1.01	97	- 83	- , 96	. 53	.01	- , 78	. 28	1.28	. 25	58	. 25	1.10	.03
V (FPS)	0.00	0.0		0.00	00.00	0.00	- , 05	04		0.00	02	.02	04	.02	0.4	- , 23	=	02	. 08	.17	. 24	.13	5.5	5.0	- 57	.37	03	06 -	. 73	-1,16	02	- , 29	19	- 43	- , 36	. 18	. 69	- , 56	.27	. 60	. 88	-1.02	. 32	23	.91
DEPTH (FT)	12		1 0	10	.26	.56	1.14	1.09	1.35	1.02	.81	.30	40	. 24	.31	1.62	.47	. 25	. 89	.71	. 55	. 15	52.1	00.		11.	1.08	2.10	Ξ.	2.15	1.25	1.54	1.37	1.71	1.59	1.45	.10	1.82	1.83	1.49	. 10	2.00	1.89	1.61	.10
STANCE (FT)	21	20	ה מי ע ק		99	75	96	8	51	69	87	97	8	33	35	40	40	69	73	91	4	02			- #	2.6	50	34	44	6 3	45	96	8	54	58	57	95	13	30	33	94	66	66	91	38

TABLE C-6 SANGAMON RIVER CROSS SECTION 6 AUGUST 6 1985 - (CONCLUDED)

W (FPS)	- 18 . 06
V (FPS)	0.00
DEFTH (FT)	.37
DI STANCE (FT)	77.02 77.75 80.81

TABLE C-7 SANGAMON RIVER CROSS SECTION 8 AUGUST 8 1985

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F H H H H H H H H H H H H H	PI PS)	51	5		61	33	.37	.36 1	.03	08	. 34	.89	.12	68	53	04	.06	.10	60.	.60	.90	.57	.12	00. 07		10	. 06	.11 2	96	60	 10	66	83	198	10	. 13	- 16 -	. 12	11	14	.15	. 95	
	H V (FPS) (FF					- 04	. 04	- 02 -	0.00		- 13	- 16 -	60	7	- 14			. 60	. 20	- 34	937	4 .04	.10		20 20			563 -1.	3 .35 .	· . 01	 			40	- 01	- 01	44	35	- 01	- 01	- 02	. 22	

TABLE C-7 SANGAMON RIVER CROSS SECTION 8 AUGUST 8 1985 - (CONCLUDED)

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(FPS)	8 8 9 9 9 9 9 9 9 9 9 9 9 9 9
(FPS)	8 2 8 8 2
DEPTH (FT)	800 10 10 10 10 10 10 10 10 10
DISTANCE (FT)	60.60 60.60 60.60 60.60 60.60 60.60 60.60 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60.00 60