

1 **Vibration induced dynamical weakening of pyroclastic flows:**
2 **Insights from rotating drum experiments**

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Abstract

Pyroclastic flows are characterized by their high mobility, which is often attributed to gas-fluidization of the usually fine and/or low density particles. However, the physical mechanism that might drive sustained fluidization of pyroclastic flows over extraordinarily long runout distances is elusive. In this letter it is proposed that a powerful mechanism to weaken the frictional resistance of pyroclastic flows would arise from the prolonged and intense mechanical vibrations that commonly accompany these dense gravitational fluid-particle flows. The behavior of fine powders in a slowly rotating drum subjected to vibrations suggests that fluid-particle relative oscillations in granular beds can effectively promote the pore gas pressure at reduced shear rates. Dynamical weakening, as caused by the enhancement of pore fluid pressure, may be a powerful mechanism in any geophysical process that involves vibrations of granular beds in a viscous fluid. This is particularly relevant for granular flows involving large amounts of fine and/or light particles such as pyroclastic density currents.

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6 I. INTRODUCTION

7 Understanding the dynamics of pyroclastic flows (PF) triggered by the collapse of lava domes
8 and explosive eruptions is crucial for hazard mitigation in populated areas around volcanoes. A
9 main characteristic of PF is their high mobility leading to astonishingly long runout distances even
10 on subhorizontal slopes [3, 8, 12, 33, 34, 36]. PF commonly contain large amounts of fine ash
11 (particle size $d_p \sim 1\text{-}100 \mu\text{m}$ and density $\rho_p \sim 2500 \text{ kg/m}^3$) [28] and light pumice rock fragments
12 ($\rho_p = 500 \text{ kg/m}^3$, $d_p \sim 10 \text{ mm}$) [6] whose dynamics is essentially influenced by the gas-solid hy-
13 drodynamic interaction [9, 28, 32–34, 36]. Thus, an excess of pore gas pressure above atmospheric
14 pressure may easily lead to a fluidization state in which the drag force exerted by the interstitial
15 gas on the particles counterbalances their weight and frictional forces become negligible. In the
16 fluidization regime, the granular flow acquires the behavior of a low viscosity fluid lacking any
17 resistance to shear. Fluidization has long been considered a key process to explain the enhanced
18 mobility of PF [33, 36]. Lab-scale observations on fluidized granular flows in horizontal flumes,
19 where a high pore gas pressure is artificially imposed by continuously injecting gas through the
20 substrate, demonstrate that a state of sustained fluidization leads to essentially infinite runouts
21 [34]. Fluidized PF would propagate as low viscosity fluids over most of their emplacement [33].
22 Accordingly, ignimbrite deposits from prehistoric PF show the signatures of negligible friction
23 and suppressed turbulence at the depositional boundary layer, which is indicative of a fluidization
24 governed dynamics [8].

25 Several mechanisms have been proposed as responsible for the enhancement of pore gas pres-
26 sure leading to fluidization such as the hindered settling of fine particles from the initially flu-
27 idized bed, exsolution of gas from juvenile clasts, engulfing of air at the avalanche front, released
28 gas from rough substrates and high shear rates [3, 12, 34, 36]. Yet, it is uncertain whether the
29 mechanisms considered so far would be intense enough to sustain fluidization of PF over runout
30 distances up to 100 km and across topographic obstacles as inferred from some PF deposits [8].
31 Once the pore gas pressure within a finely grained mixture (with typically low hydraulic perme-
32 ability and high porosity compressibility) is increased by any mechanism it could be maintained
33 for long durations due to retarded pore pressure diffusion, which could explain the considerable
34 runout distances of PF [28, 32]. The key question is then what mechanism plays the major role on
35 rising the pore gas pressure.

36 A further important characteristic of PF is the generation of intense mechanical vibrations usu-

ally by collisions of pyroclasts onto the mountain slope or other obstacles [15, 40]. Seismic signals associated with PF are generally distinguished by significantly long durations and large amplitudes proportional to the volume of PF. Remarkably, seismic signals of similar features are generated by snow avalanches, which also exhibit long runout distances [37]. In this letter we explore the possibility that mechanical vibrations excited by the propagation of PF promote their mobility by enhancing the pore gas pressure. To this end, we have observed the behavior of two types of powders in a rotating drum subjected to mechanical vibrations. The rotating drum setup has been already used in previous studies within the context of PF to assess the role of particle size and density on fluidization [4] and the generation of ash by abrasion of volcanic rock fragments [21]. Here, we use the rotating drum setup to shed light on the effect of mechanical vibrations on the dynamical behavior of PF.

II. MATERIALS AND METHODS

Two diverse types of granular materials have been tested in our work. On one hand, we have observed the behavior of cornstarch (particle size $d_p \simeq 15 \mu\text{m}$ and density $\rho_p \simeq 1550 \text{ kg/m}^3$) as representative of an easily fluidizable fine powder. Cohesiveness of the cornstarch powder is reduced by mixing it with 0.5% by weight of flow control additive (Aerosil[®] from Evonik) [13]. When subjected to a gas flow, a bed of this fine powder reaches a highly expanded fluidized state characterized by a nonbubbling and liquid-like frictionless behavior [41]. On the other hand, we have tested the behavior of glass beads (particle size $d_p \simeq 100 \mu\text{m}$ and density $\rho_p \simeq 2500 \text{ kg/m}^3$). Fluidization of these glass beads is characterized by the development of large gas bubbles just beyond the onset of fluidization, which hampers bed expansion. This is the characteristic behavior of Geldart B granular materials [18, 41].

Figure 1 illustrates a schematic representation of the experimental setup. In our work we have used a cylindrical Plexiglas drum (4.5 cm internal radius and 2cm depth), which is driven by a motor that allows a maximum angular velocity of 100 revolutions per minute (rpm) around its horizontal axis. The drum is rotated by a shaft supported on the base of an electromagnetic vibration exciter through a pair of bearings. The vibrator is driven by a signal generator that provides sinusoidal, vertical vibrations of controlled amplitude ξ_1 , and frequency f (in the range 25-200 Hz). The shaft is fitted at the other end to the motor axis by means of an elastic cardan. Peak vibration velocity $u_1 = \xi_1 2\pi f$ is monitored using a piezoelectric accelerometer. A CCD camera interfaced to a computer for image processing records the profile of the powder as affected

68 by rotation and vibration.

69 When a bed of a cohesive powder is tilted the avalanche angle depends generally on the length
70 of the slope since cohesion leads to a coherence length of size not negligible as compared to the
71 system size for small scale systems [30, 45]. The smaller the length of the slope is, the larger is
72 the angle it can sustain. In the case of the glass beads used in our study cohesion is negligible
73 and therefore effects on the avalanche angle related to the finite size of the drum are not expected.
74 On the other hand, cohesion of the fine powder used in our work has been artificially reduced,
75 which serves to minimize the dependence of the maximum slope angle on its length. As reported
76 in previous studies [45] for a powder similar to the one used in the present work, the avalanche
77 angle of a tilted bed becomes roughly independent of the length of the slope for lengths above
78 $\simeq 8$ cm. Since the diameter of the drum used here is 9 cm, it may be expected that effects on the
79 avalanching behavior due to the finite size of the drum are neither relevant for this material.

80 III. RESULTS AND DISCUSSION

81 As reported in a previous study [9], the two types of materials used in the present work display
82 also contrasting behaviors with increasing drum rotating velocity (see Fig. 2). The fine powder
83 is progressively fluidized by the air that becomes engulfed in the powder by each avalanche. The
84 extent of fluidization increases with the rotation velocity and the whole bed becomes fluidized at
85 a rotation velocity $\Omega \simeq 90$ rpm (Fig. 2(a2)). At this point the granular bed loses any frictional
86 resistance to shear. As may be seen in Fig. 2(a2), the gas-solid mixture acquires an expanded
87 state with a nearly horizontal slope. Particle volume fraction is decreased from $\phi \simeq 0.4$ in the
88 settled (solid) state to $\phi \simeq 0.2$ in the fully expanded (fluidized) state. Using $L \sim 1$ cm for the
89 characteristic size of avalanches, the shear rate is estimated to vary between $\dot{\gamma} \sim \Omega R/L \sim 1$ s⁻¹
90 at the lowest rotation velocity ($\Omega \simeq 4.8$ rpm) and $\dot{\gamma} \sim 40$ s⁻¹ in the full fluidization state ($\Omega \simeq 90$
91 rpm). According to field observations, the thickness of natural PF is usually on the order of tens
92 to several hundred meters whereas PF velocities are in the range ~ 10 ms⁻¹-300 ms⁻¹ [8, 20].
93 Thus, the maximum expected values of shear rate in gas-fluidized PF would be $\dot{\gamma} \sim 10$ s⁻¹ as in the
94 recent Merapi eruption where fine ash PF around 10 m thick propagated at velocities of about 125
95 m/s [20]. Bearing in mind that comparing the observations of small-scale experimental flows and
96 large-scale natural phenomena presents both temporal and spatial scaling problems [8], our results
97 suggest that shear rates in PF are not high enough to cause full fluidization of fine powders.

98 The behavior of fluidizable fine powders in a rotating drum has been analyzed in detail in

99 [10]. Experimental results for several materials and using drums of different diameters show that
 100 the interstitial velocity of the air that continuously escapes from the fluidized powder while the
 101 drum is rotating scales on average proportionally to the tangential velocity of the drum $u_i \sim \alpha \Omega R$,
 102 where $\alpha \sim 0.01$. At the smallest rotation velocity used in the experiments described in the present
 103 paper ($\Omega = 4.8$ rpm), $u_i \sim 0.02$ cm/s, which is not large enough as to fluidize the powder. Thus,
 104 the material displays a plastic behavior. Fluidization starts to occur when the rotation velocity is
 105 further increased leading to interstitial gas velocities similar to the minimum fluidization velocity,
 106 which is $u_{mf} \sim 0.1$ cm/s for the powder analyzed in this work [13]. Figure 2(a3) shows that
 107 the addition of vibration to the slowly rotating drum ($\Omega = 4.8$ rpm) has a relevant effect on the
 108 dynamical behavior of the fine powder. As seen when full fluidization occurs at large rotation
 109 velocities (Fig. 2(a2)), the slowly rotated bed loses progressively its frictional strength as the
 110 intensity of vibrations is increased. For sufficiently strong vibrations, the angle of the slope drops
 111 to zero (Fig. 2(a3)) even at the relatively small shear rates ($\dot{\gamma} \sim 1$ s⁻¹) corresponding to the
 112 lowest rotation velocity. Note however that powder expansion is not observed as in the case of full
 113 fluidization in the rapidly rotating drum (Fig. 2(a2)).

114 In contrast to the fine powder behavior in the rotating drum, the avalanching dynamics of the
 115 glass beads is determined by inertial stresses [9]. In this case, the average angle of the slope is
 116 increased with the rotation velocity as the centrifugal acceleration builds up (Fig. 2(b2)). However,
 117 vibration is seen to weaken the frictional resistance of the glass beads in the slowly rotating drum
 118 too (see Fig. 2(b3)). The minimum vibration intensity to reduce the angle of the slope to zero
 119 depends on the vibration frequency and the material. Thus, it is $u = 4.2$ mm/s at $f = 50$ Hz for
 120 the fine powder (Fig. 2(a3)) whereas, by extrapolating the experimental results (shown below), it
 121 may be estimated that a vibration velocity of about $u \simeq 10$ mm/s would be needed to completely
 122 nullify the frictional resistance of the avalanching glass beads.

123 If a powder bed is sheared it yields plastically when the shear stress reaches a critical value
 124 τ_c (yield strength) usually related to the normal stress σ by means of the Coulomb friction law
 125 $\tau_c = \mu_s \sigma$, where $\mu_s = \tan \theta_s$ is the static friction coefficient and cohesion is neglected [2]. After
 126 failure, the granular bed acquires a state of higher porosity and continues to deform at a slightly
 127 decreased shear stress $\tau_d < \tau_c$. The coefficient of dynamic friction $\mu_d = \frac{\tau_d}{\sigma}$ is thus somewhat
 128 smaller than the static friction coefficient. Assuming that μ_d is a constant which depends only
 129 on the material, its value may be estimated from the average angle of the slope. In the slowly
 130 rotated drum and in the absence of vibrations we obtain $\mu_d \simeq \tan \theta_d \simeq 0.48$ for the fine powder and

131 $\mu_d \simeq 0.54$ for the glass beads . Data on the relative variation of the dynamic friction coefficient $\frac{\Delta\mu_d}{\mu_d}$
132 (estimated from the average angle of the slope) are plotted in Fig. 3a as a function of the relative
133 increase of velocity $\frac{\Delta u}{u_0}$, being u either the tangential rotation velocity (in the absence of vibrations)
134 or the peak vibration velocity in the slowly rotating drum ($\Omega = 4.8$ rpm fixed). Remarkably, the
135 data from both type of experiments for the fine powder conform to a common linear trend, which
136 suggests that the effect of vibration obeys also to the enhancement of the gas-solid hydrodynamic
137 interaction. On the other hand, vibration and rapid rotation yield contrasting effects for the glass
138 beads. Inertial stresses prevail in the rapidly rotating drum for the glass beads, which leads to
139 dynamical strengthening as the centrifugal acceleration increases. Contrarily, the vibrated rotating
140 bed loses progressively frictional resistance as the intensity of vibrations is increased. Thus,
141 one might wonder whether gas-solid hydrodynamic interactions might play also a role on the
142 dynamical weakening of vibrated granular beds of relatively large inertia particles, which would
143 otherwise reach a bubbling fluidization regime only when subjected to large gas-solid relative
144 velocities ($u \gtrsim 0.5$ m/s) [9].

145 The question on the role of gas effects on vibrated granular beds of large inertia grains is
146 longstanding. The interested reader may find a recent review on this subject in [42]. It dates
147 back to Faraday [17] who already observed the onset of convective currents within the bulk of
148 thick layers of large inertia sand grains subjected to vertical vibrations. A strong indication of the
149 relevant role of gas effects was that convection disappeared when air was pumped out and appeared
150 again as the air was readmitted as more recently confirmed by other works [24, 29, 42]. Convective
151 currents in a vibrated granular bed give rise to the formation of a surface heap along which particles
152 avalanches down to be subducted into the bed at its lowest point. In close analogy with our
153 observations, the angle of this slope is smaller than the dynamic friction angle of the material
154 (in the absence of vibrations) and decreases further as the intensity of vibrations is increased.
155 Moreover, chemical engineering studies have long reported that the gas-solid drag can be notably
156 promoted by oscillations even with no net fluid flow [16, 22, 39]. For example, the settling of
157 large inertia beads is substantially slowed down by vertical oscillations of the surrounding fluid
158 and eventually stopped [16, 39]. The oscillating surrounding fluid generates an additional drag
159 that retards settling of the beads and the observed retardation was much greater than that expected
160 from the fluid-solid drag under steady conditions. Thus, it seems likely that an enhancement of
161 the fluid-particle hydrodynamic interaction due to oscillations in dense gravitational flows could
162 lead to dynamical weakening.

163 Our observations demonstrate that dynamical weakening occurs as the rotating velocity is in-
 164 creased due to fluidization in the case of the fine powder [10]. Fluidization is caused by the
 165 increase of pore gas pressure over atmospheric pressure Δp , which leads to a reduction of the
 166 effective friction coefficient μ_{ef} with the rotation velocity. The rise of the pore gas pressure acts by
 167 decreasing the effective normal stress $\sigma_{ef} = \sigma - \Delta p$, where $\sigma = \rho_p \phi g L$ is the powder weight per
 168 unit area, ϕ is the particle volume fraction and L is the typical thickness of the bed. Thus, the bed
 169 would exhibit an effective dynamic friction coefficient given by $\mu_{ef} \sigma = \mu_d (\sigma - \Delta p)$. The relative
 170 decrease of μ_d can be thus obtained from the pressure drop per unit length: $\frac{\Delta \mu_d}{\mu_d} = -\frac{\Delta p}{\sigma} = \frac{1}{\rho_p \phi g} \frac{\Delta p}{L}$.

171 When a fluid flows steadily across a granular bed at low Reynolds numbers, the Carman-Kozeny
 172 equation [7, 31] applies:

$$\frac{\Delta p}{L} = E \frac{\phi^2}{(1-\phi)^3} \frac{\eta}{d_p^2} u_s = \Lambda n_0 F_S \quad (1)$$

173 where u_s is the superficial gas velocity defined as the gas flow rate per unit area. E is an empirical
 174 constant ($E \simeq 180$ for spheres), $d_p = 2R$ is particle size, $n_0 = \frac{3\phi}{4\pi R^3}$ is the number of particles
 175 per unit volume, $F_S = 6\pi\eta R u_s$ is the Stokes drag force, η is the fluid's dynamic viscosity and
 176 $\Lambda \simeq \frac{10\phi}{(1-\phi)^3}$ is the factor that takes into account the hydrodynamic interactions within the bed
 177 [42]. In the case of a vibrated bed, gas-particle relative oscillations (oscillatory flow) would be
 178 established instead of a steady flow. The drag force on a sphere of radius R undergoing oscillations
 179 in a viscous fluid can be calculated in the limit of either small oscillations amplitude ($\frac{\xi_1}{R} < 1$) or
 180 small Reynolds number ($Re_1 = \frac{u_1 \rho R}{\eta} = \frac{\xi_1}{R} \left(\frac{R}{\delta}\right)^2 < 1$) as

$$F_1(t) = 6\pi\eta R \left(1 + \frac{R}{\delta}\right) u_1(t) + 3\pi R^2 \sqrt{\frac{2\eta\rho}{\omega}} \left(1 + \frac{2R}{9\delta}\right) \frac{du_1}{dt} \quad (2)$$

181 where $u_1(t)$ is the instantaneous oscillation velocity, $u_1 = \xi_1 2\pi f$ is the peak oscillation velocity,
 182 ρ is the fluid density, and $\delta = \sqrt{\frac{\eta}{\rho\omega}}$ is the thickness of the Stokes boundary layer surrounding the
 183 sphere across which the fluid flow becomes irrotational [23]. If the interaction between the Stokes
 184 boundary layers developed around neighbor spheres is neglected, the Carman-Kozeny equation
 185 can be adapted to oscillatory flows through granular beds [44], which leads to a root mean square
 186 (rms) pressure drop per unit length

$$\frac{\Delta p'_1}{L} = \Lambda \Upsilon n_0 F'_S \quad (3)$$

187 where $F'_S = 6\pi\eta R u'_1$, u'_1 is the rms oscillation velocity ($u'_1 = \frac{u_1}{\sqrt{2}}$) and

$$\Upsilon = \left[\left(1 + \frac{R}{\delta}\right)^2 + \left(\frac{R}{\delta}\right)^2 \left(1 + \frac{2R}{9\delta}\right)^2 \right]^{1/2} \quad (4)$$

188 Thus, we can write

$$\frac{\Delta\mu_d}{\mu_d} = -\beta\Upsilon u'_1 \quad (5)$$

189 where $\beta = \frac{6\pi\eta R\Lambda n_0}{\rho_p\phi g}$.

190 For a randomly packed bed ($\phi \simeq 0.6$) of the noncohesive glass beads used in our tests $\beta \simeq$
 191 120. On the other hand, the analysis of the fine powder is more subtle. Fine cohesive particles
 192 are usually agglomerated due to the prevalence of the attractive force between the particles as
 193 compared to their weight [11]. By considering agglomerates as effective spheres, the volume
 194 fraction filled by the agglomerates ϕ^* is simply related to the particle volume fraction ϕ by $\phi^* =$
 195 $\phi \frac{k^3}{N}$, where N is the average number of particles per agglomerate and $k = \frac{d^*}{d_p}$ is the ratio of the
 196 agglomerates size to the size of the individual particles. Fluidized agglomerates screen the external
 197 flow field and can be treated as effective particles of density $\rho^* = \rho_p \frac{N}{k^3}$ [46]. Settling tests reported
 198 elsewhere [43] yield $N \simeq 9.1$, $k \simeq 2.44$ for the powder used in the present work. Thus we may
 199 use $d^* \simeq 37 \mu\text{m}$ and $\rho^* \simeq 990 \text{ kg/m}^3$ for the size and density of the agglomerates as effective
 200 particles, respectively. The packing density measured for this powder under flow conditions [43]
 201 is in the range $\phi \simeq 0.2 - 0.3$, which yields $\phi^* \simeq 0.3 - 0.5$. Using these numbers, we would predict
 202 $\beta \simeq 200 - 800$.

203 Experimental data of $\frac{\Delta\mu_d}{\mu_d}$ versus $\Upsilon u'_1$ for the slowly rotating drum with added vibrations of
 204 increasing intensity are plotted in Fig. 3b. The best linear fits yield $\beta \simeq 116$ for the glass beads
 205 and $\beta \simeq 285$ for the fine powder, which are close to the theoretically expected values as caused
 206 by an increase of the rms pore air pressure. Data obtained for the fluidizable fine powder in the
 207 non-vibrated drum with increasing rotation velocity are plotted also in Fig. 3b. In this case $\Upsilon = 1$
 208 and u'_1 is replaced by the superficial gas velocity u_s , which is related to the interstitial gas velocity
 209 u_i by $u_s = (1 - \phi^*)u_i$. Remarkably, the data fits to the same trend of the vibrated drum data
 210 for $u_s = 0.0065\Omega R$, which is consistent with the scaling of the interstitial gas velocity with the
 211 tangential rotation velocity of the drum reported elsewhere ($u_i \sim 0.01\Omega R$) [10].

212 As a general comment, it must be noticed that the diameter of the drum used in our experiments
 213 is small. Ideally, only drums of very large size would allow dismissing any effect of boundary
 214 curvature on powder flow [4]. However, the application of vibrations to heavy drums would be

215 technically difficult using our setup. Nevertheless, the experiments carried out in our work serve
216 to capture the essential effect of vibrations on dense granular flows, which can be explained from
217 a simple physical model where boundary effects are neglected. Additional experimental work
218 should be pursued in future studies to address the possible role of surface flow curvature. A key
219 issue in the interpretation of our results is that high pore gas pressure was generated within the
220 granular media by vibrations, which leads to a reduction of the dynamic friction angle. It should
221 be remarked however that pore gas pressure was not measured in our experiments and this would
222 be a further interesting subject for future work. On the other hand, recent experimental results do
223 show that the pore gas pressure in a granular bed subjected to a gas flow is enhanced when high
224 intensity sound vibrations are applied [47]. As a result, the minimum gas fluidization velocity
225 is decreased in a similar way that fluidization is observed in our rotating drum experiment at re-
226 duced rotation velocities when mechanical vibration is applied. Application of either mechanical
227 or sound vibrations would equivalently promote the pore gas pressure due to the enhancement of
228 drag by gas-solid relative oscillations. Finally, it must be stressed that vibration induced dynamical
229 weakening could be promoted by other physical mechanisms that might prevail in some particular
230 situations as proposed elsewhere [1, 24, 26, 27]. For example, according to the *acoustic flu-*
231 *idization* mechanism proposed by Melosh [26, 27], high-amplitude and high-frequency vibrations
232 would be also capable of generating transient mechanical stresses intense enough as to facilitate
233 failure under a reduced shear stress in geophysical granular flows (regardless of the presence of
234 interstitial fluids) although the consistency of this theory was later questioned [35].

235 IV. CONCLUSIONS

236 Our observations suggest that vibrations may promote fluidization of dense granular flows by
237 enhancing the pore fluid pressure. Thus, the intense and prolonged mechanical vibrations that can
238 arise in nature as dense granular flows propagate on terrains with irregular surfaces [48] could
239 contribute to the sustained fluidization of these flows over long runout distances. Vibration en-
240 hanced fluidization would be a main mechanism specially in the case of pyroclastic flows mostly
241 consisting of a matrix of easily fluidizable fine ash and light pumice particles. In particular, if the
242 ratio $\frac{R}{\delta}$ is very large, vibrations would lead to a notable enhancement of the rms pore gas pressure
243 at small shear rates (by a factor $\Upsilon \simeq \frac{2}{9} \left(\frac{R}{\delta}\right)^2 \gg 1$). This can be the case of large and light pumice
244 rock fragments found in ignimbrite deposits of pyroclastic flows with very large runout distances,
245 which are easily fluidizable due to their very low density [8] (using $R = 2$ cm, it is $\Upsilon \sim 100$ for

246 an oscillation frequency of 10 Hz). Dynamical weakening would be further promoted for high
247 vibration frequencies that may be produced by the scattering of waves in heterogeneities as it is
248 believed to occur in fault gouges [27].

249 The enhancement of pore gas pressure by oscillatory flows could be also a potentially relevant
250 mechanism on the dynamical weakening of granular materials observed in other geological pro-
251 cesses such as seismic faulting [5], landslides [25], detachment faulting [19] and impact crater
252 formation [26], although pore pressure generation in gas-particle systems by oscillatory flows is
253 likely to be efficient only for small and/or light particles. Dynamical weakening by oscillatory
254 flows would be more pronounced if the interstitial fluid is a liquid as might occur in dense gravi-
255 tational liquid-particle flows (lahars) usually triggered by heavy rains on unconsolidated materials
256 such as ash, sand, or gravel, which are accompanied by mechanical vibrations of very long dura-
257 tion and high frequencies [48].

258 In the present study it has been assumed that pore gas compressibility can be ignored, which
259 is justifiable for small scale systems at ambient temperature [28]. However, gas compressibility
260 cannot be neglected in the case of large scale gas-particle mixtures such as pyroclastic flows and
261 snow avalanches, where it would arguably have a relevant effect on the pore pressure diffusion
262 process [28]. In addition, a temperature gradient across a porous solid may lead to a great ampli-
263 fication of acoustic oscillations as gas parcels within the bed are compressed and expanded by the
264 oscillating pressure. Likewise, intense pore gas oscillations may create a significant temperature
265 gradient across the material. This is the so-called thermoacoustic effect, which is at the basis of
266 thermoacoustic engines and refrigerators [38]. An open issue that would deserve further study is
267 whether temperature gradients commonly expected in pyroclastic flows [14] could favor sustained
268 fluidization by enhancing acoustic oscillations of the pore fluid.

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