

# The landscape model: a model for exploring trade-offs between agricultural production and the environment

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<sup>2</sup> The Landscape Model: a model for exploring trade-offs

- <sup>3</sup> between agricultural production and the environment.
- 4
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- 15
- 16 HIGHLIGHTS
- Understanding trade-offs between yield and environment is essential for SI
- The Landscape Model aids the understanding of crop-soil-water interactions
- Model validated against 50 years of data from two long-term experiments
- Model validated against spatially-explicit data from the North Wyke farm platform
- The model simulated wheat yield, grain N and grain P particularly well

22

#### 24 ABSTRACT

We describe a model framework that simulates spatial and temporal interactions in agricultural 25 landscapes and that can be used to explore trade-offs between production and environment so 26 27 helping to determine solutions to the problems of sustainable food production. Here we focus on models of agricultural production, water movement and nutrient flow in a landscape. We 28 validate these models against data from two long-term experiments, (the first a continuous 29 30 wheat experiment and the other a permanent grass-land experiment) and an experiment where water and nutrient flow are measured from isolated catchments. The model simulated wheat 31 32 yield (RMSE 20.3-28.6%), grain N (RMSE 21.3-42.5%) and P (RMSE 20.2-29% excluding the nil N plots), and total soil organic carbon particularly well (RMSE 3.1 - 13.8%), the 33 simulations of water flow were also reasonable (RMSE 180.36 and 226.02). We illustrate the 34 use of our model framework to explore trade-offs between production and nutrient losses. 35

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#### 43 **1. Introduction**

Increasingly, agricultural production is being compelled to look not just at its externalities such as the environmental pollution or depletion of natural resources but also at the provision of wider ecosystem services such as biodiversity. Schemes to monitor or assess land for all of these factors are prohibitively expensive and yet there is a need to analyse modern agricultural systems for the purposes of policy, planning or management. Not surprisingly therefore, computer simulation models have a role to play in filling the large gaps between what we need to know and what is available from measurements.

Simulation models of agricultural systems abound, some focussing on specific aspects 51 such as soil organic matter dynamics (Coleman et al., 1997), crop growth (Semenov and 52 Stratonovitch, 2015), water movement (Addiscott and Whitmore, 1991), emissions (Rolston et 53 54 al., 1984), competing organisms (Andrew and Storkey, 2017), and some integrating to agricultural management systems (Brisson et al., 2003; Keating et al., 2003). Others focus on 55 56 the natural systems, tracing biodiversity often quite specifically (Andam et al., 2008; Koh et al., 2010). Some models, particularly agricultural ones, focus on field (Bell et al., 2012; Parton 57 et al., 1994) or farm scales (Del Prado et al., 2011). Biodiversity models often focus on larger 58 scales and water management models are naturally focussed on river basins or catchments 59 (Whitehead et al., 2014). 60

Many models simulate fields or regions, some simulate particular fluxes, say water from land to rivers. It is rarer to find models that try to integrate several of the impacts of farming in the landscape, and those that do adopt a relatively empirical, data-driven approach (Jackson et al., 2013; Tilman et al., 2001) that makes it difficult to explore the interactions between components of that landscape that might be better managed with a more holistic overview. It is rarer still to find models that make explicit spatial and temporal linkage between

67 adjacent fields and integrate all aspects of the managed farm environment up to the catchment level. Such a model would be useful to understand the spatial interactions and impact of the 68 natural (weeds, pest and diseases) as well as management (irrigation, fertilizer and application 69 70 of pesticides) events on an agricultural landscape. Our aim is to develop a spatially explicit model that can simulate the essential processes of soil, water, crop growth and biodiversity for 71 agricultural landscapes in the UK. This model can then be used to understand the trade-off 72 73 between farm management practices on farm economy and the environment. The ability to quantify such trade-offs is critical to our management of the landscape and underpins many 74 75 sustainability frameworks including the three pillars of sustainability (environmental, economic and social), the UN Sustainable Development goals which includes several targets 76 77 that relate to agricultural landscapes (Gil et al., 2017), and water-energy-food nexus approaches 78 that aim to consider the use of all of these resources. While tradeoff models exist (e.g. see 79 Sharps et al., 2017) they usually operate at large scales, not accounting for the field or farm scale at which land management decisions are often made. These models are often focussed on 80 81 land-use options within in GIS-based systems, operate on annual time-scales and can be focussed on policy. Our approach, and ultimate aim, is to simulate interactions between the 82 multiple processes that take place in agricultural fields and the farmed landscape with a view 83 to uncovering strategies for development and improvement of agri-environmental systems, 84 85 beyond the current envelope (Fig 1). By working on a daily time-step we can simulate the 86 processes and inform the decisions that someone who manages land will have to take.

Here we report the first version of our model that integrates agricultural production, water movement and nutrient flow in a landscape. The model combines aspects of several published model [RothC (Coleman and Jenkinson, 2014), LINTUL (Wolf, 2012), SUCROS (van Laar et al., 1997), and Century (Parton et al., 1994)], but also includes novel factors that have been implemented to capture potential improvements in yield that result from 92 management actions. These include coupling the RothC model to include the dynamics of N 93 and P and responses to changes in bulk-density that result from changes in soil organic matter. 94 We evaluate the model against data on crop growth and nutrient uptake for cereals and for 95 grass, and the integration in space of water and nutrients leaving agricultural fields. We then 96 illustrate how our model can be used to explore trade-offs between production and environment 97 with a scenario based on a wheat crop grown in conditions typical of arable England.





Fig1: Representation of an environmental-economic production possibility frontier. The blue diamonds are independent outcomes of management that optimises both yield and environmental quality at the same time. A decision along this line is a matter for policy. The orange squares within the envelope are inefficient in the sense that either production or environmental quality could be improved without impacting the other. This is the region for extension. Beyond the envelope is a zone where outcomes are currently infeasible and this is

the area which research addresses. An origin placed over any point (for example the cross shown in the figure on the middle of the envelope), facilitates the definition of the envelope algorithmically: if another point can be found in the first quadrant (North East) then the first point in not on the envelope.

110

#### 111 **2. Methodology**

Our intention was to build a model system capable of exploring the multiple interactions 112 113 between components of a simple landscape and to take into account both within and between field movement of components such as water and nitrate. Nonetheless, because we wished to 114 build a system that can be used on a reasonably large landscape comprising many fields and 115 116 boundaries, we based our system on simple but adequate descriptions of the processes involved. Here we report on interactions and differences between single or adjacent but joined fields and 117 focus our discussion on productivity and loss of water and nitrogen to water courses and the 118 atmosphere. To do so we describe an integrated model of crop, water and soil processes that 119 runs on a daily time step. We validate this using data from the Broadbalk and Park Grass long-120 121 term experiments at Rothamsted Research, in Harpenden, SE England, and spatial interactions are tested on data from the more recently established North Wyke Farm Platform, at 122 Rothamsted Research, near Okehampton, SW England (Orr et al., 2016). 123

124

125 2.1 Spatial structure

We impose a grid on the landscape where, dependant on size, each field is represented by one or more grid cells. Soil properties are set in each cell and initial values are given for bulk density, pH and soil water. Within each cell we model crop growth, the dynamics of soil water, total soil organic carbon (TOC), changes in bulk density and nutrient flows on a daily time step. In cases where fields are made up of several cells, water and nutrients can move laterally between cells, as well as vertically though the soil profile. This model structure allows us to explore both temporal and spatial interactions. Cell edges can be designated as ditches (into which water and nutrients may flow), hedgerows or field margins.

134

#### 135 2.2 *Soil water*

The soil water model uses a capacity based approach (Addiscott and Whitmore, 1991; 136 Van Ittersum et al., 2003; van Laar et al., 1997). The soil is divided into three layers. This 137 choice is a compromise between capturing the heterogeneity of the soil profile (which would 138 require multiple layers in the simulation) and minimising complexity to enable fast run-times 139 140 which are important when coupling models with optimisation algorithms over large spatial scales. In our study each layer was initially set to 230mm. The capacity of each of the soil 141 layers is calculated with van Genuchten (1980) soil water release curves determined using the 142 HYPRES pedo-transfer functions (Wösten et al., 1999). These functions use texture, soil 143 organic matter and bulk density to derive the water release curves. For the topsoil, these release 144 145 curves are updated daily to take into account changes in bulk density, for example, when farmyard manure (FYM) is added (see section 2.6). 146

Infiltrating water fills the soil layers to field capacity (-10 kPa), and starting from the top layer, excess water drains to the layer below, with water draining from layer 3 becoming drainage. In addition to percolation, water is lost by runoff and evaporation from the soil surface, and transpiration by the growing crop. The water available for crop uptake at any time is equal to the quantity of water stored above wilting point (-1500 kPa) in the rooted soil profile. A detailed description of the soil water model can be found in van Laar et al. (1997), with our modifications described in section 2.7. The change in water content in each layer is derived

154 from the balance between inputs from precipitation, and outputs from drainage, runoff,155 evaporation and transpiration.

Working at the water catchment scale Bell et al. (2007) developed a simple algorithm for estimating the total surface water leaving a sloping (i.e. not uniform in the vertical dimension) region. The storage capacity (*S*) of high zones is reduced in relation to the topographic gradient according to

160 
$$S = \left(1 - \frac{g}{g_{\max}}\right) S_{\max}$$
 1

161 where  $S_{\text{max}}$  is the maximum storage capacity,  $\bar{g}$  is the average gradient in the cell and  $g_{\text{max}}$  is 162 the upper limit on the gradient. By adopting this strategy on a grid cell basis, we increase the 163 flow of water out of each cell compared to that if it were flat. Runoff moves from the highest 164 cell to the lowest by moving between cells with neighbouring boundaries. The proportion of 165 runoff allocated in each direction is determined by the relative magnitude of the downward 166 slopes. Dissolved substances such as nitrate, move in proportion to the water.

167

#### 168 2.3 Soil total organic carbon, nitrogen and phosphorus

The soil total organic carbon (TOC) model is based on the Rothamsted carbon model, RothC, (Coleman and Jenkinson, 2014). Soil total organic carbon is split into four active compartments and a small amount of inert organic matter (IOM). The four active compartments are Decomposable Plant Material (DPM), Resistant Plant Material (RPM), Microbial Biomass (BIO) and Humified Organic Matter (HUM). Each compartment decomposes by a first-order process with its own rate constant. The IOM compartment is resistant to decomposition. Decomposition of each of the four active pools is modified by rate modifying factors for temperature, moisture and plant retainment. Full details of the model can be found in (Colemanand Jenkinson, 2014).

The dynamics of the soil organic nitrogen (SON) and phosphorus (SOP) are modelled 178 in a similar way to the TOC dynamics, both SON and SOP have the same pool structure as the 179 active TOC pools. To determine initial values for each TOC pool, the model is run to 180 181 equilibrium so that the modelled TOC matches the initial measured TOC. The initial values of each of the SON and SOP pools are then determined using the TOC values, and the C:N and 182 C:P ratios of each pool. The C:  $N_{\text{Bio}}$  and C:  $N_{\text{Hum}}$  ratios are both fixed at 8.5 (Bradbury et al., 183 1993), whereas  $C: N_{\text{DPM}}$  and  $C: N_{\text{RPM}}$  ratios vary over time depending on the carbon inputs to 184 soil from the crop or the addition of organic amendments. The  $C: P_{Bio}$  and  $C: P_{Hum}$  ratios are 185 fixed at 50.0 and 100.0 respectively, like nitrogen the  $C: P_{\text{DPM}}$  and  $C: P_{\text{RPM}}$  ratios vary over 186 time depending on the carbon inputs to the soil from the crop or the addition of organic 187 188 amendments.

#### 189 The N in pool i that is mineralised or immobilized is given by

190

$$M_i = \frac{\Delta_i}{\rho_i} - \frac{B_i}{\rho_{\text{Bio}}} - \frac{U_i}{\rho_{\text{Hum}}}$$

where  $\Delta_i$  is the change in pool *i* from day *t* to *t* + 1,  $B_i$  is the amount of pool *i* transformed to biomass from day *t* to *t* + 1,  $U_i$  is the amount of pool *i* transformed to humus from day *t* to t + 1,  $\rho_i$  is the C:N ratio for pool *i*, and  $\rho_{Bio}$  and  $\rho_{Hum}$  are the C:N ratios for the biomass and humus pools respectively. The sum of  $M_i$  across the four pools gives the net mineralisation or immobilisation, if the sum of  $M_i$  is negative immobilisation occurs and mineral N is removed from the soil, if the sum of  $M_i$  is positive mineralisation occurs and mineral N is added as NH<sup>4</sup><sub>4</sub> to the soil. If there is not enough soil mineral N (NO<sup>-</sup><sub>3</sub> and NH<sup>+</sup><sub>4</sub>) on a particular day, then decomposition of TOC does not happen. If there is enough soil mineral N, then N is removed from the  $NH_4^+$  pool in preference to  $NO_3^-$  pool.

The P mineralisation or immobilization of each SOP pool is calculated in a similar way to the mineralisation N, where in Equation (2),  $\rho_i$  is the C:P ratio for pool *i*, and  $\rho_{Bio}$  and  $\rho_{Hum}$ are the C:P ratios for the biomass and humus pools respectively. When P is mineralised 80% is added to the available P pool, and the remaining 20% is added to the non-available P pool. For P 80% of mineralised P is added to or subtracted from the available P pool, similarly when immobilization of P occurs 80% is taken from the available P pool, and the remaining 20% is taken from the non-available P pool (see section 2.5).

207

#### 208 2.4 Soil Mineral Nitrogen

In the model, soil mineral N consists of N in ammonium (NH<sub>4</sub><sup>+</sup>) and nitrate (NO<sub>3</sub><sup>-</sup>). Inputs of N through atmospheric deposition ( $N_{AtDep}$ ) were set to 35 kg N yr<sup>-1</sup> (Anon, 1998) for the UK in 1966, decreasing linearly to 20 kg N yr<sup>-1</sup> in 2012 (pers. comm. Goulding). Like, Sundial (Anon, 1998) it was distributed evenly throughout the year as nitrate. Nitrogen applied as fertilizer enters the NH<sub>4</sub><sup>+</sup> or NO<sub>3</sub><sup>-</sup> pools depending on the type of fertilizer applied. When organic amendments are added, N enters the soil inorganic nitrogen pools by mineralisation (see section 2.3).

Rainfall runoff mixes in the model with the water and minerals in the top 20 mm of the soil profile. The amount of mineral nitrogen ( $NH_4^+$  and  $NO_3^-$ ) in runoff from the the top 20 mm of soil ( $N_{Run}$ ) is given by (Sharpley, 1985)

219 
$$N_{\rm Run} = \frac{N_{\rm Surf} W_{\rm Run}}{W_{\rm Run} + W_{\rm Surf}}$$
 3

where the surface water ( $W_{Surf}$ ) is given by difference in the volumetric water content at saturation and air dried, multiplied by 20 to give the water (mm) in the top 20 mm,  $W_{Run}$  is the water runoff (mm) and the surface N ( $N_{Surf}$ ) is given by

223 
$$N_{\text{Surf}} = \frac{20}{\delta(1)} (N_{NH4} + N_{NO3})$$
 4

where  $\delta(1)$  is the depth of the first layer.

Any nitrate in the soil can potentially move down the soil profile with the water. The concentration of NO<sub>3</sub><sup>-</sup> in layer *l*, ( $\gamma_{NO3}(l)$ ) is given by:

227 
$$\gamma_{\rm NO3}(l) = \frac{N_{\rm NO3}(l)}{W(l)}$$
 5

where  $N_{NO3}(l)$  is the NO<sub>3</sub><sup>-1</sup> (kg N ha<sup>-1</sup>) in layer  $l, l = 1 \dots 3$ , and W(l) is the water content of layer l.

The amount of  $NO_3^-$  (kg N d<sup>-1</sup>) that moves down each layer *l* is given by

231 
$$F_{NO3}(l) = \max(0, \min\{N_{NO3}(l), \gamma_{NO3}(l)F_W(l+1)\})$$
 6

where  $F_W(l)$  is the water that flows from layer *l* to layer l + 1. The nitrate that moves down from layer 3,  $F_{NO3}(3)$ , is N leached out of the profile.

Nitrification is an aerobic process whereby the NH<sub>4</sub><sup>+</sup> in the soil is oxidised to form NO<sub>3</sub><sup>-</sup> and N<sub>2</sub>O. Our models are based on Milne et al. (2005) and Parton et al. (2001). The rate of nitrification depends on the soil properties, such as water filled pore space  $\theta/\theta_{\text{Sat}}$ , soil temperature (T), soil moisture (M), and pH ( $S_{\text{pH}}$ ). In the model the amount of N<sub>2</sub>O (kg N ha<sup>-1</sup> day<sup>-1</sup>) produced from a given amount of NH<sub>4</sub><sup>+</sup> ( $N_{\text{NH4}}$  (l)) in layer l is given by

239 
$$N_{\rm N2O}(l) = k_{\rm N2O} N_{\rm NH4}(l) S_{\rm pH}(l) \left(1 - \frac{\theta}{\theta_{\rm Sat}(l)}\right), \qquad 7$$

where  $k_{N20}$  is a constant that takes the value 0.0001. The amount of nitrate (kg N ha<sup>-1</sup> day<sup>-1</sup>) produced from soil NH<sub>4</sub><sup>+</sup> is given by

242 
$$N_{\text{NO3}}(l) = \max \left[ (N_{\text{NH4}}(l) - N_{\text{N2O}}(l) - N_{\min})(1 - e^{-k}) f(T(l)) g(M(l)), 0 \right]$$

where  $N_{\min}$  is the minimum amount of NH<sub>4</sub><sup>+</sup> that must be in the soil for nitrification to occur (we assume  $N_{\min} = 0.05$ ), *k* is a constant for nitrification which is set at 0.15, and f(T) and g(M(i)) are functions that describe the effect respectively of temperature and moisture on nitrification, for details see Godwin and Allan Jones (1991).

Denitrification is an anaerobic process whereby the NO<sub>3</sub><sup>-</sup> in the soil is reduced to nitrous oxide and nitrogen. The amounts of these gases produced depends on the soil conditions, most notably the nitrate in the soil ( $N_{NO3}$ , kg ha<sup>-1</sup>), the water filled pore space ( $\theta/\theta_{Sat}$ ), soil temperature (T, °C), soil organic carbon (c) and pH ( $S_{pH}$ ) (Del Grosso et al., 2000; Milne et al., 2011; Nömmik, 1956). The effect of soil organic carbon on emissions is felt indirectly as a result of the temperature function g(T). We assumed the following simple model to describe N<sub>2</sub>O emissions (kg N ha<sup>-1</sup> day<sup>-1</sup>)

N<sub>2</sub>O = 
$$aN_{NO3}f(\theta/\theta_{Sat})g(T)$$
 9

where a is a constant. We took the functional forms of  $f(\theta/\theta_{\text{Sat}})$  and g(T) from the literature 255 and then fitted the model parameters to data from field experiments from around the UK where 256 nitrate, soil temperature, water filled pore space, and N<sub>2</sub>O emissions (kg N ha<sup>-1</sup> day<sup>-1</sup>) were 257 measured. Similar to other empirical or semi-empirical models, these parameter values can 258 only be assumed to hold for the range of conditions for which they were fitted, and outside of 259 this range further validation would be required. Nitrous oxide is linearly related to nitrate 260  $(N_{NO3})$  and we used the function defined by Lark and Milne (2016) to describe the effect of 261 water-filled-pore space on N<sub>2</sub>O emissions. That is 262

263 
$$f(w) = \exp\left[-0.6151\left(\log\left\{\frac{\theta/\theta_{Sat}(l)}{1 - \theta/\theta_{Sat}(l)}\right\} - 1.19\right)^2\right]$$
 10

where  $\theta/\theta_{\text{Sat}}(l)$  is the water filled pore space in each layer *l*. Data from Nömmik (1956) suggested that the relationship between temperature and N<sub>2</sub>O emissions should follow a normal distribution with mean 23.65 and standard deviation 5.53. However, data from the Defra project AC0116 (http://www.environmentdata.org/archive/ghgno:676) which we used to relate average temperature to emissions, did not conform to the standard deviation given by Nömmik (1956). Therefore, we assumed the same mean but fitted the standard deviation to our field data. Our fitted model was

271 
$$N_2 O = 0.000735 N_{NO3} \exp[-0.6151(\theta/\theta_{Sat}(l) - 1.19)^2] \exp[-0.00045(T - 23.65)^2]$$
 11

which we apply in only the top two layers of our model as there is not sufficient biologicalactivity for denitrification to occur in the bottom soil layer.

When water filled pore space increases, the soil becomes more anaerobic and so the amount of  $N_2$  produced increases. A similar relationship holds for temperature (Nömmik, 1956). We used the following model and fitted the parameters so that our model gave proportions of  $N_2O$  to  $N_2$  similar to those observed in Colbourn (1988)

278 
$$N_2 = \frac{0.0052 \text{ N}_{NO3}}{(1 + e^{-0.14975(T+4.0)})(1 + e^{-12.0(\theta/\theta_{\text{Sat}}(l) - 0.62)})}.$$
 12

The nitrogen taken up by the crop each day is taken from the nitrate pool with an upper limit of 6 kg N ha<sup>-1</sup> day<sup>-1</sup> (Semenov et al., 2007).

281

#### 282 2.5 Mineral Phosphorus

In the model, mineral phosphorus is split into two pools: available P (which includesphosphorus in soil solution and loosely adsorbed to the clay surface) and non-available P.

Eighty percent of the fertilizer P enters the available P pool and the remaining 20% enters thenon-available P pool (Wolf et al., 1987).

287 Similar to the N model, a proportion of the available P contained in the top 20 mm of288 soil can be lost through runoff.

289 
$$P_{\text{Run}} = \frac{P_{\text{Surf}} W_{\text{Run}}}{W_{\text{Run}} + W_{\text{Surf}}}$$
13

290 where the surface  $P(P_{Surf})$  is given by

 $P_{\text{Surf}} = \frac{20}{\delta(1)} P_{\text{M}}$  14

where  $P_{\rm M}$  is the mobile (dissolved and particulate) P which we assume to be 10% of  $P_{\rm AV}$ . We set solution P to 1% of the available P (pers. comm. Paul Poulton). This can potentially be leached when water flows down the profile.

The soil organic P that is mineralised is added to the available P pool. Mineral P may also be immobilised, in which case it is taken from the available P pool first and then from the non-available P pool.

Available P ( $P_{Av}$ ) is converted to non-available P ( $P_{NonAv}$ ) by reversible processes which reduce its extractability. In the model, the P content for each soil layer (available and non-available P), which we define  $P_{Tot}$ , is calculated in mg kg<sup>-1</sup> soil. The release to fixation variable, V(l), for layer l is given by

302 
$$V(l) = \begin{cases} \frac{\alpha_b P_{\text{Tot}}(l) + \beta_b}{P_{\text{Tot}}(l)}, & P_{\text{Tot}}(l) > \frac{\beta_a - \beta_b}{\alpha_b - \alpha_a} \\ \frac{\alpha_a P_{\text{Tot}}(l) + \beta_a}{P_{\text{Tot}}(l)}, & P_{\text{Tot}}(l) \le \frac{\beta_a - \beta_b}{\alpha_b - \alpha_a} \end{cases}$$
15

303

were  $\alpha_b$  and  $\beta_b$  are the slope and intercept, of value 0.113 and -49.3 respectively, for the linear relationship between  $P_{Av}$  and  $P_{Tot}$ . For small values of  $P_{Tot}$ , an alternative set of 306 coefficients  $\alpha_a$  and  $\beta_a$ , of value 0.0201 and -5.1, are used (see supplementary Fig. 1). The 307 ratio of release to fixation is given by

308 
$$R_{\rm RF}(l) = \frac{V(l)}{1.0 - V(l)}$$
 16

309 The transfer of P from the non-available to the available pool  $P_{NA \rightarrow Av}$ , and the reverse transfer 310  $P_{Av \rightarrow NA}$  in layer *l* on day t + 1 are given by

311 
$$P_{\mathrm{NA}\to\mathrm{Av}} = R_{\mathrm{RF}}(l) P_{\mathrm{Av}}(l;t) f_{\mathrm{pH}}\left(S_{\mathrm{pH}}(l)\right)$$
17

312 
$$P_{\mathrm{Av}\to\mathrm{NA}} = \lambda P_{\mathrm{NonAv}}(l;T)R_{\mathrm{RF}}(l)f_{\mathrm{pH}}\left(S_{\mathrm{pH}}(l)\right), \qquad 18$$

313 and so

314 
$$P_{\text{NonAv}}(l;t+1) = P_{\text{NonAv}}(l;t) + P_{\text{Av}\to\text{NA}} - P_{\text{NA}\to\text{Av}}$$
 19

315 
$$P_{Av}(l;t+1) = P_{Av}(l;t) - P_{Av \to NA} + P_{NA \to Av}$$
20

The constant  $\lambda$  determines the rate of re-equilibration between  $P_{Av}$  and  $P_{NonAv}$  following the addition of mineral P, and is set to 0.01 giving a half-life of approximately 65 days. The values of coefficients,  $\alpha$ ,  $\beta$  and  $\lambda$  were established for a silty clay loam soil at Rothamsted. The rate modifying function  $f_{pH}$  linearly increases from 0.0 to 1.0 as pH increases from 0 to 7, and then linearly decreases back to zero as pH increases from 7 to 14. The P required by the crop is taken from the available P pool, up to a limit of 2 kg P ha<sup>-1</sup> day<sup>-1</sup>.

322

323 2.6 Bulk density

To take into account changes in depth caused by changes in bulk density as a result of, for example, the addition of FYM, we used the Rawls (1983) nomogram to estimate bulk density in relation to sand, clay and organic carbon contents of soil. The depth of the topsoil is modified to reflect the change in bulk density (changes in depth and bulk density only occur in
the top soil). Because of the changes in depth and bulk density in the top soil, we modify water
properties, such as the water content at saturation, field capacity, and wilting point, daily (see
section 2.2). Modelling bulk density dynamically in this way has been described previously by
Whitmore et al. (2011).

332

#### 333 *2.7 Crop model*

Our crop model is a generic plant growth model, which uses a light use efficiency (LUE, g dry matter MJ<sup>-1</sup>) based approach to calculate the biomass production (Monteith, 1990; Monteith and Moss, 1977). The rate of biomass (*B*<sub>crop</sub>) produced each day is given by

$$\frac{dB_{\rm crop}}{dt} = Q \ \varepsilon \ W_{\rm rf} \ N_{\rm NI} \ P_{\rm NI}$$
 21

where Q is the intercepted PAR (MJ PAR  $m^{-2}$  surface area) which depends on the solar 338 radiation and canopy leaf area,  $\varepsilon$  is the crop specific LUE, which for grass, changes with 339 development stage see Schapendonk et al. (1998),  $W_{rf}$  is the transpiration reduction factor, 340  $N_{\rm NI}$  and  $P_{\rm NI}$  are the nitrogen and phosphorus nutrition indices, which range from zero to one. 341 For grass, LUE is reduced for higher radiation levels (Schapendonk et al., 1998). In our model 342 343 LUE is reduced by a factor  $R_{LUE}$  which decreases from 1.0 to 0.33 when radiation increases from 10 to 40 MJ m<sup>-2</sup> d<sup>-1</sup>. Schapendonk et al. (1998) also modified LUE, by the temperature 344 factor  $T_{LUE}$ , which in this study increases linearly from 0.0 to 1 between 6.0 and 9.0 °C. The 345 346 biomass formed is partitioned between roots, stem, leaves and storage organs based on the development stage (DVS) (Boons-Prins et al., 1993; Wolf, 2012). 347

348 The transpiration reduction factor ( $W_{rf}$ ) is defined as the ratio of actual transpiration 349 (mm day<sup>-1</sup>) to potential transpiration (mm day<sup>-1</sup>) and is calculated

350 
$$W_{\rm rf} = \frac{\sum_{l=1}^{3} A_{\rm Tran}(l)}{P_{\rm Tran}}$$
22

351 where  $P_{\text{Tran}}$  is the daily potential transpiration which is calculated as in Lintel (Wolf, 2012).

352 The amount of the actual transpiration coming out of layer (l) is given by

353 
$$A_{\text{Tran}}(l) = \frac{P_{\text{Tran}}(l)W_{S}(l)^{2}F_{\text{RL}}(l)}{W_{S}(1)F_{\text{RL}}(1) + W_{S}(2)F_{\text{RL}}(2) + W_{S}(3)F_{\text{RL}}(3)}.$$
 23

Here  $F_{\text{RL}}$  is the fraction of root in each layer and  $W_S$  is the impact of water content on the water stress function. This follows the approach of Li et al. (2001). This impact of water content is based on the method described in Feddes et al. (1976) given by

358  

$$W_{S} = \begin{cases} \frac{\theta_{s} - \theta}{\theta_{s} - \theta_{a}}, & \text{for } \theta > \theta_{a} \\ 1 & \text{for } \theta_{a} \ge \theta > \theta_{d} \\ \frac{\theta - \theta_{w}}{\theta_{d} - \theta_{w}}, & \text{for } \theta_{d} \ge \theta > \theta_{w} \end{cases}$$
24

where  $\theta$  is the volumetric water content,  $\theta_s$  is the water content at saturation,  $\theta_a$  is the water content at -5 kPa,  $\theta_d$  is the water content at -40 kPa, and  $\theta_w$  is the water content at wilting point (-1500 kPa). Water stress affects grass less than arable crops (per comms J. Storkey). In simulations, when the soil is saturated grass does not suffer water stress. When the volumetric water content falls below  $\theta_d = -40$  kPa the water stress factor  $W_s$  decreases linearly between  $\theta_d$  and  $\theta_w$  to 0.4.

364

365 The proportion of root  $(F_{RL})$  in each layer *l* is given by

366 
$$F_{RL}(l) = \frac{R_{\text{Len}}(l)}{R_{\text{Len}}(1) + R_{\text{Len}}(2) + R_{\text{Len}}(3)}$$
25

367 where  $R_{\text{Len}}$  is the root length per unit area (mm mm<sup>-2</sup>).

The root depth  $(d_{root})$  increases by 12.0 mm per day to a maximum root depth which depends on the crop being modelled. The root length per unit area within each layer, calculated according to an adaptation of the method of Gerwitz and Page (1974), is given by

371 
$$R_{\text{Len}}(l) = -\frac{R_0}{a \left(e^{-a Z_2(l)} - e^{-a Z_1(l)}\right)}$$
 26

where  $R_0$  is the root length density at the soil surface (mm mm<sup>-3</sup>) the value of which is nonessential to the model as it cancels out in Equation (25),  $z_1(l)$  and  $z_2(l)$  are the upper and lower horizon depth (mm) of layer l, and a is given by

$$a = -\frac{\ln(1 - F_{\rm r})}{d_{\rm root}}$$
<sup>27</sup>

where  $F_{\rm r}$  is the fraction (arbitrarily defined as 0.98) of the root length that is present above  $d_{\rm root}$ .

378 The uptake of plant nutrient (N and P) is determined by the crop demand and the supply of these nutrients by soil. The total nutrient demand of the crop is the sum of the nutrient 379 demand from its individual organs (i.e. roots, stems and leaves excluding storage organs, for 380 381 which nutrient demand is met by translocation from the other organs). Nutrient demand of the individual organs is calculated as the difference between maximum and actual organ nutrient 382 contents. The maximum nutrient content is defined as a function of canopy development stage. 383 The total nutrient uptake of the crop takes place before anthesis. Sub-optimal nutrient 384 availability in the soil leads to nutrient stress in the crop. A detailed description of crop nitrogen 385 386 dynamics is reported by Shibu et al. (2010) and P dynamics follows N in a similar way.

387 Nitrogen stress in the plant growth model is expressed as nitrogen nutrition index ( $N_{\rm NI}$ ) 388 and is calculated by:

389

390 
$$N_{\rm NI} = \max\left[0, \min\left(1, \frac{N_{\rm leaf} + N_{\rm stem} - N_{\rm Res}(\Omega_{\rm leaf} + \Omega_{\rm stem})}{\Omega_{\rm leaf}N_{\rm MaxPropleaf} + \Omega_{\rm stem}N_{\rm MaxPropstem} - N_{\rm Res}(\Omega_{\rm leaf} + \Omega_{\rm stem})}\right)\right] 28$$

where  $N_{\text{leaf}}$  and  $N_{\text{stem}}$  are the *N* in the leaf and stem respectively,  $\Omega_{\text{leaf}}$  and  $\Omega_{\text{stem}}$  are the weights of the leaf and stem respectively,  $N_{\text{MaxPropleaf}}$  and  $N_{\text{MaxPropstem}}$  are the maximum proportion of N in the leaf and stem respectively. The residual N ( $N_{\text{Res}}$ ) is the fraction of N which is part of the cell structure and was fixed at 0.004 for wheat (Wolf, 2012) and 0.01 for
grass (Bouman et al., 1996). For wheat, the maximum N in the leaf is given by:

396 
$$N_{\text{MaxPropleaf}} = 0.046 \exp(-1.7D) + 0.014$$
 29

where *D* is the development stage of the crop which is calculated using thermal time modified by a vernalisation factor and the photosensitivity of the crop (see Wolf (2012), and references therein). For grass we set  $N_{\text{MaxPropleaf}}$  to 0.0425. The maximum N in the stem is given by  $N_{\text{MaxPropStem}} = 0.5 N_{\text{MaxPropLeaf}}$ , (see Wolf (2012).

401 The phosphorus nutrition index  $(P_{\rm NI})$  is calculated by:

$$402 \quad P_{\rm NI} = \max\left[0, \min\left(1, \frac{P_{\rm leaf} + P_{\rm stem} - (\Omega_{\rm leaf} P_{\rm ResLeaf} + \ \Omega_{\rm stem} P_{\rm ResStem})}{\Omega_{\rm leaf} P_{\rm MaxPropleaf} + \Omega_{\rm stem} P_{\rm MaxPropstem} - N_{\rm Res}(\Omega_{\rm leaf} P_{\rm ResLeaf} + \ \Omega_{\rm stem} P_{\rm ResStem})}\right)\right]3$$

where  $P_{\text{leaf}}$  and  $P_{\text{Stem}}$  are the P in the leaf and stem respectively, and  $P_{\text{MaxPropleaf}}$  and 403  $P_{\text{MaxPropstem}}$  are the maximum proportion of P in the leaf and stem respectively. For wheat the 404 residual P in the leaf is  $P_{\text{ResLeaf}} = 0.0003$  and in the stem  $P_{\text{ResStem}} = 0.00018$ . For grass both 405  $P_{\text{ResLeaf}}$  and  $P_{\text{ResStem}}$  are set to 0.001 (Wolf et al., 1987). For wheat the maximum P in the 406 leaf reduces with development stage. From development stages 0 to 0.7 it reduces linearly from 407 0.0066 to 0.0036 and then from 0.0036 to 0.0009 from development stage 0.7 to 1, after which 408 it holds the value of 0.0009. For grass the maximum P in the leaf is fixed at 0.0035 (Bouman 409 et al., 1996). 410

Processes leading to the aboveground litter formation and carbon turnover below ground are similar for both crops and grass but their rates are different. We assume that 50% of the dead leaves become litter on a daily basis and the remainder is left on the stem. The rate at which the roots die is a function of growth stage. In case of crops, the root death happens towards the latter part of the growing season (DVS >1.5) at a rate of 0.02 per day. In case of 416 grass, once the root system has been established (3-6 months after sowing, DVS=0.01), root 417 death becomes continuous at a rate of 0.01 per day. The root exudates are considered to be a 418 part of root death, so are not modelled seperately . The leaf death rate is a function of heat 419 stress, nitrogen stress and shading as described in Schapendonk et al. (1998). All C, N, and P 420 from dead roots and litter is returned to the soil.

421 The grass model differs somewhat from the crop model as grass has indeterminate growth and is not allowed to flower (so always has a DVS always < 1.0) as it can be cut or 422 grazed in the model (unlike the crop which completes its life cycle in a given growing season). 423 424 Grass is a perennial crop that grows for one or more seasons before being reseeded. Cut grass and grazed grass is removed from the modelled system. The amount removed is such that the 425 remaining biomass cannot fall bellow 50 g m<sup>-2</sup>. Livestock deposit nutrients into the system as 426 427 manure. When animals are on the field, we set the deposition of C and N for each animal type based on data from (Cottrill and Smith, 2007), for each beef animal this was 4.03 kg C of 428 manure per day containing 0.22 kg N, for each dairy cow this was 6.45 kg C per day containing 429 0.35 kg N, and for each sheep this was 0.45 kg C per day as fresh deposit, containing 0.02 kg 430 N per day. These rates are multiplied by the stocking rate to give the rate of deposit per hectare. 431

432

#### 433 2.8 *Data requirements*

For each layer of the soil, the model requires initial values for soil depth, clay, silt, TOC, bulk density, available P, non available P, soil NH<sub>4</sub>, soil NO<sub>3</sub>, soil pH. Initial values for elevation and latitude are also needed. The model runs with a daily time-step and so for each simulated day weather data (minimum and maximum temperature, rainfall, radiation, vapour pressure and windspeed) are needed. For each season and where relevant to the crop, sowing dates, fertilizer application timing, type and dose and dates when the grass is cut are required.

#### 441 *2.9 Case studies*

To test our model, we used data from two long-term agricultural experiments and one 442 more recent grass-livestock experiment. These were: The Broadbalk wheat experiment, and 443 the Park Grass permanent grassland experiment at Rothamsted Research, Hertfordshire, UK 444 (51.8° N, 0.37° W), and the more recent North Wyke farm platform at Rothamsted Research, 445 near Okehampton, UK (50.77° N, 3.92° W), which has spatially integrated data from livestock-446 bearing grassland in a sloping terrain. We used a suite of statistical metrics (including the mean, 447 standard deviation, root mean square error, and sample correlation coefficient, r) to quantify 448 449 the performance of our model (see Smith et al., 1997).

450

#### 451 2.9.1 *Broadbalk*

The Broadbalk wheat experiment has been running since 1843, and wheat has been 452 sown and harvested on all or part of the experiment every year since then. The original aim of 453 the experiment was to test the effects of various combinations of inorganic fertilizers and 454 organic manures on the yield of winter wheat. The experiment was divided into different strips 455 456 given a range of fertilizer applications, which extended the whole length of the field. In 1926 the experiment was divided into five Sections, crossing the fertilizer treatments at right angles, 457 where each section was bare fallowed one year in five to control weeds. In 1968 the experiment 458 459 was further divided into 10 Sections, so that the yield of wheat grown continuously could be compared with that grown in rotation after a two-year break. The plots receive management 460 consistent with standard practice for the time. The soil is clay loam to silty clay loam, 461 462 predominately Batcombe series (Avery and Catt, 1995), FAO classification: Chromic Luvisol (or Alisol), U.S. Soil Taxonomy: Aquic (or Typic) Paleudalf. The site is thought to have been 463

464 in arable cropping for many centuries before the start of the experiment. Further details are465 available from http://www.era.rothamsted.ac.uk/Broadbalk

The plots from the continuous wheat sections (Section 1 and 9), selected for this study, 466 receive a range of fertilizer and FYM applications (see Table 1). Wheat has been grown every 467 year on these Sections, since 1966. Modern, short-strawed high yielding varieties were 468 introduced in the 1967–1968 season and it is from this date that we test the model. Most of the 469 data available from the electronic Rothamsted Archive 470 are (eRA http://www.era.rothamsted.ac.uk). Periodic measurements of TOC were made on all plots 471 472 (Watts et al., 2006; Pers. comm. P. Poulton for later data), measurements of volumetric water content on plot 8 in 2007 (Pers. Comm, C. Watts) and measurements and estimates of N 473 leaching were made between 1990 and 1998 (Goulding et al., 2000). Grain N was measured 474 475 1968-2012, and grain P from 1968-2011 (except 1976-1985), Section 1 only.

Table 1. The fertilizer and manure treatments applied annually to the Broadbalk experimentplots used in the simulations.

	Treatments										
Plot	up to 1967	1968 - 1984	1985 - 2000	2001 - 2004	2005 - 2012						
3	Nil	Nil	Nil	Nil	Nil						
5	P K Na Mg	P K Na Mg	P K Mg	K Mg	K Mg						
6	48N P K Na Mg	48N P K Na Mg	48N P K Mg	48N K Mg	48N K Mg						
7	96N P K Na Mg	96N P K Na Mg	96N P K Mg	96N K Mg	96N K Mg						
8	144N P K Na Mg	144N P K Na Mg	144N P K Mg	144N K Mg	144N K Mg						
9	48N* P K Na Mg	192N P K Na Mg	192N P K Mg	192N K Mg	192N K Mg						
15	96N P K Na Mg	192N P K Na Mg	240N P K Mg	240N K Mg	240N K Mg						
16	96N* P K Na Mg	96N P K Na Mg	288N P K Mg	288N K Mg	288N K Mg						

2.1	FYM since 1885	FYM 96N	FYM 96N	FYM 96N	FYM 144N
2.2	FYM	FYM	FYM	FYM	FYM

The values of N are in kg N ha<sup>-1</sup>, applied as ammonium sulphate 1843-1967, as calcium
ammonium nitrate between 1968-1985, and as ammonium nitrate thereafter. Treatments with
\* were applied as sodium nitrate. Farmyard manure (FYM) was applied at 35 t ha<sup>-1</sup> fresh
weight, and contains approximately 250kg N ha<sup>-1</sup>. Other elements were applied at 35 kg P
ha<sup>-1</sup>, 90 kg K ha<sup>-1</sup>, 16 kg Na ha<sup>-1</sup> until 1973 and 12 kg Mg ha<sup>-1</sup> respectively. P has not been
applied since 2001, due to high levels of plant available P in the soil. For more details see
http://www.era.rothamsted.ac.uk/Broadbalk

We ran the model to simulate the plots listed in Table 1 using weather data from the Rothamsted meteorological station from 1966 to 2012. Comparisons were made between measured and simulated values of crop yield, content of N and P in the grain, TOC, volumetric water content and nitrate leaching.

490

491 2.9.2 *Park Grass* 

The Park Grass experiment is the oldest experiment on permanent grassland in the 492 493 world. Started by Lawes and Gilbert in 1856, its original purpose was to investigate ways of improving the yield of hay by the application of inorganic fertilizers and organic manure. 494 Within 3 years it became clear that these treatments were having a dramatic effect on the 495 496 species composition of what had been a uniform sward. The continuing effects of the original treatments on species diversity and on soil function, together with later tests of liming and 497 interactions with atmospheric inputs and climate change (Storkey et al., 2015), has meant that 498 Park Grass has become increasingly important to ecologists, environmentalists and soil 499

scientists. The soil is silty clay loam, predominately Hook series, with areas more typical of
the Batcombe series (Avery and Catt, 1995), FAO Classification: Chromic Luvisol (or Alisol),
U.S. Soil Taxonomy: Aquic (or Typic) Paleudalf. The site is known to have been in permanent
pasture for at least 100 years before the start of the experiment. For further details see
http://www.era.rothamsted.ac.uk/Park

The plots are cut in mid-June, and made into hay. A second cut is usually taken in the autumn, except in a few years, when there was insufficient herbage to sample. Since 1960, yields have been estimated from strips cut with a forage harvester. The remainder of the plot is still mown and made into hay, continuing earlier management. For the second cut, the whole of each plot is cut with a forage harvester. The experiment is never cultivated, and the site was in permanent grassland for at least 100 years before the experiment began. Further details are available from http://www.era.rothamsted.ac.uk/Park

512 Here we simulated two plots, Plot 3a and 14/2a, with contrasting fertilizer treatments. Plot 3a has received no inorganic fertilizer or manure since 1856. Plot 14/2a has received 96 513 kg N ha<sup>-1</sup> in the spring, and 35 kg P in the autumn each year since 1858, plus K, Na and Mg. 514 In 1965 the plots were divided into four subplots, given different amounts of chalk to maintain 515 516 soil at pHs of 7, 6 and 5 (sub-plots a, b and c, respectively). The fourth sub plot (d) receives no chalk. We have selected sub-plot 'a' for this simulation, with a pH of 7. We use yield data from 517 518 1966-2012, with two cuts each year except in 2003, when no second cut was taken, with weather data from the Rothamsted meteorological station. 519

We chose Plot 14/2a over the other N fertilizer plots because N is applied as sodium nitrate, whereas in most other plots N is applied as ammonium sulphate, which has an acidifying effect on the soil and so a dramatic effect on species composition and the decomposition of soil organic matter (see <u>http://www.era.rothamsted.ac.uk/Park</u>).

#### 525 2.9.3 The North Wyke Farm Platform

The North Wyke Farm Platform, near Okehampton, SW England was established as a UK 526 National Capability for collaborative research, training and knowledge exchange in agro-527 environmental sciences related to beef and sheep production in lowland grasslands (Orr et al., 528 2016). The soils on the farm platform are predominately Halstow, (Pelo-stagnogley soils, 529 Avery, 1980), FAO Classification: Stagni-vertic cambisol, U.S. Soil Taxonomy: Typic 530 haplaquept. For more details see Harrod and Hogan (2008). A system based on permanent 531 pasture was implemented on three 21-ha farmlets to obtain baseline data on hydrology, nutrient 532 cycling and productivity for 2 years. Since then, two of the farmlets have been modified by 533 534 either (i) planned reseeding with grasses that have been bred for enhanced sugar content or 535 deep-rooting traits or (ii) sowing grass and legume mixtures to reduce nitrogen fertilizer inputs. The third farmlet continued under permanent pasture. The quantities of nutrients that enter, 536 cycle within and leave the farmlets are recorded using sensor technologies alongside more 537 traditional field study methods. Here we simulated the water and nutrient flows from October 538 2012 to 25<sup>th</sup> December 2013 from catchment 4 (Golden Rove) and catchment 5 (Orchard 539 Dean), two of the un-modified permanent grassland catchments, that had contrasting 540 topologies. The North Wyke data that we used for this study are available from 541 http://www.rothamsted.ac.uk/farmplatform). 542

543

544 2.10 *Trade offs* 

545 We coupled the simulation model with an optimisation algorithm to determine Pareto 546 optimal fronts between multiple objectives defined in terms of outputs from the model. The 547 optimised Pareto fronts describe the trade-offs between objective variables such as yield and nitrate leaching. To illustrate how these can be identified, we used the fertiliser application time and amount as two management variables that the optimisation algorithm could vary in order to affect three objectives: the yield of a wheat crop, nitrate leaching and N<sub>2</sub>O emissions. Simulations used the soil properties and weather data from plot 9 of the Broadbalk experiment for the years 1968–1978. For this period the mean measured yield was 5.4 t ha<sup>-1</sup> at 85% dry matter.

Initially the algorithm, which combines non-dominated sorting (Deb et al., 2002) with 554 differential evolution (Storn and Price, 1997), randomly selects a number of possible 555 556 management variables, implements these management options in the simulation model and records the effect on each of the multiple objectives. Non-dominated sorting then identifies the 557 management options that result in the 'best' objectives, i.e. those that are non-dominated. A 558 559 point is said to be dominated by another if it is worse for every single objective. For example, if we aim to maximise two objectives, point A (Fig. 2) is dominated by point B because the 560 value of both objectives is greater at B than A. Points B and C, however, are both non-561 dominated with respect to one another because whilst objective 1 is higher for B, objective 2 562 is higher for C. The non-dominated sorting algorithm performs a series of pairwise 563 564 comparisons in order to identify all of the management options that lead to non-dominated sets of objectives. The differential evolution algorithm then combines aspects of the management 565 566 options that led to non-dominated objectives to identify new management options that could 567 potentially perform even better. The process is iterated in directions that the differential evolution algorithm suggests will be an improvement, , until the results converge and produce 568 569 a similar Pareto front with each iteration.



Fig.2 Example of how a Pareto front is identified from a number of points simulated by the model with the aim to improve multiple objectives (1 & 2) simultaneously. Point B is selected over point A because B scores better for both objectives. It can be seen that neither of points B or C dominates the other, because point B does better at objective 1 whilst point C improves on objective 2. Consequently, both are retained. The Pareto front (line) can be identified by connecting together all of the non-dominated points.

577

#### 578 **3. Results**

#### 579 *3.1 Broadbalk*

The simulated and measured grain yields for the plots listed in Table 1 are shown inFig. 3. The model captures the differences between the plots well and this is quantified by the

overall correlation between modelled and measured (Pearson correlation, r = 0.86). The plot means for the modelled and measured yields are similar, as are the variances, although the variance for the modelled yield in plots with little fertilizer N applied are smaller than the observed (Table 2). The model reflects the year-to-year fluctuations in yield, although notably under-predicts the 1995 yield from the plots with larger N applications (9, 15, 16, 2.1 and 2.2).



Fig. 3 Measured (black) and modelled (red) grain yields for ten plots from the Broadbalk
long-term wheat experiment, 1968-2012, continuous wheat (sections 1 and 9). The measured
values were averaged over Sections 1 and 9 (see 2.9.1).

593 The model replicates the plot-to-plot and year-to-year variation in grain N, grain P and TOC (see Figs 4, 3 and 6, and Tables 3, 4 and 5), although we note that year-to-year variation 594 in TOC is minimal. The correlations across all plots between modelled and measured grain N, 595 grain P, and TOC are 0.88, 0.84 and 0.99 respectively. The model reproduces the pattern in the 596 variation of volumetric water content for plot 8, following one of the observed realisations 597 closely (Fig. 7). Note that measurements with such probes are sometimes biased towards drier 598 measurements because instrument range is short and if contact is lost between the access tube 599 and soil then the soil can appear drier than it actually is. 600



Fig. 4: Measured (black) and modelled (red) grain N content for ten plots from the Broadbalk
long-term wheat experiment, 1968-2012, continuous wheat. The measured values were from
Section 1 only (see 2.9.1).



Fig. 5: Measured (black) and modelled (red) grain P content for ten plots from the Broadbalk
long-term wheat experiment (1968–1975 and 1986–2011), continuous wheat. The measured
values were from Section 1 only (see 2.9.1)



Fig. 6: Measured (black circles) and modelled (red) soil total organic carbon (TOC) for ten
plots from the Broadbalk long-term wheat experiment. The measured values were averaged
over Sections 1 and 9.

![](_page_36_Figure_0.jpeg)

Fig. 7: Measured for three replicates (black) and modelled (red) volumetric water content insoil from plot 8 of the Broadbalk experiment.

- 622 Table 2 Summary statistics for measured and simulated grain yields (at 85% dry matter),
- 623 19682012 for the Broadbalk wheat experiment. The measured values for yield in each year

	Measured			Simulated		
Plot	Mean	Standard	Mean	Standard	RMSE	Correlation
no.	t ha <sup>-1</sup>	deviation/	t ha <sup>-1</sup>	deviation/	(%)	
		t ha <sup>-1</sup>		t ha <sup>-1</sup>		
3	1.16	0.5	1.26	0.22	42.56	0.28
5	1.37	0.42	1.27	0.22	33.41	0.16
6	3.4	0.67	2.86	0.51	28.61	0.07
7	4.99	1.02	4.62	0.7	24.03	0.16
8	5.71	1.18	6	1.01	24.27	0.25
9	6.23	1.06	6.66	1.32	23.01	0.36
15	6.32	1.28	6.58	1.37	23.29	0.4
16	6.34	1.41	6.12	1.64	20.88	0.64
2.1	7.16	1.35	6.75	1.41	20.28	0.49
2.2	5.67	1.12	5.49	1.13	22.84	0.35

624 were averaged over Sections 1 and 9 (see 2.9.1).

625

Table 3 Summary statistics for measured and simulated grain N content, 1968–2012, for the
Broadbalk wheat experiment. The measured values for grain N were from Section 1 only (see
2.9.1).

	Measured		S	imulated		
Plot	Mean	Standard	Mean	Standard	RMSE	Correlation
no.	kg N ha <sup>-1</sup>	deviation/	kg N ha⁻	deviation/	(%)	
		kg N ha <sup>-1</sup>	1	kg N ha <sup>-1</sup>		
3	16.33	7.23	19.76	3.1	42.54	0.57
5	18.48	6.36	20.07	3.15	31.58	0.47
6	46	9.12	47.69	5.79	23.42	0.03
7	76.69	16.21	80.44	9.03	23.58	0.11
8	99.03	21.31	110.25	15.84	25.44	0.29
9	117.91	20.91	126.27	23.92	23.32	0.32
15	122.99	28.03	124.05	25.96	25.17	0.35
16	121.34	37.12	113.62	32.82	22.98	0.71
2.1	128.08	23.56	129.28	27.52	21.27	0.44
2.2	86.83	18.58	98.09	17.27	27.04	0.34

- Table 4 Summary statistics for measured and simulated P in the grain, 1968–1975 and 1986–
- 634 2011 for the Broadbalk wheat experiment. The measured values for grain P were from
- 635 Section 1 only (see 2.9.1).
- 636

	-	Measured	5	Simulated		
Plot	Mean	Standard	Mean	Standard	RMSE	Correlation
no.	kg P	deviation/	kg P ha⁻	deviation/	(%)	
	ha <sup>-1</sup>	kg P ha <sup>-1</sup>	1	kg P ha <sup>-1</sup>		
3	3.04	1.49	5.36	0.43	89.47	0.29
5	4.05	1.47	5.4	0.43	48.33	0.26
6	9.89	2.23	10.35	0.86	22.5	0.26
7	14.11	2.89	14.58	1.66	20.73	0.29
8	15.38	3.85	17.31	2.38	29.01	0.23
9	16.49	3.59	18.64	3.06	27.8	0.26
15	17.47	4.73	18.77	3.27	28.34	0.33
16	17.33	4.3	18.42	3.57	25.47	0.42
2.1	21.49	4.09	19.36	3.26	20.42	0.47
2.2	17.14	3.57	15.8	2.5	20.24	0.49

- 639 Table 5 Summary statistics for measured and simulated total soil organic carbon (TOC),
- 640 when measured between 1967–2012 Broadbalk wheat experiment. The measured values for
- TOC were averaged over Sections 1 and 9 (see 2.9.1).
- 642

	Measured		S	Simulated		
Plot	Mean	Standard	Mean	Standard	RMSE	Correlation
no.	t C ha⁻	deviation/	t C ha <sup>-1</sup>	deviation/	(%)	
	1	t C ha <sup>-1</sup>		t C ha <sup>-1</sup>		
3	22.95	1.37	21.01	1.6	11.51	0.28
5	24.84	1.05	21.78	1.76	13.83	0.47
6	27.96	0.88	26.82	0.84	6.07	-0.08
7	29.57	0.83	30.39	0.3	3.88	0.24
8	29.73	1.12	31.69	0.53	7.88	-0.1
9	29.97	1.47	30.36	1.18	3.11	0.82
15	29.45	1.84	29.74	1.33	4.42	0.72
16	30.75	1.96	30.55	1.06	4.3	0.78
2.1	68.9	4.76	68.97	2.61	5.46	0.62
2.2	75.18	3.27	72.36	1.06	5.61	0.28

645	The measured (Goulding et al., 2000) and modelled N leached for each plots are shown
646	in Fig. 8. The model predictions match the N leached from the mineral fertilized plots
647	reasonably well, although the model consistently overestimates N leached from plots receiving
648	the most N (plots 15 and 16 and the FYM plots 2.1 and 2.2) and in the driest years (1991/2,
649	1996/7 and 1997/8). The variances for measured leaching are larger than the modelled for all
650	but plot 2.1 (Table 6). Note that measurements were not determined for every plot in every
651	year.

![](_page_42_Figure_0.jpeg)

Fig. 8: Estimated and modelled N leached from study plots on the Broadbalk wheat experiment
1990–1998. Measurements are from Section 9 only. The black open circle indicates that no
measurement was taken.

657	Table 6 Summary	statistics for	measured and	simulated	nitrate l	eached	(kg N	ha <sup>-1</sup> y	<sup>1</sup> ) between
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658 1990 and 1998, Broadbalk wheat experiment. Measurements are from Section 9 only.

	Measured		S	Simulated		
Plot	Mean	Standard	Mean	Standard	RMSE	Correlation
no.	kg N	deviation/	kg N ha⁻	deviation/	(%)	
	ha <sup>-1</sup> y <sup>-1</sup>	kg N ha <sup>-1</sup> y <sup>-1</sup>	<sup>1</sup> y <sup>-1</sup>	kg N ha <sup>-1</sup> y <sup>-1</sup>		
3	13.00	8.51	11.94	5.99	33.29	0.85
5	11.71	8.92	12.86	6.21	58.56	0.60
6	11.88	9.76	18.24	7.38	111.01	-0.02
7	15.00	10.38	20.68	7.01	81.13	0.17
8	22.00	16.55	22.73	6.47	65.70	0.36
9	30.00	22.44	32.74	12.83	57.82	0.58
15	42.38	33.31	53.57	17.51	79.79	0.36
16	47.57	47.02	77.51	22.69	107.32	0.23
2.1	76.86	36.19	130.33	50.46	85.84	0.38
2.2	59.00	50.10	105.98	37.58	103.82	0.45

#### 660 3.2 Park Grass

- 661 The model captures the differences between the plots and between the first and second cuts
- well (Figure 9 and Table 7). The first cut, usually taken in June, is normally higher than the 662
- second cut which is usually taken in November). 663

![](_page_44_Figure_4.jpeg)

![](_page_44_Figure_5.jpeg)

Figure 9 Simulated (red) and measured (black) yields for plots 3a and 14/2a Park Grass 668 669 permanent grassland experiment, showing both cuts each year.

Table 7 Summary statistics for measured and simulated yield 1966-2012, Park Grass

671 experiment, (47 years, n = 93).

	]	Measured	Simulated			
Plot	Mean	Standard	Mean	Standard	RMSE	Correlation
no.	t ha <sup>-1</sup>	deviation/	t ha <sup>-1</sup>	deviation/	(%)	
		t ha <sup>-1</sup>		t ha <sup>-1</sup>		
За	1.61	0.78	1.79	0.43	49.91	0.28
14/2a	3.32	1.67	2.91	1.24	34.35	0.77

672

673

### 674 3.2 North Wyke Farm Platform

The simulation of water flow rates  $(m^3 day^{-1})$  for catchments 4 and 5 reflect those measured

676 (Fig. 10 and Table 8). This is quantified by the correlations between modelled and measured

677 (Pearson correlation, r = 0.57 and r = 0.55 respectively). The modelled water flow rate and

678 variation are slightly smaller than the measured in each case.

![](_page_46_Figure_0.jpeg)

Fig. 10 Simulated (red) and measured (black) flow rates  $(m^3 day^{-1})$  for catchment 4 and 5 of

<sup>684</sup> the North Wyke Farm Platform

Table 8 Summary statistics for measured and simulated flow and nitrate (kg N per catchment per day) in the drains, North Wyke Farm Platform

688 Catchments 4 and 5.

	Flow (m <sup>3</sup> day <sup>-1</sup> )					Nitrate (kg N catchment <sup>-1</sup> day <sup>-1</sup> )						
	Meas	sured	Simu	lated			Meas	ured	Simu	lated		
Catchment	Mean	Std	Mean	Std	RMSE	Correlation	Mean	Std	Mean	Std	RMSE	Correlation
		dev		dev	(%)			dev		dev	(%)	
4	213.60	457.01	147.83	315.90	180.36	0.57	0.13	0.27	0.47	1.64	1287.69	0.21
5	114.10	281.82	122.88	258.19	226.02	0.55	0.13	0.29	0.01	0.08	248.45	-0.01

689

The simulation of nitrate in the drainage water over estimates nitrate for catchment 4 and under estimates it for catchment 5, but the peaks of nitrate after May 2013 broadly correspond to that which was measured (see Fig. 11 and Table 8).

![](_page_48_Figure_1.jpeg)

Fig. 11 Simulated (red line) and measured (black line) log nitrate (kg N /catchment) for
catchment 4 and 5 of the North Wyke Farm Platform. The black discs show when nitrogen
fertilizer was applied. For details see <a href="http://www.rothamsted.ac.uk/farmplatform">http://www.rothamsted.ac.uk/farmplatform</a>.

700 3.3 *Trade offs* 

By allowing an optimisation algorithm to vary the timing and amount of a single fertilizer 701 application, we identified the trade-offs between yield, nitrate leaching and N<sub>2</sub>O emissions for 702 703 an illustrative example (Fig. 12). The results show that as the yield increases (due to changes in fertilizer application) the lowest possible N<sub>2</sub>O emissions that could be achieved 704 simultaneously increases non-linearly. The range of fertilizer N applied to achieve these Pareto 705 optimal objectives was  $0 - 210 \text{ kg N} \text{ ha}^{-1} \text{ y}^{-1}$ . The N<sub>2</sub>O emissions reduce as a result of applying 706 less fertilizer later in the growing season. As yield approaches its maximum, both the N<sub>2</sub>O 707 708 emissions and the nitrate leaching increase substantially with increasing amounts of fertilizer for an increasingly marginal improvement in yield. Nitrate leaching and N<sub>2</sub>O emissions are 709 synergistic throughout most of the range, however a trade-off appears as the emissions reach 710 711 their minimum value, as this also results in an increase in leaching. This illustrates how an optimisation approach (e.g. minimising  $N_2O$ ) could have unitended consequences for another 712 process (nitrate leaching), if both objectives are not considered simultaneously. The 713 optimisation algorithm does not identify a single fertilisation strategy, but highlights 714 715 nonlinearities thus identifying where a small reduction in one objective could have a large 716 benefit to another. Here, for example, the simulation indicates that the fertilizer application 717 conditions which correspond to a moderate yield, reduce the nitrate that is available to leach 718 from the soil substantially compared to those required for the most yield.

![](_page_50_Figure_0.jpeg)

720

Fig. 12 Illustrative example of use of the model to identify trade-offs between multiple objectives such as maximising yield, minimising nitrate leaching and minimising  $N_2O$ emissions. As maximising or minimising any one of these objectives affects the others, the optimisation identifies points on a multi-dimensional frontier with Pareto optimality. On this frontier no objective can be improved upon without a detrimental effect on at least one of the other objectives. This frontier therefore represents the best trade-offs that can be achieved.

727

#### 729 **4. Discussion**

We have built and used a model framework to simulate spatial and temporal interactions 730 731 in agricultural landscapes. The framework allows us to explore trade-offs between production and environmental outcomes to determine strategies that could contribute to sustainable food 732 production. It is important that the models reflect the important mechanisms that relate to 733 734 production and the environment. It is also essential that the models are parsimonious and run quickly so that a large range of scenarios can be tested, perhaps in conjunction with an 735 optimisation algorithm. Our simulations are within 25% of all the observations across multiple 736 years and plots and this is good evidence that the model is robust and that we can use it with 737 confidence to explore trade-offs relevant to farm and environmental management. 738

739 Simulation of wheat yields from the Broadbalk experiment and grass yields from the 740 Park Grass experiment reproduced both the differences between plots caused by the various fertilizer rates ( $\rho > 0.78$ ) and the observed year-to-year variation (RMSE ranging between 741 742 20.3 and 28.6% for the mineral N and FYM plots on Broadbalk and 34.3% for Park Grass, correlations were up to 0.77). According to the RMSEs, the model performed less well for the 743 plots that received no fertilizer (plots 3 and 5 on Broadbalk and plot 3a on Park Grass) where 744 the RMSEs were 42.6, 33.4% and 49.9% respectively. The larger values for the RMSE on the 745 lower-yield plots to some extent result from the form of this statistic which is scaled by the 746 747 reciprocal of the mean observation (i.e. the sum of the squared difference for the lower-yielding plots are scaled by larger values than the higher plots). Over the 46 years that we 748 simulated Broadbalk, the model tended to under predict yield between 1994 and 1996 for plots 749 750 with higher rates of N fertilizer applied (plots 8, 9, 15, 16, 21, 22) (Fig. 3). This is likely to be a result of excessive water stress when there was no N limitation. It was drier than normal in 751 the three months before harvest in 1994, 1995 and 1996, this led to higher water stress during 752 those months, and so a reduction in dry matter production. 753

The predictions of the variation in grain N for the Broadbalk plots were also good, with 754 the RMSE ranging from 21.3 to 42.5% (Fig. 4, Table 3), and again illustrated the differences 755 between plots receiving different rates of fertilizer N. For P uptake by the crop, the model 756 757 performed well for most plots with RMSE between 20.2 - 29.0% for all plots except 3 and 5 which had RMSE of 89.5% and 48.3% respectively (Fig. 5 and Table 4). In the experiment 758 applications of P stopped in 2001 due to large amounts of plant-available P in the soil, and the 759 P measured in the grain declines noticeably in plots with larger applications of fertilizer but 760 this is not exhibited in the model. However, this does not affect the measured grain yields (Fig 761 762 3). The variations in simulated yield, grain N and P are approximately 50% smaller than the observed for plots 3 and 5 (for other plots the variation is proportionally more similar). This 763 764 suggests that the nitrogen stress function maybe over-damping the simulated response to 765 variation in the weather.

The modelled total soil organic carbon for the Broadbalk plots fits the measured datawell with the RMSE ranging from 3.1 to 13.8% (Fig. 6 and Table 5).

The model simulations of N leached from the Broadbalk plots were compared with 768 estimates of leaching from (Goulding et al., 2000), based on nitrate concentrations in drainage 769 770 and soil water and calculations of drain flow. The measured concentrations of nitrate in soil water were subject to the usual large spatial variation with typical CVs of 50-90%. The 771 simulations reflected the differences in leaching between the different amounts of N, although 772 they tended to overestimate N leached at the largest N rates and in the driest years (Fig. 8 and 773 774 Table 6). IPCC guidelines (IPCC, 1997; Del Grosso et al., 2005) assume that 30% of applied N is leached or runs off into groundwater or surface waters and this accords with our simulations 775 of Broadbalk where approximately 31.7% of N applied is lost through leaching. 776

777 The simulation of water flow from the two North Wyke Catchments matches the pattern in the variation of water flow but the average water flow over the simulated periods was larger 778 than that simulated, as was the variation. This suggests that our model system is buffering the 779 780 water through-put in the catchment and that too much is being taken by the crop or evaporating from the system. The simulations of nitrate in drainage water on the North Wyke plots appeared 781 to be poorer than the simulations of N losses for Broadbalk. Although the timing of peaks in 782 783 nitrate towards the end of the simulation were determined well, little nitrate was simulated in the first part of the simulated time period. This was because there was very little nitrate left in 784 785 the model soil profiles at the beginning of the simulated run, and during the summer period (May 2013 – September 2013) there was very little simulated discharge (see Fig. 10). An 786 addition of nitrate on 5<sup>th</sup> March 2013 to catchment 4 increased the nitrate levels in the soil and 787 788 a peak in nitrate followed. Further additions of nitrate fertilizer kept the soil nitrate in this 789 simulation at a larger concentration than that in the catchment 5 simulation, which despite having similar levels of nitrate applied, retained less nitrate in the soil. The difference in the 790 791 simulated soil nitrate between the two catchments manifests as differences in the nitrate in the drainage water in the autumn and winter of 2013 where the nitrate leached was greater for 792 793 catchment 4 than for catchment 5. The simulated nitrate in the drainage water is larger than that measured for catchment 4 yet smaller for catchment 5. This suggests problems with the 794 795 modelled uptake of nitrate by the grass and retention in the soil in this case, but we have no 796 explanation for the counter-intuitive discrepancy between the measurements on the two plots. Quantifying the fate of nitrate is notoriously difficult (Senapati et al., 2016). Recently 797 calculated field level budgets of N from the North Wyke platform show unaccounted for losses 798 of between 30 and 60 kg N ha<sup>-1</sup> (Misselbrook pers. comm.). This highlights the need for more 799 research on the processes that control N transformations from micro-scale to field scale, and 800 larger-scales. Facilities such as the North Wyke farm platform are ideally placed to support 801

this kinds of research. Models such as the one described here can help to identify the parts of the processes where understanding is incomplete and so can help to inform the design of experiments as well as benefit from any new understanding obtained.

805 Others have explored trade-offs using empirical data. For example Phalan et al. (2011) compared the effects of land sparing and land sharing on crop yields and the densities of tree 806 807 and bird species across the UK, while Lamb et al. (2016) explored the need to cut greenhouse gas emissions, while increasing agricultural yields to meet the rapidly rising food demand 808 through land sparing. Eory et al. (2013) examined the trade-offs and synergies between 809 greenhouse gas mitigation measures and other environmental pollutants. The limitation of such 810 empirical studies is that there is a lack of data and so it is often not possible to consider more 811 than two factors at a time. Whilst models should always be used with caution, they do allow us 812 813 to consider multiple interactions under a large range of management strategies. Used appropriately, models such as the one we present here should allow sound conclusions to be 814 drawn on the relative impact of management strategies and might highlight unintended 815 consequences of certain actions. Whilst the complexity of agricultural systems across the 816 landscape could warrant a complex model, a simpler model that runs more quickly but still 817 818 captures the key processes can be coupled more easily to an optimisation algorithm. This then provides the opportunity to identify the form of the synergies and trade-offs between multiple 819 820 objectives at a broad and often neglected scale. Here, for example, we observe that objectives 821 that are largely synergistic such as nitrate leaching and  $N_2O$  emissions still exhibit a trade-off as the N<sub>2</sub>O emissions approach the minimum. The non-linearity in the leaching and emissions 822 as yield increases is also clear, indicating a strong trade-off. 823

In order to generate frontiers such as the ones we did here (Fig. 12) an optimisation algorithm must be chosen and a set of management options that the optimisation algorithm can manipulate identified. Within an agricultural landscape, management options are numerous. 827 For example, even considering only fertiliser applications, the timing, amount and type of multiple applications could all be included in the set of management options to be optimised. 828 This set of options will constrain the frontier, thus care must be taken to identify a reasonable 829 830 range of options, whilst keeping the number of variables that the algorithm can manipulate to a minimum. Even so, the set of options is likely to represent a complex optimisation problem, 831 involving multiple control variables, with the risk that the algorithm may be trapped in local 832 833 minima. The optimisation algorithm must be chosen and implemented to minimise this risk. In this case we chose to use non-dominated sorting combined with differential evolution. Whilst 834 835 the non-dominated sorting allowed us to consider multiple-objectives, which is critical to our aim of generating trade-off curves, the differential evolution combines a genetic algorithm and 836 a gradient based search to allow a complex control space to be explored efficiently. 837

838 Our framework includes models of crop growth, the dynamics of soil conditions and water and nutrient flows in order to quantify the trade-offs between agricultural production and 839 environmental factors. It could be expanded to include volatilisation and biological N fixation 840 (which should improve the simulation for certain grass and crop types). Our framework is 841 distinct from alternative models of the agricultural landscape because it simulates multiple 842 843 functions simultaneously and distinct from other models of ecosystem services (e.g. Sharps et al., 2017) because it focuses on scaling up the effect of field and farm scale management 844 845 practices to landscape scale. Additional environmental factors are also relevant to the 846 agricultural landscape and to include these the model could be expanded to include weeds, pests and diseases and aspects of biodiversity. For each new component there will be feedbacks 847 into existing models that alter the dynamics of yield accumulation and soil nutrient status. For 848 849 example, weed population dynamics will depend on the crop and the soil conditions, but in turn weeds will have a competitive effect on the crop, primarily for light, that will affect both yield 850 and to some extent soil nutrient status (Kropff and van Laar, 1993). Our model framework is 851

spatially explicit and simulates interactions between cells, in particular it describes the lateral
flows of nutrients and water from cell to cell based on relative elevation and slope of model
cell. The movement of insect pests, for example, is somewhat different as choice of destination
are influenced by host plant distribution and the dispersal characteristics of the species in
question. It will be straightforward to include these dispersal mechanisms within the landscape
framework, see Milne et al. (2015).

858

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