

Base map compiled by Illinois State Geological Survey from digital data provided by the United States Geological Survey. Topography by photogrammetric methods from aerial photographs taken 1963. Field checked 1966.

North American Datum of 1927 (NAD 27)
Projection: Transverse Mercator
10,000-foot ticks: Illinois State Plane Coordinate system, east and west zones (Transverse
Mercator)
1,000-meter ticks: Universal Transverse Mercator grid system, zone 16

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BASE MAP CONTOUR INTERVAL 10 FEET NATIONAL GEODETIC VERTICAL DATUM OF 1929

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ILLINOIS DEPARTMENT OF NATURAL RESOURCES



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Geology based on field work and data analysis by W.J. Nelson and J.A. Devera, 1994-1995; revised 1997 and 2005.

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IGQ Mt. Pleasant-BG Sheet 1 of 2

Structural Geology

The Mt. Pleasant Quadrangle is located along the southern margin of the Illinois Basin, east of the Ozark Dome and a few miles north of the Mississippi Embayment. Bedrock strata at and near the surface regionally dip northeast toward the center of the Basin. The average dip in the Mt. Pleasant Quadrangle (shown by structure contour lines on the map) is about 140 feet per mile or $1\frac{1}{2}^{\circ}$. Local attitudes, indicated by strike and dip symbols on the map and by variable dip of mapped contacts, commonly deviate from the regional dip. Local dips may result from differential compaction around lenticular rock bodies, soft-sediment deformation caused by slumping and loading, solution collapse, and other non-tectonic causes.

Five small faults that strike north-south to slightly east of north have been mapped. All have the east side downthrown. No fault surfaces are exposed; all faults are inferred from indirect evidence. (1) Near the north edge of the quadrangle (SE¹/₄, Sec. 1, T12S, R1E), marker beds in the Clore Formation are displaced 20 to 30 feet down to the east. North-south jointing is prominent near the inferred fault and in line with it to the south. (2) On the south side of the Cache River along the border of Sections 14 and 23, T12S, R1E, sandstone of the Waltersburg Formation exposed in a roadbed dips 35° east. A linear north-south valley is just east of the dipping sandstone outcrop. (3) On the north side of Buck Run just east of the center of the quadrangle, the Waltersburg Formation is juxtaposed with the Tar Springs Formation and Glen Dean Limestone on the west. The indicated displacement is as large as 100 feet. Sandstone near the inferred fault contains closely spaced orthogonal joints that strike north-south and east-west. Weller and Krey (1939) previously mapped this fault. (4) On the south side of Buck Run, a fault is inferred along the boundary between R1E and R2E. Evidence includes a linear valley, well-developed north-south jointing, and displacement of the Tar Springs-Glen Dean contact 30 to 40 feet down to the east. (5) Southwest of the center of the map, the top of the Cypress Formation is at least 50 feet higher on the west side of a linear valley. Steeper than normal dips (4° to 6° northeast) were measured on sandstone outcrops near the northeast end of this valley.

Additional north-trending faults occur in the nearby Makanda (Jacobson and Weibel 1993), Anna (J.A. Devera and W.J. Nelson, unpublished map), Vienna (Nelson et al. 2004), Creal Springs (Trask and Jacobson 1990), and Jonesboro Quadrangles (Devera and Nelson 1995). All of these are highangle normal faults that displace Pennsylvanian and older rocks. Striations, drag folds, and fracture patterns imply dip-slip movements. There is no indication that any north-south faults have been active since late Paleozoic time.

A northeast-trending fault is inferred to underlie Quaternary alluvium along Buck Run in the east-central part of the quadrangle. Our interpretation of this fault is nearly the same as that of Weller and Krey (1939). The Glen Dean Limestone and Tar Springs Formation southeast of Buck Run are uplifted roughly 100 feet relative to the Waltersburg Formation and

Vienna Limestone northwest of the stream. This fault is in line with a small northeast-trending fault near the west edge of the adjacent Vienna Quadrangle (Nelson et al. 2004).

The Mt. Pleasant Quadrangle lies along a northeast-trending zone of magnetic and gravity highs known as the Commerce Geophysical Lineament (CGL), which extends from northeastern Arkansas to near Vincennes, Indiana. In southern Illinois, a series of small, discontinuous northeast-trending faults line up with the CGL. The fault along Buck Run is one of these faults. The CGL has been interpreted as a series of dense, highly magnetic igneous bodies possibly emplaced during Precambrian time along a rifted margin or plate boundary (Langenheim and Hildenbrand 1997; Hildenbrand et al. 2002). Holocene faulting has been documented along the CGL in Missouri just across the Illinois border, and the CGL is implicated in ongoing seismic activity within Illinois (Harrison and Schultz 1994; Harrison et al. 1999; Hildenbrand et al. 2002).

Groundwater

In the northeast corner of the quadrangle (T12S, R2E) all water wells are relatively shallow (100 to 200 feet) and yield 1.5 to 3 gallons of water per minute. Aquifers are sandstone in the Palestine Formation and the Tygett Sandstone Member of the Clore Formation and limestone at the top of the Cora Member of the Clore Formation.

Most wells in the central part of the quadrangle were completed in sandstone in the lower half of the Cypress Formation. These wells commonly yield 30 to 60 gpm at depths as great as 600 feet. Some water from the Cypress reportedly is salty. Shallower sandstones in the Tar Springs and Hardinsburg Formations generally yield less than 10 gpm. A few wells were finished in limestone aquifers of the Menard, Glen Dean, and upper Golconda formations; these wells have outputs of a few gallons per minute.

In the southwestern part of the quadrangle, where the Cypress is at the surface or eroded, wells are drilled as deep as 870 feet in search of water in the Ste. Genevieve and St. Louis (not shown) Limestones. Some of these wells failed to find potable water; others were plugged back for shallow production of a few gallons per minute from the Paoli Limestone. The Aux Vases Formation apparently is not an aquifer in this area.

Water production from limestone is sporadic, depending on fracture systems and solution cavities. Large quantities of water may flow in solution channels, but such water is vulnerable to contamination from surface runoff.

Quaternary valley fill along the Cache River, Lick Creek, and Cypress Creek is dominantly silty clay. This material has low permeability and generally does not yield useful amounts of water.

Mineral Resources

Stone

No commercial stone quarries are active in the Mt. Pleasant Quadrangle. Several small, abandoned quarries in sandstone of the bluff-forming lower portion of the Cypress Formation were observed during mapping. These quarries probably furnished stone for local construction projects.

Lamar (1959) reported that the Menard, Glen Dean, Golconda, Downeys Bluff/Renault (Paoli) (not shown), and Ste. Genevieve formations contain commercially valuable limestone in southern Illinois. Most of these units contain numerous shale interbeds in the study area, rendering them undesirable for most purposes. The Glen Dean contains limestone as thick as 25 feet with little shale or chert and could support small quarries. The Glen Dean currently is being mined underground north of Chester, Illinois.

The best unit for quarrying in the area is the Ste. Genevieve Limestone. Large quarries in the Ste. Genevieve are currently active in the nearby Anna and Cypress Quadrangles. Stone from these quarries is used for aggregate, agricultural lime, and road surfacing. The Ste. Genevieve crops out at the southwest corner of the Mt. Pleasant Quadrangle and extends beneath the entire quadrangle in the subsurface.

Oil and Gas

Only two oil and gas test holes are on record in the quadrangle. The Henry Leschen No. 1 Schenker test, in the NE¹/₄ NW¹/₄ NW¹/₄, Sec. 14, T12S, R1E, was drilled to a total depth of 2,220 feet. The date of drilling was not recorded, no samples are on file, and the only log available is a driller's log that provides scanty geologic information. A show of oil was reported in limestone (Middle Devonian?) near the bottom of the hole.

The Maloney No. 1 Osman test hole was drilled in 1970 to a total depth of 815 feet in the St. Louis Limestone (Mississippian). The well site was SE¹/₄ SW¹/₄ SE¹/₄, Sec. 33, T12S, R1E. A gamma-ray log and two sample studies are on public file at the ISGS. The driller recorded "good shows of oil" in the Ste. Genevieve Limestone and "dead oil" in the Aux Vases Formation and St. Louis Limestone (not shown). The Osman well was dry and abandoned.

The nearest producing oil fields are in Williamson County, approximately 20 miles north of the study area.

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SYSTEM SERIES	GROUP	FORMATION	MEMBER	GRAPHIC COLUMN	THICKNESS (FEET)	UNIT DESCRIPTION	A Clay, silt, sand, gravel, and peat Valley fill along larger streams is brownish gray, mottled silty clay that contains scattered lenses of sand and a thin gravel layer at the base. Organic matter is common; peat lenses are as thick as 10 feet. These sediments represent the Cahokia Formation (Holocene), lacustrine deposits of the Equality Formation (Wisconsinan), and possibly older alluvium, but the Cahokia and Equality are difficult to differentiate. Alluvium along small upland streams is a mixture of clay, silt, and rock fragments.	 P Limestone Limestone is mostly brownish gray, medium- to coarse-grained, crinoid-bryozoan packstone and grainstone. Light gray, cross-bedded oolitic grainstone is common near the top. Thin beds of silty to argillaceous, dolomitic lime mudstone and wackestone are present. Most fossils are disarticulate. The lower contact was not observed. Q Shale with thin limestone interbeds The shale is medium to dark gray, greenish gray, and
QUATERNARY	CENE U	Cahokia Equality loess			0–65 0–20	A B	B Silt Silt is yellowish, brownish, and reddish gray, massive, and rooted. Loess blankets uplands; it is thickest near the Cache River and in the southwest corner of the quadrangle. At least three loesses are present: the Peoria Silt (late Wisconsinan), the Roxana Silt (early Wisconsinan), and the Loveland Silt (Illinoian). Loess is not shown on the man	olive-gray, platy and fissile, and slightly calcareous. It is mostly clay shale, but some is silty. Lime- stone is dark reddish to brownish gray, coarse, crinoid-brachiopod packstone and grainstone. A bed of very sandy limestone or calcareous sandstone 2 to 3 feet thick is widespread at the base of the unit. The lower contact is gradational.
PIAN		Kinkaid Ls. Degonia	Negli Creek Ls.		8 exposed 30–40	C D	C Limestone Limestone is dark gray, argillaceous lime mudstone (Dunham 1962) in hummocky beds 2 to 14 inches thick. Fossils include bellerophontid gastropods, small brachiopods, and <i>Chaetetes</i> .	R Sandstone, siltstone, and shale Fine- to medium-grained, thick-bedded to massive sand- stone at the base grades upward to thin-bedded and laminated siltstone and shaly sandstone at the top. The thin-bedded sandstone contains interference ripples, small load casts, and casts of plant stems and bark. Trace fossils are common, including <i>Lockeia</i> , <i>Aulichnites</i> , <i>Uchirites</i> , <i>Sclara</i> -
		Clore	Ford Station Tygett Sandstone		20–30 () () () () () () () () () ()	E F G H	D Siltstone, sandstone, shale, and mudstone Siltstone and very fine sandstone are olive gray to greenish gray and have planar laminations separated by shale partings. In adjacent quadrangles the Degonia contains shale that is dark gray to greenish and olive-gray and, at the top, an interval of red and green variegated, blocky mudstone. Only small outcrops of siltstone and shale were found in the Mt. Pleasant Quadrangle.	<i>tuba missouriensis</i> , and <i>Olivellites</i> . The lower contact is sharp and locally erosional. S Limestone with shale interbeds Limestone is mostly brownish gray skeletal packstone to wackestone that is thin- to medium-bedded and interlayered with soft greenish gray shale. Whole calices of <i>Pentremites</i> and <i>Phanocrinus</i> are common along with diverse brachiopods and bryozo-ans. The crinoids <i>Zeacrinus</i> and <i>Pterotocrinus</i> also are present. The lower contact is gradational
		Palestine	Cora		<u>හි</u> 20–55 47–75	J	E Limestone and shale Limestone is medium to dark gray (weathering olive-gray to yellowish gray) argillaceous lime mudstone that is massive to nodular. Only a few small exposures were found. Shale, exposed in adjacent quadrangles, is dark gray to olive-gray, calcareous, and fissile.	and intertonguing. T Shale with limestone interbeds Shale is gray to greenish gray, calcareous, and fossilifer- ous; limestone is similar to that in the overlying unit. This unit is poorly exposed, and well logs lack detail.
		Menard Limestone			108–140	L	 F Sandstone, Shale, and slitstone Gray slity shale at the base of the unit grades upward to laminated siltstone and thin-bedded sandstone in the middle. Thick-bedded sandstone at the top of the unit contains stigmarian root casts. In places, the upper sandstone cuts downward into the lower shale with an erosional lower contact. Cross-bedding in this channel-fill sandstone dips southwest. The lower contact is sharp. G Limestone and shale Two limestone beds 2 to 4 feet thick are separated by 4 to 5 feet of 	 U Mudstone and shale The upper part is blocky mudstone that is variegated green, red, and gray and is probably a paleosol. The lower part is gray shale containing lenses and laminae of sandstone. This unit is equivalent to the Big Clifty Sandstone Member. The lower contact is gradational. V Limestone and shale The limestone is mostly dark gray skeletal packstone and wackestone
		Waltersburg			41–88	M	shale. Limestone is dark gray argillaceous lime mudstone to skeletal wackestone that contains rugose corals, spiriferid brachiopods, and echinoderm fragments. Shale is olive-gray and contains lenses of siltstone and nodules of limestone. The lower contact is rapidly gradational.	that weathers yellowish orange. The shale is gray to greenish gray and olive-gray and calcareous. Limestone and shale both are highly fossiliferous: brachiopods dominate, but bryozoans, crinoids, blastoids, and rugose corals also are common. The lower contact probably is gradational.
		Vienna Ls.			5–29	N	H Shale, siltstone, and sandstone Dark gray to olive-gray clay shale at the base grades up- ward through ripple- and planar-laminated siltstone and sandstone in the middle, to thick-bedded sandstone at the top. The upper sandstone contains stigmarian root casts and the trace fossils	W Shale Shale is medium to dark gray and calcareous and contains fenestrate bryozoans. Thin lenses and interbeds of limestone are present. This unit is mostly concealed by alluvium along Cypress Creek
		Tar Springs			90–120	0	 Planolites and Rhizocorallium. The lower contact is sharp. I Limestone Limestone is dark gray, argillaceous lime mudstone to skeletal wackestone that weathers yellowish gray to olive-gray. The upper part has thin shale interbeds. Thin siliceous bands in the lower part stand out as parallel ridges on weathered surfaces. Fossils include spiriferid brachiopods, rugose corals, and bryozoans. The lower contact is sharp. 	X Limestone Limestone is medium to dark gray skeletal wackestone, packstone, and grain- stone. In outcrops near the center of the N½, Sec. 3, T13S, R1E, dark gray argillaceous wacke- stone at the base grades to light gray, coarse-grained crinoidal grainstone at the top. Brach- iopods and fenestrate bryozoans are common. The limestone forms wavy beds a few inches to 1½ feet thick. Sample logs from wells indicate that some of the limestone is sandy. The lower con-
		Glen Dean Limestone			02 -45 	PQ	J Shale with thin limestone interbeds Shale is dark gray to greenish gray and varies from fissile clay shale to blocky mudstone. Limestone is mottled gray to yellowish orange (when weathered) and highly fossiliferous; it contains trepostome and fenestrate bryozoans, echinoderm fragments, trilobites, and the brachiopods <i>Spirifer increbescens</i> , <i>Composita subquadrata</i> , <i>Diaphragmus alagana</i> , <i>Chapataa and and productide</i> . At the base a lasticular dark	Y Sandstone, siltstone, and shale The unit is mostly very fine-grained, laminated to thinly bedded sandstone that contains interbeds and laminae of siltstone and silty shale. Planar laminations, interference ripples, and ladderback ripples are present. Some layers are intensely biotur-
MISSISSIP	POPE	Hardinsburg			53–90	R	 <i>phragmus elegans, Cliothyridina</i> sp., <i>Chonetes</i> sp., and productids. At the base a lenticular, dark reddish gray to brown, very sandy limestone grades laterally to calcareous, burrowed sandstone. The lower contact is gradational and intertonguing on a small scale. K Sandstone, siltstone, and shale The unit commonly comprises two upward-coarsening sequences of roughly equal thickness. A rooted mudstone (paleosol) and thin coal occur locally. 	Z Sandstone Sandstone is white to light gray, very fine- to fine-grained quartz arenite, forming cliffs and ledges. The lower part has prominent low-angle wedge-planar and tabular planar cross-bedding, along with ripple and planar lamination. Paleocurrents were mostly toward the west and porthwest and less commonly toward the south and southwest. In one area, bidirectional north-

	Haney Limestone	30-60	S T	sequences of roughly equal thickness. A rooted mudstone (paleosol) and thin coal occur locally at the top of the lower sequence. Sandstone at the top of the upper sequence locally fills small channels; cross-bedding dips south or southwest. A thin ripple-laminated, calcareous sandstone at the base of the Palestine intertongues on a small scale with sandy limestone at the top of the Menard	northwest and, less commonly, toward the south and southwest. In one area, bidirectional north- east-southwest cross-bedding was observed. The upper 5 to 30 feet of the sandstone has perva- sive slumped and distorted lamination and "healed" planar fractures that strike N5-20°W and dip steeply. These features possibly were produced by earthquakes during sedimentation. The lower
Golconda	Fraileys Shale		U	L Limestone and shale The limestone is mostly medium to dark gray lime mudstone and skel-	AA Shale with thin limestone interbeds Shale is olive-gray to dark gray, platy to fissile clay
	Beech Creek	20–40	W	Bedding is tabular to hummocky, and most beds are a few inches to about 2 feet thick. Scattered beds of light gray crinoidal packstone and grainstone are present. Shale interbeds are dark gray	gray to reddish gray, shaly to sandy, coarse-grained skeletal packstone and wackestone. Con- glomerate composed of red, gray, green, and ochre mudstone clasts in a sandy limestone matrix
	Limestone	5–10 50–90	Y	to olive-gray and calcareous and contain limestone lenses. The uppermost 10 to 15 feet of the Menard consists of nodular-bedded limestone with many shale interbeds. The middle 40 to 50 feet is dominantly limestone, less argillaceous than above and containing fewer shale interbeds. The lower part of the Menard is covered by alluvium and is known only from well records.	was seen in several places. The Ridenhower (Butts 1917) is a shale-dominated unit that grades laterally to sandstone east of the study area and is equivalent to part of the Paint Creek Lime-stone west of the study area (Cole and Nelson 1995). The lower contact appears to be grada-tional.
Cypress		50–100	Z	M Shale and sandstone This poorly exposed unit is composed mostly of medium to dark gray, soft and platy clay shale. Sandstone at the top weathers orange-brown; it is very fine grained, thinly bedded, and contains horizontal burrows and trails. Locally, a thin sandstone occurs at the base of the Waltersburg.	BB Limestone and shale Limestone is variable in lithology. Light gray skeletal and oolitic grain- stone and packstone are most widespread. Yellowish to greenish gray crinoidal wackestone and packstone contain pink crinoid grains and rip-up clasts of red and green shale. Fossils include bryozoans, brachiopods, crinoids, and blastoids. Cross-bedded oolitic and crinoidal grainstone is at the base. Thin interbeds and lenses of greenish gray, calcareous and fossiliferous shale are
Ridenhower Shale		60–70	AA	N Limestone Limestone in the middle and upper part of the unit is largely dark gray lime mud- stone to skeletal wackestone that is very siliceous and cherty. It weathers to a porous, punky residuum with angular blocks of chert. Fenestrate bryozoan fronds are abundant both in fresh limestone and residuum. The lower part of the Vienna is composed of light gray crinoid-bryozoan	present. The Downeys Bluff Limestone, Yankeetown Formation, and Renault Limestone were not differentiated in this area. This unit is equivalent to the Paoli Limestone in Indiana (Droste and Carpenter 1990). The lower contact appears to be gradational and intertonguing.
Paoli Limestone		50–120	BB	packstone and grainstone, some of which is oolitic and cross-bedded. This lower limestone con- tains little chert and is not siliceous. The lower contact is sharp.	CC Sandstone and limestone Sandstone is gray to greenish gray, very fine- to fine-grained, and commonly glauconitic and calcareous. It is mostly thin-bedded and contains laminae of green and purple shale. Onlites, calcareous bioclasts, and burrows are prominent. Limestone is light to
Aux Vases		90–140	сс	lenses of thin, shaly coal. Below this is mostly ripple-laminated and cross-bedded sandstone that forms bluffs up to 30 feet high. Paleocurrent directions are highly variable, and tidal rhythmites are present. Trace fossils include <i>Planolites, Lockeia</i> , and <i>Conostichus</i> . The lower Tar Springs contains up to 50 feet of shale that is medium-dark gray with an olive or greenish cast, platy and fissile, and contains small limppite lenses. A lenticular conditione up to 10 feet thick is at the base	 stone intertongue in a complex fashion within this unit and at the contact with the underlying unit. Outcrops are few; descriptions are based partly on well records and outcrops in adjacent quadrangles. DD Limestone Limestone is white to medium brownish gray colitic and skoletal grainstone with
Ste. Genevieve Limestone		60 exposed; 190 total	DD	of the Tar Springs. Interference ripples, load casts, and ball-and-pillow structures are common; shale rip-up clasts and plant fossils (<i>Sigillaria</i>) also are present in the basal sandstone. The lower contact is sharp or rapidly gradational.	thinner interbeds of medium to dark gray, cherty lime mudstone and wackestone. Outcrops are few; descriptions are mainly based on well records and outcrops in adjacent quadrangles.

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