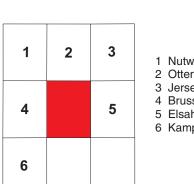




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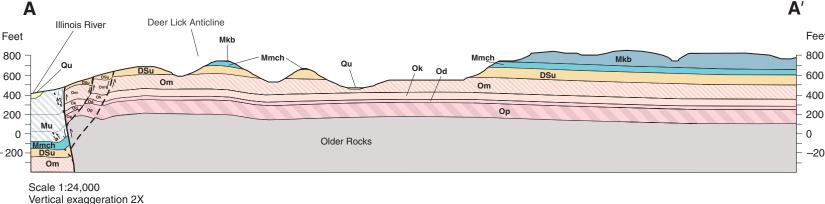
4 Brussels 5 Elsah 6 Kampville

ADJOINING 7.5-MINUTE QUADRANGLES

Released by the authority of the State of Illinois: 2002

UTM GRID AND 1995 MAGNETIC NORTH DECLINATION AT THE CENTER OF THE SHEET

This geologic map was funded in part by the USGS National Cooperative Geologic Mapping Program. It is one of a series prepared for the USGS 7.5minute Grafton Quadrangle (Illinois portion) by a multidisciplinary team of geologists from the Illinois State Geological Survey (ISGS). This series will haracterize surface landscapes and surface, bedrock, and engineering geology and will delineate coal and sand and gravel resources. This map was significantly improved through review, suggestions, and comments by the following individuals: Dennis R. Kolata (ISGS), David A. Grimley (ISGS), Fred Marshall (Principia College), Rodney D. Norby (ISGS), W. John Nelson (ISGS), Zakaria Lasemi (ISGS), B. Brandon Curry (ISGS), Jonathan H. boodwin (ISGS), and Tom Miller (IEPA). Digital cartography by Pamella K. Carrillo, F. Brett Denny, and Barbara J. Stiff. Photography by Joel M. Dexter.



ION		GRAPHIC COLUMN	THICKNESS (feet)		DESCRIPTION UNIT
m			0-100		A
is ne	Chert Breccia Zone		0-75		В
m one			0-90		С
aw le			0-90		D
(eokuk nes			190-200		E
			0-3	0	F
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<b>`</b>				40-55	Н
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)	Phosphatic zone		100 -140		N
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# Description

A. Alluvium: *clay, silt, sand, gravel, and cobble.* The upland sediments are composed of a mixture of clay, silt, sand, gravel, and cobbles. The clay, silt, and fine sand fraction is colluvium derived from loessal deposits that thickly mantle the upland areas. Most of the gravel originated from the underlying bedrock, but some glacially derived cobbles are basalt, granite, and metamorphic rocks weathered from diamicton that occurs above the bedrock. Small, reworked geodes filled with calcite and quartz crystals were observed. These geodes weather out of the Mississippian bedrock and are more abundant in the western half of the quadrangle. Alluvium in the bottomland is composed of a thick sequence of clay, silts, sand, and gravel that ranges from Pleistocene to Holocene. Slack-water lake deposits composed of gray, laminated silt with wood fragments were observed in some of the valleys at elevations near 450 feet (Grimley 1999).

B. St. Louis Limestone: limestone, limestone breccia, siltstone, and shale. Light gray to medium gray dense limemudstone with fossil wackestones. Part of the unit contains quartz sand and subangular limestone breccia clasts. Brecciation is attributed to ancient karstification of gypsum and anhydrite (Saxby and Lamar 1957). Oolitic grainstones, greenish oncolitic packstones, peloidal grainstones, stromatolitic boundstones, and carbonate intraclastic conglomerates make up a highly variable mix of microfacies. Acrocyathus floriformis, a colonial coral, occurs in the upper part of the basal portion of this formation. A. *floriformis* is widespread near the base of the unit. Yellowish dolostone beds are also present in this formation. Gray to dark gray chert occurs as nodules and stringers. Siltstones are calcareous and greenish. The shales are greenish gray and reddish brown, calcareous, soft, and non-fissile. This unit was only exposed along a fault slice in the extreme southwest corner of the Illinois portion of the quadrangle.

C. Salem Limestone: limestone, dolomite, chert, and siltstone. Limestones are tan-brown to light gray and contain laminated tidallites, wackestones to grainstones composed of rounded and broken fossils and coated grains. Bedding styles range from tabular to undulatory. Cross-beds are present in grainstone facies. The unit has a dirty gray-brown grainy appearance. The diagnostic character of this formation is alternating beds of laminated, fine-grained (calcisiltite) facies with coarse bioclastic, peloidal to oolitic grains in shoalingupward cycles. Dolomites are brown and have moldic porosity. Cherts are light gray and may be bioclastic and weather with a porous rind. Cherts occur between grainstones and laminated beds as elliptical nodules containing concentric rings that spall off like egg shells when weathered. Siltstones are brown to light gray and thinly bedded, typically less than 1 inch thick. Oolitic beds are rare. The foraminifera, *Globoedothyra baileyi*, is an index fossil for this unit. Other microfossils include calcareou algae, conodonts, and ostracodes. Fossil invertebrates include spiriferid and productid brachiopods, rugose corals, conularids, and crinoids. Ramose, fenestrate, encrusting, and bifoliate bryozoans are also present. The contact with the underlying unit

is gradational. **D.** Warsaw Shale: dolomitic limestone, siltstone, and *mudstone*. Medium-gray, crinoidal, bryozoan wackestones and packstones that contain a few brachiopods. In the limemudstone beds Archimedes sp. are preserved with coil and fronds attached. Dolostone beds are gray-brown, thinly bedded and contain chlorite-rich shale clasts with some small quartz geodes. The upper half of the unit is dominated by shaly limestone and dolostone beds. The lower half contains bluish gray mudstones up to 20 feet thick interbedded with thin limemudstones. Conularids and gastropods occur in the shaly portion of this unit. Siltstones are calcareous and fossiliferous and thinly bedded in the lower part. Quartz geodes are common in the shaly sequences. The geodes weather out and are very common in the drainages in the lower portion of the unit. The basal contact is poorly exposed but thought to be sharp and conformable with the underlying carbonate beds.

E. Burlington and Keokuk Limestone: limestone, siltstone, and shale. Light gray to white crinoidal grainstones dominate and are interbedded with nodular and bedded light gray to black

bioclasts of crinoids and brachiopods. Sandy limestones weather light brown, are cross-bedded, and contain brachiopod and crinoid molds. The unit is characterized by alternating layers of light gray to white crinoidal grainstones with beds of argillaceous and sandy limestones. This cyclic sequence of crinoidal limestone over sandy cross-bedded limestone is common in the lower part of the unit. Large spirifers are common along with crinoids, bryozoans, and corals. Siltstones are dark gray with a greenish tint and are calcareous. Calcite and quartz-filled vugs from 0.5 to 2 inches in diameter were observed. The unit is weathered on the upland surface where cherty residuum is 20 feet thick. The unit is conformable with the underlying unit.

cherts. The cherts are white when weathered, and some have

F. Fern Glen Formation: limestone, siltstone, and shale. On the eastern side of the Illinois portion of the quadrangle, the limestone is greenish gray, thin-bedded, and argillaceous; it contains small calcite geodes and crinoid stems. Green and red shaly calcareous siltstones are diagnostic and are well exposed on the river bluffs near Chautauqua, Illinois. The cherts are greenish gray, nodular, and fossiliferous. In Dagett Hollow and westward, the unit is dominantly thin, irregularly bedded, lime mudstone with cherty, crinoidal wackestone and packstone facies; these facies are indistinguishable from the lower part of the overlying unit E. Yellowish dolostone facies are also present. The basal part is gradational with underlying formation.

**G.** Meppen Limestone: *dolomitic limestone*. Light gray to tan, massive dolomite limestone, containing small (0.5 to 1.0 inch diameter) calcite geodes, light gray chert nodules, and locally calcareous siltstones. The unit is less than 12 feet thick and normally forms a small resistant weathered face that is fairly well exposed in most drainages. A minor unconformity exists between this unit and the underlying formation. On the river bluff just west of Chatauqua, an angular relationship between the underlying unit and this formation can be observed (fig. 1).

**H.** Chouteau Limestone: *limestone and siltstone*. Light brown to greenish gray irregular to wavy, thin beds of lime mudstone with thin beds of silty dolostone. Calcite geodes with diameters from 0.5 to 2 inches are common. Some of the calcite geodes are replaced with quartz. Chert nodules are locally abundant and typically are dark gray with light gray rims. Crinoidal wackestones containing fenestrate bryozoans and brachiopods occur in southwest dipping beds near Chatauqua. The unit appears to be gradational with the underlying unit.

**I.** Hannibal Shale: *shale, mudstone, and siltstone.* The upper portion may interfinger with the overlying argillaceous limestone Non-fossil erous gray fissile silty shale to laminated siltstone having brown iron oxide and manganese on fractures of weathered surfaces. The lithology is dominated by a soft, greenish to light gray mudstone, silty at base and fines upward to a non-fissile mudstone that fractures conchoidally. A thin, black, silty shale near the base of the unit (NE1/4, NE 1/4) Section 1, T6N, R13W, had a very strong petroliferous odor. Typically, the mudstone is not well exposed; however, good sections were observed at the head of Graham, Dagett, Distillery, and Jerseyville hollows. On the western side of the quadrangle, this unit is conformable with the underlying unit.

J. Glen Park Formation: *limestone*. Identified only in the eastern portion of the quadrangle. The limestone is poorly exposed and was only identified at two locations. At one location it was composed of an argillaceous lithographic limestone with small concretions (1/8 inch diameter) of pyrite and some glauconite. At the second location it was composed of a fossiliferous and oolitic limestone. The fossils were mostly chonetid and spiriferid brachiopods with crinoid fragments in an organic-rich reddish brown shale oolitic grainstone. The limestone interfingers with the overlying the dominant fossils are strophom shales. It is unconformable with the underlying unit.

**K.** Cedar Valley Limestone: *limestone and sandstone*. Thin and discontinuous fossil packstone with quartz sand. The lowest alternating fossiliferous shales and unit is a brownish gray sandstone overlain by fossiliferous and base. sometimes argillaceous limestone. It is gray where fresh and

## Data Type

Ouaternary

Devonian

and

Silurian

Ordovician

Line symbols are solid where observed, dashed where inferred, dotted where concealed

dashed wh	iere inferred, dotted whe
	Contact
• •	Fault: bar and ball on downthrown side
	Landslide failure plan
	Strike-slip fault: arrow direction of movemen
<u>A A</u> '	Line of cross section
400_	Structure contour top top of the Cedar Valle

le failure plane ip fault: arrow indicates of movement cross section contour top of the op of the Cedar Valley/Joliet

Formation (contour interval 20 feet) Anticline: direction of plunge indicated by the arrows

Strike and dip of bedding: number indicates degree of dip

Horizontal bedding

- Vertical joints
- Water well: number indicates depth of boring
- (feet)
- $-\phi_{50}$ Oil well: number indicates depth of boring (feet)
- Engineering or ISGS boring:
- letters indicate ISGS number
- Abandoned quarry Location of photograph
- Elevation (feet) of benchmark or survey point

## **Cross Section A-A' Looking Northwest**

## **Economic Geology of the Grafton Quadrangle**

Several quarries once mined Silurian dolomites in the quadrangle. Currently, none of these operations are active. According to local residents, most of the quarries were operated for local supply of aggregate and building stone.

The Burlington and Keokuk Formations in the Grafton area are nearly identical in lithologic character and were mapped as a single unit. Both units are composed in part of calcium-rich limestone. Portions of these units contain white crinoidal grainstones to packstones, which commonly are high in calcium carbonate. Hindering the quarrying in this unit are the cherty intervals located above and below the high-calcium zones. Relatively thick chert-free beds of economically important limestone are present locally.

### Oil

Two oil tests have been drilled in the quadrangle. According to an oil well report written by consulting geologist Lawrence Bengal, the first well had a show of oil, and the second well drilled in 1984 (Section 2, T6N, R12W) was interpreted by the geologist as intersecting a fault and repeating the Devonian and Silurian units (Bengal 1984). This interpretation would require at least 150 feet of vertical displacement. No field evidence to support a fault of this size was identified. The chance for economic oil production in this quadrangle is marginal because of the shallow depth to the pay zones along the anticlines and because of faulting in the area along the Cap au Grès. Neverthless, the shales and the limestone of the Ordovician Decorah contain hydrocarbons. Qualitative distillation tests in the area have reported the Decorah to produce between 15 to 25 gallons of crude oil per ton (Rubey 1952).

Sand and Gravel Sand and gravel are available in the alluvial deposits of the Illinois and Mississippi Rivers (see Grimley 1999).

Structural Geology of the Grafton Quadrangle The major structural feature of the quadrangle is the Cap au Grès Faulted Flexure (Keyes 1894). The Cap au Grès is the southeastern extension of the Lincoln Fold, which extends over 165 miles into northeast Missouri (Nelson 1995). The axis of the Lincoln Fold follows a general northwesterly trend but turns easterly at its southernmost exposures. The south-easternmost portion of this structure in Missouri and Illinois is called the Cap au Grès. In this quadrangle, the Cap au Grès is a faulted monocline with dips averaging less than  $30^{\circ}$  to the southwest and less than  $4^{\circ}$  to the northeast. The faulted blocks strike N80° W and dip between  $40^{\circ}$  to  $80^{\circ}$  to the south. The fault in this quadrangle juxtaposes Silurian age dolomites with Mississippian carbonates. Geologic reconstruction of the flexure indicates as much as 950 feet of vertical displacement may be present along this fault zone at the Deer Plain Ferry location. Evidence for the structure can be observed in road cuts along Illinois Highway 100 just west of Graham Hollow north of the Deer Plain Ferry landing. At this location the Mississippian St. Louis Limestone dips to the south up to  $70^{\circ}$ .

Along the bluffs west of Grafton, several large blocks of Silurian dolomite are exposed that are interpreted to be rotational slumps. Several more slump blocks were observed along a drainage on the west side of Graham Hollow. The basal failure plane of these slumps occurs in the underlying Maquoketa Shale. The failure plane was not observed because of the lack of bedrock exposures of the Maquoketa in the immediate area of the slump blocks. Rubey (1952) mapped the Hardin and Brussels 15-minute quadrangles to the west of the Grafton Quadrangle. He identified an oval-shaped uplift located in Deer Lick Hollow just west of the Grafton Quadrangle, which he named the Deer Lick Dome. We traced this feature into the Grafton Quadrangle and determined that it was more accurately described as a

plunging anticline. The Deer Lick Anticline is closely related to the Dome described by Rubey and may be considered the eastern limit of the Deer Lick Dome. A second anticline was identified on the structure contours to the northeast of the Deer Lick Anticline. This second anticline roughly parallels the Deer Lick Anticline but the dips of both limbs are less than 4°. This anticline also plunges to the southeast where it is concealed by the

alluvial sediments of the Illinois and Miss Dome observed on the Bedrock Geologic Minute Quadrangle and Report (Harrison tion of the second unnamed anticline in the Two northeasterly trending strike-slip faul half of the quadrangle. The first was locat The fault zone was 10 feet wide and conta and secondary calcite veinlets. No vertica Silurian dolomite. The fractures were near mullion-like planes, which had general tre movement was probably horizontal, but the right or left was not readily apparent. A se fied along Babbs Hollow (1,100 feet WL, R11W). The fault zone was less than 5 fee central breccia zone down-dropped less t N30° E to N40° E and is located in the bas Keokuk Limestone. Indications are that th lateral. A third fault is concealed under all was inferred from the 10 to 15 feet elevation Devonian and Silurian units across the Ho

These faults and the anticlines suggest a n principal stress direction. The most plausil Grès feature was postulated by Rubey (19 Lumm (1985), and Nelson (1995) who di seated reverse basement fault. The Cap au folds found on the Colorado Plateau that f overlying reactivated basement faults (Han (1985) compared the Cap au Grès flexure Rocky Mountains and Colorado Plateau, overlie faults in the Precambrian crystalli The timing of the Cap au Grès/Lincoln Fo stratigraphic relationships. Facies variatio Kinderhookian through lower Valmeyera near the Cap au Grès indicates that these l apparently thin toward the structure. There was active starting in Late Devonian and c earliest Pennsylvanian. Outliers of Pennsy occur at only slightly different elevations in Calhoun County (Rubey 1952). The St. unit significantly displaced, which indicat place on the structure between Valmeyera The eastern quarter of the quadrangle is r Cap au Grès and has a regional easterly dir

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1:24,000.

Rubey, W.W., 1952, Geology and mineral Brussels Quadrangles (in Illinois): Un Professional Paper 218, 179 p. Saxby, D.B., and J.E. Lamar, 1957, Gyps Illinois State Geological Survey, Circu

weathers to a brown tint and cont Paraspirifer sp. brachiopods, rug gastropods. In the quarry east of was found in the sandy limestone unconformable with and may dow underlying formation.

L. Joliet Formation: doloston dolostone is yellowish brown to b weathered surface; sucrosic textu common. The upper part of the d overlying strata. Bedding planes beds are typically several feet thi Dagett Hollow the upper surface cracks. Chert occurs as nodules s unit. The thin shales have a gree Sthenarocalymene celebra is con Grafton. A cheirurid trilobite was part of the dolostone. Dolomitiz formations within the Silurian ma separately difficult. This unit is un underlying units.

M. Kankakee Formation: dol dolostone is yellowish brown to b characteristic pitted weathered sur from thick to thin. Bedding plane are separated by thin green clay la nodules sporadically throughout gray tint. Glauconite is present w include brachiopod molds, straig This unit is unconformable with places.

N. Edgewood Limestone: da dolostone is brown to buff gray. I argillaceous, and sandy in places. sporadically throughout the unit. tint. Glauconite was present in a f were observed. This unit is uncon units

**O. Maquoketa Formation:** *a* This unit is poorly exposed and f well vegetated. The lower part of and grades upward into bluish gr interbedded with bluish gray muc buff-gray to greenish gray and ha hales. A thin black shale having fragments, and abundant dissen the base of the unit. This black sh location in Section 13 (SW NF formation was not observed, but unconformable with the underly Subsurface only (described from a

P. Kimmswick Limestone–Tre subsurface: limestone, dolostone gray, coarsely crinoidal grainston formation. Other fossils include R Isotelus gigas (trilobites), brachie commonly broken in the cross-be formation. Shales are calcareous are not very common and are wh When cracked, the limestones ha The basal contact is a distinct ha

Q. Decorah Formation: lime to greenish limestone or lime mu

R. Plattin Limestone: limesto brown to chocolate brown sublith

Referen Bengal, L.E., 1984, Well report on the O'

Illinois: Illinois State Geological Surv Grimley, D.A., 1999, Surficial geology m Portion), Jersey and Calhoun Countie Survey, Illinois Geological Quadrangle

Harrison, R.W., 1997, Bedrock geologic r Minute Quadrangle and report: United

Miscellaneous Field Studies, 22 p. Keyes, C.R., 1894, Crustal adjustments in Geological Society of America Bulleti Nelson, W.J., 1995, Structural features in

Survey, Bulletin 100, 144 p. Nelson, W.J., and D.K. Lumm, 1985, Ste and Illinois: U.S. Nuclear Regulatory

*ucrospirifer* sp. and als, and platycerid , an arthrodire tooth indy limestone is everal feet into the

inor shale. The y and has a pitted molds of fossils are is truncated by to wavy in places, and an be thinly bedded. In s polygonal mud lly throughout the tint. The trilobite the quarries east of llected from the upper his and lower pping the Silurian units mable with the

*ind shale*. The y and has a olostone beds range it to wavy in places and Chert occurs as Shales have a greenish s formation. Fossils plopods, and trilobites. rlying units in some

nd minor shale. The thick to thin, occurs as nodules have a greenish gray he beds, and few fossils le with the underlying

siltstone, mudstone. ntle hill slopes that are nation is calcareous a calcareous siltstones The upper part is shaly iminated silts and atic nodules, fossil rite was identified near only observed at one 3W The base of this s indicate that it is

s and reports).

mestone in the *inor shale*. White to dominant facies in this culites sp., Illaenus sp., nd gastropods that are arse bioclastics of the y contain pyrite. Cherts slight yellow tones. nt petroliferous odor. d omission surface.

*l shale*. Light brownish nterbedded with herts are dark gray, and achiopods.

stone, and shale. Light ic limestone with limestones near the

Rivers. The Florissant *f the St. Louis*  $30 \times 60$ appears to be a continuaton Quadrangle.

e identified in the eastern ie bluff just east of Grafton reccia fragments, gouge, could be observed in the ical with smooth wavy N50° E. The direction of e of displacement to the trike-slip fault was identiet SL; Section 7, T6N, and appeared to have the et. The fault zone strikes tion of the Burlingtonement is probably right at Rice Hollow. This fault erence on the top of the

st to southwest maximum planation for the Cap au arrison (1997), Nelson and ne possibility of a deepesembles monoclinal drape in sedimentary strata 1993). Nelson and Lumm aramide monoclines in the folds in sedimentary cover ment.

nt is constrained by the structure occur in the ssion. Detailed mapping Aississippian rocks *we* suggest that the structure ied sporadically through the n Carbondale Formation er side of the Cap au Grès Formation is the youngest major movement took Desmoinesian times.

ificantly influenced by the ughly 50 feet per mile.

l well, Jersey County, ublished report, County

fton Quadrangle (Illinois is: Illinois State Geological IGQ Grafton-SG,

the St. Louis  $30 \times 60$ Geological Survey, per Mississippi Valley:

p. 231–242. : Illinois State Geological

ieve Fault Zone, Missouri ission, 1985–3. ces of the Hardin and ates Geological Survey,

anhydrite in Illinois: 5, 26 p.