

1 **Abstract**

2

3 The effects of the polychaetes *Marenzelleria* sp. (Polychaeta, Spionidae), nonindigenous,
4 rapidly increasing species in the Baltic Sea, on benthic nutrient fluxes, denitrification and
5 sediment pore water nutrient concentration were studied in laboratory experiments using
6 a flow-through setup with muddy sediment from coastal regions of the Gulf of Finland.
7 In addition, different forms of sediment phosphorus (P), separated by chemical
8 fractionation, were studied in three sediment layers. At a population density
9 corresponding to about half the highest measured in the northern Baltic Sea,
10 *Marenzelleria* sp. increased the fluxes of P and ammonium to the water column. No
11 effect could be recorded for denitrification. Since the previously dominant species of the
12 area, *Monoporeia affinis*, can enhance denitrification and lower the amount of dissolved
13 P in the pore water, the replacement of *M. affinis* with *Marenzelleria* spp. may lead to
14 increased P flux to the water column and decreased denitrification, further increasing the
15 ammonium flux to the water column. However, sediment reworking by *Marenzelleria*
16 spp. also oxidizes the surface sediment in the long run, improving its ability to retain P
17 and support nitrification. Therefore, the impact of *Marenzelleria* spp. on sediment
18 nutrient release may not be as drastic as the initial reactions seen in our experiments
19 suggest.

20

21 **Keywords**

22 Benthic nutrient fluxes, bioturbation, denitrification, *Marenzelleria* spp., phosphorus
23 fractionation.

24

1 **Introduction**

2

3 Mineralization of the organic matter in shallow bodies of water mainly occurs in
4 sediments. Environmental conditions at the bottom, such as temperature and oxygen (O_2)
5 concentration, vary seasonally and yearly, affecting the structure of microbial
6 communities, the rate at which mineralization occurs and the pathways that are used.
7 Both nitrogen (N) and phosphorus (P) mineralization and the fate of the mineralized
8 nutrients are dependent largely on the O_2 conditions. Microbial denitrification is an
9 anaerobic process converting nitrite and nitrate ($NO_x = NO_2^- + NO_3^-$) to gaseous
10 dinitrogen (N_2), thereby decreasing the amount of fixed N in the water ecosystem. In
11 recent years, another natural process removing fixed N from the water ecosystem,
12 anaerobic ammonium oxidation (anammox), was discovered in marine sediments
13 (Dalsgaard and Thamdrup, 2002; Thamdrup and Dalsgaard, 2002; Trimmer et al., 2003).
14 In this process, microbes oxidize ammonium (NH_4^+) with NO_2^- to form N_2 . The NO_x
15 used in the processes can originate either from the water column, or from an aerobic
16 nitrification process in the oxic layer of the sediment. Since nitrification is dependent on
17 the supply of O_2 , anoxic denitrification can also proceed only as long as there is enough
18 O_2 for the nitrifiers to produce NO_x . During anoxia, NH_4^+ accumulates in the bottom
19 water, from which it can return to the productive water layers in mixing or upwelling
20 events. The P release from the sediment is also commonly associated with the depletion
21 of bottom water O_2 (e.g. Mortimer, 1941, 1942). Anoxia leads to reduction of iron (Fe)
22 compounds (hydrated oxides) in sediments and in suspended particles, and to consequent
23 release of Fe-bound P to the bottom water. In addition to the Fe compounds, P in the
24 sediment also has several other binding forms, the binding strengths and reactivities of
25 which differ (e.g. Boström et al., 1982).

26

1 Benthic animals enhance microbial activities in several ways. They mix newly
2 sedimented material to the deeper sediment layers, transport oxidized and reduced
3 compounds in their burrows, concentrate organic matter in faecal pellets and reduce the
4 size of particles, thereby increasing particle surface area for microbial colonization (e.g.
5 Kristensen, 1988). The animals mix oxic bottom water into the surface sediment,
6 increasing the volume of oxygenated sediments. In addition, the channels of the
7 burrowing animals increase transport of oxic bottom water, both by diffusion and
8 physical transport by the animals, to the deeper anoxic sediment layers. The
9 mineralization products and the metabolites of the burrowing animals are effectively
10 removed from the channels by irrigation, increasing the fluxes out of the sediment by
11 keeping the concentration gradient between the sediment pore water and burrow water
12 high. The shape, location and ventilation frequency of the channels in the sediment vary
13 according to the species responsible for bioturbation and have major implications for the
14 nutrient dynamics within and around the burrows (e.g. Welsh, 2003).

15

16 Bioturbation enhances nitrification activity by increasing surfaces with access to both
17 NH_4^+ and O_2 (Pelegri and Blackburn, 1994; Tuominen et al., 1999; Svensson et al.,
18 2001), although reduction in nitrification, caused by partial digestion of nitrifiers by
19 insect larvae in organic-poor (but not organic-rich) sediment, has also been reported
20 (Altmann et al., 2004). Due to enhanced nitrification, the denitrification rate also
21 generally increases in the presence of benthic animals (Pelegri and Blackburn, 1994;
22 Svensson and Leonardson, 1996; Hansen and Kristensen, 1998; Bartoli et al., 2000;
23 Svensson et al., 2001), although this effect is not always clear (Tuominen et al., 1999).

24

25 Bioturbation enhances P binding to the sediment, since Fe compounds in oxygenated
26 sediments retain P more effectively (e.g. Andersen et al., 1991; Tuominen et al., 1999).
27 However, bioturbation can also temporarily increase sediment P release when

1 bioturbation reaches deeper, reduced sediment layers with dissolved P in the pore water
2 (e.g. Hansen et al., 1998). If the oxic sediment layer is thin, or the binding sites for P are
3 already highly saturated, there may not be enough binding sites for all the released P. In
4 larger channels, P can also be transported by advection in addition to diffusion (e.g.
5 Kristensen, 1988), and the flux of pore water P can occur so quickly that some of the P
6 released will reach the bottom water. The compounds, excreted by the benthic animals,
7 that line the walls of the burrow channels, may also hinder the adsorption of dissolved P
8 to the oxides of Fe and aluminium (Al) on sediment particles. In addition, the general
9 increase in microbial activity caused by bioturbation leads to increased release of P in
10 mineralization of organic matter (e.g. Andersen and Jensen, 1991; Hansen et al., 1998).
11 However, if the bottom water is oxic, the P released is trapped back on the suspended
12 particles and the sediment surface.

13

14 The invasive polychaete worm, identified as the North American *Marenzelleria viridis*
15 (Verrill) (Polychaeta, Spionidae), was first recorded in the southern Baltic Sea in 1985
16 (Bick and Burckhardt, 1989), in 1990 in the coastal Gulf of Finland (Stigzelius et al.,
17 1997), and in the northernmost arm of the Baltic, the Gulf of Bothnia, in 1996 (Stigzelius
18 et al., 1997). Later the taxonomy of *Marenzelleria* species was revised by Sikorski and
19 Bick (2004) who described a new species, *M. neglecta*, in the Baltic Sea. More recently,
20 the occurrence of three different species in the Baltic Sea, namely *M. viridis*, *M. neglecta*
21 and *M. arctia* (Chamberlin) was confirmed by molecular methods (Bastrop and Blank,
22 2006). The species are morphologically very similar and likely to have only minor, if
23 any, functional (e.g. bioturbation) differences. The species status in previous Baltic Sea
24 studies cannot be fully confirmed and their current distribution is under study (Blank et
25 al., in prep.).

26

1 *Marenzelleria* spp. now commonly occur in the coastal northern Baltic Sea (Laine et al.,
2 2003a; 2003b; Perus and Bonsdorff, 2004), and high abundances and biomasses were
3 recorded both in the southern Baltic (Kube et al., 1996; Zettler, 1996) and in the Gulf of
4 Riga (Cederwall et al., 1999). Recently, high densities and biomass values of up to 4000
5 ind. m⁻² and 80 g m⁻² wet weight, respectively, have been observed in the deep open Gulf
6 of Bothnia (Finnish Institute of Marine Research (FIMR), unpublished data). The
7 previously abundant and dominant amphipod (*Monoporeia affinis* (Lindström))
8 populations have declined strongly during recent decades in the coastal areas of the
9 northern Baltic Sea (Laine et al., 2003a, 2003b; Perus and Bonsdorff, 2004) and also in
10 the open Bothnian Sea (Norkko et al., 2007). Concomitantly, the importance of
11 *Marenzelleria* spp. is expected to increase in the soft-bottom ecosystem. *Marenzelleria*
12 *viridis* is an infaunal species, forming highly branched burrow networks, that feeds with
13 palps in the sediment-water interface by collecting either material deposited on the
14 sediment surface or suspended particles in the near-bottom water (Dauer et al., 1981). In
15 the southern Baltic, the L- or J-shaped burrows of *Marenzelleria* spp. extend downwards
16 25-35 cm into the sediment (Zettler et al., 1994), which is much deeper than for any of
17 the native Baltic soft-bottom macrofaunal species. Thus, in the Baltic Sea ecosystem
18 these species appear to have occupied an open niche and have the potential for affecting
19 the sediment nutrient dynamics through deep-reaching bioturbation (Olenin and
20 Leppäkoski, 1999). Despite the small, slim morphology of the animals, they have the
21 potential for affecting conditions on the seafloor, due to the large number of animals, the
22 deep burrows they dig and their clear resilience to challenging environmental conditions,
23 such as low O₂ concentrations (Schiedek, 1997).

24

25 We studied the effect of *Marenzelleria* sp. (most probably *M. arctia*, see Material and
26 methods) population on its new environment. Laboratory experiments were performed
27 using a flow-through setup with coastal Gulf of Finland muddy sediment. We measured

1 benthic nutrient fluxes, denitrification, O₂ penetration depth in the sediment and sediment
2 pore water nutrient concentration at low and high densities of *Marenzelleria* sp. as well
3 as in control sediments without animals. In addition, different forms of sediment P,
4 separated by chemical fractionation, were studied in three sediment layers with and
5 without high densities of *Marenzelleria* sp.

6

7 **Materials and methods**

8

9 **Experimental setup**

10 The *Marenzelleria* sp. individuals were collected on October 18, 2004 in the Åland Sea
11 at a 285-m-deep station (F64; 60°18.0' N, 19°15.0' E). This area hosts a dense
12 *Marenzelleria* sp. population, e.g. in June 2004 a density of 4000 ind. m⁻² was recorded
13 at this site. At the time of sampling the species was assumed to be *M. viridis* but samples
14 taken at the same site in 2005 and identified with molecular methods revealed that the
15 population consisted of *M. arctia* only (Blank et al., submitted). Thus, the worms used in
16 this study most probably belong to *M. arctia* but since the species identity in the present
17 material was not confirmed, we refer to them as *Marenzelleria* sp. Sediment was
18 collected using a van Veen grab and gently mixed with water. Swimming *Marenzelleria*
19 spp. were caught in a 1-mm sieve. The animals were then placed in aerated transport
20 boxes with 5 cm of sieved local sediment and kept at +5 °C until transport to the
21 laboratory.

22

23 The sediment used in the experiments was collected from a coastal station (Tvärminne,
24 Storfjärden, northern Gulf of Finland, 59°51.3' N, 23°15.8' E), representing a
25 characteristic, outer archipelago accumulation bottom consisting of soft mud. The water
26 depth at the sampling station is 33 m. A box corer was used to collect the deeper anoxic
27 sediment and an Ockelman sledge to collect the oxic surface sediment. The mud was

1 gently sieved (1 mm for the deeper sediment and 0.5 mm for the surface slurry) to
2 remove macrozoobenthos. The experiment was conducted at +5 °C in a coldroom,
3 corresponding to *in situ* temperature. Round polyacrylic aquariums (30 aquariums with
4 inner diameter 14 cm, height 16 cm) were packed with 8 cm of deeper mud that was
5 covered with 2 cm of surface mud. The aquariums were filled with filtered seawater of
6 the same salinity (6.1) and temperature as the sampling station, taken at the beginning of
7 the experiment from a nearby location, filtered (0.2 µm) and kept in the coldroom for the
8 length of the experiment. The aquariums were covered with 0.5-mm wire mesh covers.
9 After letting the aquariums stabilize for 12 days, the first samples (time zero
10 measurements) were taken and animals were distributed in the aquariums so that nine
11 aquariums received no worms (control units, C), nine received five worms each (low-
12 density units, LOW, corresponding to 325 ind. m⁻²) and nine received 30 worms each
13 (high-density units, HIGH, corresponding to 1950 ind. m⁻²). The mesh covers were
14 equipped with hypodermic needles that were attached to tubing, bringing filtered water
15 from the container through a peristaltic pump at a rate of 1.5 ml min⁻¹. Excess water was
16 allowed to overflow.

17

18 **Sampling and analyses**

19 Samples were taken at time zero (before adding the worms) and then on days 2, 6 and 14
20 after connecting the flow-through system. The sampling was randomized and three
21 aquariums of each treatment were sacrificed at each sampling time (except on day 14,
22 when one LOW, one HIGH and two C units were discarded due to problems in
23 waterflow).

24

25 The samples for nutrient fluxes (NH₄⁺, NO_x, phosphate (PO₄³⁻) and Fe) were collected
26 from the aquariums using a peristaltic pump, collecting outflowing water at the same rate
27 as water was pumped into the aquariums (1.5 ml min⁻¹). In addition, the inflowing water

1 from the container was sampled and the fluxes were calculated from the difference in
2 concentrations, using the residence time and sediment area. All nutrients were analysed
3 using standard methods (Koroleff, 1983). All units were photographed prior to sediment
4 sampling. The sediment samples for all the different measurements were taken
5 simultaneously so that all the tube cores needed were inserted in the aquarium at the
6 same time, ensuring undisturbed samples.

7

8 Denitrification was measured using the isotope pairing technique (Nielsen, 1992). Three
9 replicate samples were taken in clear plastic cores (diameter 2.6 cm, height 9 cm) so that
10 about half of the core was filled with the sediment and half with the water from above.

11 The samples were enriched with K^{15}NO_3 (98% labelling, Cambridge Isotope
12 Laboratories, Andover, MA, USA) to a final concentration of $100\ \mu\text{M}$ of $^{15}\text{NO}_3^-$ in the
13 overlying water and incubated, with a magnetic stirrer on the lid of the cores, at *in situ*
14 temperature in darkness for 3-4 hours. Incubation was terminated with ZnCl_2 , samples
15 were mixed and subsamples were sent in gastight 12-ml vials (Exetainer; Labco, High
16 Wycombe, Buckinghamshire, UK) to the National Research Institute, Silkeborg,
17 Denmark, for analysis of N_2 isotopic composition.

18

19 The O_2 profiles were measured in undisturbed sample cores, similar to those used in the
20 denitrification measurements, using Clark-type oxygen microelectrodes (100- μm tips,
21 OX-100; Unisense, Aarhus, Denmark) that gave a spatial resolution of about 200 μm . For
22 pore water nutrient analyses, one sample per aquarium was taken using a corer with an 8-
23 cm inner diameter. The sample was sliced in 2-cm strata from the surface to and 8-cm
24 depth in the sediment, and slices from the same depth from replicate aquariums (three of
25 each treatment) were combined. The sediment redox potential was measured in each
26 combined slice with an electrode (SenTixORP), corrected by temperature and converted
27 to an E_h value. The pore water was then extracted by squeezing the samples with a

1 Millipore Zero Headspace Extractor, (Millipore, Billerica, MA, USA) using 0.45- μm
2 filter (Nuclepore; Whatman), under N_2 atmosphere.
3
4 Two replicate sediment samples for P fractionation studies were collected from two C
5 units and two HIGH units on days 2 and 14. The samples were taken into small plastic
6 corers (see denitrification) and the water above the sediment was removed (siphoned)
7 immediately. The samples were kept in the corers at +5 °C, capped and protected from
8 light. Each sediment core was subsampled the following day in an N_2 atmosphere (O_2
9 content below 5%) in a glove box, separating three sediment layers: 0-2 cm, 2-4 cm and
10 4-6 cm. The subsamples were stored in plastic bottles at +5 °C until analyses (for about 3
11 weeks). The various chemical forms of sediment P were determined using a P
12 fractionation method slightly modified from that described in Jensen and Thamdrup
13 (1993) (detailed description in Lukkari et al., submitted). The method separates six P
14 pools (Jensen et al., 1995): loosely adsorbed and pore water P (extracted with sodium
15 chloride, NaCl; referred to as NaCl-iP), a redox-sensitive fraction of P bound to hydrated
16 oxides of reducible metals (mainly those of Fe) (sodium dithionite, $\text{Na}_2\text{S}_2\text{O}_4$, in
17 bicarbonate buffer, NaHCO_3 , at pH 7; NaBD-iP), P bound to oxides of Al and
18 nonreducible Fe (sodium hydroxide, NaOH; NaOH-iP), apatite-P (hydrochloric acid,
19 HCl; HCl-iP) and residual, mainly organic, P (extracted with HCl after combustion; Res-
20 iP). In addition to these five fractions, the pool of mobile organic P (nonreactive P, NRP)
21 was determined as the difference between total P (TP) and dissolved inorganic P (iP),
22 summarized from the first three steps. The iP was determined from filtered (Nuclepore
23 polycarbonate membranes, pore size 0.4 μm) extracts with a UV-VIS spectrophotometer
24 (Genesys 10uv Thermo Spectronic; Genesys, Daly City, CA, USA) with a 50-mm flow-
25 injection cuvette and TP with a spectrophotometer after acid persulphate digestion
26 (Koroleff, 1983).

27

1 The sediment dry matter (DM) content was first determined using a moisture analyser (a
2 balance equipped with a halogen lamp dryer, Ohaus MB45; Ohaus, Pine Brook, NJ,
3 USA), and the amount of fresh sediment extracted was determined as a fresh weight
4 corresponding to 0.3 g dry mass. The volume of the extracts was always 30.0 ml yielding
5 a sediment DM-to-solution ratio of 1:100. The sediment subsamples were not sieved, but
6 visible animals or their remains were removed.

7

8 The effects of *Marenzelleria* spp. and time on denitrification and benthic nutrient fluxes
9 were tested using two-factor analysis of variance (ANOVA) and when significant
10 differences ($\alpha = 0.05$, $p < 0.01$) were found, treatments were further tested against the C
11 units, using Dunnett's test ($\alpha = 0.05$). Differences in pore water nutrient profiles between
12 the treatments and sampling occasions, based on concentrations in different layers, were
13 tested using the multivariate analysis of similarities (ANOSIM) procedure (Clarke and
14 Warwick, 2001). The differences in sediment P fractions between the C and HIGH
15 treatments were tested using t-tests assuming unequal variances ($\alpha = 0.05$, $p < 0.05$) and
16 the effects of time within treatment were tested using paired t-tests ($\alpha = 0.05$, $p < 0.05$).
17 The significances of all correlations (denitrification, fluxes, concentrations of dissolved
18 O₂ and nutrients) were analysed using Pearson's correlation test ($\alpha = 0.05$, $p < 0.01$).

19

20 **Results**

21

22 In the C units, the two-layer structure basically persisted until the end of the experiment.
23 By day 6 some darkening of the initially light brown colour in the lower parts of the
24 surface layer was observed. By day 14, there was an even more clear difference and only
25 the uppermost 5 mm in the C units had retained the original colour, whereas the lower
26 part of the surface layer was clearly darker, with black spots, indicating reduction and

1 sulphide formation. Microbial growth on the sediment surface was indicated by light-
2 coloured patches and increase in surface roughness by day 14.

3

4 The worms began to burrow immediately after they had been transferred to the
5 aquariums. Burrows were observed from day 2 onwards, both in the LOW and HIGH
6 units. Sometimes worm heads projected out of the sediment and the worms were lashing
7 with their palps to feed but disappeared quickly when disturbed. The burrows formed a
8 dense network in the vertical and horizontal directions, especially in the HIGH density
9 units. Active burrowing was observed in the uppermost 5 cm and single burrows
10 penetrated to more than 8-cm depths in the sediment. Occasionally a thin and transparent
11 membrane lining was observed in the burrows. The deeper parts of the upper sediment
12 layer became darker in the same way as in the C units. However, the actively used worm
13 burrows could be distinguished even in the deep sediment by a thin light-coloured,
14 apparently oxidized layer that was similar to the colour of the sediment surface.

15

16 Burrow openings and faecal pellets were visible on the sediment surface. The faecal
17 pellets were up to 7 mm in length and formed radial cones around the openings. In the
18 HIGH units the sediment surface was almost completely covered by pellets by day 14,
19 whereas in the LOW units roughly half of the surface was covered by pellets and the
20 other half resembled the surface of the C units. After day 14 some additional units that
21 were not used for the experiment were sieved. The worms in these units were all alive
22 and actively swimming when transferred to the water.

23

24 The O₂ penetration depth ranged from 0.3 to 1.5 mm and was 0.5 mm at 17 out of the 28
25 times measured, with no differences between the treatments. The O₂ consumption
26 likewise did not differ between the treatments, since the O₂ concentration 0.5 cm above
27 the sediment surface was similar in all aquariums. There were no clear differences in

1 redox profiles between the different units. However, the redox potential decreased in the
2 surface layer (0-2 cm) of all units during the experiment from an initial level of -10 to
3 +105 mV to a level of -130 to -30 mV in the end. In the deeper sediment the redox
4 potential varied less and was mostly between -150 and -100 mV. Despite the dense
5 burrow networks that developed especially in the HIGH units, no obvious change was
6 detected in porosity between the treatments. However, the surface sediments could be
7 distinguished from the deeper sediments by the porosity difference.

8

9 **Nutrient fluxes**

10 The nutrient fluxes at the sediment-water interface fluctuated considerably during the
11 experiment (Figure 1). The NO_x flux was out of the sediment, except on the last sampling
12 day when it was directed into the sediment in all treatments. There were significant
13 differences between the C and HIGH (but not the C and LOW) units on day 2, when the
14 C units released nine times more NO_x out of the sediment than did the HIGH units. No
15 differences between the treatments were found on day 6, and the significance of the
16 differences on day 14 could not be tested due to the lack of replicates of the C treatment.
17 The NH_4^+ fluxes increased throughout the experiment, were always directed out of the
18 sediment and were always highest in the HIGH, and lowest in the C units. Significant
19 differences between treatments were found on day 2, when the HIGH units released nine
20 times more NH_4^+ than did the C units, but not on day 6. On day 14 the difference was
21 even greater, but the significance could not be tested (see above). The PO_4^{3-} flux was
22 likewise directed out of the sediment throughout the experiment and the highest values
23 were always recorded in the HIGH units and lowest in the C units. The differences
24 between the HIGH (but not the LOW) and C units were significant on days 2 and 6. They
25 were even higher on day 14, but the significance could not be tested (see above). The Fe
26 flux was always out of the sediment and increased in the course of the experiment. On

1 day 2, the flux in the HIGH (but not the LOW) units was significantly higher than in the
2 C units.

3

4 The NO_3^- flux correlated positively (Pearson correlation, Table 1) with the O_2 penetration
5 depth, NO_3^- concentration and denitrification, and negatively with the denitrification
6 potential (D15). The NH_4^+ flux was negatively correlated with total denitrification (Dtot)
7 and coupled nitrification-denitrification (Dn), and positively with D15. The PO_4^{3-} flux
8 was positively correlated only with the Fe flux, which was negatively correlated with the
9 O_2 penetration depth.

10

11 **Denitrification**

12 No effect of *Marenzelleria* spp. was detected on the denitrification rates in the
13 experiment. The Dtot and Dn values decreased significantly in all aquariums in the
14 course of the experiment, independent of the treatment (Figure 2). The percentage of
15 denitrification that was based on water column NO_x (Dw) likewise differed significantly
16 between sampling days, but instead of a general decrease, it showed very high values on
17 day 6 in all treatments, decreasing again to very low values by day 14. No interactions
18 between the treatments and time were found. In contrast to the ^{14}N based natural
19 denitrification (Dtot), the denitrification potential (D15), based on the added ^{15}N , tended
20 to increase throughout the experiment in all treatments (Figure 2). In addition to the time,
21 the treatments also significantly affected the D15 rates. On days 2 and 6, the D15 rates in
22 the HIGH, but not in the LOW, units were significantly higher than in the C units. The
23 effect disappeared by day 14, when no differences between units with and without
24 animals were found.

25

26 The Dtot rates correlated positively (Table 1) with Dn, O_2 penetration depth and NO_3^-
27 concentration and flux, and negatively with the D15 rate and NH_4^+ flux. The Dn level

1 was positively correlated with the O₂ penetration depth and negatively with the D15 rate
2 and NH₄⁺ flux. The Dw level correlated positively with the NO₃⁻ concentration and flux.
3 The D15 rate was positively correlated with the NH₄⁺ flux and negatively with the NO₃⁻
4 concentration.

5

6 **Pore water nutrient profiles**

7 In all units, the pore water NH₄⁺ concentration increased almost linearly with sediment
8 depth (Figure 3). The concentrations in the surface layer (0-2 cm) varied from about 220
9 to 380 μmol l⁻¹ at the beginning of the experiment and increased to about 300-425 μmol
10 l⁻¹ on day 14. In the bottom layer (6-8 cm) the concentration was more stable, about 820-
11 1050 μmol l⁻¹. No differences were found between the C, LOW and HIGH units. Only on
12 day 14 were the NH₄⁺ concentrations in all layers of the HIGH units slightly lower than
13 in the C and LOW units. The sediment pore water NO_x concentration was measured only
14 in the surface layer (0-2 cm) and was very low throughout the experiment (mean 1.9
15 μmol l⁻¹).

16

17 The PO₄³⁻ profiles were the most variable of the measured pore water parameters (Figure
18 3). In the surface layer, the concentrations were below 65 μmol l⁻¹ except on day 14 in
19 the C unit. The maximum PO₄³⁻ concentrations in the deeper layers exceeded 200 μmol l⁻¹
20 ¹. On days 6 and 14, the PO₄³⁻ concentrations in the deeper layers of the HIGH units were
21 lower than in the C and LOW units. The Fe concentration varied only in the uppermost
22 layer (Figure 3). In the C units the surface concentration of Fe always exceeded 15 μmol
23 l⁻¹. In all units the concentrations at 2-8-cm depths were very low, < 5 μmol l⁻¹, and did
24 not vary during the experiment. In the HIGH units the surface layer Fe concentrations fell
25 to low levels, similar to those observed in the deep sediments on days 6 and 14. No
26 significant differences were detected in any of the pore water nutrient profiles between
27 the treatments or sampling occasions.

1

2 **Sediment P fractions**

3 The difference in P fractions between the treatments and as a function of time could not
4 be tested statistically at each data point, because there were only two replicate samples in
5 the C units on day 14. The concentration of total extractable P (TP_{extr}) was about 60 and
6 $40 \mu\text{mol g}^{-1}$ DM in the surface (0-2 cm) and bottom (2-6 cm) sediment layers,
7 respectively, with NaBD-iP and HCl-iP forming the major fractions. NaCl-iP and NaBD-
8 iP were the fractions most affected by the treatment, while Res-P was the most stable P
9 form (Figure 4).

10

11 NaCl-iP constituted only a minor percentage of TP_{extr} (< 1%). Its concentration was
12 highest in the surface layer (0-2 cm) and did not change markedly in the C units between
13 days 2 and 14. The increase in NaCl-iP in the surface layer from day 2 to day 14 in the
14 HIGH units was significant. On day 2, the HIGH units had significantly lower NaCl-iP at
15 0-2 cm than the C units. On day 14, no marked differences between the treatments were
16 found.

17

18 NaBD-iP formed 11-35% of the TP_{extr} . Its percentage of the sediment P was higher in the
19 0-2-cm layer than in the other two layers and increased during the experiment, especially
20 in the surface layers and in the HIGH units. NaBD-iP in the HIGH units was significantly
21 higher on day 14 than on day 2 in the two uppermost layers. The evident increase with
22 time in the C units (Figure 4) could not be tested statistically (only two replicates on day
23 14). On day 2, NaBD-iP was lower in the HIGH units than in the C units at 0-2 cm and 2-
24 4 cm; the concentration difference was larger in the former, although statistically
25 significant only in the latter. On day 14, the HIGH units had lower NaBD-iP
26 concentrations at 2-4 cm and 4-6 cm than the C units.

27

1 NRP constituted 20-31% of the TP_{extr} and its percentage of the sediment P decreased in
2 the C units and increased in the HIGH units (not significantly) during the experiment.
3 The decrease in NRP in the surface layer of the C units is evident (Figure 4), while that in
4 the 2-4-cm and 4-6-cm layers is within standard deviation of the replicate results. On day
5 2, NRP was lower in all depth layers in the HIGH units compared with those in the C
6 units, but significantly only in the surface layer. On day 14, the NRP concentration was
7 higher in the two deepest layers of the HIGH units than in the C units (Figure 4).

8

9 NaOH-iP formed 4-7% of the TP_{extr} in the sediment and its percentage did not change
10 during the experiment. However, possibly due to the very small deviation in the results,
11 the relatively small ($< 1 \mu\text{mol g}^{-1} \text{DM}$) changes in NaOH-iP concentrations were
12 statistically significant. NaOH-iP decreased in the 0-2-cm layer of the C units between
13 days 2 and 14. On day 2, the NaOH-iP concentration was lower in the 4-6-cm layer of the
14 C units than of the HIGH units, and on day 14 it was higher in the 2-4-cm layer of the C
15 units compared with the HIGH units.

16

17 The percentage of HCl-iP was 17-35% of the TP_{extr} . No statistically significant
18 differences were found in the concentrations of HCl-iP between days and treatments. The
19 Res-P fraction formed 15-27% of the TP_{extr} . Its percentage of TP increased slightly in the
20 C units and decreased in the HIGH units during the experiment. None of the changes
21 were statistically significant, but the increase in Res-P in the surface layer of the C units
22 from day 2 to day 14 is evident (Figure 4). In addition, despite the high deviation in the
23 HIGH units (especially on day 14), the results suggest a decrease in Res-P from day 2 to
24 day 14. On day 2, Res-P was higher in the HIGH units compared with the C units and
25 vice versa on day 14.

26

27

1 Discussion

2

3 The redox conditions in the surface layer of the aquariums gradually deteriorated in all
4 treatments during the experiment. This could be seen both visually, as a darkening of the
5 light brown surface layer from below, and as a decreased redox potential (Mortimer
6 1941). Reduced substances diffused from the deep sediment layer upwards and at the
7 same time the mineralization processes in the upper sediment layer used up the O₂
8 storage, resulting in a decreasing volume of the oxidized sediment. Despite active burrow
9 formation, the *Marenzelleria* spp. did not counteract the decrease in redox potential by
10 transporting more O₂ to the sediment, when the entire sediment volume was considered.
11 They also did not affect the porosity, which may have been caused by different water and
12 sediment relocation in the HIGH units in comparison to C units. Burrow formation could
13 have caused sediment packing in the interspace between the burrows with no overall
14 change in the water content. The coarse sampling procedure may also have failed to
15 detect small-scale differences in the loose surface layer where the water content was
16 highest.

17

18 The animals affected the nutrient fluxes shortly after they were introduced to the
19 sediment, but the effect may have been transient, since no significant differences were
20 found in the rates on day 6 (except for the PO₄³⁻ flux). The differences could not be
21 tested on day 14, but appeared to increase from day 6 for NH₄⁺ and PO₄³⁻, although there
22 were wide variations. No significant differences were detected in the pore water nutrient
23 profiles, although some effects could be seen at the higher population density. Similarly,
24 Karlson et al. (2005) found no significant effect of *M. viridis* (or *Macoma balthica* (L.) or
25 *Monoporeia affinis*) on pore water nutrients in their recolonization experiment, using
26 reduced natural Baltic Sea sediments. The general increase in the surface layer NH₄⁺
27 concentration indicates diffusion from the deeper layers. The slightly lower NH₄⁺

1 concentrations in the HIGH units, compared with the C and LOW units, at the end of the
2 experiment is in accordance with the observed higher efflux of NH_4^+ that could have
3 been caused by the more intensive burrow formation at the higher worm density.

4

5 **Nitrification and denitrification**

6 The high flux of NO_3^- out of the sediment confirms that the nitrification rate was rapid in
7 the sediment. On day 2, the HIGH units released significantly less NO_3^- in the water than
8 did the C units. The Dn rate was, however, similar in all treatments, indicating that
9 nitrification saturated the NO_3^- demand of the denitrifiers in all treatments. At the same
10 time the NH_4^+ flux was significantly higher in the HIGH units, suggesting either a lower
11 nitrification rate in the HIGH units, simultaneous dissimilatory NO_3^- reduction to NH_4^+
12 (DNRA) or simply more rapid transport of NH_4^+ out of the sediment. A lower
13 nitrification rate was not likely, because the animals enhance nitrification by increasing
14 the oxic-anoxic interfaces in the sediment (Pelegri and Blackburn, 1994; Tuominen et al.,
15 1999; Svensson et al., 2001). DNRA is usually considered to be of minor importance in
16 natural sediments and to occur mainly in organically enriched environments, such as fish
17 farm sediments (Hattori, 1983; Christensen et al., 2000). However, Karlson et al. (2005)
18 found that DNRA represented a major pathway of NO_3^- removal in laboratory
19 experiments similar to ours, in which reduced Baltic Sea sediments were used. Yet
20 another explanation for the low NO_3^- and high NH_4^+ fluxes is the shorter residence time
21 of NH_4^+ , produced in the mineralization of organic matter in the sediments with animals,
22 due to the enhanced transport from the burrows into the overlying water. Similarly,
23 increase in NH_4^+ flux was observed in deep-burrowing polychaete *Nereis virens* (Sars)
24 (Henriksen et al., 1980). The situation on day 14, when the NO_3^- concentration in the
25 water was very low and the flux was towards the sediments, represents typical steady-
26 state conditions in which denitrification consumes NO_3^- at the rate it is produced. That
27 the Dn rate did not increase as the flux turned into the sediments suggests that other

1 processes, such as DNRA, contributed to the NO_3^- reduction as well. This is also
2 supported by the high NH_4^+ flux out of the sediments.
3
4 Macrofauna stimulated denitrification in several studies (e.g. Kristensen and Blackburn,
5 1987; Hansen and Kristensen, 1998; Pelegri and Blackburn, 1994; Svensson and
6 Leonardson, 1996; Gilbert et al., 1998; Bartoli et al., 2000; Svensson et al., 2001).
7 However, in the only study using *M. viridis*, no significant effects on denitrification rate
8 were found (Karlson et al., 2005). Similarly, in our study the *Marenzelleria* spp. did not
9 affect the denitrification rates that decreased in all aquariums during the experiment,
10 independent of the treatment. No explanation was found for the decline. The
11 denitrification process is generally regulated by temperature, O_2 concentration (both
12 directly and indirectly via nitrification), NO_3^- concentration and the availability of
13 organic carbon. In these experiments, the temperature was stable and the concentration of
14 O_2 dropped from 75% to 50% saturation from day 0 to day 2, after which it remained
15 constant in all aquariums. The O_2 penetration depth likewise did not change. The NO_3^-
16 availability clearly regulated the Dw rate that fluctuated together with the NO_3^-
17 concentration, but the Dn rate was independent of the water column NO_3^- concentration.
18 Dn is controlled by the rate of nitrification, which in turn is regulated by the availability
19 of O_2 and NH_4^+ , neither of which were in short supply in the aquariums, although the
20 availability of NH_4^+ increased during the experiment. The availability of organic carbon
21 in the aquariums could have been too low to support the heterotrophic denitrifying
22 community. However, both the denitrification rates and the concentrations of labile
23 organic carbon peak in October-December in the study area (Hietanen and Kuparinen, in
24 press). In addition, the D15 rates nearly doubled from day 2 to day 14 in all aquariums,
25 indicating that enough carbon was available for the denitrifiers. The increasing D15 rates
26 also verify that the denitrifying bacteria were not grazed faster than they grew. While the
27 animals did not affect D14 (denitrification based on the natural $^{14}\text{NO}_3^-$), the D15 rate was

1 significantly enhanced on days 2 and 6 in the HIGH units, although the effect
2 disappeared by day 14. The D15 rate reflects D_w in unlimited NO_3^- concentrations. The
3 increase in D15 in the HIGH units could have been caused by expansion of the
4 denitrifying zone, since the animals create oxic-anoxic interfaces in their burrows below
5 the primary oxidized layer. That this effect was not significant in the D_w rates could be
6 explained by the slightly lower NO_3^- concentration in the HIGH units – the D_w rates
7 were similar in all treatments, although less NO_3^- was available in the HIGH units.
8 However, time was also the most significant factor affecting the D15 rate, since no
9 differences between the rates could be seen on day 14 any more. The increase in D15 in
10 the C units also indicates that the sieved and repacked sediment layers had not fully
11 stabilized by the beginning of the experiments, despite the 12-day interval between filling
12 the aquariums and the first measurements. The same effect could also be seen in the NO_x
13 flux that reached the typical steady-state pattern only by day 14. In experiments using a
14 similar setup and northern Baltic Sea sediments, Tuominen et al. (1999) found that the
15 pore water NO_x profile already resembled that of the intact sediment 2 days after
16 sediment repacking, indicating rapid recovery of the sensitive nitrifiers after sediment
17 mixing. In these experiments the cores were fully stabilized after 2 weeks. However, due
18 to the slow growth of the nitrifying bacteria compared with the denitrifying bacteria, it
19 may also take considerably longer for the system to stabilize (Welsh, 2003).

20

21 **Dissolved phosphorus**

22 The PO_4^{3-} profiles in the C units reflect the different redox states of the two sediment
23 layers. The oxidized surface layer retains PO_4^{3-} in the Fe compounds, decreasing its
24 concentration in the pore water, while in the more reduced bottom layer PO_4^{3-} remains in
25 the pore water in dissolved form, diffusing upwards driven by the concentration
26 difference between the layers (Mortimer 1941). The maximum formed at the 2-4-cm
27 layer, immediately below the interface between the surface and bottom layers (see also

1 Lewandowski and Hupfer, 2005). The differences in the pore water PO_4^{3-} between the
2 treatments already appeared on day 2 and by day 14 the PO_4^{3-} concentration in the HIGH
3 treatment was clearly lower than in the C and LOW treatments. This may have resulted
4 from the more intensive mixing and subsequent higher initial release of PO_4^{3-} into the
5 pore water in the HIGH treatments, also seen as the higher flux of PO_4^{3-} out of the
6 sediment on all sampling occasions. A similar effect was noted by Lewandowski and
7 Hupfer (2005) in experiments using chironomids (*Chironomus plumosus* (L.)) and
8 oligochaetes (*Tubifex tubifex* (Müller) and *Limnodrilus hoffmeisteri* (Claparède)). In
9 addition, organic substances excreted by burrowing animals (e.g. Kristensen, 1988) and
10 the organic molecules produced in the enhanced degradation or transformation of organic
11 matter may to some extent block the binding surfaces for P in the burrow channels,
12 leading to higher efflux out of the sediment. In the LOW treatments, the PO_4^{3-} flux
13 increased to the same level as in the HIGH units on the last sampling occasion,
14 suggesting that the less extensive burrow formation at the lower animal density delayed
15 the release of the dissolved P from the bottom layer. The high and increasing
16 concentration of Fe in the pore water at the surface of the C units is related to the
17 reduction and upward diffusion in the deeper layers (Mortimer 1941). In contrast, the
18 surface pore water Fe concentration decreased throughout the experiment, in the presence
19 of *Marenzelleria* spp., especially at the high density, indicating higher release from the
20 sediment (seen also in the Fe flux to the overlying water), or precipitation as
21 oxyhydroxides to the surface sediment. Similar observations were made by Lewandowski
22 and Hupfer (2005) in experiments using chironomids and oligochaetes.

23

24 **Sediment P fractions**

25 The high density of *Marenzelleria* sp. clearly affected the distribution of P into different
26 forms in the sediment. At the beginning of the experiment, the activity of *Marenzelleria*
27 sp. released loosely bound or pore water P (NaCl-iP) to the overlying water from the

1 surface sediment, which was reflected in the high PO_4^{3-} flux out of the sediment in the
2 HIGH units. Although small, such changes in NaCl-iP may be important, because even P
3 concentrations near the detection limit of the common analytical methods used can affect
4 algal growth (Baldwin et al., 2003). The effect of *Marenzelleria* sp. was most
5 pronounced in the redox-sensitive P fraction (NaBD-iP). The increase in NaBD-iP in the
6 surface sediment of the HIGH units was probably related to sediment oxidation by
7 animal reworking, as was also noted by Lewandowski and Hupfer (2005). The increase
8 in redox-sensitive P in the surface layer of the C units may have been caused by a slow
9 reduction of the bottom sediment and consequent release, upward diffusion and
10 entrapment of P to the Fe compounds in the oxic sediment.

11

12 Part of the transformable organic P (NRP) may have been degraded in the surface
13 sediment layer during the 12-day preincubation, since NRP also contains easily
14 hydrolyzable compounds (Ahlgren et al., 2005). The lower percentage of NRP in the
15 bioturbated units 2 days after introducing the animals may have resulted from enhanced
16 mineralization of the easily degradable organic molecules (Hansen et al., 1998), although
17 it is not clear whether this could have occurred within such a short time. Pure physical
18 mixing of the surface and bottom sediment layers by the animals does not solely explain
19 the lower NRP levels, because the same effect was also seen in the two deeper layers.
20 The increase in NRP towards the end of the experiment in the HIGH treatment, on the
21 other hand, may have resulted from the enhanced degradation of the more resistant
22 organic matter affected by sediment reworking, and the congruent decrease in the Res-P
23 fraction supports this conclusion. Kristensen et al. (1992) reported that bioturbation
24 caused loss of detritus and that the relatively refractory organic matter was affected more
25 than the labile material. Since NRP extracted with NaOH also includes inorganic P
26 compounds, e.g. polyphosphates and pyrophosphates that are produced in biological
27 transformation processes (degradation of organic matter), these probably also increased

1 the percentage of this fraction (Hupfer et al., 1995; Ahlgren et al., 2005). Defecation of
2 *Marenzelleria* sp., visible as piles of faecal pellets on the sediment surface, may also
3 have increased the share of NRP.

4

5 As expected, the P bound to oxides of nonreducible metals (NaOH-iP) was not greatly
6 affected during the experiment, since release from this fraction requires increase in pH or
7 presence of competing anions (Scheffer and Scheffer, 1984; Ryden et al., 1987). Benthic
8 animals can indirectly liberate NaOH-iP by mixing the surface sediment with bottom
9 water of higher pH (Drake and Heaney, 1987; Lewandowski et al., 2005). In our
10 experiment, the sediment probably acted as a buffer against pH changes, but some
11 microscale changes may have occurred, explaining the minor changes observed. These
12 can also be artifacts, despite the statistical significance, since the experimental units had
13 not fully stabilized by the beginning of the experiments.

14

15 Apatite P (HCl-iP), which is mainly detrital, is a resistant form of sediment P and did not
16 change markedly during this experiment. Res-P, however, showed marked changes
17 despite its theoretically resistant nature. Res-P contains primarily organic P that is
18 resistant to degradation, but may release P under favourable conditions, e.g. with
19 enhanced microbial degradation. This may explain the decrease in Res-P in the
20 *Marenzelleria* sp.-treated units during the experiment. The total organic P also decreased
21 in the bioturbated units during the experiment, suggesting that the high density of
22 *Marenzelleria* spp. enhanced degradation of the organic P in the sediment (Hansen et al.,
23 1998).

24

25 **Ecological aspects of *Marenzelleria* spp. invasion**

26 Introduced species alter the community structure and may also cause significant large-
27 scale changes in the function of the recipient ecosystem (e.g. Carlton, 1996; Leppäkoski

1 et al., 2002; Vanderploeg et al., 2002). The soft-bottom macrozoobenthos in the Baltic
2 Sea is characterized by very low species richness and a low number of functional groups
3 (Bonsdorff and Pearson, 1999; Laine, 2003). Different and even contrasting effects by
4 macrofauna on nutrient fluxes were explained by species-specific differences in the mode
5 of sediment mixing and structure building, irrigation behaviour and burrowing depth
6 (Welsh, 2003; Mermillod-Blondin et al., 2004; Michaud et al., 2006). Thus, invasions by
7 nonindigenous species, if they result in a change in the community structure, also have a
8 high potential for causing changes in biogeochemical cycles.

9

10 The recent invasion of *Marenzelleria* spp. in the northern Baltic Sea now also extends to
11 open sea areas in the Gulf of Bothnia. This change has also caused a complete shift in
12 community dominance. Some areas that previously were highly dominated by the
13 amphipod *Monoporeia affinis* (e.g. Andersin et al., 1977) are currently dominated by
14 corresponding numbers of *Marenzelleria* spp. At several sites, densities corresponding to
15 and exceeding (up to 4000 ind. m⁻²) the densities used in our experiment developed in
16 only 1-2 years after the first occurrence of the species (FIMR unpublished data). This
17 could also have caused profound changes in the nutrient fluxes, since these two species
18 may affect nutrient cycling in different ways. *Monoporeia affinis* can stimulate
19 denitrification and decrease the amount of dissolved P in the pore water of sediments in
20 the Gulf of Finland (Gran and Pitkänen, 1999; Tuominen et al., 1999). Karlson et al.
21 (2005) also found higher denitrification rates in sediments affected by *M. affinis*,
22 compared with *M. viridis*, but no difference in the P fluxes. However, in the less
23 eutrophied Gulf of Bothnia, no increase in the denitrification rate due to bioturbation of
24 the amphipods could be detected (Tuominen et al., 2003), probably because the oxidized
25 sediment layer is thicker than in the Gulf of Finland and bioturbation by *M. affinis*,
26 therefore, does not reach deep enough to increase the volume of the oxidized sediment.
27 The different bioturbation effects of *Monoporeia affinis* and *Marenzelleria viridis* are

1 probably caused by their functional differences. *Monoporeia affinis* is a surface and
2 subsurface deposit feeder that dwells mostly in the uppermost sediment surface layer
3 (Hill and Elmgren, 1987; Karlson et al., 2005). Its active sediment mixing at high
4 densities causes homogenization of sediment without permanent burrow structures,
5 whereas *Marenzelleria* builds more permanent burrows that may enhance the solute
6 fluxes between the deeper sediment layers and the overlying water. However, in natural
7 communities with mixed species compositions, the patterns identified experimentally for
8 single species may not emerge similarly, due to the complex interactions in the
9 mineralization processes (Welsh, 2003). Therefore, *in situ* studies are needed to predict
10 the actual impact of *Marenzelleria* in the benthic processes of the Baltic Sea.

11

12 **Conclusions**

13

14 In our 2-week experiments *Marenzelleria* sp. increased the fluxes of P and NH_4^+ to the
15 water column. No effect could be recorded for denitrification. Since the previously
16 dominant species of the area, *Monoporeia affinis*, can enhance denitrification and lower
17 the amount of dissolved P in the pore water, the replacement of *M. affinis* with
18 *Marenzelleria* spp. may lead to increased P flux to the water column and decreased
19 denitrification, further increasing the NH_4^+ flux to the water column. Nutrient release to
20 the bottom water instead of burial and removal from the water ecosystem by
21 denitrification further accelerates eutrophication, the main problem in the Baltic Sea.
22 However, in the long run sediment reworking by *Marenzelleria* spp. also oxidizes the
23 surface sediment, improving its ability to retain P and support nitrification. Therefore, the
24 impact of *Marenzelleria* spp. on sediment nutrient release may not be as drastic as the
25 initial reactions seen in our experiments suggest.

26

27

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9

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14

1 **Tables**

2

3 Table 1. Pearson correlation matrix ($\alpha = 0.05$, $p < 0.01$). Dtot total denitrification, Dw
 4 denitrification based on water column NO_x , Dn coupled nitrification-denitrification, D15
 5 denitrification potential, O_2 depth O_2 penetration depth into the sediment, Conc
 6 concentrations and Flux fluxes of O_2 and nutrients. Significant correlations marked in
 7 bold.

8

| | Dtot | Dw | Dn | D15 | O_2 depth | Conc NO_x | Conc O_2 | Flux NH_4^+ | Flux NO_x | Flux PO_4^{3-} |
|-------------------------|--------------|-------------|--------------|--------------|-----------------------|-----------------------|----------------------|-------------------------|-----------------------|----------------------------|
| Dw | 0.51 | | | | | | | | | |
| Dn | 0.88 | 0.04 | | | | | | | | |
| D15 | -0.66 | -0.05 | -0.74 | | | | | | | |
| O_2 depth | 0.72 | 0.50 | 0.56 | -0.41 | | | | | | |
| Conc NO_x | 0.77 | 0.79 | 0.46 | -0.52 | 0.81 | | | | | |
| Conc O_2 | 0.25 | 0.10 | 0.23 | 0.01 | 0.44 | 0.26 | | | | |
| Flux NH_4^+ | -0.70 | -0.36 | -0.61 | 0.68 | -0.38 | -0.52 | 0.11 | | | |
| Flux NO_x | 0.72 | 0.80 | 0.40 | -0.47 | 0.82 | 1.00 | 0.26 | -0.48 | | |
| Flux PO_4^{3-} | -0.35 | -0.37 | -0.21 | 0.20 | -0.07 | -0.32 | 0.29 | 0.48 | -0.28 | |
| Flux Fe | -0.51 | -0.41 | -0.36 | 0.37 | -0.39 | -0.55 | 0.00 | 0.52 | -0.52 | 0.81 |

9

10

11

1 **Figure legends**

2

3 Figure 1. Nutrient fluxes between sediment and water ($\text{mmol m}^{-2} \text{d}^{-1}$, avg \pm sd).

4 Significant differences from the control treatment are marked with an asterisk.

5

6 Figure 2. A) Coupled nitrification-denitrification (D_n , white columns) and denitrification
7 based on water column nitrate (D_w , dark columns) ($\text{mmol N m}^{-2} \text{d}^{-1}$, avg \pm sd) B)

8 Denitrification potential (D_{15}) ($\text{mmol N m}^{-2} \text{d}^{-1}$, avg \pm sd). Significant differences from
9 the control treatment are marked with an asterisk.

10

11 Figure 3. Pore water nutrient concentrations (μM). Diamonds C, squares LOW, triangles
12 HIGH treatments.

13

14 Figure 4. P fractions ($\mu\text{mol P g}^{-1} \text{DM}$, avg \pm sd) in C and HIGH units on days 2 and 14.

15 White bars 0-2 cm, grey bars 2-4 cm, black bars 4-6 cm.