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# Dual wave farms for energy production and coastal protection

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# Dual wave farms for energy production and coastal protection

## Abstract

The synergetic application of wave farms, i.e., arrays of wave energy converters (WECs), for protecting the coast in addition to their main objective of generating carbon-free energy can place this renewable resource as a major element in ocean and coastal management. In particular, their ability to mitigate coastal erosion by reducing the amount of wave power reaching the coast will be put to use – and this is the motivation for this work. We propose a new approach in which the wave farm has a dual purpose: to generate carbon-free energy and to contribute to coastal erosion management. We illustrate this approach by means of a case study: a dual-purpose wave farm off Xago, a beach-dune system in Asturias (N Spain) subject to severe erosion – manifested dramatically in the retreat of the dune – and located in the area earmarked for the first wave farm in Spain. The objective of this work is to establish whether or not the wave farm may be useful to counter the erosion of the beach-dune system. To this end a wave propagation model is coupled with a state-of-the-art coastal processes model and applied to analyse the response of the system under storm conditions in two scenarios: with and without the farm. The efficiency of the wave farm in mitigating erosion is determined by comparing the results in both scenarios by means of a series of coastal indicators defined *ad hoc*. We find that the farm reduces storm-induced erosion particularly where it is most acute, in the dune front, and thus contributes to alleviate the current erosive trends. This opens up exciting possibilities of using dual wave farms in lieu of, or as a complement to, coastal structures or beach nourishment. As wave energy develops into a major renewable energy source in the coming decades, dual wave farms are poised to constitute a breakthrough in coastal erosion management.

**Keywords:** wave energy; coastal management; coastal erosion; beach morphology; sediment transport.

28 The current status of wave energy is similar to that of wind energy in the early 80's. With a vast  
29 resource and a very active R&D community, wave energy is expected to become a major renewable in  
30 the coming decades, with wave farms deployed in a number of coastal regions throughout the world  
31 (Bernhoff et al., 2006; Cornett, 2008; Folley and Whittaker, 2009; Guedes Soares et al., 2014; Iglesias  
32 and Carballo, 2009, 2010b; Pontes et al., 1998; Veigas and Iglesias, 2013, 2014; Vicinanza et al.,  
33 2013). In previous work it was established that the extraction of wave energy by a nearshore wave  
34 farm results in a milder wave climate in its lee (Carballo and Iglesias, 2013; Iglesias and Carballo,  
35 2014; Mendoza et al., 2014; Millar et al., 2007; Palha et al., 2010; Ruol et al., 2011; Smith et al.,  
36 2012; Veigas et al., 2014a; Veigas et al., 2014b; Vidal et al., 2007; Zanuttigh and Angelelli, 2013).  
37 The scientific hypothesis of this work is that this reduction in wave energy can be used for coastal  
38 erosion management, in particular in the case of a beach-dune system. To test the hypothesis, a case  
39 study is carried out on Xago Beach, in the area proposed by FAEN (*Fundación Asturiana de la*  
40 *Energía*, Asturian Energy Foundation) for the deployment of the first wave farm in Spain. Previous  
41 studies of relevance for this work include the characterisation of the wave resource in the region  
42 (Iglesias and Carballo, 2010a) and the geological and geotechnical study for wave farm development  
43 (Flor-Blanco et al., 2011), in which two areas off the beaches of Xago and Llumeres (Figure 1) were  
44 recommended.

45 The Xago beach-dune system constitutes an ideal case study, for it has experienced significant erosion  
46 in recent years. This is revealed particularly by the dune toe, which receded up to 11.5 m over a  
47 relatively short period of time, 2011-2014 (Figure 2) (Flor-Blanco et al., 2013; Flor et al., 2015). The  
48 conventional approach to defending the coast against flooding and erosion involves coastal structures:  
49 stone-armour or concrete-unit revetments, seawalls, groynes, detached breakwaters, etc. – this is the  
50 so-called “hard engineering” approach. The downsides of this approach are well known: it results in  
51 armoured coastlines, which bear little resemblance to their natural counterparts. Structures such as  
52 seawalls tend to have high wave reflection coefficients (far higher than those of beaches), which

53 implies larger wave heights in front of the structure and often loss of sediment. Moreover, in the  
54 current context of climate change and transition coasts, the inability of structures to adapt to sea-level  
55 rise poses a problem. Indeed, there have been recently many cases of coastal structures failing to cope  
56 with the increased pressures of climate change (Castelle et al., 2015; Kendon and McCarthy, 2015;  
57 Senechal et al., 2015). These examples of failures of coastal structures – due to either structural  
58 collapse or excessive overtopping – expose the dramatic consequences of the inadequacy of many of  
59 the existing structures in the current transition scenario. The conventional approach to solving this  
60 problem entails upgrading the existing structures or building new ones, in both cases at a large cost.

61 On these grounds, nearshore wave farms present three main advantages relative to conventional  
62 coastal structures. First, by providing renewable, carbon-free energy, wave farms contribute to  
63 decarbonising the energy supply and thereby combatting the manmade causes of climate change.  
64 Second, the environmental impact of wave farms on the littoral – the single most sensitive  
65 environment in the planet – is considerably lower than that of coastal structures. Last, but not least,  
66 wave farms consisting of floating wave energy converters (WECs) – e.g., WaveCat, WaveDragon,  
67 DEXA – adapt naturally to sea level rise, and therefore can cope well with the main impact on the  
68 littoral of climate change.

69 Thus, rather than resorting to the conventional approach (more structures) to fix obsolete,  
70 underperforming structures, deploying wave farms to generate carbon-free energy as their main  
71 purpose and, in synergy with it, defend the coastline against erosion and flooding is a new alternative  
72 that warrants consideration. Incidentally, their application to coastal defence would enhance the  
73 economic viability of wave energy through the savings achieved in conventional defence schemes.

74

## 75 2. MATERIALS AND METHODS

### 76 2.1 STUDY SITE

77 Xago (Figure 2) is a ~2 km sandy beach with a flat intertidal area. The sedimentology is characterised  
78 by siliciclastic sands of medium size. Their grain size distribution – a prerequisite for the coastal  
79 processes model – was established based on sediment samples, and the values of the most relevant  
80 metrics (D50 and D90, which are the intercepts for 10%, 50% and 90% of the cumulative mass) were  
81 obtained by means of the GRADISTAT model (Blott and Pye, 2001). The tidal regime is semidiurnal,  
82 with maximum and mean tidal ranges of 4.98 m and 2.66 m, respectively (Flor-Blanco et al., 2013) –  
83 a macro-tidal system close to the transition to meso-tidal. The beach is exposed primarily to waves  
84 from the IV quadrant (NW).

85

## 86 2.2 WAVE PROPAGATION MODEL

87 SWAN (Simulating WAVes Nearshore) is a third-generation spectral wave model that solves the  
88 equation of conservation of wave action considering the relevant wave generation and dissipation  
89 processes, such as shoaling, refraction due to current and depth, whitecapping, bottom friction and  
90 depth-induced wave breaking. The deep water boundary conditions were obtained from WaveWatch  
91 III (WWIII), a third-generation offshore wave model consisting of global and regional nested grids  
92 with a resolution of 100 km (Tolman, 2002). The model was validated over a twelve-month period  
93 using data from the wave buoy off Salinas Beach (~1 km to the west of Xago Beach) in conjunction  
94 with data from node #3085039 of the SIMAR-44 dataset (off Xago Beach), kindly provided by  
95 Spain's State Ports (*Puertos del Estado*).

96 In the twelve-month period considered for model validation purposes (January 2010 - December  
97 2010), the average values of significant wave height ( $H_s$ ), mean period ( $T_m$ ) and wave direction ( $\theta$ )  
98 were: 1.40 m, 6.02 s and 317.1°, respectively. Storm waves are also typically from the NW; for  
99 instance, during the storm from 7<sup>th</sup> November 2010 to 16<sup>th</sup> November 2010 the average deepwater  
100 wave conditions were:  $H_s = 3.72$  m,  $T_m = 7.49$  s and  $\theta = 299.9^\circ$ . This period was selected for the  
101 assessment of the effects of the wave farm as it presents storm clustering, one of the most relevant

102 phenomena in coastal erosion (Dissanayake et al., 2015). Indeed, in this period up to four storm peaks  
103 occur within approximately a week.

104 Deepwater wave data from the nearby Avilés offshore wave buoy and hindcast wave data with a  
105 three-hourly frequency, along with tidal data from the port of Gijón (20 km away) with an hourly  
106 frequency, were used to force the wave propagation and coastal processes models.

107 As regards wind conditions, in the twelve-month validation period the highest probability of  
108 occurrence (22.4%) corresponded to southerly winds (from  $157.5^\circ$  to  $202.5^\circ$ ). More importantly,  
109 however, the strongest winds (with wind speeds,  $u_{10}$ , exceeding  $20 \text{ ms}^{-1}$ ) were associated with  
110 northwesterly directions (from  $292.5^\circ$  to  $337.5^\circ$ ). Three-hourly values of wind speed and direction  
111 obtained from the Global Forecast System (GFS) weather model were input into the wave propagation  
112 model.

113 High-resolution bathymetric and topographic data, obtained in *ad hoc* surveys, were used as input for  
114 the coastal processes and wave propagation numerical models (Figure 3). Importantly, the dataset  
115 covered not only the submarine beach but also the subaerial beach, including the dune system, with  
116 elevations ranging from  $-20 \text{ m}$  to  $+15 \text{ m}$  (with reference to the Spanish National Geodetic Vertical  
117 Datum).

118 In order to locate the WECs accurately and simulate their effects on the nearshore wave conditions,  
119 two computational grids with different resolutions were defined (Figure 5): (i) a coarse grid ( $50 \times 50$   
120 m), which extended  $25 \times 25 \text{ km}$  and covered part of the Avilés submarine canyon system (including  
121 the Avilés Canyon itself, with water depths over  $900 \text{ m}$ ); and (ii) a high-resolution nested grid ( $12 \times$   
122  $15 \text{ m}$ ), which extended  $5.4 \text{ km}$  offshore and  $4.5 \text{ km}$  from east to west, covering the area of interest.

123 For the purposes of the modelling in the present research work, the area selected for the wave farm  
124 was situated off Xago Beach, at a water depth of  $\sim 30 \text{ m}$  (Figure 6). Following previous work  
125 (Carballo and Iglesias, 2013), the WECs were laid out with a spacing of  $2.2D$ , where  $D = 90 \text{ m}$  is the  
126 distance between the twin bows of an individual WaveCat WEC. Their interaction with the wave field

127 was modelled based on the wave transmission coefficients obtained in laboratory tests (Fernandez et  
128 al., 2012). The results of the wave propagation model provided boundary conditions for the XBeach  
129 model grid covering the study area (Figure 5).

### 130 2.3 COASTAL PROCESSES MODEL

131 XBeach is a process-based model that predicts the response of the beach under storm conditions  
132 considering wave propagation, sediment transport and seabed updates. Wave propagation takes into  
133 account the coastal processes in the nearshore area by means of the time-dependent wave action  
134 balance coupled to the roller energy equations and the nonlinear shallow water equations of mass and  
135 momentum. Sediment transport is modelled applying the depth-averaged advection diffusion equation  
136 on the scale of wave groups based on different equilibrium concentration and the Van Rijn–Van Thiel  
137 formulation (Van Thiel de Vries, 2009). The complete description of XBeach is given by Roelvink et  
138 al. (2006) or Roelvink et al. (2009). XBeach has been successfully applied to predict storm-induced  
139 erosion in sandy beaches (Abanades et al., 2014b; McCall et al., 2010; Pender and Karunarathna,  
140 2013; Villatoro et al., 2014) and gravel beaches (Jamal et al., 2014; McCall et al., 2014; Williams et  
141 al., 2012).

142 The computational grid extended some 1.7 km alongshore and 2 km offshore, from the dune system to  
143 water depths of approximately 20 m, with a resolution of  $7.5 \times 5$  m. Spectral wave parameters (wave  
144 height, period, direction and spreading) from the SWAN runs were used to prescribe the offshore  
145 boundary conditions.

146 Regarding coastal morphodynamics, the beach may be divided into three main sections: (i) the west  
147 section, which experiences significant storm-induced erosion; (ii) the middle section, characterised by  
148 some erosion on the dune front and deposition in the intertidal area of materials eroded in the west  
149 section; and (iii) the east section, subject to intense erosion of the foredune (Flor-Blanco et al., 2013).  
150 It is also noteworthy how the evolution of the foredune front has changed over the last years relative  
151 to previous decades: whereas the dune limit advanced from 1970 to 2011 (Figure 4), it retreated from  
152 2011 to 2014 (Figure 2). This change from progradation to recession, which may well be related to the

153 severe winter gales experienced in the 2011-2014 period (Flor-Blanco et al., 2013), would highlight  
154 the sensitivity of the system to storm-induced erosion, and thus the interest of mitigating the latter  
155 from a coastal management perspective.

156 The response of the Xago beach-dune system to storm conditions was compared in two scenarios,  
157 with and without the wave farm, in order to establish the effects of the wave farm on the system and,  
158 on these grounds, test the scientific hypothesis of this work.

159

## 160 2.4 COASTAL INDICATORS

161 Coastal indicators were used to quantify the effects of the wave farm on the beach-dune system, as  
162 follows. The effects on the nearshore wave conditions were analysed through Reduction in Significant  
163 wave Height (*RSH*) (Abanades et al., 2015), a dimensionless indicator defined as

$$164 \quad RSH(x, y) = H_{s,b}(x, y)^{-1}(H_{s,b}(x, y) - H_{s,f}(x, y)) , \quad (2)$$

165 where  $H_{s,f}$  and  $H_{s,b}$  are the significant wave height with the wave farm and without (baseline),  
166 respectively, at a generic point of coordinates  $(x,y)$ . This indicator shows the reduction of the  
167 significant wave height induced by the wave farm in terms of its magnitude in the baseline scenario.

168 The effects of the wave farm on coastal morphodynamics were studied by means of the coastal  
169 indicators developed by Abanades et al. (2014a): (i) Bed Level Impact (*BLI*), (ii) beach Face Eroded  
170 Area (*FEA*) and (iii) Non-dimensional Erosion Reduction (*NER*).

171 The *BLI* indicator, with SI units of m, represents the difference in seabed level at a point of the beach  
172 between the two scenarios, with and without the farm:

$$173 \quad BLI(x, y) = \zeta_f(x, y) - \zeta_b(x, y), \quad (3)$$



174 where  $\zeta_f(x,y)$  and  $\zeta_b(x,y)$  are the seabed level with the farm and without it (baseline), respectively, at a  
 175 generic point of the beach designated by its coordinates  $(x,y)$ , and at the point in time considered. The  
 176  $y$ -axis is directed along the general coastline orientation, with the coordinate increasing towards the  
 177 eastern end of the beach, and the  $x$ -axis is directed along the beach profiles, with the coordinate  
 178 increasing towards the landward end. A positive or negative value of  $BLI$  signifies that the presence of  
 179 the wave farm resulted in a higher (accretion) or lower (erosion) seabed level relative to the baseline  
 180 (no farm) scenario, respectively.

181 The  $FEA$  indicator, with units of  $m^2$ , represents the volume of the beach face (the section of the profile  
 182 exposed to wave uprush) eroded per unit length of beach relative to an initial condition – typically, a  
 183 point in time before the beginning of the storm considered. It is defined in both scenarios, baseline  
 184 ( $FEA_b$ ) and with the wave farm ( $FEA_f$ ):

$$185 \quad FEA_b(y) = \int_{x_1}^{x_{max}} [\zeta_0(x,y) - \zeta_b(x,y)] dx, \quad (4)$$

$$186 \quad FEA_f(y) = \int_{x_1}^{x_{max}} [\zeta_0(x,y) - \zeta_f(x,y)] dx, \quad (5)$$

187 where  $\zeta_0(x,y)$  is the seabed level at the point of coordinates  $(x,y)$  at the initial condition (typically, a  
 188 point in time before the storm), and  $x_1$  and  $x_{max}$  are the values of the  $x$ -coordinate at the seaward end of  
 189 the beach face and landward end of the profile, respectively. The  $FEA$  indicator is a profile function,  
 190 and hence depends only on the  $y$ -coordinate, so it allows the alongside analysis of the erosion in the  
 191 beach.

192 Finally, the  $NER$  indicator is also a profile function, in this case non-dimensional, defined as

$$193 \quad NER(y) = 1 - (x_{max} - x_1)^{-1} \int_{x_1}^{x_{max}} [\zeta_0(x,y) - \zeta_f(x,y)] [\zeta_0(x,y) - \zeta_b(x,y)]^{-1} dx. \quad (6)$$

194 It expresses the variation in the eroded area of a generic profile ( $x$ ) caused by the wave farm as a  
 195 percentage of the total area eroded between the initial condition (typically, a point in time before the

196 storm) and the point in time considered. Thus, a positive or negative value implies a reduction or  
197 increase in the eroded area as a result of the wave farm, respectively.

198

### 199 3. RESULTS AND DISCUSSION

200 First, the model was validated using the data from the wave buoy off Avilés (with available data from  
201 1<sup>st</sup> March 2010 – 1<sup>st</sup> September 2010) and the point SIMAR44-3085039 (1<sup>st</sup> January 2010 – 1<sup>st</sup>  
202 January 2011). The model results are in excellent agreement with the observations (Figures 7 and 8),  
203 as corroborated by the root mean square error and coefficient of determination (Table 1).

204 The effects of the wave farm on beach morphodynamics were analysed by comparing the response of  
205 the beach with and without the wave farm to the storm from 7 November 2010, 12:00 UTM to 16  
206 November 2010, 06:00 UTM (delimited by dashed lines in Figures 7 and 8). In the significant wave  
207 height patterns at 18:00 on 9<sup>th</sup> November 2010 (Figure 9) an area of wave energy concentration is  
208 apparent. During the propagation of waves towards the coast their properties change as a result of  
209 their interaction with the seabed (refraction, shoaling, friction). Over an irregular bathymetry (Figure  
210 6) this interaction often leads to areas of energy concentration, also known as nearshore hotspots  
211 (Iglesias and Carballo, 2010b) – such as the area selected for the deployment of a wave farm off  
212 Xago.

213 The effects of the wave farm on wave heights were substantial directly behind the wave farm (Figure  
214 9), with a reduction in significant wave height (*RSH*) of over 50%. This reduction decreased towards  
215 the coastline due to the energy diffracted from both sides of the farm into its wake; importantly,  
216 however, it was still significant near the coastline, with *RSH* values exceeding 15% along the 10 m  
217 contour. The farm was not directly in front of the beach but somewhat to the east (Figure 9), which  
218 reduced its impact on the conditions in the western part of the beach. This is apparent in the  
219 significant wave height values along the 20 m contour (Figure 10) and the resulting *RSH* values  
220 (Figure 11). (The 20 m contour itself is depicted in Figure 6). Strictly speaking, *RSH* was non-zero

221 over a 3500 m stretch alongshore. More importantly, however, relevant *RSH* values (above 30%)  
222 extended over some 1300 m alongshore – and this, it may be argued, is the alongshore extent of the  
223 wave farm wake for practical purposes. Owing to the position of the farm to the east of the beach, the  
224 wake extends some distance east of the beach (Figure 11). The maximum *RSH* values within the wake  
225 were a hefty ~50%.

226 As regards beach morphodynamics, in the baseline scenario (without the wave farm) the storm  
227 produced acute erosion in the subaerial beach, in particular in front of the dune in the eastern part of  
228 the beach, with a fall in the beach level of up to 2.5 m (Figure 12). In general, some erosion in front of  
229 the dune occurred throughout the beach, in line with the general recessionary trend of the dune toe in  
230 the period 2011-14 (Figure 2). The low-tide terrace, for its part, experienced accretion in the west and  
231 mid-sections – which would appear to be a consequence of the deposition of the sand eroded from the  
232 subaerial beach – and some erosion in the east section of the beach. Further offshore areas with some  
233 degree of erosion are interspersed with areas of little erosion.

234 For the analysis of the effects of the wave farm, the coastal indicators defined in Section 2.4 were  
235 applied. Based on the *BLI* values after the storm (Figure 13) three main areas can be distinguished: the  
236 dune front along the entire beach, the low-tide terrace in the east section, and the low-tide terrace in  
237 the west and mid-section of the beach. *BLI* values are positive in the first two areas, negative in the  
238 latter. Considering also the results shown in Figure 12, these values indicate that the wave farm  
239 reduced storm-induced erosion in the dune front and the east section of the low-tide terrace, and  
240 reduced storm-induced accretion in the west and mid-sections of the low-tide terrace.

241 The greatest *BLI* values (over 2 m) were found in the first area, the dune front, and especially in the  
242 east section of the beach, which experienced the greatest erosion in the absence of wave farm (Figure  
243 12). In the west and middle sections *BLI* values were also significant, between 1-1.5 m. In sum, the  
244 wave farm contributed significantly to mitigate storm-induced erosion on the dune front.

245 In the second area, the low-tide terrace in the east section of the beach, the maximum *BLI* values were  
246 smaller than those on the dune front but nevertheless relevant – with storm-induced erosion

247 decreasing by up to 1 m over a large area. Thus, in the east section of the beach, both on the dune  
248 front and the low-tide terrace, erosion is significantly mitigated, which can be explained by the fact  
249 that this area is directly in the lee of the wave farm, with consequential reductions in significant wave  
250 heights (*RSH* values of ~50%, Figure 11).

251 Finally, in the third area, the west and middle sections of the low-tide terrace, negative *BLI* values  
252 occur. The sediment eroded from other areas is deposited in part here, so the storm actually results in  
253 accretion in this area (Figure 12). In this context, the negative values of *BLI* indicate that the wave  
254 farm reduces accretion.

255 In sum, the wave farm, by reducing the amount of wave energy that reaches the beach during the  
256 storm, dampens the morphodynamics of the system: it reduces both erosion (in the first and second  
257 areas) and accretion (in the third area).

258 The wave farm was most effective at countering erosion on the dune front. This is no mean feat given  
259 that it was precisely the dune front that experienced the greatest erosion in the baseline scenario.  
260 Thanks to the absorption of wave energy by the wave farm the landward reach of erosion was  
261 displaced offshore by over 10 m along the three profiles considered, representative of the three  
262 sections of the beach (Figure 14): P1 (west), P2 (middle) and P3 (east). In the east section (P3) this  
263 displacement reached a very substantial 25 m, which is indicative of the efficiency of the wave farm  
264 in mitigating storm-induced erosion in the area directly in its lee.

265 To investigate in particular erosion on the beach face, two coastal indicators were applied:  $FEA_b(y)$   
266 and  $FEA_f(y)$ , which represent the area eroded in the beach face at the  $y$  profile in the baseline and the  
267 wave farm scenarios, respectively (Figure 15). The largest values in both scenarios correspond to the  
268 east section of the beach, although significant erosion occurred throughout. The reduction achieved by  
269 the wave farm is apparent by comparing the curves of  $FEA_b$  and  $FEA_f$ , and particularly so in the east  
270 section – again, the wave farm is most effective where erosion is most pronounced.

271 Finally, the *NER* indicator represents the reduction in the eroded area caused by the wave farm as a  
272 proportion or percentage of the total eroded area in the baseline scenario (Figure 16). The wave farm  
273 was found to reduce erosion ( $NER > 0$ ) on most of the beach, with the highest values in the east  
274 section ( $NER > 60\%$ ). As indicated, this is precisely the area where the erosion of the beach face was  
275 more pronounced in the baseline scenario (Figure 15) – hence the interest for coastal management. In  
276 any case, this drastic reduction in erosion in the east section, for all its relevance, must not obscure the  
277 significant reductions elsewhere, with an average *NER* value of 17.64%.

278 These results indicate that a wave farm off Xago Beach would have contributed to mitigating the dune  
279 erosion over the last decade at the very least during the storm events (short-term analysis). Although  
280 further analysis in the long-term would undoubtedly be useful – possibly using behavioural models  
281 rather than process-based models – the present results can be of particular relevance in cases with  
282 assets (promenades, buildings, roads, railways, etc.) close to the beach and at risk from storm-induced  
283 scour at the toe of their foundations, for they indicate that a wave farm deployed off the affected  
284 section of coast would contribute to preventing storm-induced toe scour, which may lead to structural  
285 failure.

286 However, the effects of wave farms on the coast do not lend themselves to general statements, for  
287 they will depend on the characteristics of the area in question (wave energy resource, wave climate  
288 and grain size distribution, among others), of the WECs and their layout. In this sense, *ad hoc* studies  
289 are necessary for determining the viability of such projects in an area, considering not only the  
290 effectiveness of the wave farm in mitigating coastal erosion but also any other effects. In certain  
291 cases, these may be negative, e.g. in an coastal area popular with surfers. The reduction in wave  
292 power and, consequently, wave height near shore might have a negative impact on the tourism and the  
293 economy of the area; in the latter the deployment of the WECs could affect the fisherman's activity  
294 and the lower resource could reduce the nutrient flow. On the other hand, in beaches like Xago, that  
295 are experiencing a drastic reduction of the sand volume, the installation of a wave farm can contribute  
296 for the production of renewable energy and mitigate coastal protection.

297

298 4. CONCLUSIONS

299 This work posited the hypothesis that wave farms can serve for a dual purpose, production of carbon-  
300 free energy and coastal erosion management, and tested this hypothesis through a case study: the  
301 Xago beach-dune system in N Spain – a beach backed by a dune field which has experienced  
302 substantial erosion in recent years, and is located in the area proposed for the deployment of the first  
303 wave farm in Spain.

304 To establish the effectiveness of the wave farm in mitigating storm-induced erosion, a series of  
305 coastal indicators were applied to the results of two numerical models – a coastal processes model,  
306 XBeach, fed by a wave propagation model, SWAN. The wave propagation model was used to  
307 determine the effects of the wave farm on the nearshore wave conditions. Wave energy extraction by  
308 the WECs was found to have relevant nearshore effects, with reductions in the significant wave height  
309 (*RSH*) of up to 50% along the 20 m contour.

310 For the purpose of coastal erosion management the crux of the matter is of course the effects of the  
311 wave farm on the morphodynamics of the beach-dune system. These may be summarised as follows:  
312 by reducing the amount of wave energy available at the beach, the wave farm acted as a dampener of  
313 coastal processes. The areas which experienced erosion under storm conditions (dune front, low-tide  
314 terrace in the east part of the beach) saw their erosion mitigated thanks to the wave farm. Similarly,  
315 the areas which experienced accretion under storm conditions (low-tide terrace in the west and middle  
316 parts of the beach) saw their accretion reduced.

317 Importantly, the farm reduced storm-induced erosion of the dune front. The scarp, or landward  
318 extreme reached by erosion, was displaced seaward by up to 25 m in the east part of the beach, i.e. in  
319 the section most affected by storm-induced erosion. Had the wave farm been operating, this would  
320 have been of great practical significance, given the unrelenting retreat of the dune toe experienced by  
321 the beach in the period 2011-14.

322 Further research is certainly warranted to validate our hypothesis, and analysis at the local scale will  
323 always be necessary. In any case, the results so far are promising as to the potential of dual-purpose  
324 wave farms to serve as a coastal erosion management tool, replacing, or at least complementing,  
325 conventional coastal structures (detached breakwaters, groynes, etc.) which tend to have far greater  
326 visual impact and, unlike offshore floating wave farms, do not adapt naturally to sea level rise – an  
327 important aspect in the current context of climate change.

328 For this new approach to wave energy, the dual-purpose wave farm, to materialise, its benefits in  
329 terms of coastal management must be translated into incentives for the developers of the wave farms;  
330 in economic terms, the positive externalities (reduced storm-induced erosion, reduced visual impact,  
331 etc.) must be internalised through appropriate mechanisms (Astariz and Iglesias, 2015a; Astariz and  
332 Iglesias, 2015b; Astariz and Iglesias, 2016). If correctly applied, the potential of this new approach is  
333 immense.

334 In sum, with the advent of large scale wave energy exploitation over the coming decades, a new,  
335 potent tool for coastal erosion management will arise.

336

### 337 **AKNOWLEDGEMENTS**

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341 Engineering of the University of Plymouth (UK). The authors are grateful to *Cluster de Energía,*  
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344

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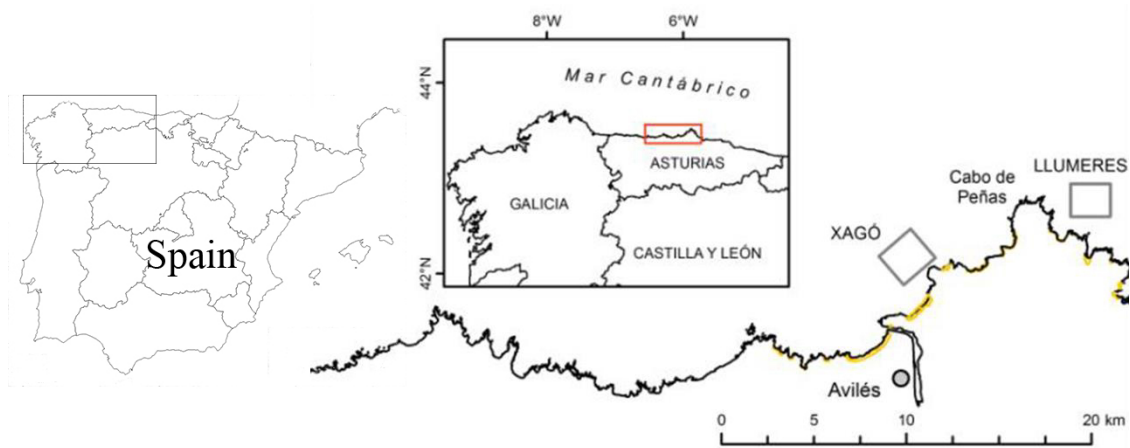
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452

453 **FIGURES**



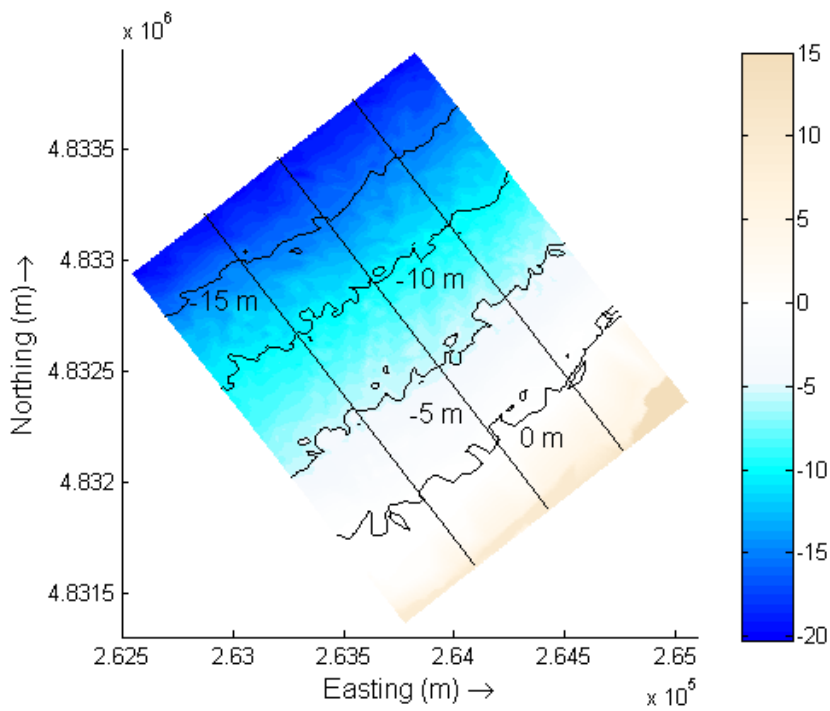
454

455 Figure 1: Location of the Xago beach-dune system in Asturias, N Spain. The squares on the right-  
456 hand side of the figure delimit the areas selected for the deployment of wave farms: Xago and  
457 Llumeres (Flor-Blanco et al., 2011).



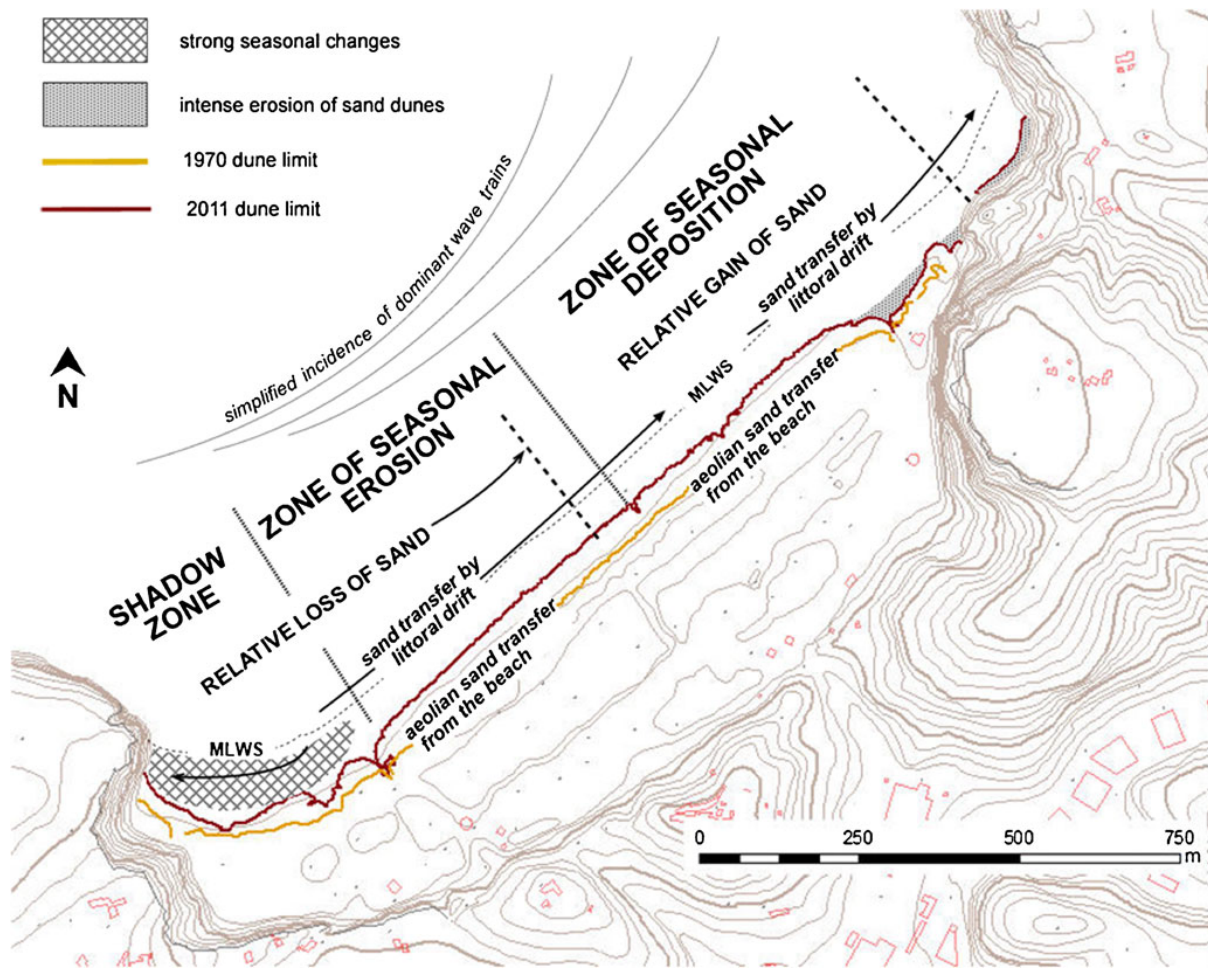
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459 Figure 2: Dune toe recession at Xago.



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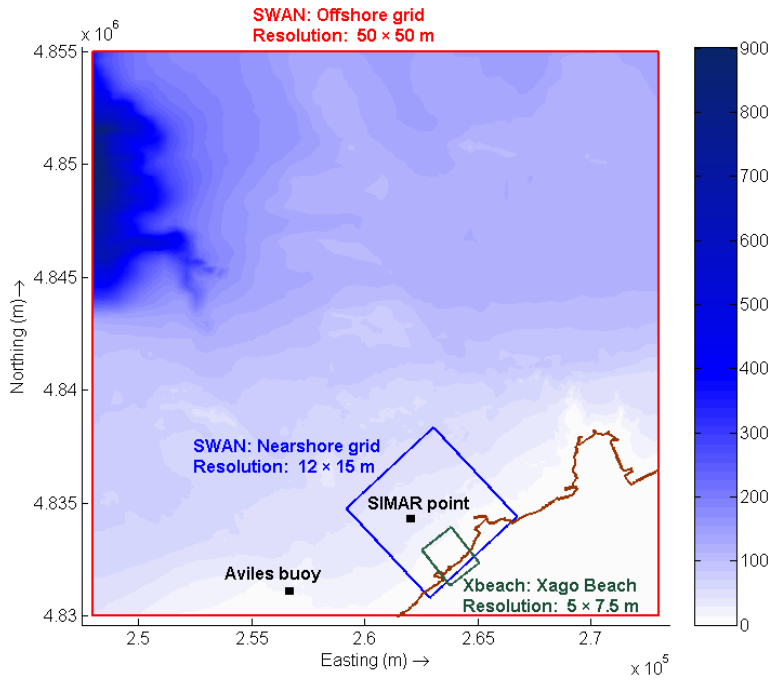
461 Figure 3: Bathymetry for the coastal processes model, with Profiles P1, P2 and P3 from left to right.  
 462 [Water depths in m].



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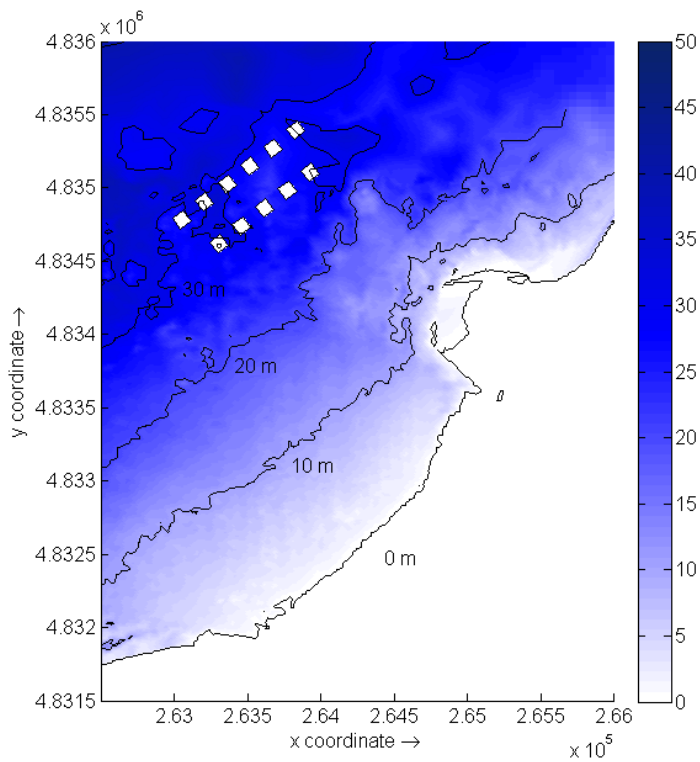
465 Figure 4: Simplified dynamic and sedimentary model of Xago (Flor-Blanco et al., 2013).

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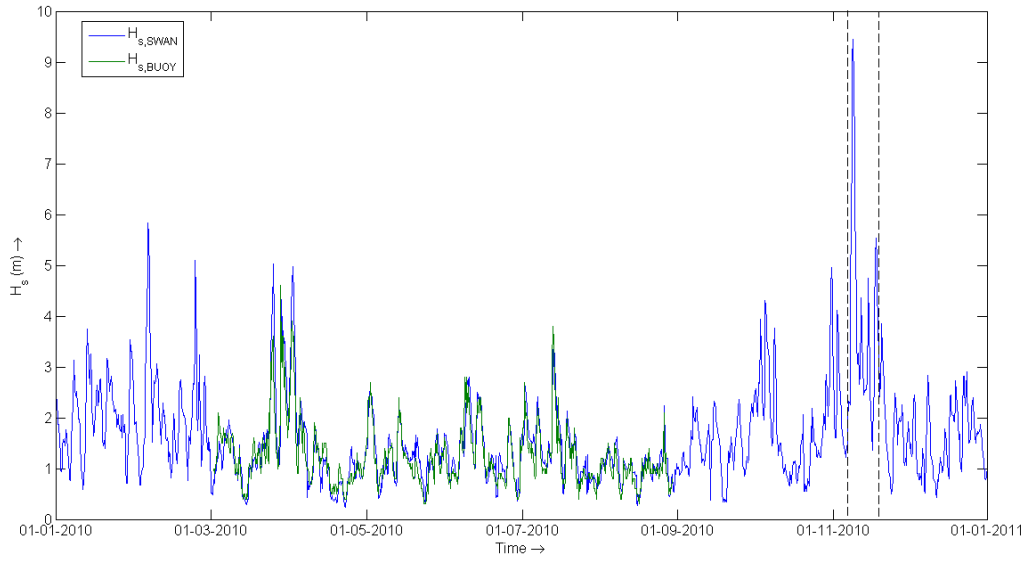
467

468 Figure 5: Computational grids for the SWAN and XBeach models, and wave buoy locations [Water  
 469 depths in m].



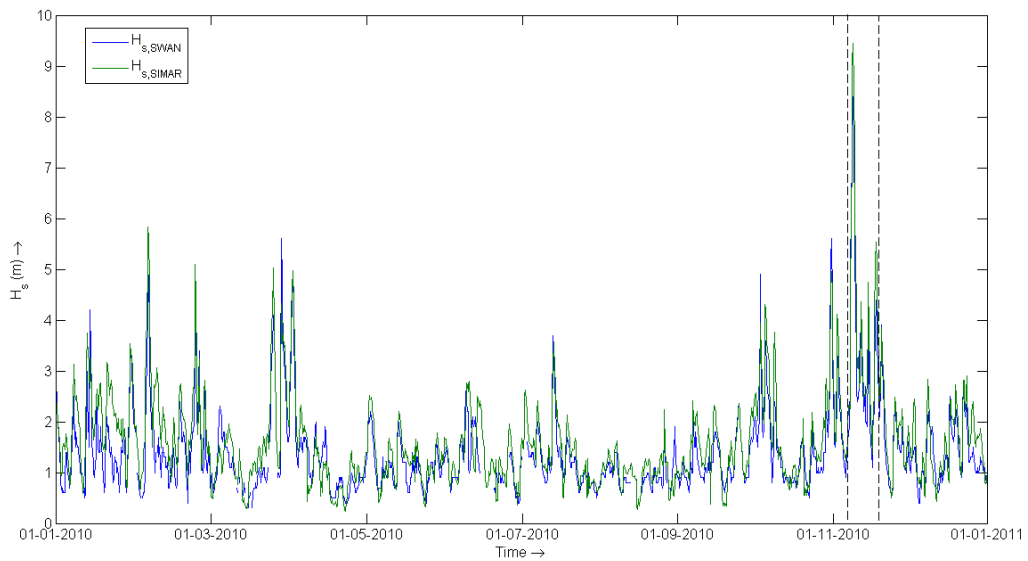
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471 Figure 6: Wave farm layout off Xago [Water depths in m].



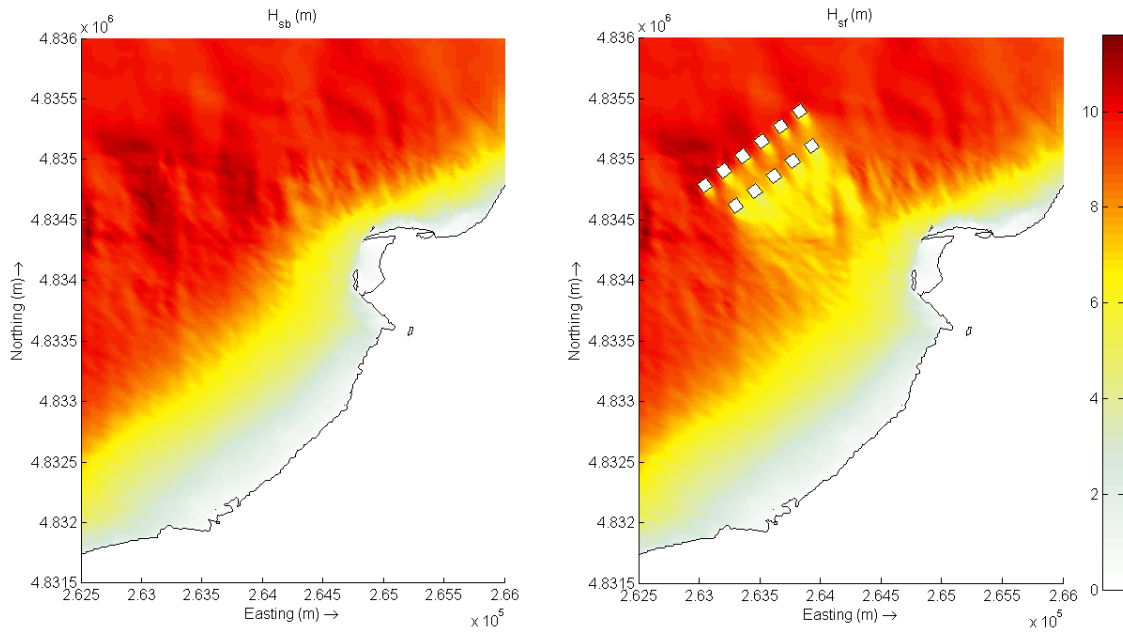
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473 Figure 7: Validation of the SWAN wave propagation model with the Aviles wave buoy data: observed  
 474 ( $H_{S, BUOY}$ ) vs. calculated ( $H_{S, SWAN}$ ) time series of significant wave height. The dashed lines delimit the  
 475 storm period used for the study.



476

477 Figure 8: Validation of the SWAN wave propagation model with the point SIMAR44-3085039 off  
 478 Xago Beach: observed ( $H_{S, BUOY}$ ) vs. calculated ( $H_{S, SWAN}$ ) time series of significant wave height. The  
 479 dashed lines delimit the storm period used for the study.

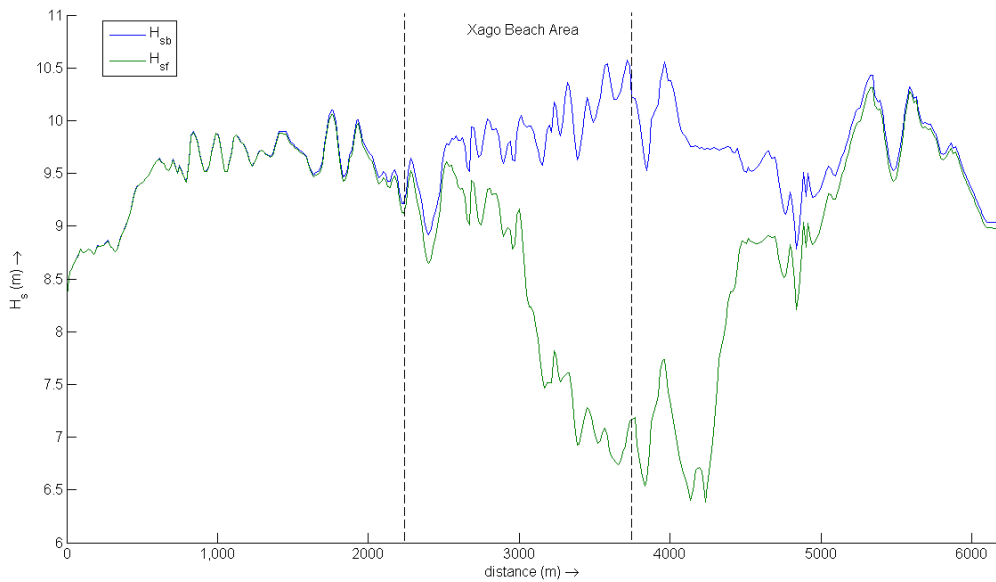


480

481 Figure 9: Significant wave height in the baseline scenario ( $H_{sb}$ ) and with the wave farm ( $H_{sf}$ ) on 9 Nov  
 482 2010, 18:00 UTC. [Deep water wave conditions:  $H_{s0}=10.28$  m,  $T_p = 15.64$  s,  $\theta_p = 268.45^\circ$ ].

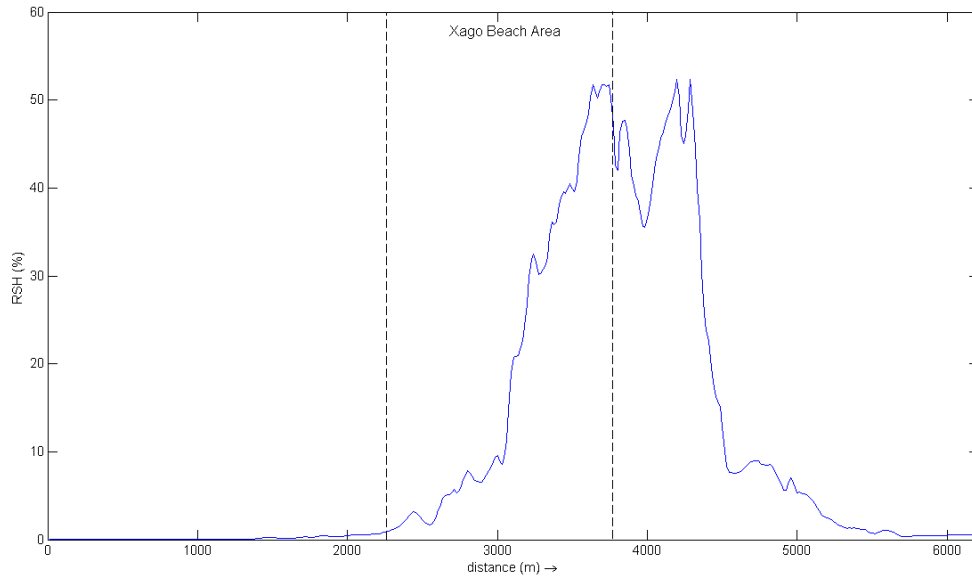
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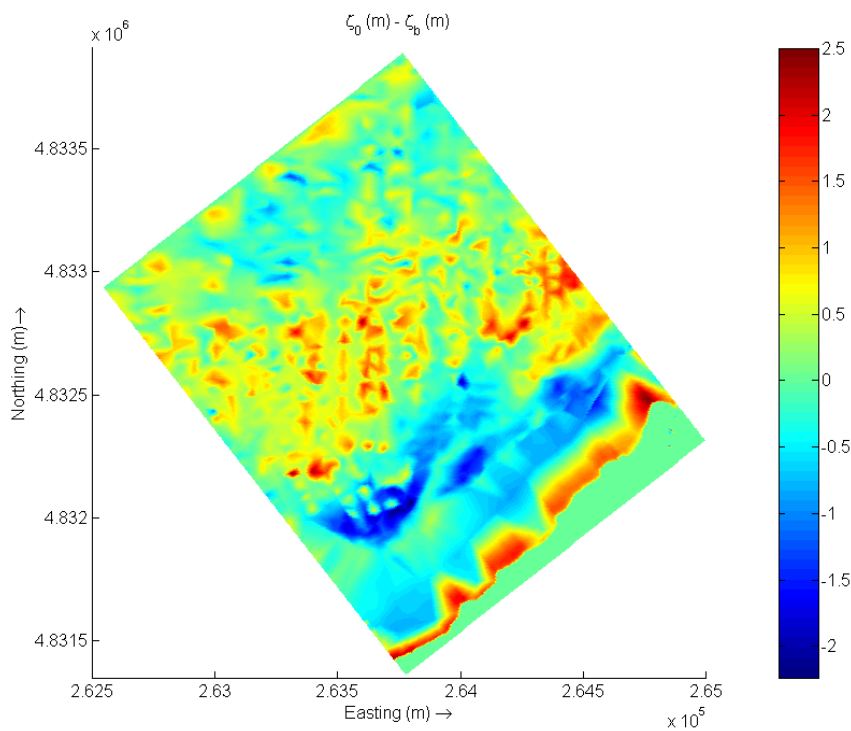
485

486 Figure 10. Significant wave height in the baseline scenario ( $H_{sb}$ ) and with the wave farm ( $H_{sf}$ ) along  
 487 the 20 m water depth contour on 9 Nov 2010, 18:00 UTC. [Deep water wave conditions:  $H_{s0}=10.28$   
 488 m,  $T_p = 15.64$  s,  $\theta_p = 268.45^\circ$ ].



489

490 Figure 11: Reduction of the significant wave height (*RSH*) parameter along the 20 m water depth  
 491 contour on 9 Nov 2010, 18:00 UTC. [Deep water wave conditions:  $H_{s0}=10.28$  m,  $T_p = 15.64$  s,  $\theta_p =$   
 492  $268.45^\circ$ ].



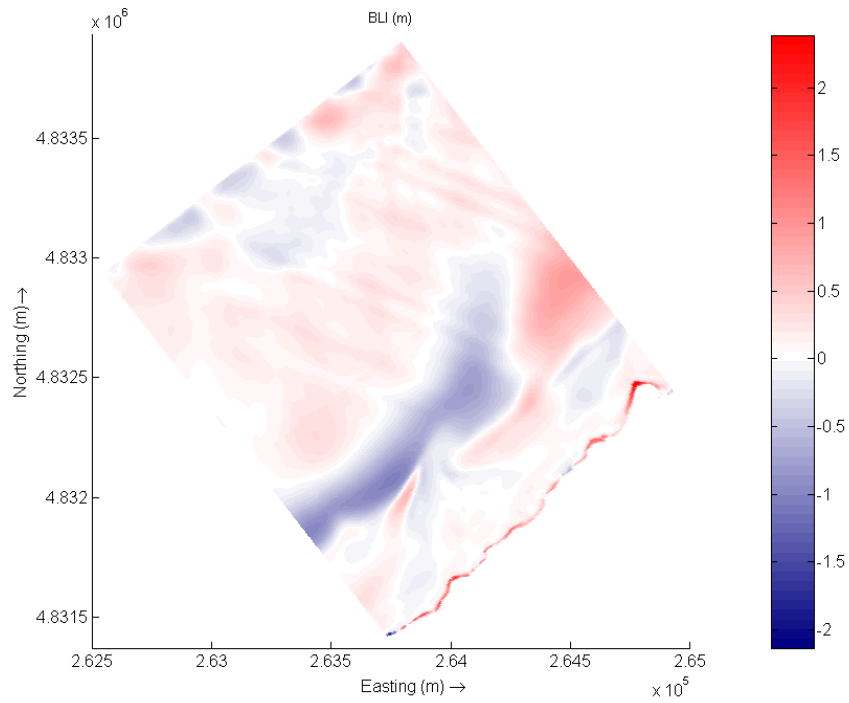
493

494 Figure 12: Fall in bed level after the storm (without the wave farm). Positive and negative values  
 495 indicate erosion and accretion, respectively.



496

497

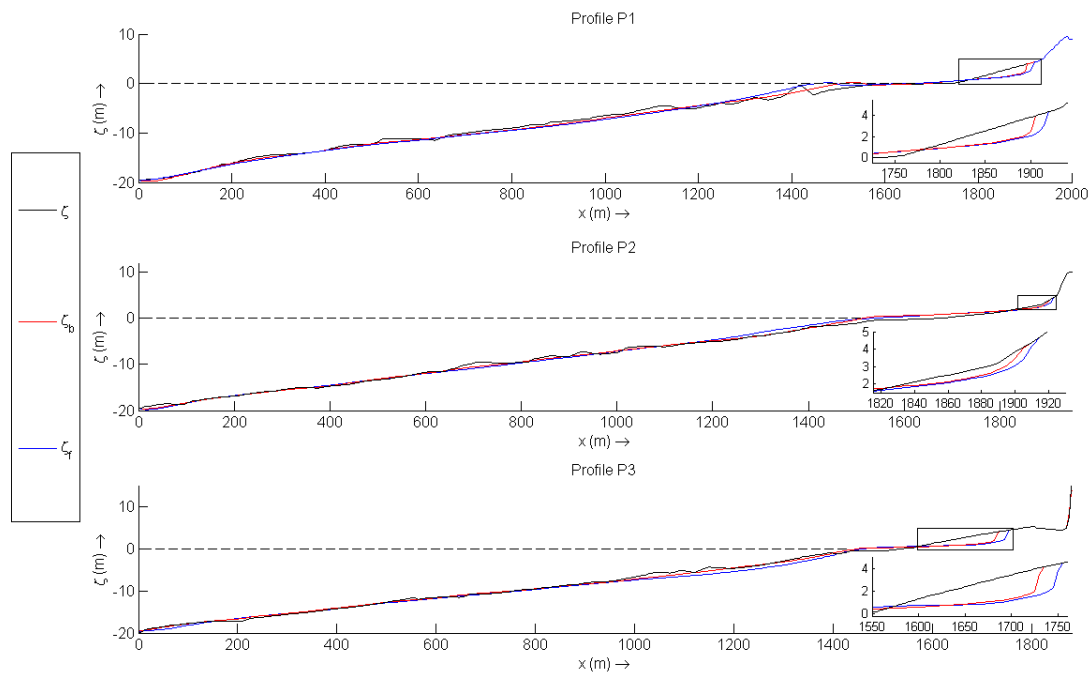


498

499 Figure 13: Bed level impact (*BLI*) at Xago after the storm [16 Nov 2007, 06:00 UTC]. Positive values  
500 indicate reduction in erosion due to the wave farm.

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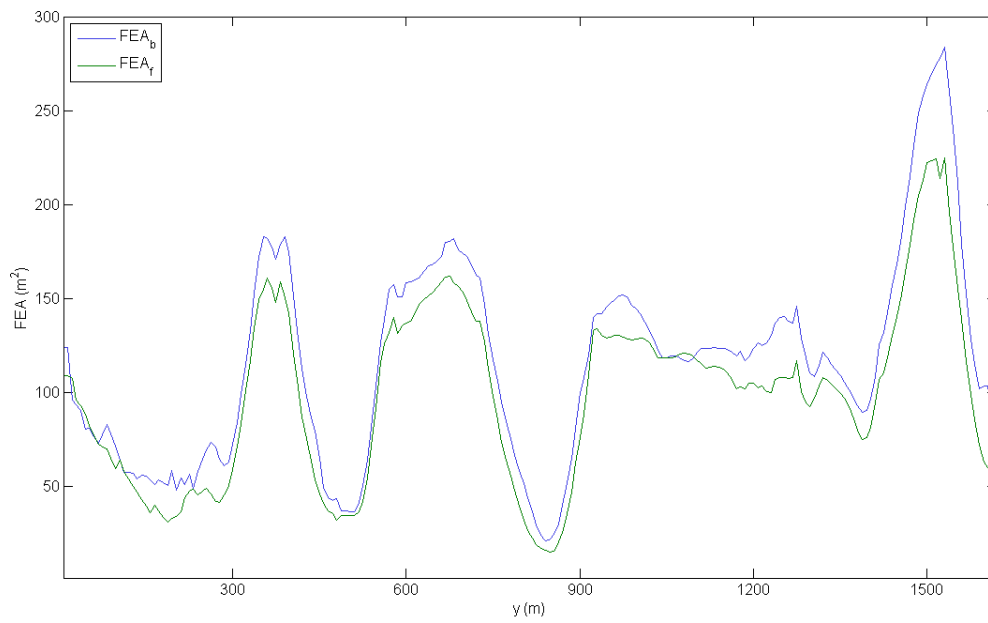
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503

504 Figure 14: Bed level at profiles P1, P2 and P3: before the storm ( $\zeta_0$ ) [07 Dec 2007, 12:00 UTC] and  
505 after the storm [16 Nov 2007, 06:00 UTC] in the baseline scenario ( $\zeta_b$ ) and with the wave farm ( $\zeta_f$ ).  
506 [Profiles P1, P2 and P3 are delimited in Figure 3].

507

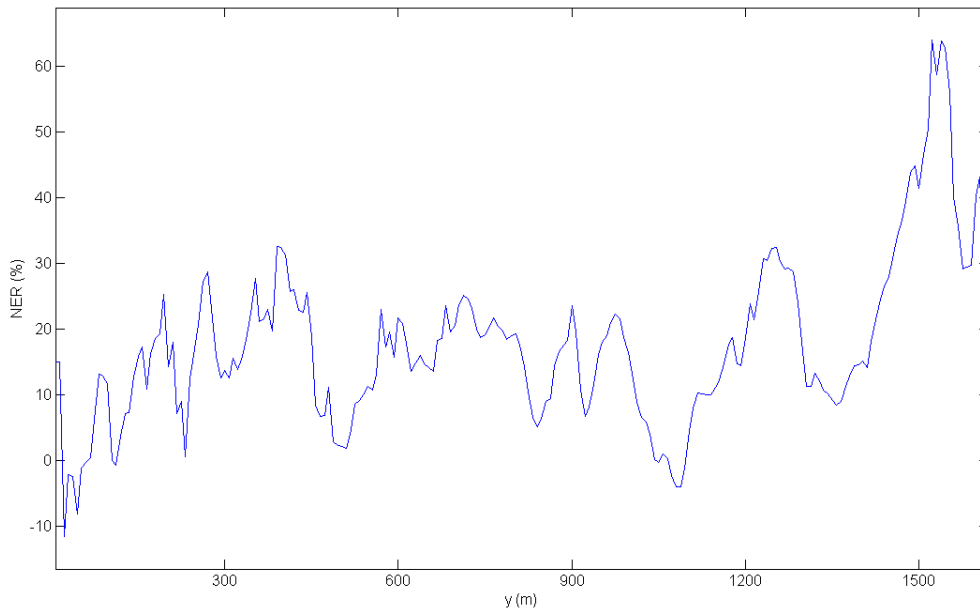


509

510 Figure 15: Beach face eroded area at the end of the time period studied [16 Nov 2007, 06:00 UTC] in  
511 two scenarios: baseline ( $FEA_b$ ) and with the wave farm ( $FEA_f$ ). The  $y$ -coordinate is the alongshore  
512 coordinate, with  $y$  increasing eastwards.

513

514



515

516 Figure 16: Non-dimensional erosion reduction (NER) on the beach face at the end of the time period  
517 studied [16 Nov 2007, 06:00 UTC]. The y-coordinate is the alongshore coordinate, with y increasing  
518 eastwards.

519

520

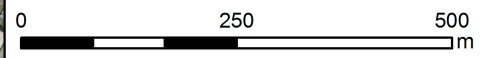
521 Table 1: Wave data used to validate the wave propagation model and values of the error statistics:  
522 Root Mean Square Error (*RMSE*) and coefficient of determination ( $R^2$ ).

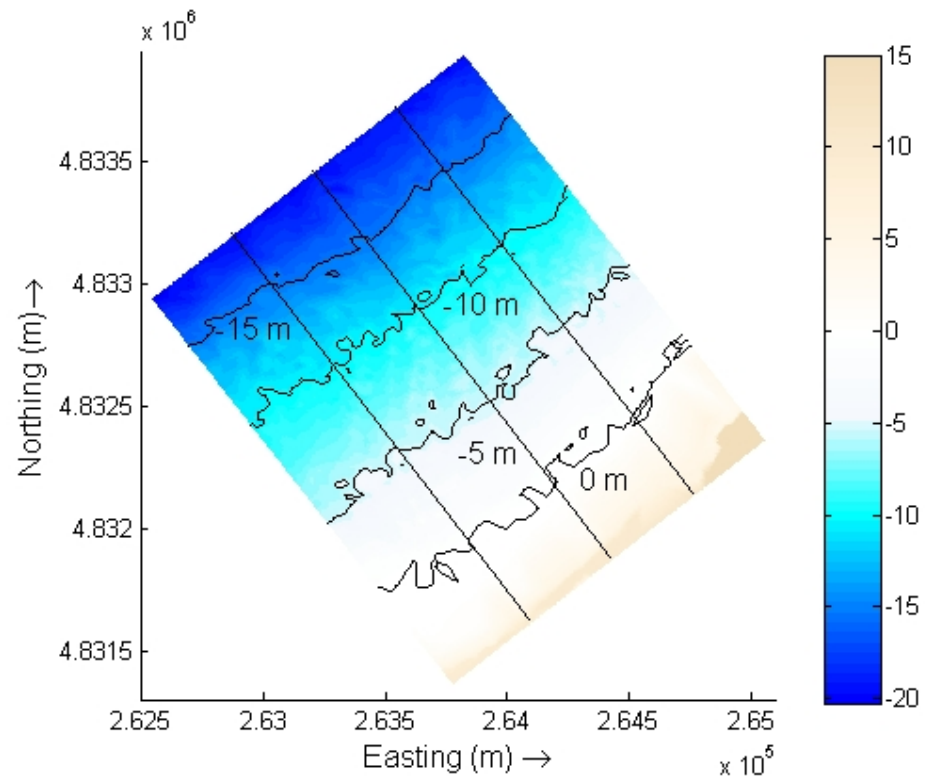
Data	Data available	Error statistics	
		Root Mean Square Error ( <i>RMSE</i> ) [in m]	Coefficient of determination ( $R^2$ ) [-]
Wave buoy data off Avilés	1 <sup>st</sup> March 2010 – 1 <sup>st</sup> September 2010	0.33	0.89
SIMAR44-3085039 off Xago	1 <sup>st</sup> January 2010 – 1 <sup>st</sup> January 2011	0.45	0.92

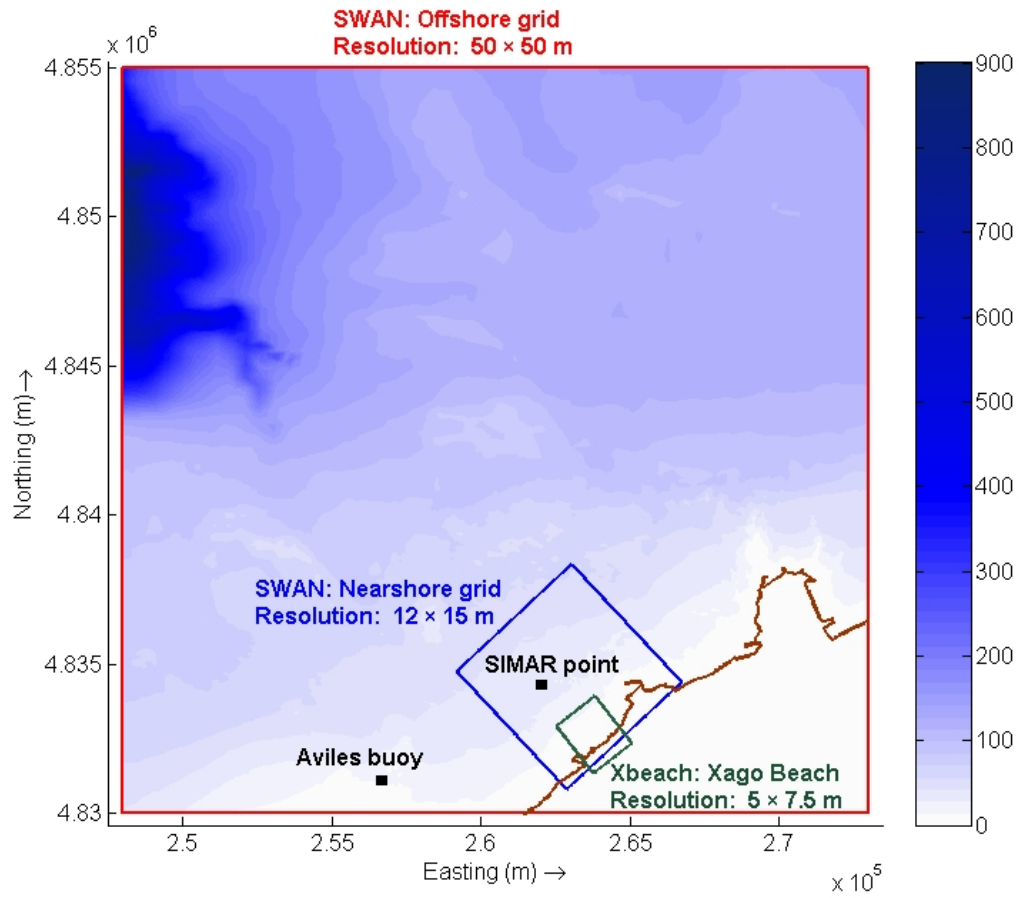
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**Legend**

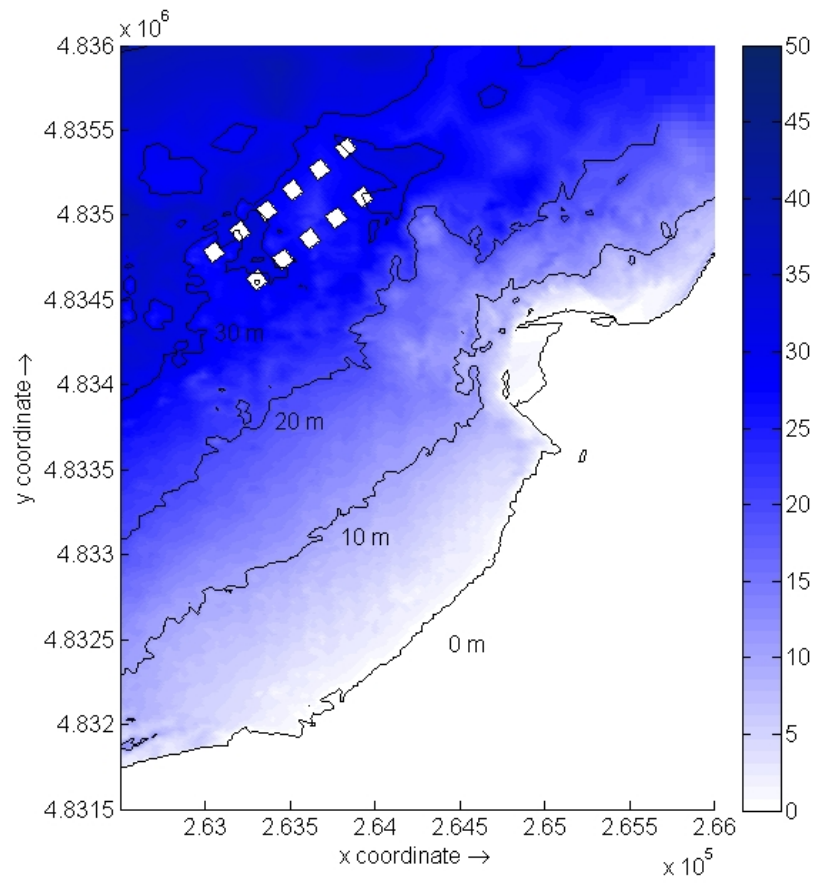
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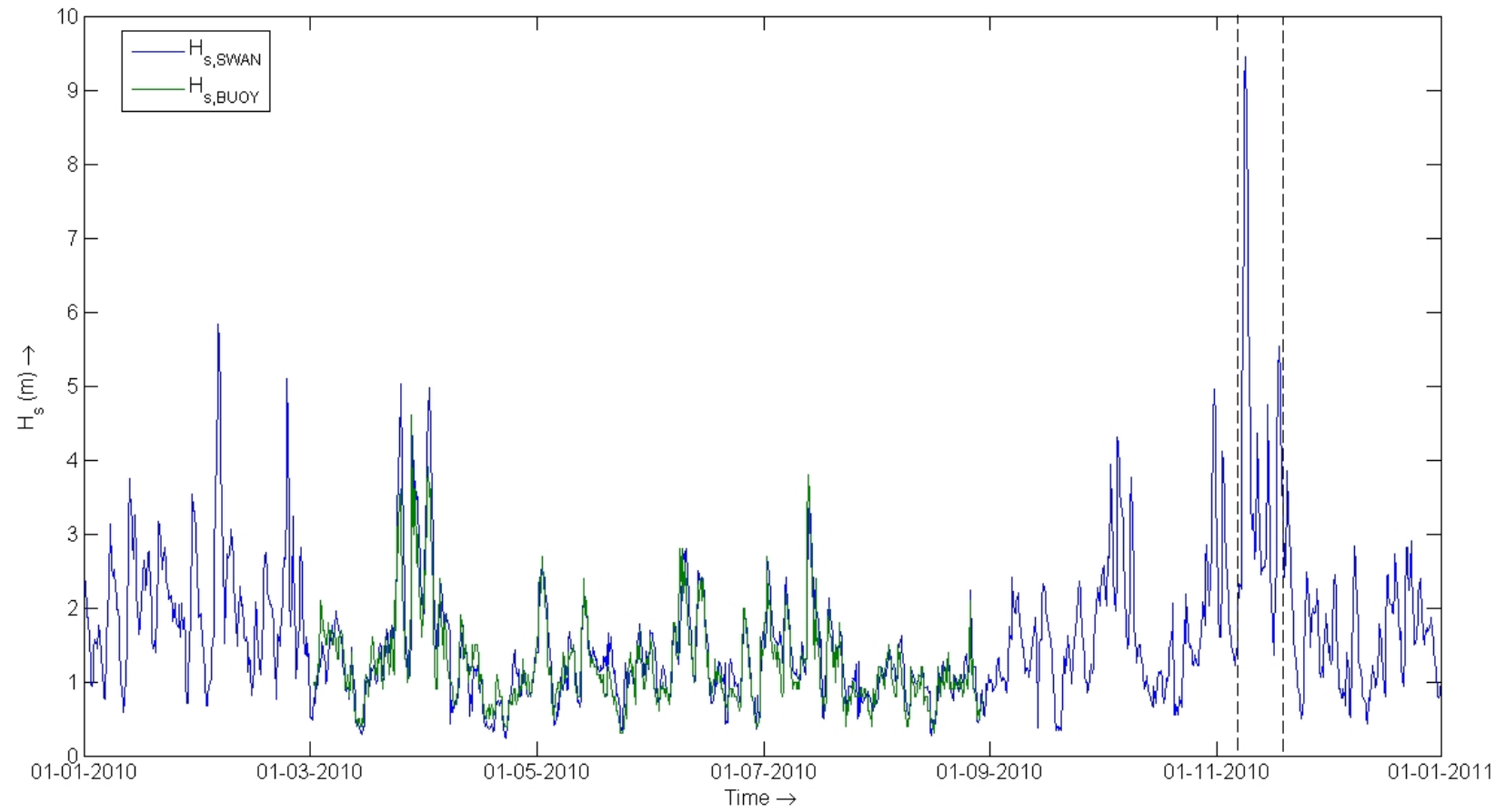


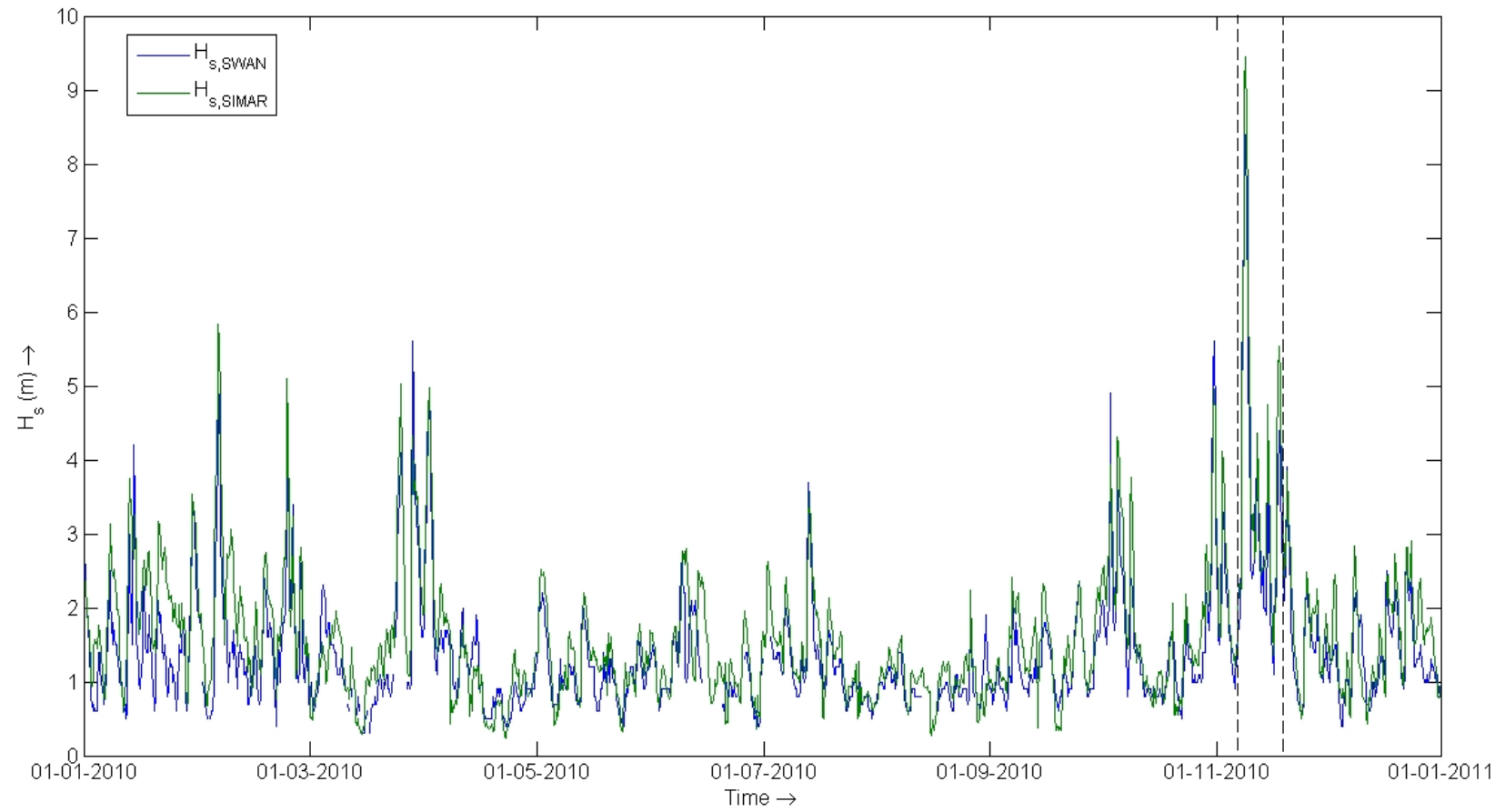


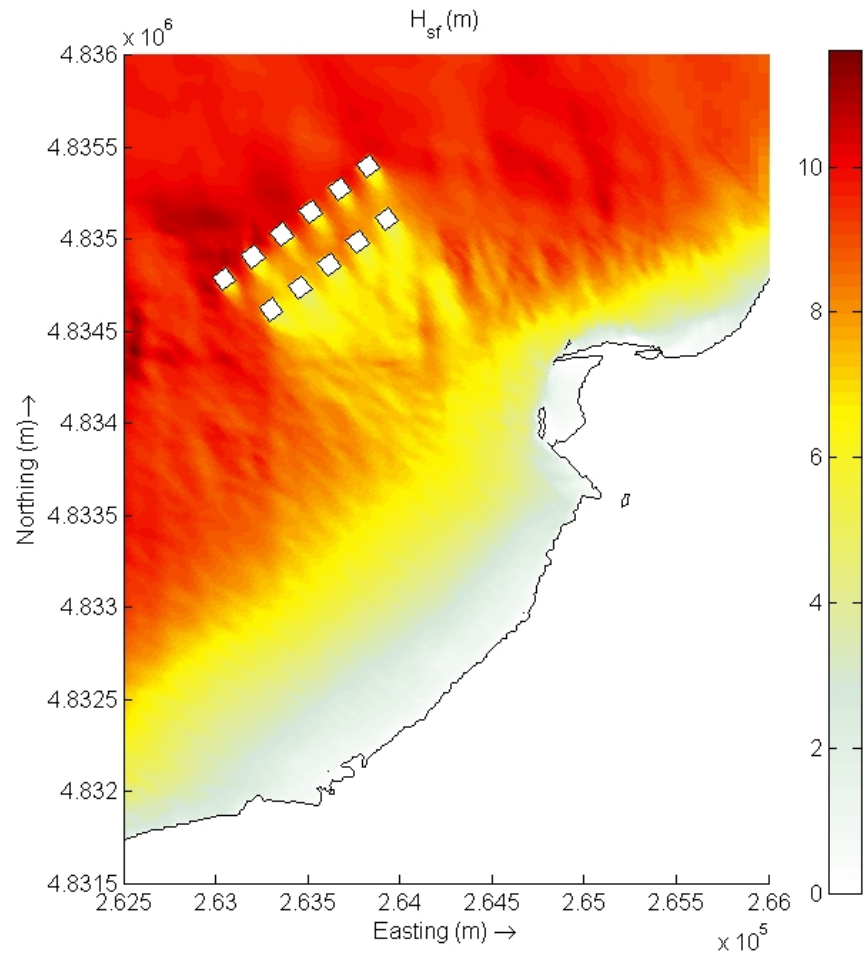
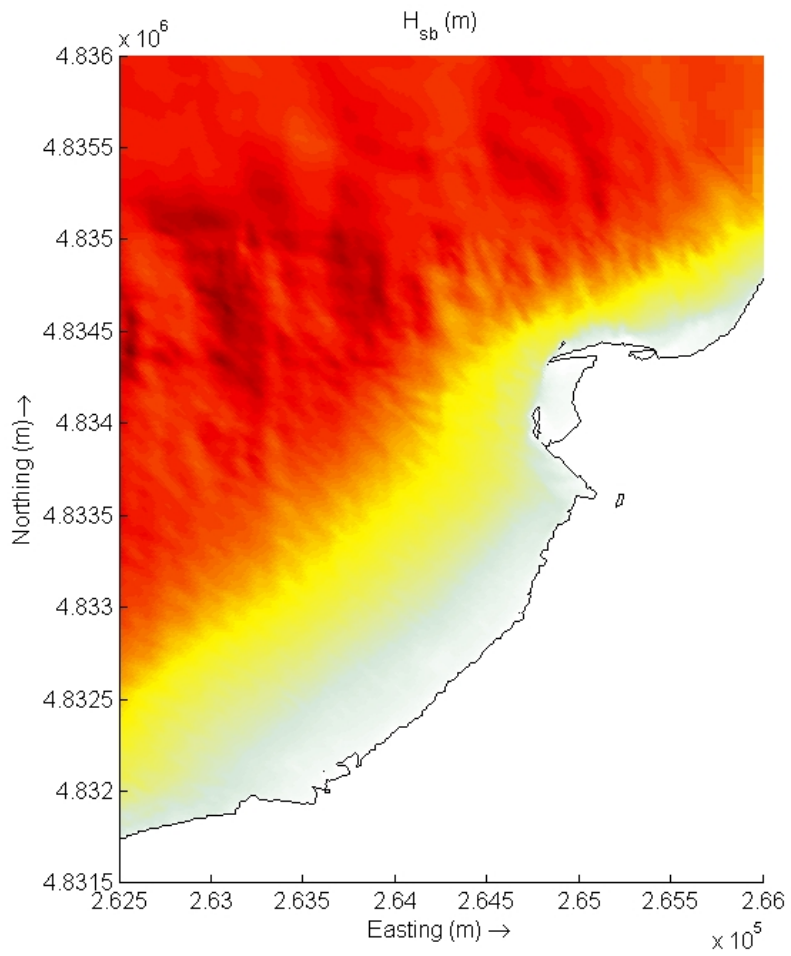


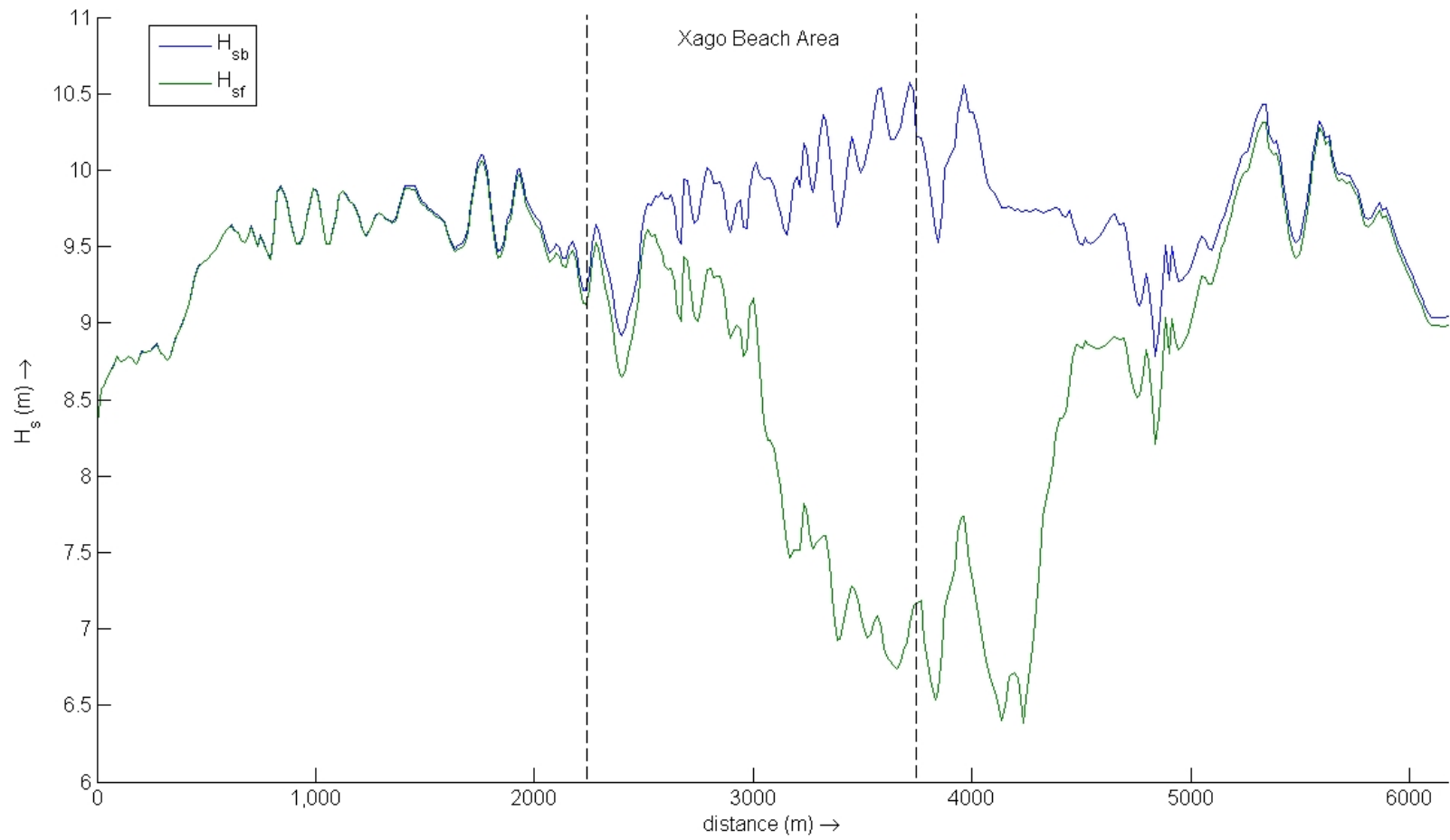


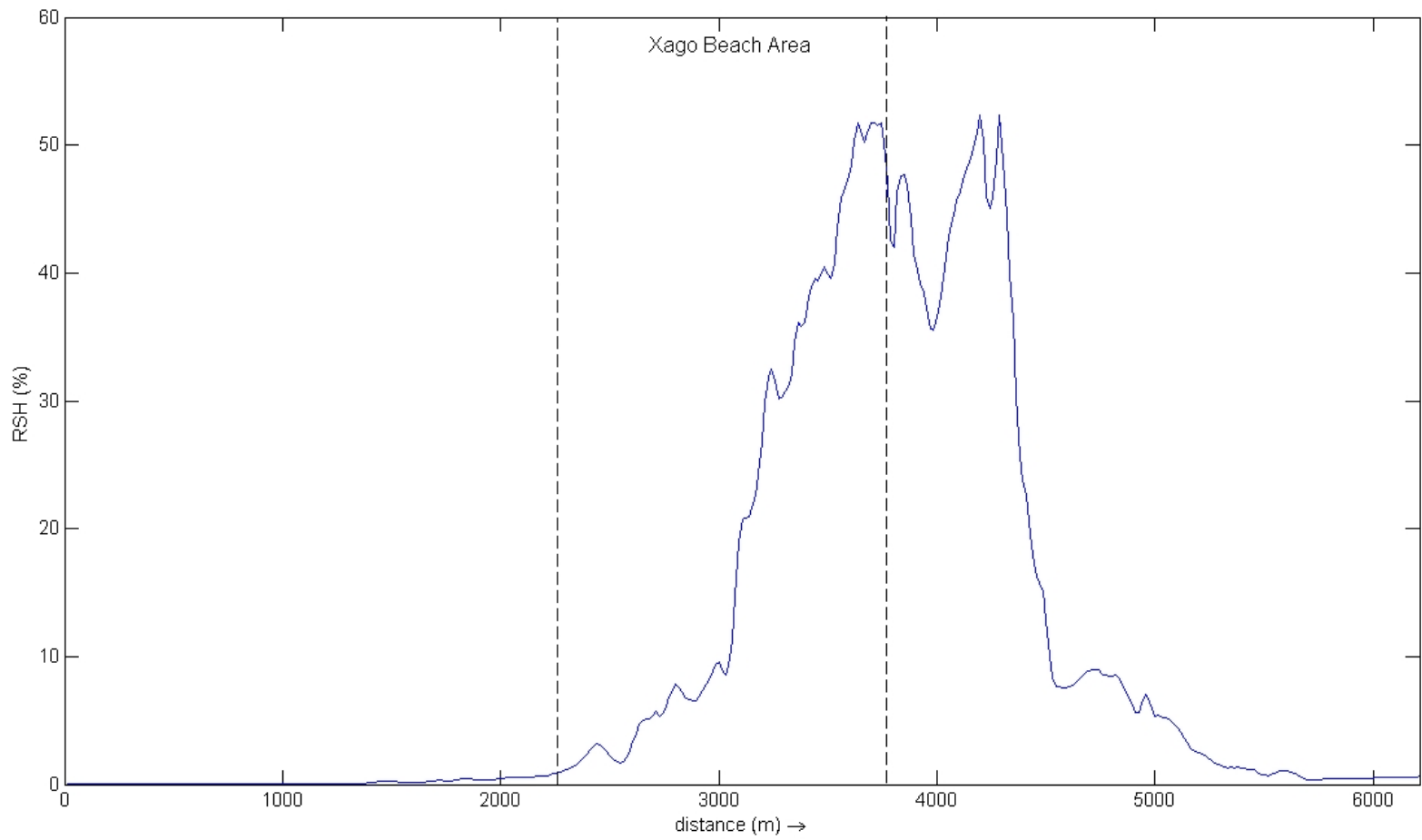


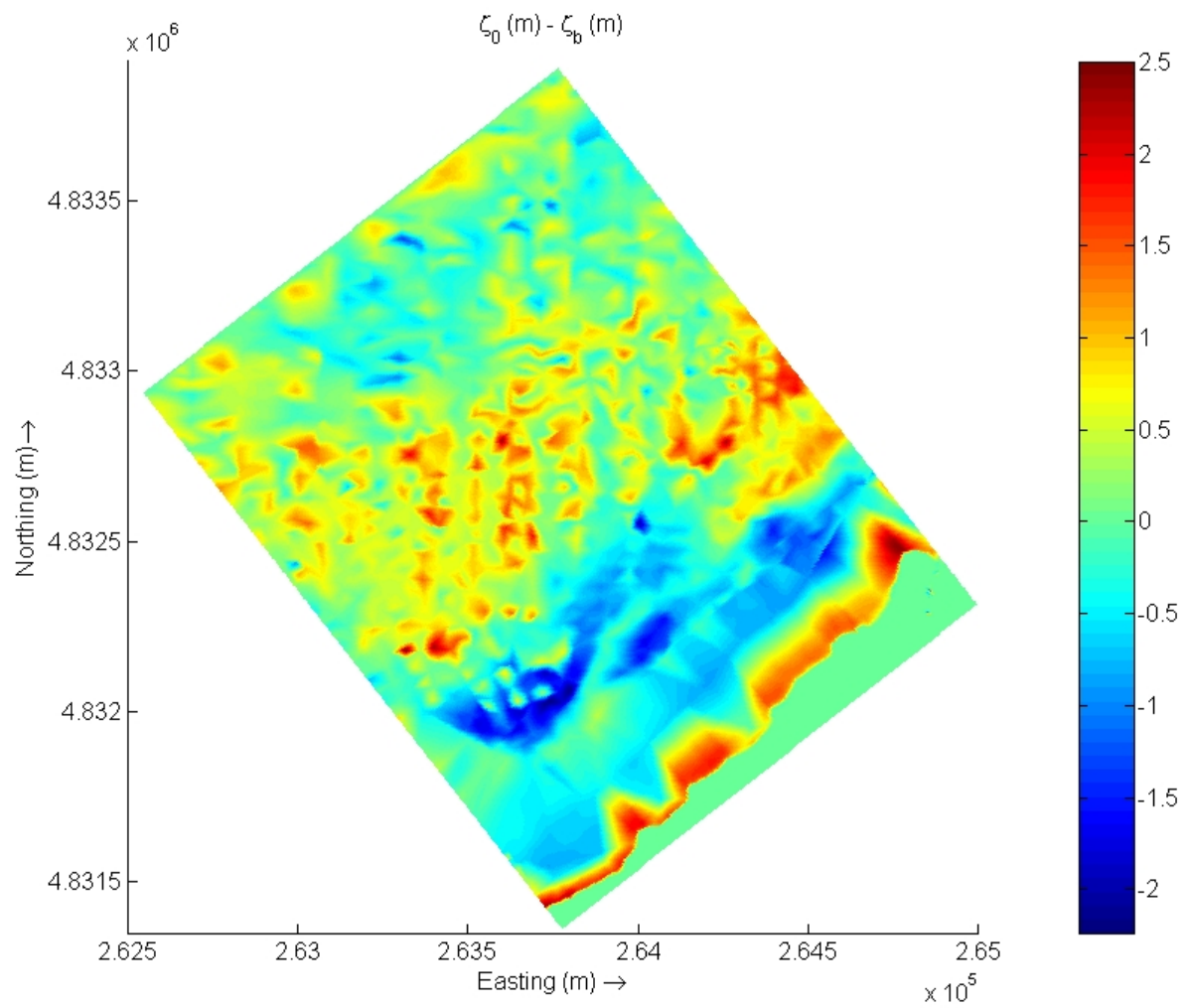


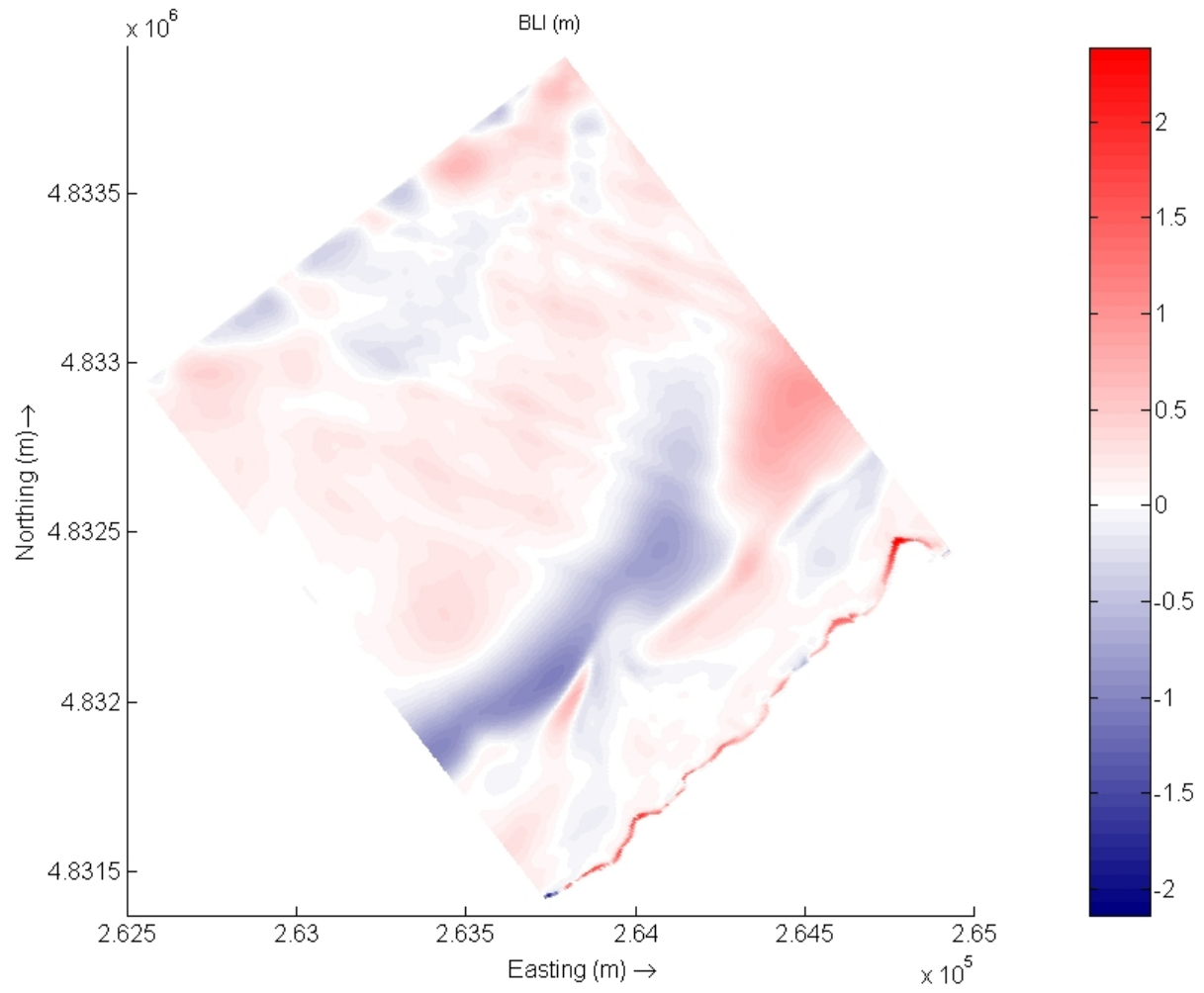




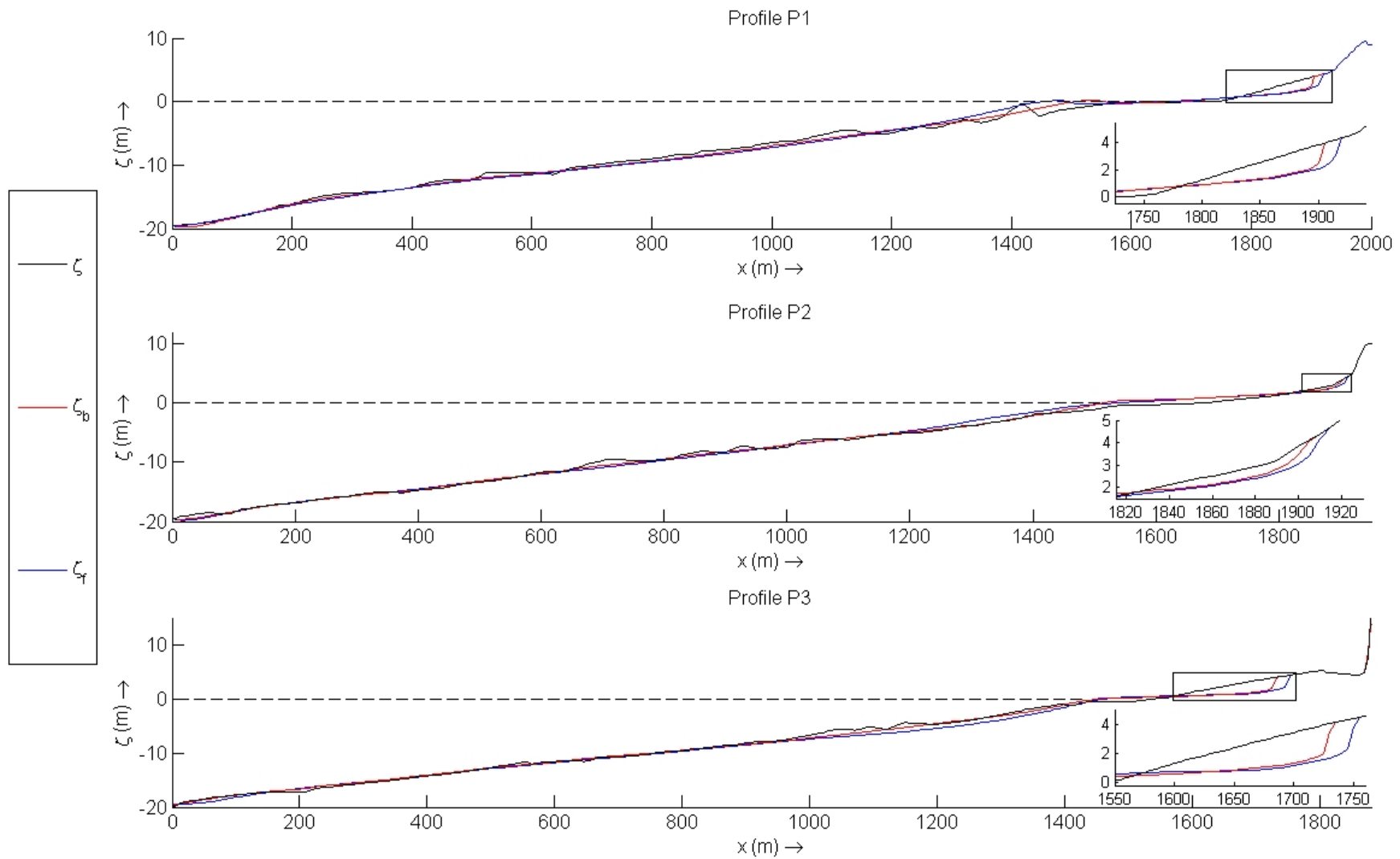


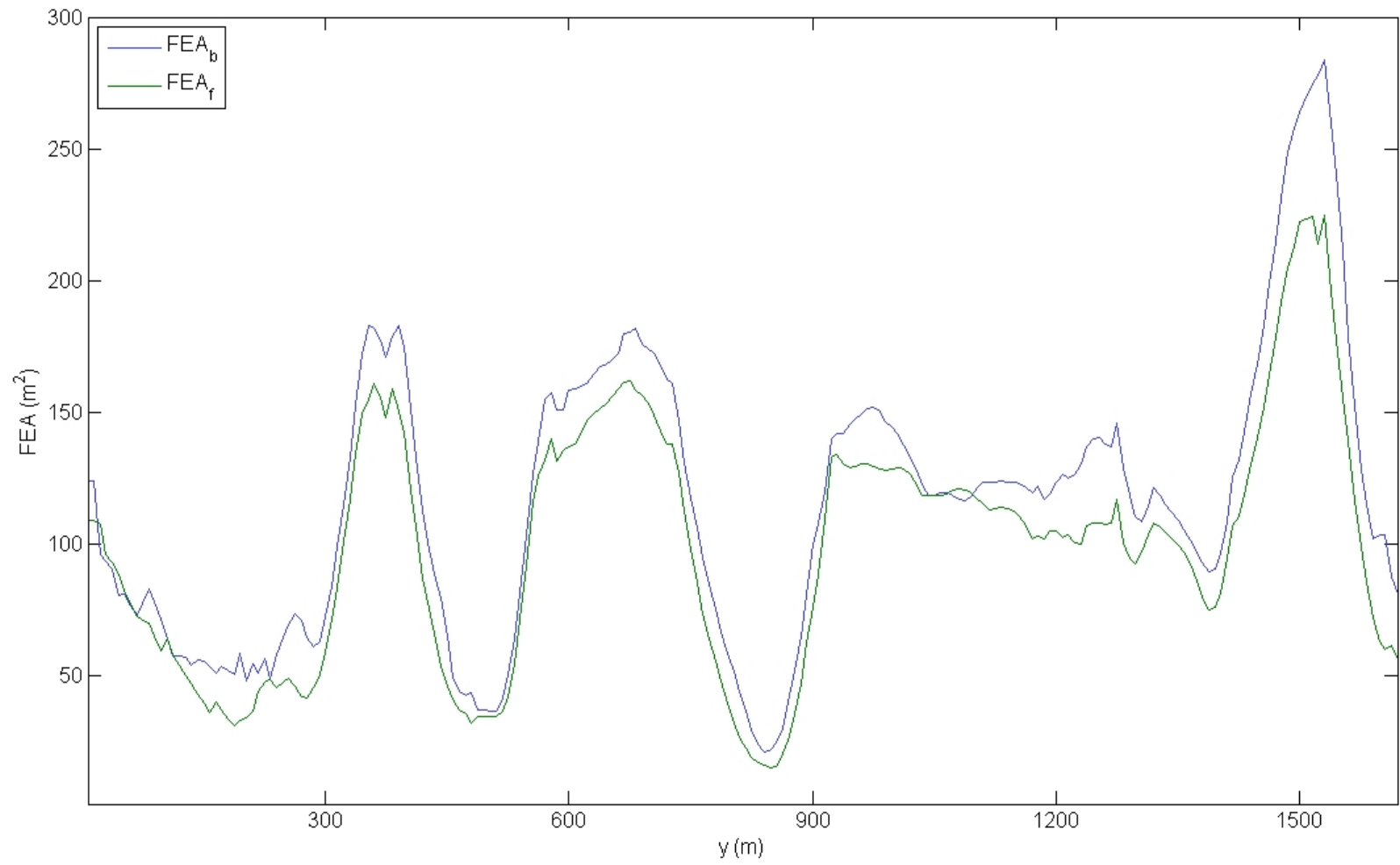


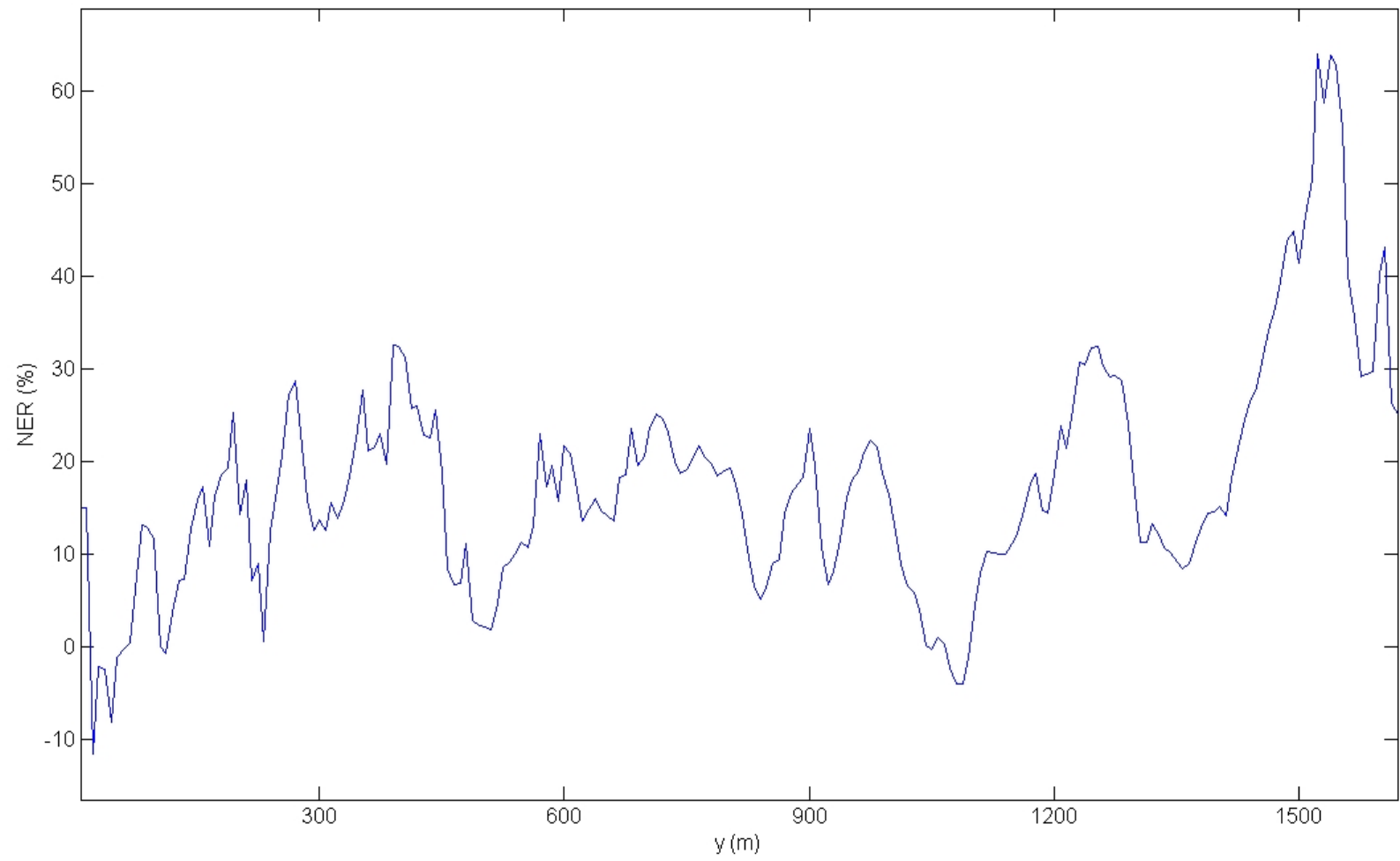












Data	Data available	Error statistics	
		Root Mean Square Error ( <i>RMSE</i> ) [m]	Coefficient of determination ( <i>R</i> <sup>2</sup> ) [-]
Wave buoy data off Avilés	01/03/2010 – 01/09/2010	0.33	0.89
SIMAR44-3085039 off Xago	01/01/2010 – 01/01/2011	0.45	0.92

Table 1: Wave data used to validate the model and values of the error statistics: Root Mean Square Error (*RMSE*) and coefficient of determination (*R*<sup>2</sup>).