

Organised turbulence over mobile and immobile hydraulically rough boundaries

Rui M. L. Ferreira⁽¹⁾, Mário J. Franca⁽¹⁾, João G. B. Leal⁽²⁾ and António H. Cardoso⁽¹⁾

⁽¹⁾ CEHIDRO-Instituto Superior Técnico, TULisbon, Av. Rovisco Pais, 1049-001 Lisboa, Portugal; ruif@civil.ist.utl.pt

⁽²⁾ Dept. Civil Engr. and IMAR-CMA, University of Coimbra, Pólo II, 3030-788 Coimbra, Portugal; mfranca@dec.uc.pt

⁽³⁾ Dept. of Civil Engineering, Faculty of Sciences and Technology, UNL, Quinta da Torre, 2829-516 Caparica, Portugal; e-mail: jleal@fct.unl.pt

⁽⁴⁾ CEHIDRO-Instituto Superior Técnico, TULisbon, Av. Rovisco Pais, 1049-001 Lisboa, Portugal, ahc@civil.ist.ut.pt

ABSTRACT

The present work is aimed at the study of near-bed organised turbulence over mobile and immobile, porous, hydraulically rough boundaries. The bed was permeable and composed of non-cohesive sediments. Two data sets were analysed, characterised by the same u_* . The mobile bed data featured generalised sediment transport for all size fractions smaller than the d_{90} . The fixed bed was obtained as result of an armouring process. Comparison of these data sets reveals differences on such parameters of the bursting cycle as the maximum shear stress and the transported momentum. These results point to a reorganization of turbulence, in the near-bed region, when the bed is mobile. The impacts of these reorganization are discussed, namely in what concerns the third order moments of the distributions of the velocity fluctuations.

Keywords: Turbulent coherent structures, Laser Doppler Anemometry, mobile bed.

1 INTRODUCTION

Bursting phenomena form a class of coherent structures, i.e., organised turbulence, responsible for generating turbulent energy and shear stress (Grass 1971, Nakagawa & Nezu 1977). The complete cycle is composed by an ejection event and a sweep event, mediated by outward or inward interactions.

Accurate visualization techniques and non-intrusive measuring capabilities allowed for more sophisticated detection techniques and for advances in phenomenological descriptions of coherent turbulence (Franca & Lemmin 2006, Hurther et al. 2007) and fluid-particle interactions (*cf.* e.g. Nelson *et al.* 1995, Ferreira et al. 2002). In spite of these conceptual progresses, the increased mobility of coarse size fractions in the presence of finer fractions in the bedload (MacAuley & Pender (1999) and Ferreira (2005), pp. 139-151, among others) remains to be fully understood.

Ferreira et al. (2007) investigated the role of particle arrangement and found that the competent velocity for the entrainment of coarse fractions is dependent on the presence of fine sediments in the bed mixture. One complementary hypothesis is that, under generalized sediment transport conditions, a reorganization of the near-

bed turbulence promotes a greater transfer of momentum to the bed and, thus, increases the probability of entrainment of coarse particles.

The present work pursues this line of research and is aimed at the identification of the differences exhibited by organized turbulence in the inner layer of open-channel flows over mobile and immobile beds. The latter is porous and hydraulically rough. In particular, it is sought to evaluate the statistical characteristics of sweeps and ejections over mobile and immobile beds, both in the roughness dominated layer and in the upper region of the logarithmic layer.

To accomplish these objectives, two sets of instantaneous velocity data, corresponding to mobile and immobile beds, were obtained with LDA instrumentation in an open-channel. Nakagawa & Nezu's (1977) quadrant threshold analysis was modified to detect sweep and ejection events. Relevant statistics of the events were computed and analysed. The two sets of data have, approximately, the same value of the friction velocity, u_* , in order to overcome normalization errors.

2 LABORATORY FACILITIES AND INSTRUMENTATION

The laboratory data were obtained at a 12 m long and 40 cm wide recirculating and tilting flume with glass side-walls. The bed was permeable and composed of cohesionless sediments. Two sets of data were obtained. Set E1 was obtained under uniform flow and recirculation conditions. E1D was an armoured bed test (details in Ferreira 2005, pp. 68-73). To maintain approximately the same mean bed shear stress in both data sets, the slope was slightly adjusted.

Two orthogonal components of the instantaneous flow velocity were measured with a monochromatic He-Ne forward scatter Laser Doppler Anemometer operated at 240 Hz. A laser bed-profiler was employed to evaluate the bed texture and the water depth and the bed elevation were measured with a 0.5 mm precision point-gage. Sediment discharge was computed from dry-weighted samples. For both data sets the flow discharge was $Q = 13,5$ l/s, the initial bed slope was $i_0 = 0.0025$. The remaining main features of the two experimental data sets are shown in Table 1.

Table 1. Main characteristics of the data sets.

Set	h_0 (m)	u_* (m/s)	B/h_0 (-)	Fr (-)	θ_{90} (-)	θ_{c90} (-)	d_{90b} (mm)	d_{90BL} (mm)	k_s (mm)	$\phi \times 10^{-3}$ (-)
E1	0.069	0.042	5.8	0.59	0.021	0.031	5.49	2.86	5.6	3.9
E1D	0.072	0.043	5.6	0.56	0.021	0.032	5.87	-	7.3	-

In Table 1 h_0 is the water depth, u_* is the friction velocity, B/h_0 is the aspect ratio, Fr is the Froude number, θ_i is the Shields parameter for the size fraction represented by the diameter d_i , $\theta_{ci} = \theta_c \xi_i$, is the critical Shields parameter computed from the initial bed mixture (incorporates an hiding-exposure coefficient, ξ_i , Ferreira *et al.* 2007) where $\theta_c = 0.05$, k_s is the roughness scale (computed as in Ferreira *et al.* 2008); ϕ_{BL} , the non-dimensional sediment transport. The subscripts “b” and “BL” stand for “bed surface” and “bedload”, respectively.

3 DATA ANALYSIS

3.1 Conditional sampling and detection criteria

A conditional sampling technique was employed to identify the events of the bursting cycle. The proposed detection criterion was based on the u' vs. v' threshold quadrant technique, chosen for its conceptual simplicity. In its standard version (Nakagawa & Nezu 1977), a sweep occurs when, in quadrant IV, $|u'v'| > \sigma_h^{(IV)}$ and an ejection is identified when, in quadrant II, $|u'v'| > \sigma_h^{(II)}$, where $\sigma_h^{(IV)}$ and $\sigma_h^{(II)}$ stand for the detection thresholds (u' and v' are the fluctuations of the longitudinal and vertical velocities, respectively).

This procedure may lead to faulty characterization: a given turbulent event, seen by direct visualisation, may not be detected as so but as a train of separated smaller events. This is illustrated in Figure 1 where a clear sweep event, occurred between $t = 6.6$ s and $t = 6.7$ s, features a persistent positive value of u' and an *essentially* negative value of v' . Because of the fluctuation in the value of v' , this event would be recognised as train of 5 events of smaller duration resulting in underestimation of the values of the mean of the duration and of the period.

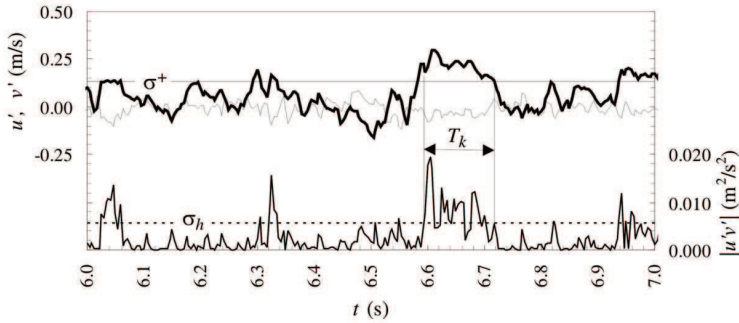


Figure 1. Top: detail of u' (—) and v' (- - -) series. Bottom: detail of a $|u'v'|$ series. Data series belonging to the E1 data set.

A simple method to enhance the results of the quadrant threshold method is based on the observation that the persistence of u' is well correlated with the persistence of the event. The detection of the initiation of ejections or sweeps remains that of the standard quadrant threshold method but the event is considered to last while the absolute value of u' is above a given threshold, σ^+ for sweeps and σ^- for ejections. The domains of occurrence of each event, once detected, are:

$$\text{i) outward interactions, } Q_{out} = \left\{ u', v' \in : u' > 0 \wedge v' > \frac{\sigma_h^{(I)}}{u'} \wedge u' < \sigma^+ \right\};$$

$$\text{ii) ejections, } Q_{ej} = \left\{ u', v' \in : u' < 0 \wedge \left\{ v' > \frac{\sigma_h^{(II)}}{|u'|} \vee |u'| > \sigma^- \right\} \right\};$$

iii) inward interactions, $Q_{in} = \left\{ u', v' \in : u' < 0 \wedge |v'| > \frac{\sigma_h^{(III)}}{|u'|} \wedge |u'| < \sigma^- \right\}$;

iv) seeps, $Q_{sw} = \left\{ u', v' \in : u' > 0 \wedge \left\{ |v'| > \frac{\sigma_h^{(IV)}}{u'} \vee u' > \sigma^+ \right\} \right\}$.

After a survey of the data, it was adopted $\sigma_h^{(I)} = \sigma_h^{(II)} = \sigma_h^{(III)} = \sigma_h^{(IV)} = \sigma_h$ with $\sigma_h = H u_{rms} v_{rms}$ where u_{rms} and v_{rms} are the turbulent intensities associated to the longitudinal and normal velocities, respectively, and H is the hole size. The value of the latter, $H = 1.7$, was determined heuristically and was found to be close to the half-value threshold level (Nezu & Nakagawa 1993, p. 182). The values of σ^+ and σ^- depend on the value of σ_h . It was adopted $\sigma^+ = 2.5 u_{rms}$ and $\sigma^- = -2.5 u_{rms}$.

The domains of persistence of each of the events can be seen in Figure 2 for two sets of data taken at $y/k_s = 0.8$, where y is the normal (upwards) coordinate.

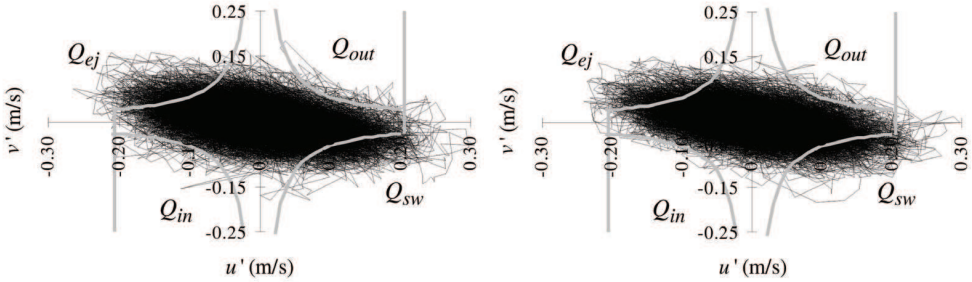


Figure 2. Quadrant plot of the instantaneous velocity data with thresholds. Left: data set E1 (mobile bed). Right: data set E1D (immobile bed). Gray lines (—) stand for the $|u'v'|$ and $|u'|$ thresholds.

The data plotted in Figure 2 was treated to eliminate spurious points – values that differ more than 600% from the average of neighbouring points. In data sets of 12000 entries, bad data were no more than 22 points.

3.2 Event statistics

Four parameters will be used to characterize the sweep and the ejection events of the bursting cycle: the duration, T_k , the transported momentum, M_k , the maximum shear stress, A_k (in absolute value) and the period, P_k , computed as seen in Figure 3.

For each data series, a record of the detected events is created and analysed. All events with a duration smaller than $2dt$, where $dt = 1/f_{LDA}$ is the sampling interval, were eliminated. The N_j valid events in a given data series j are used to compute the average values $\bar{X}^{(j)}$.

Three data series were collected in E1 and in E1D. The ensemble averages of the parameters that characterize the events are $\langle \bar{X} \rangle = \frac{1}{3} \sum_{j=1}^3 \bar{X}^{(j)}$, where $\langle \bar{X} \rangle$ stands

for $\langle \bar{T} \rangle$, $\langle \bar{M} \rangle$, $\langle \bar{A} \rangle$ or $\langle \bar{P} \rangle$. To simplify the notation, the brackets and the overbar will be dropped.

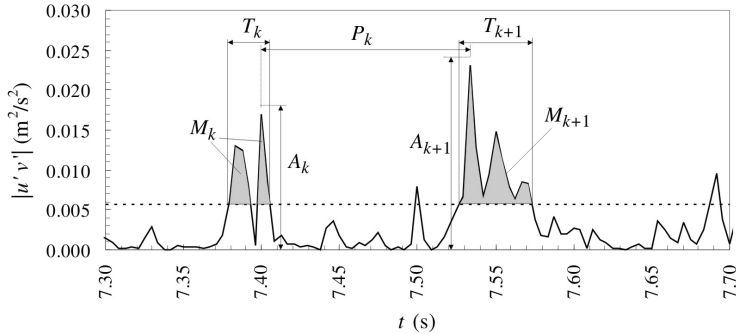


Figure 3. Detail of a $|u'v'|$ series with definition of the parameters that characterize sweep and ejection events.

4 EXPERIMENTAL RESULTS

To compare the turbulent structure of the flow for sets E1 and E1D the statistical distributions of the parameters that characterize events sweep and ejection (see Figure 3), in both mobile and immobile bed conditions, are computed and compared. The comparative analysis is made at two reference positions within the water column: in the roughness layer ($y/k_s = 0.80$) and in the upper region of the logarithmic layer ($y/h = 0.19$). Emphasis will be given to the results in the roughness layer. Figures 4 and 5 present the statistics of sweep and ejections events, respectively, measured at $y/k_s = 0.80$ for the mobile and the immobile data sets.

The visible effect of a hypothetical reorganization of the turbulent structure due to the presence of mobile sediment occurs mainly on the sweep domain and concerns the transported momentum and the extreme value of the events. In the mobile bed case, larger values of M and of A are more frequent: the average value of M is $0.099 \text{ m}^2/\text{s}$ and $0.075 \text{ m}^2/\text{s}$ for the mobile bed and the immobile bed, respectively. The average value of A , it is $0.012 \text{ m}^2/\text{s}^2$ and $0.010 \text{ m}^2/\text{s}^2$ for the mobile bed and the immobile bed, respectively.

The statistics of the extremes shows even more discrepant values: the extreme maximum value of A is $0.028 \text{ m}^2/\text{s}^2$, for the mobile bed data set, and $0.023 \text{ m}^2/\text{s}^2$, for the immobile bed data set. The extreme values of the transported momentum are $0.819 \text{ m}^2/\text{s}$ and $0.400 \text{ m}^2/\text{s}$ for the mobile and immobile bed data sets, respectively. This configures an increased tendency to anisotropy near the bed in the mobile bed case. This claim is corroborated by the quadrant plot of the data (*cf.* Figure 2) where a higher dispersion of the data cloud is observed in the mobile bed data set.

For the ejection event, the histogram of A shows a deviation to the right, indicative of a higher frequency of extreme events. It was found that the extremes were $0.030 \text{ m}^2/\text{s}^2$ and $0.026 \text{ m}^2/\text{s}^2$ for the mobile and the immobile bed data sets, respectively. This tendency is, however less pronounced than in the sweep event case: the average values of A are $0.011 \text{ m}^2/\text{s}^2$ and $0.010 \text{ m}^2/\text{s}^2$ for the mobile and the immobile bed data sets, respectively.

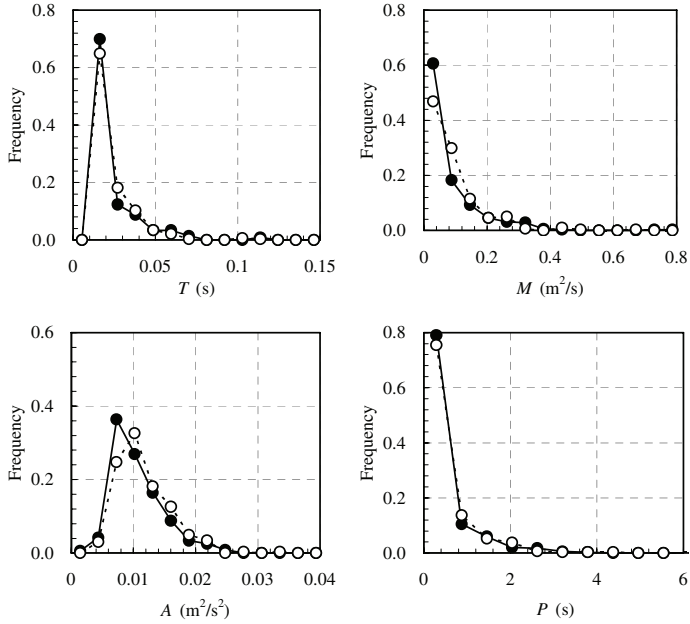


Figure 4. Histograms of duration (T), transported momentum (M), maximum shear stress (A) and period (P) of sweep events at $y/k_s = 0.80$. Comparison of data sets E1, mobile bed, (--○--) and E1D, immobile bed, (—●—).

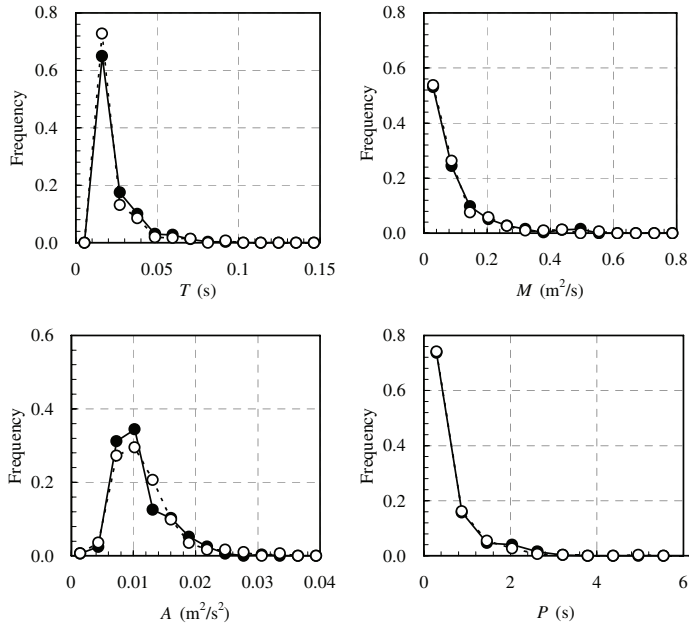


Figure 5. Histograms of the parameters characterizing the ejection events at $y/k_s = 0.80$. Remaining caption as in Figure 4.

Since the shear stress for both data sets is approximately the same, the higher

frequency of extreme sweeps and ejections in the mobile bed case must be compensated by stronger interactions. Observing the data clouds depicted in Figure 2, it is concluded that that is the case: the mobile bed data is characterized by a greater dispersion in all quadrants.

At the measuring position $y/h = 0.19$, at the upper region of the logarithmic layer, no appreciable differences are found on the distributions of any of the parameters that characterize sweep events and the ejection events. For comparison, the transported momentum is shown in Figure 6.

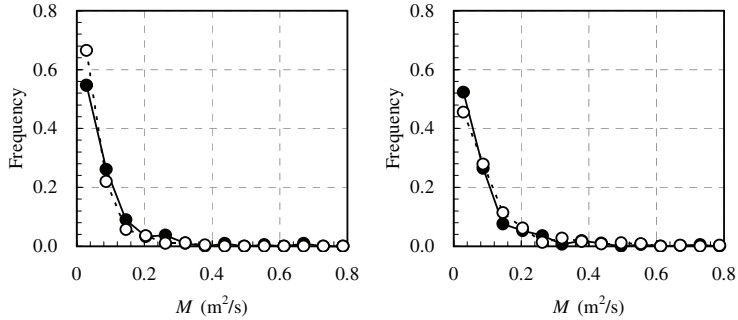


Figure 6. Histograms of the transported momentum characterizing the sweep (left) and ejection (right) events at $y/h = 0.19$. Remaining caption as in Figure 4.

It thus appears that the effects of the mobile bed are confined to the near-bed region and no significant reorganisation of turbulence occurs in the upper regions of the logarithmic region.

It is worthy of notice that only the transported momentum and the maximum value of the shear stress associated to the events seem to be affected by the existence of mobile sediment near the bed. The duration of the events and the period between events seem to remain unaltered. Only the values of the skewness of the period are slightly different (3.11 for the mobile bed data set and 2.11 for the immobile bed data set), indicating that the spacing of extreme ejection events may be greater in the mobile bed case.

5 DISCUSSION AND CONCLUSIONS

The analysis of the data sets E1 and E1D reveals that the ensemble average of the turbulent intensities u_{rms} and v_{rms} is approximately insensitive to sediment mobility (Table 2).

The second moments of the instantaneous velocity distributions are essentially not affected by the apparent reorganization of near-bed turbulence. Hence, these effects must be felt in higher order moments. Observing Table 2, and keeping in mind that the quality of the estimates of third-order moments is lower than that of standard deviation, it may be the case that the effect of the turbulence reorganization is felt on skewness of the longitudinal velocity distribution. The larger positive skewness displayed by data set E1 (mobile bed) signals the fact that larger instantaneous velocities are felt near the bed. This would, in turn, be responsible by increased transfer of momentum to mobile particles, namely to coarser grain sizes.

Table 2. Ensemble averages of 2nd and 3rd order moments at $y/k_s = 0.8$

	Data set E1	Data set E1D
$u_{\text{rms}}/u_* (-)$	1.96	1.93
$v_{\text{rms}}/u_* (-)$	0.95	0.91
Skewness(u') (-)	0.110	0.073
Skewness(v') (-)	0.062	0.065

It is not clear, however, if the increased mobility of coarse sediment is a cause or an effect of the increased intensity of the coherent events. Further experimental work will be undertaken to clarify this issue.

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