1	Prebiotic Chemistry and Atmospheric Warming of Early Earth By An Active
2	Young Sun
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11	Nitrogen is a critical ingredient of complex biological molecules [1]. Molecular nitrogen,
12	however, which was outgassed into the Earth's early atmosphere [2], is relatively chemically
13	inert and nitrogen fixation into more chemically reactive compounds requires high temperatures.
14	Possible mechanisms of nitrogen fixation include lightning, atmospheric shock heating by
15	meteorites, and solar ultraviolet radiation [3,4]. Here we show that nitrogen fixation in the early
16	terrestrial atmosphere can be explained by frequent and powerful coronal mass ejection events
17	from the young Sun - so-called superflares. Using magnetohydrodynamic (MHD) simulations
18	constrained by Kepler space telescope observations, we find that successive superflare ejections
19	produce shocks that accelerate energetic particles, which would have compressed the early
20	Earth's magnetosphere. The resulting extended polar cap openings provide pathways for
21	energetic particles to penetrate into the atmosphere and, according to our atmospheric chemistry
22	simulations, initiate reactions converting molecular nitrogen, carbon dioxide and methane to the
23	potent greenhouse gas nitrous oxide as well as hydrogen cyanide, an essential compound for life.

Furthermore, the destruction of  $N_2$ ,  $CO_2$  and  $CH_4$  suggests that these greenhouse gases cannot explain the stability of liquid water on the early Earth. Instead, we propose that the efficient formation of nitrous oxide could explain a warm early Earth.

27 MAIN TEXT

Here we develop a new concept of the rise of prebiotic chemistry on early Earth that suggests 28 abiotic nitrogen fixation mediated by the energy flux from paleo solar eruptive events. The flare 29 statistics of *Kepler* data suggests that the frequency of occurrence of superflares with energies > 30  $5 \times 10^{34}$  erg observed on G-type dwarfs follows a power-law distribution with spectral index 31 between ( $\alpha$ =- 2.0), which is comparable to those observed on dMe stars and the Sun [5,6]. If the 32 occurrence rate of superflares on young solar-like  $\sim 0.1$  events/day [6], then, the frequency of 33 super Carrington-type flare events with  $E \sim 10^{33}$  ergs on the early Sun ( $\leq 0.5$  Gyr) is expected to 34 be  $\sim 250$  events per day! Current data suggest that powerful solar flares (over X5 type) are 35 usually associated with fast ( $\geq 1000 \text{ km s}^{-1}$ ) wide ( $\theta > 100^{\circ}$ ) coronal mass ejections (CMEs) and 36 high-fluence solar energetic particle (SEP) events with kinetic energies up to 10<sup>33</sup> ergs [7-9]. 37 Tree ring data have recently provided evidence in favor of past superflares from the Sun [10,11]. 38 Their energy is a factor of 2-3 greater than that suggested for the famous Carrington-type CME 39 event [12]. Recent direct measurements of surface longitudinal magnetic fields on young solar-40 type stars imply that our young Sun had generated at least 10 times greater magnetic flux than 41 that observed in the current Sun [13]. The stronger magnetic flux produces frequent and 42 energetic flares, fast and wide coned CMEs and associated energetic SEP events with energies 43 up to  $10^{36}$  ergs. Our calculations suggest that the probability of CME striking the Earth is about 44 5% [14]. In the "perfect" magnetospheric storm, when the incoming cloud magnetic field,  $B_z$ 45 component is sheared with respect to the Earth's magnetic field, the frequency of CME impacts 46

is > 1 event per day! To model a CME event and its effects on the Earth's magnetosphere, we 47 used the Space Weather Modeling Framework (SWMF) available through the Community 48 Coordinated Modeling Center (CCMC) (see Supplementary Material). We assumed a steady 49 state paleo solar wind at 0.7 Gyr with the mass loss rate of  $1.7 \times 10^{-12} M_{sun}/yr$  and the wind speed 50 of 700 km/s as obtained from the 3D MHD young Sun's wind model [15] and a Carrington-type 51 CME cloud propagating at the radial speed of 1800 km/s with the total energy of 2 x  $10^{33}$  erg 52 [12]. Figure 1 presents a 2D map of the steady-state plasma density superimposed by magnetic 53 field lines for the magnetospheric configuration in the Y=0 plane corresponding to the initial 30 54 minutes of the simulations, when the Earth's magnetosphere was driven only by dynamic 55 pressure from the paleo-solar wind. The left panel of Figure 1 shows the steady state paleo solar 56 wind compresses the Earth's magnetosphere to  $\sim 9 R_E$ . The right panel of Figure 1 shows the 57 state of the magnetosphere two hours later when the CME cloud hits the Earth's magnetosphere 58 (also see the movies in Supplementary Material). At this time, the solar wind dynamic pressure 59 and the magnetic reconnection between the southward directed CME's cloud magnetic field and 60 northward Earth's dipole field pushing the dayside magnetosphere earthward reducing the 61 magnetopause stand-off distance from 9 to ~1.5 Earth's radii. The CME drives large field aligned 62 current distributions and produces significant disturbance of the magnetospheric field shifting the 63 boundary of the open-closed field shifts to 36° in latitude and producing a polar cap opening to 64 70% of the Earth's dipole magnetic field. In the current version, we used the dipole magnetic 65 field of the current Earth, however paleomagnetic studies of the Earth's ancient rocks suggest 66 that the field was weaker [16]. This suggests that the fraction of the open field used in our model 67 represent only a lower bound. Energetic particles accelerated in shocks driven by successive 68 flare/CME events (see for example [17]) can then efficiently penetrate the early terrestrial 69

atmosphere through the expended polar cap region. 70

We applied the Aeroplanet model [18] to simulate the atmospheric chemistry of the 71 nitrogen-dominated (80% N<sub>2</sub>, 20% CO<sub>2</sub> and 0.03% CH<sub>4</sub>) primitive Earth's atmosphere [19]. The 72 upper boundary of the atmosphere at 100 km is exposed to the steady state XUV flux with the 73 spectrum reconstructed for the early Sun at 0.7 Gyr [20] and to energetic protons with the energy 74 flux of 5 x  $10^{11}$  protons/cm<sup>2</sup>/MeV at 0.1 MeV with the spectral index of the energy spectrum of -75 2.15 representative of the Jan 20, 2005 SEP event and the energy range within 1 GeV [21]. The 76 model calculates photoabsorption of the EUV-XUV flux from the early Sun (see Figure 2) and 77 particle (electron and proton) fluxes to compute the corresponding fluxes at the atmospheric 78 altitudes between 200 km and the surface. These fluxes are used to calculate the photo and 79 particle impact ionization/dissociation rates of the atmospheric species producing secondary 80 electrons due to ionization processes. Then, using the photon flux and the photoionization-81 excitation-dissociation cross-sections, the model calculates the production of ionized and excited 82 state species and as a result, photoelectrons. In our steady-state model of the early Earth's 83 atmosphere, energetic precipitating protons from an SEP event impacted the middle and low 84 atmosphere and produce ionizations, dissociations, dissociative ionizations, and excitations of 85 atmospheric species and as a result, secondary electrons. The model includes 117 neutral 86 chemical reactions. The destruction of  $N_2$  into reactive nitrogen,  $N(^2D)$  and  $N(^4S)$  and the 87 subsequent destruction of CO2 and CH4 produces NO, CO, CH and NH in the polar regions of 88 the atmosphere as shown in Figure 3 (see Supplementary Material). 89

Our model predicts the formation of abundant NO and NH molecules and efficient 90 formation of N<sub>2</sub>O through NO + NH  $\rightarrow$  N<sub>2</sub>O + H with the major sink through the reaction N<sub>2</sub>O + 91  $H \rightarrow OH + N_2$  (see the pathway diagram in Figure 3). Photolysis of N<sub>2</sub>O via the reaction 92

pathway  $N_2O + hv \rightarrow N_2 + O(^1D)$  is not an efficient loss channel for  $N_2O$ , because of absorption 93 of solar flux shorter than 2300 Å by CH<sub>4</sub>. Atmospheric N<sub>2</sub>O steady-state density reaches a 94 concentration with the mixing ratio of 2 and 20 ppbv at 30 km in the 1 PAL (present atmospheric 95 level) atmosphere with 100% (solid line) and 10% (dashed line) of the maximum photochemical 96 destruction rate, as shown in Figure 4a. The derived value at 100% of the photodestruction rate 97 should be considered as a lower bound, because our model does not account for a number of 98 factors including the eddy diffusion and convection effects, the effects of Rayleigh scattering of 99 solar EUV radiation in the atmosphere and formation of hazes that significantly reduces the 100 photo-destruction rate of nitrous oxide, and therefore increases the production of N<sub>2</sub>O. Thus, the 101 model with 10% of the maximum photo-destruction rate probably better represents the density 102 profiles when all factors are accounted for. The steady-state density of N<sub>2</sub>O reaches 20 to 3000 103 ppbv in the 2 PAL model with 100% (solid line) and 10% (dashed line) of the maximum 104 photochemical destruction rate, as shown in Figure 4b. The choice of 2 PAL in Figure 4b is 105 consistent with compelling evidence that the atmospheric pressure of early Earth was enhanced 106 by a factor of 2-3 [22]. Another factor affecting the equilibrium mixing ratios of Figure 4 is the 107 representative energy of SEP events, which could be greater than that assumed in the model. 108 Laboratory experiments report the production of nitrogen oxides and N<sub>2</sub>O when N<sub>2</sub>-CO<sub>2</sub> mixture 109 that simulates the early Earth atmosphere was exposed by lightning and coronae discharges [23]. 110 Enhanced production of nitrous oxide in the lighting experiments are caused by energetic 111 electrons accelerated in the discharge and UV emission. Other evidence for the role of energetic 112 particles in N<sub>2</sub>O production comes from direct observations of its enhancement by 3% associated 113 114 with thunderstorm events [24].

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The efficient production of N<sub>2</sub>O in our model offers a solution to warming the early

Earth. The 0.7 Gyr old Sun was 25-30% fainter than the present-day Sun [25], which would be 116 insufficient to support liquid water on the early Earth contrary to geological evidence of that time 117 [26]. Current models of atmospheric warming offer solutions of this problem, commonly known 118 as Faint Young Sun (FYS) paradox due to a large atmospheric concentration of CO<sub>2</sub>, H<sub>2</sub>O, CH<sub>4</sub> 119 or N<sub>2</sub> and H<sub>2</sub> [27]. However, as our model implies, these molecules will be efficiently dissociated 120 due to photo-collisional processes driven by SEPs from the young Sun, which is consistent with 121 the recent mineralogical data [28]. Instead, the production of CH, NH and NO sets stage for the 122 formation of N<sub>2</sub>O, HCN and other N-containing species in the lower parts of the atmosphere. 123 HCN concentration reaches up to tens ppmv in the lower atmosphere. The calculated production 124 rate of HCN in the low atmosphere is driven by the following major reactions: NO + CH  $\rightarrow$ 125  $HCN + O, CH_2 + N(^4S) \rightarrow HCN + H, CH_3 + N(^4S) \rightarrow HCN + H + H, CH + CN \rightarrow HCN + H,$ 126  $N_2O + CH \rightarrow HCN + NO$ . Organic molecules may subsequently rain out into surface reservoirs 127 and engage in higher order chemistry producing more complex organics. For example, further 128 HCN polymerization is known to produce various amino acids, the building blocks of proteins 129 [29]. Production of other types of soluble N-containing species (NH<sub>3</sub>, HNO, NO) by particles 130 may have provided a massive dose of nitrogen "fertilizer" to early surface biology on terrestrial 131 planets. 132

Thus, our concept implies that early Sun's activity provided a window of opportunity for prebiotic life on Earth. The proposed model also redefines the conditions of habitability not just in terms of a "liquid water zone", but as a biogenic zone (BZ), within which the stellar energy fluxes are high enough to ignite reactive chemistry that produces complex molecules crucial for life. As a by-product, this chemistry forms greenhouse gasses that may efficiently keep the atmosphere warm for liquid water to exist. The model predictions can tested by observing broad



Figure 1. The initial (panel a) and the final state (panel b) magnetic field lines in white) and the plasma pressure (in color) of the Earth's magnetosphere due to the CME event

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*Figure 3. The pathway diagram of abiotic production of odd nitrogen and nitrogen-bearing* 

163 compounds including nitrous oxide and hydrogen cyanide due to photo and collisional

*dissociation and ionizations caused by XUV solar flux and SEP particle flux.* 





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Figure 4. Radial profiles of the steady-state mixing ratios of various species produced by
incoming flux of primary protons and secondary electrons for 10% (dotted lines) and 100%
(solid lines) of the maximum photo-destruction rate at 1 PAL (top figure, a) and at 2 PAL
(bottom figure, b).

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### 270 **AUTHOR CONTRIBUTIONS**:

V. Airapetian conceived and designed the numerical models, analyzed the data, contributed
 materials and wrote the manuscript. A. Glocer and G. Gronoff contributed to the

- 273 development and execution of codes and data analysis. E. Hébrard contributed to the
- chemistry model and data analysis, W. Danchi contributed to data analysis and proofreading
- of the paper.

### 276 **Competing financial interests**.

277 The authors declare no competing financial interests.

## 278 FIGURE CAPTIONS.

**Figure 1**. The initial (left panel, a) and the final state (right panel, b) magnetic field lines in

white) and the plasma pressure (in color) of the Earth's magnetosphere due to the CME event

281 [14].

Figure 2. XUV flux of the young Sun at 0.7 Gyr [20].

**Figure 3**. The pathway diagram of abiotic production of odd nitrogen and nitrogen-bearing

compounds including nitrous oxide and hydrogen cyanide due to photo and collisional

dissociation and ionizations caused by XUV solar flux and SEP particle flux.

**Figure 4**. Radial profiles of the steady-state mixing ratios of various species produced by

incoming flux of primary protons and secondary electrons for 10% (dotted lines) and 100%

(solid lines) of the maximum photo-destruction rate at 1 PAL (top figure, 4a) and at 2 PAL

289 (bottom figure, 4b).

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- 299 METHODS
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## 301 **1. SWMF DESCRIPTION**

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In this paper, we utilized the Space Weather Modeling Framework (SWMF) available at 304 Community Coordinated Modeling Center (CCMC) at NASA Goddard Space Flight Center (see 305 http://ccmc.gsfc.nasa.gov). A single-fluid, time dependent fully non-linear 3D 306 at magnetohydrodynamic (MHD) code BATS-R-US (Block-Adaptive-Tree Solar-wind Roe-type 307 Upwind Scheme) is a part of SWMF and was developed at the University of Michigan Center of 308 Space Environment Modeling (CSEM). The spine of the SWMF is the BATS-R-US code 309 [30,31], which is coupled to Rice Convection Model (RCM, [31] to model a propagation and 310 interaction of a model SCME with a magnetosphere and ionosphere of a young Earth. The MHD 311 part of the code calculates the dynamic response of the large-scale magnetospheric plasma to 312 varying solar wind conditions in a self consistent manner by using the block-adaptive wind Roe-313 type upwind scheme global MHD code [30]. The dynamics of the magnetosphere is described in 314 a Cartesian geometry by using resistive MHD equations. The electromagnetic coupling of the 315 magnetosphere to a conducting ionosphere is handled in a standard way [32]. Specifically, the 316 magnetospheric currents near the inner boundary of the MHD simulation are mapped to the 317 ionosphere where. A potential solver is then used which combines these currents with a 318 conductance map of the ionosphere (including solar and auroral contributions) to produce the 319 320 electric potential in the ionosphere. That potential is then used to set the electric field and corresponding drift at the magnetospheric simulations inner boundary. 321

The MHD approximation does not provide an adequate description of the inner magnetosphere because energy dependent particle drifts and ring current evolution become

important. Here we use the Rice Convection Model, embedded in the MHD simulation, to model 324 this important region [31]. The RCM code is a kinetic plasma model that couples plasma motions 325 in the inner magnetosphere and calculates the energy dependent particle drifts and ring current 326 evolution in the inner magnetosphere. The ring current carries the most of the energy density 327 during magnetic storms and is essential to modeling strong storms. This coupling is crucial for 328 description of solar wind effects on a magnetosphere, because the ionosphere provides closure of 329 magnetospheric currents, which is needed for realistic description of magnetospheric convection 330 and associated electric fields. Thus, we apply a dedicated inner magnetospheric model that is 331 fully coupled to the MHD code for the treatment of the inner magnetosphere. We simulate the 332 magnetospheric cavity (outer and inner magnetosphere) in a computational box defined by the 333 following dimensions  $-224R_E < x < 224R_E$ ,  $-128R_E < y < 128R_E$ ,  $-128R_E < z < 128R_E$ , where 334 RE is the radius of the Earth placed at the center of the computational box. The dipole tilt is 335 neglected in this problem. The simulations were carried out using a block adaptive high-336 resolution grid with the minimum cell size of  $1/16 R_{\rm E}$ . 337

The inner boundary is set at 1.25  $R_E$  with a density of 100 cm<sup>-3</sup>. The velocity at 338 the inner body is set to the  $\vec{E} \times \vec{B}$  velocity, where  $\vec{E}$  is determined from the ionospheric 339 potential and B is the Earth's magnetic field. The pressure is set to float. The magnetic field is set 340 in a way that the radial component is the Earth's dipole and the tangential components are 341 allowed to float. The simulation is initialized with a dipole everywhere in the computational 342 domain and a small density, zero velocity, and a finite pressure. The solar wind conditions are set 343 at the upstream boundary and some period of local time stepping is used to get an initial steady 344 state solution. We assume the solar wind input parameters including the three components of 345

interplanetary magnetic field,  $B_x$ ,  $B_y$  and  $B_z$ , the plasma density and the wind velocity,  $V_x$ , using the physical conditions associated with a Carrington- type event as discussed by [33] and [14], see Figure 1. The time evolution of the plasma pressure (in nPa) and current density (in microAmps/m<sup>2</sup>) during the extreme CME event are presented in the attached Movie 1.

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# 2. AEROPLANETS MODEL DESCRIPTION

We used our sophisticated Aeroplanets model with enhanced chemistry [18] to model the upper 351 atmospheric region (up to 200 km) in response to young Sun's XUV (X-ray and EUV) emission 352 from and precipitating electrons and protons due to an SEP event. The model calculates the 353 photo and collisional (due to protons) dissociation, ionization and photoexcitation processes in 354 the Earth's atmosphere. The primary photoelectrons are then transported along a magnetic field 355 line, and the electron impact is computed solving the stationary kinetic Boltzmann equation. This 356 results in the dissociation, ionization and excitation of the different atmospheric species. The 357 Aeroplanets code incorporates 117 chemical reactions with the rates presented in Table 1. 358

To converge to steady state chemical solution for the early Earth atmosphere described in the Main section was reached after running the code for 6 months of physical time.

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### Table 1. List of Chemical Reactions Used in Our Model

Reaction	Reaction rate (in cgs units)
H + CH -> C + H2	0.124E-09 * (T / 300) ** (0.260E+00)
H + CH2 -> CH + H2	0.220E-09 * (T / 300) ** (0.320E+00)

H + e3CH2 -> CH + H2	0.220E-09 * (T / 300) ** (0.320E+00)
H + CH3 -> e3CH2 + H2	0.100E-09 * exp(-0.760E+04 / T)
H + CH4 -> CH3 + H2	0.589E-12 * (T / 300) ** (0.300E+01) * exp(-0.404E+04 / T)
CH + H2 -> e3CH2 + H	0.310E-09 * exp(-0.165E+04 / T)
CH + CH -> C2H + H	2.00E-10
$CH + CH \rightarrow e3CH2 + C$	2.00E-11
CH + e3CH2 -> C2H2 + H	2.00E-10
CH + CH3 -> C2H3 + H	1.00E-11
CH + CH3 -> C2H2 + H + H	1.00E-10
	0.105E-09 * (T / 300) ** (-0.104E+01) * exp(-0.361E+02 /
CH + CH4 -> C2H4 + H	T)
CH2 + H2 -> CH3 + H	0.880E-10 * (T / 300) ** (0.350E+00)
CH2 + CH4 -> e3CH2 + CH4	0.310E-11 * exp(0.250E+03 / T)
CH2 + CH4 -> CH3 + CH3	0.279E-10 * exp(0.250E+03 / T)
CH2 + C2H2 -> e3CH2 + C2H2	2.30E-10
CH2 + C2H2 -> C3H3 + H	0.760E-10 * (T / 300) ** (-0.300E+00)

CH2 + N2 -> e3CH2 + N2	0.110E-10 * (T / 300) ** (0.810E+00)
e3CH2 + H2 -> CH3 + H	0.800E-11 * exp(-0.450E+04 / T)
e3CH2 + CH4 -> CH3 + CH3	0.713E-11 * exp(-0.505E+04 / T)
CH3 + H2 -> CH4 + H	0.245E-13 * (T / 300) ** (0.288E+01) * exp(-0.460E+04 / T)
N(4S) + CH -> CN + H	0.140E-09 * (T / 300) ** (0.410E+00)
N(4S) + e3CH2 -> HCN + H	0.500E-10 * (T / 300) ** (0.170E+00)
N(4S) + e3CH2 -> HNC + H	0.300E-10 * (T / 300) ** (0.170E+00)
N(4S) + CH3 -> H2CN + H	5.60E-11
N(4S) + CH3 -> HCN + H + H	6.00E-12
N(4S) + NH -> N2 + H	0.250E-10 * (T / 300) ** (0.170E+00)
N(4S) + CN -> C + N2	0.900E-10 * (T / 300) ** (0.420E+00)
N(4S) + H2CN -> N2 + e3CH2	4.00E-11
N(4S) + H2CN -> HCN + NH	5.00E-12
N(2D) -> N(4S)	2.30E-05
N(2D) + H2 -> NH + H	0.420E-10 * exp(-0.880E+03 / T)
N(2D) + CH4 -> NH + CH3	0.130E-10 * exp(-0.755E+03 / T)

N(2D) + CH4 -> CH2NH + H	0.350E-10 * exp(-0.755E+03 / T)
N(2D) + N2 -> N(4S) + N2	0.100E-12 * exp(-0.520E+03 / T)
N(2D) + NH3 -> N2H2 + H	5.00E-11
N(2D) + HCN -> CH + N2	5.00E-11
N(2D) + HNC -> CN2 + H	2.00E-11
N(2D) + HNC -> CH + N2	2.00E-11
NH + H -> N(4S) + H2	0.220E-11 * (T / 300) ** (0.155E+01) * exp(-0.103E+03 / T)
NH + CH -> HCN + H	5.00E-11
NH + CH -> HNC + H	5.00E-11
NH + e3CH2 -> H2CN + H	3.00E-11
NH + e3CH2 -> HCN + H + H	3.00E-11
NH + e3CH2 -> HNC + H2	5.00E-12
NH + CH3 -> CH2NH + H	0.130E-09 * (T / 300) ** (0.170E+00)
NH + NH -> N2 + H + H	2.00E-10
NH + NH2 -> N2H2 + H	0.100E-09 * (T / 300) ** (0.170E+00)
NH + CN -> CN2 + H	1.00E-10

NH + CN -> N2 + CH	1.00E-10
NH2 + H2 -> NH3 + H	0.209E-11 * exp(-0.428E+04 / T)
NH2 + H -> NH + H2	0.200E-10 * exp(-0.240E+04 / T)
NH2 + CH4 -> NH3 + CH3	0.399E-13 * (T / 300) ** (0.359E+01) * exp(-0.454E+04 / T)
NH2 + C2H2 -> NH3 + C2H	0.111E-12 * exp(-0.185E+04 / T)
NH2 + C2H3 -> NH3 + C2H2	2.00E-11
NH2 + C2H3 -> SOOTN + H	8.00E-11
	0.540E-10 * (T / 300) ** (-0.110E+01) * exp(-0.600E+02 /
NH2 + H2CN -> HCN + NH3	T)
NH3 + H -> NH2 + H2	0.423E-13 * (T / 300) ** (0.393E+01) * exp(-0.406E+04 / T)
	0.169E-09 * (T / 300) ** (-0.560E+00) * exp(-0.280E+02 /
NH3 + CH -> CH2NH + H	T)
NH3 + CH3 -> NH2 + CH4	0.510E-13 * (T / 300) ** (0.286E+01) * exp(-0.734E+04 / T)
CN + H2 -> HCN + H	0.412E-12 * (T / 300) ** (0.287E+01) * exp(-0.820E+03 / T)
$CN + CH \rightarrow HCN + C$	0.100E-09 * (T / 300) ** (-0.170E+00)
CN + e3CH2 -> HCN + CH	5.00E-11

CN + e3CH2 -> CHCN + H	5.00E-11
CN + e3CH2 -> C2N + H2	5.00E-11
CN + CH3 -> CH2CN + H	1.00E-10
CN + CH4 -> HCN + CH3	0.620E-11 * exp(-0.721E+03 / T)
CN + NH3 -> HCN + NH2	0.277E-10 * (T / 300) ** (-0.114E+01)
CN + HCN -> C2N2 + H	0.430E-12 * (T / 300) ** (0.171E+01) * exp(-0.770E+03 / T)
CN + HNC -> C2N2 + H	2.00E-10
HCN + CH -> CHCN + H	0.140E-09 * (T / 300) ** (-0.170E+00) * exp(-0.0 / T)
HCN + CH -> C2N + H2	0.140E-09 * (T / 300) ** (-0.170E+00) * exp(-0.0 / T)
HCN + 1C2 -> C3N + H	0.200E-09 * (T / 300) ** (0.170E+00) * exp(-0.0 / T)
HNC + H -> HCN + H	0.300E-10 * exp(-0.800E+03 / T)
H2CN + H -> HCN + H2	6.00E-11
H2CN + CH3 -> CH4 + HCN	3.00E-11
O(3P) + CH2 -> HCO + H	1.00E-11
O(3P) + CH2 -> CO + H + H	5.00E-11
O(3P) + CH2 -> CO + H2	6.00E-11

O(3P) + e3CH2 -> HCO + H	1.00E-11
O(3P) + e3CH2 -> CO + H + H	5.00E-11
O(3P) + e3CH2 -> CO + H2	6.00E-11
O(3P) + CH3 -> CO + H2 + H	2.90E-11
O(3P) + CH3 -> H2CO + H	1.10E-10
O(3P) + NH -> NO + H	6.60E-11
O(3P) + HNO -> OH + NO	3.80E-11
O(1D) + H2 -> OH + H	1.10E-10
O(1D) + CH4 -> OH + CH3	1.05E-10
O(1D) + CH4 -> CH3O + H	3.50E-11
O(1D) + CH4 -> H2CO + H2	7.50E-12
O(1D) + N2 -> O(3P) + N2	2.15E-11
OH + H2 -> H2O + H	0.280E-11 * exp(-0.180E+04 / T)
OH + CH4 -> H2O + CH3	0.185E-11 * exp(-0.169E+04 / T)
OH + N(4S) -> NO + H	4.50E-11
OH + N(2D) -> NO + H	4.50E-11

OH + CO -> CO2 + H	1.30E-13
	0.280E-10 * (T / 300) ** (-0.122E+01) * exp(-0.120E+02 /
H2O + CH -> H2CO + H	T)
H2O + N(2D) -> OH + NH	4.50E-11
H2O + N(2D) -> HNO + H	5.00E-12
CO2 + N(2D) -> CO + NO	0.100E-10 * exp(-0.100E+04 / T)
NO + CH -> HCN + O(3P)	0.100E-09 * (T / 300) ** (-0.130E+00)
NO + CH -> NCO + H	0.300E-10 * (T / 300) ** (-0.130E+00)
NO + CH -> CO + NH	0.300E-10 * (T / 300) ** (-0.130E+00)
NO + CH -> OH + CN	0.100E-10 * (T / 300) ** (-0.130E+00)
NO + e3CH2 -> HNCO + H	0.210E-11 * exp(0.554E+03 / T)
NO + e3CH2 -> CO + NH2	0.300E-12 * exp(0.554E+03 / T)
$NO + N(4S) \rightarrow O(3P) + N2$	0.400E-10 * (T / 300)
NO + N(2D) -> O(3P) + N2	0.600E-10 * (T / 300)
NO + NH -> N2O + H	0.290E-10 * (T / 300) ** (-0.300E+00) * exp(0.770E+02 / T)
NO + NH -> OH + N2	0.120E-10 * (T / 300) ** (-0.300E+00) * exp(0.770E+02 / T)

HNO + H -> NO + H2	0.310E-10 * exp(-0.500E+03 / T)
HNO + N(2D) -> NO + NH	5.00E-11
N2O + CH -> NO + HCN	0.150E-10 * exp(0.257E+03 / T)
N2O + N(2D) -> N2 + NO	0.150E-10 * exp(-0.570E+03 / T)
	(0.914E-32 * (T / 300) ** (-0.600E+00) * exp(-0.0 / T))* [M]
	/ ( 1 + (0.914E-32 * (T / 300) ** (-0.600E+00) * exp(-0.0 /
H + H -> H2	T))* [M] / (0.100E-09 * exp(-0.0 / T)))
	(0.890E-28 * (T / 300) ** (-0.180E+01) * exp(-31.8 / T))*
	[M] / ( 1 + (0.890E-28 * (T / 300) ** (-0.180E+01) * exp(-
	31.8 / T))* [M] / (0.320E-09 * (T / 300) ** (0.133E+00) *
H + CH3 -> CH4	exp(-2.54 / T)))
	(0.100E-33 * exp(-0.0 / T))* [M] / (1 + (0.100E-33 * exp(-0.0 / T)))* [M] / (1 + (0.100E-33 * exp(-0.0 / T)))* [M] / (1 + (0.100E-33 * exp(-0.0 / T)))* [M] / (1 + (0.100E-33 * exp(-0.0 / T)))* [M] / (1 + (0.100E-33 * exp(-0.0 / T)))* [M] / (1 + (0.100E-33 * exp(-0.0 / T)))* [M] / (1 + (0.100E-33 * exp(-0.0 / T)))* [M] / (1 + (0.100E-33 * exp(-0.0 / T)))* [M] / (1 + (0.100E-33 * exp(-0.0 / T)))* [M] / (1 + (0.100E-33 * exp(-0.0 / T)))* [M] / (1 + (0.100E-33 * exp(-0.0 / T)))* [M] / (1 + (0.100E-33 * exp(-0.0 / T)))* [M] / (1 + (0.100E-33 * exp(-0.0 / T)))* [M] / (1 + (0.100E-33 * exp(-0.0 / T)))* [M] / (1 + (0.100E-33 * exp(-0.0 / T)))* [M] / (1 + (0.100E-33 * exp(-0.0 / T)))
HCN + H -> H2CN	0.0 / T))* [M] / (0.980E-11 * exp(-2080.0 / T)))

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377	Code availability. BATS-R-US code as a part of SWMF used to generate the scenario shown in
378	Figure 1 can be accessed through GSFC's CCMC run-on-request web site at
379	http://ccmc.gsfc.nasa.gov. We have opted not to make the computer code, Aeroplanets,
380	associated with this paper available because we are currently in the phase of adopting it for free
381	access at Exoplanetary part of CCMC's web site. The authors declare that model data supporting
382	the findings of this study are available within the article and its supplementary information.
383	Other model related data will be available on request from the authors.