Fluid avalanches

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Simulation of avalanches: modelling and numerics - Sevilla workshop



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- Context: the dam-break problem
- Laboratory insight: flow visualization
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In engineering, dam break: sudden release of water



Teton dambreak (Idaho, 1976)

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Scientific issues





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Taum Sauk dam break (Missouri, Dec. 2005) intense erosion of the bed (down to the bed rock) and sediment transport





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Outburst flood: Lake Nostetuko (British Columbia, Canada) July 1983





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Release of a fixed volume of fluid



Questions:

- Front position over time $x_f(t)$?
- Flow depth profile h(x,t) ?
- Velocity profile within the flow (far from the sidewall) ?
- Further questions: stability, slip, influence of surface tension, etc.



Navier-Stokes equations

Dimensionless form

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 $\begin{aligned} \frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} &= 0, \\ \epsilon \operatorname{Re} \frac{\mathrm{d} u}{\mathrm{d} t} &= \phi \cos \theta \left(\tan \theta - \epsilon \frac{\partial p}{\partial x} \right) + \epsilon^2 \frac{\partial^2 u}{\partial x^2} + \frac{\partial^2 u}{\partial y^2}, \\ \epsilon^2 \operatorname{Re} \frac{\mathrm{d} v}{\mathrm{d} t} &= -\phi \cos \theta \left(1 + \frac{\partial p}{\partial y} \right) + \epsilon^3 \frac{\partial^2 v}{\partial x^2} + \epsilon \frac{\partial^2 v}{\partial y^2}, \end{aligned}$

with $\phi=\rho g H_*^2/(\mu U_*)$ a dimensionless group and $\epsilon=H_*/L_*\ll 1$ the aspect ratio.

 $Ca=\mu U_*/\gamma\gg 1$ and $Re=\rho U_*H_*/\mu\ll 1$: capillary and Reynolds numbers.

 L_* and H_* selected so that $L_*H_*=\tilde{V},$ viz, $L_*=\sqrt{\tilde{V}/\epsilon}$ and $H_*=\sqrt{\epsilon\tilde{V}}$

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Dominant balance:

- Diffusive regime. Balance between the pressure and shear stress gradients : $U_* = \rho g H_*^3 / (3 \mu L_*)$ and $\phi = 3/\epsilon$
- Advection diffusion regime. Balance between the body force and shear stress gradient + pressure gradient within the leading edge: $\epsilon = \tan \theta$, $U_* = \rho g H_*^2 \sin \theta / (3\mu)$, and $\phi = 3 / \sin \theta$
- Steep slope regime. Increasing role of the body force: $\epsilon = \tan^2 \theta$, $U_* = \rho g H_*^2 \sin \theta / (3\mu)$, and $\phi = 3 / \sin \theta$



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Diffusive regime observed for $Ca \to \infty$, Re = O(1), and $\theta \ll 1$. Dimensionless governing equation for $\theta > 0$

$$\frac{\partial h}{\partial t} + \frac{\partial h^3}{\partial x} = \frac{\partial}{\partial x} \left(h^3 \frac{\partial h}{\partial x} \right).$$

Dimensionless governing equation for $\theta=0$

$$\frac{\partial h}{\partial t} = \frac{\partial}{\partial x} \left(h^3 \frac{\partial h}{\partial x} \right).$$

No analytical solution (available), but asymptotic solutions at short or long times t with the following change of variable

$$h(x,t) = t^{-n}H(\xi,t)$$

- short-time solution: n = 1/5 (Nakaya, 1974 ; Huppert, 1982) ;
- long-time solution: n = 1/3 (Huppert, 1982 ; Lister, 1982).



Diffusive regime at $t\ll 1$

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Huppert's (or Barrenblatt-Pattle's) solution

$$h(x,t) = t^{-1/5} \left(\frac{3}{10}(\xi_f^2 - \xi^2)\right)^{1/3} \operatorname{avec} \xi_f = V_0^{3/5} \left(\frac{\sqrt[3]{\frac{3}{10}}\sqrt{\pi}\Gamma\left(\frac{1}{3}\right)}{5\Gamma\left(\frac{5}{6}\right)}\right)^{-5/3} \left(\frac{1}{5}\right)^{-5/3} \left(\frac{1}{5}\right$$



Comparison between the numerical solution and Huppert's self-similar solution at t = 1, 2, 4, 8, 16, 32, 64, 128, 256, 512, and 1024

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Huppert's solution (+ higher-order terms Ancey, JFM 2009)

$$h(x,t) = t^{-1/3} \left(\sqrt{\frac{1}{3} \frac{x}{t^{1/3}}} + K_0 \left((\xi_f - xt^{1/3}) t^{2/3} \right) - \sqrt{\frac{\xi_f}{3}} \right).$$

with the position of the front given by

$$x_f = \xi_f t^{1/3} + \left(\log 2 - \frac{1}{2}\right) \sqrt{\frac{\xi_f}{3}} t^{-1/3},$$

$$\xi_f = \left(\frac{3\sqrt{3}}{2}V\right)^{2/3}$$

This solution requires a boundary-layer treatment at the front as the diffusive effects (pressure gradient) prevail over the advection term.



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Comparison of flow depth profiles: numerical solution (solid line) and asymptotic solution (dashed line), with diffusion included (a) or not (b), for slope $\theta = 6^{\circ}$, at times t = 1, 2, 4, 8, 16, 32, 64, 128, and 256.

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Comparison of flow depth profiles for $\theta = 0^{\circ}$. For (b), we show the self-similar solution and the experimental trend $(h/t^{-1/5})^3 = \frac{9}{10}(1.3 - \xi)$. Fluid: glucose ($\mu = 345$ Pa s)



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Front position over time: experimental data (solid line) and theory (dashed line)



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Front position over time for slopes ranging from 6° to 24°



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Flow depth profiles for slopes ranging from 6° to 24° and different slopes



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Velocity profiles (u and v) in glycerol ($\mu = 1.11$ Pa s) for a 6° slope



Viscoplastic material

Simple shear constitutive law

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$$\mu \dot{\gamma}^n = \begin{cases} \tau - \tau_c & \text{ for } \tau > \tau_c, \\ 0 & \text{ for } \tau \le \tau_c, \end{cases}$$

Velocity profile for $y \leq Y_0$ or $y > Y_0$ (Liu & Mei, 1990)

$$u(x,y,t) = \begin{cases} \frac{n}{n+1} K \left(Y_0^{1+1/n} - (Y_0 - y)^{1+1/n} \right) \left(1 - \cot \theta \frac{\partial h}{\partial x} \right) \\ \frac{n}{n+1} K \left(1 - \cot \theta \frac{\partial h}{\partial x} \right) Y_0^{1+1/n}, \end{cases}$$

where $Y_0 = \max(0, h - \tau_c/(\rho g \cos \theta (\tan \theta - \partial_x h)))$ denotes the position of the yield surface and $K = \rho g \sin \theta/\mu$.



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Comparison of the velocity profiles for $\theta = 25^{\circ}$ at different distances Δx to the front. Fluid: Carbopol ultrez 10 ($\mu = 26$ Pa sⁿ, n = 0.33, $\tau_c = 33$ Pa)



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Comparison of the velocity profiles for $\theta = 15^{\circ}$ at different distances to the front. Fluid: Carbopol ultrez 10 ($\mu = 26$ Pa sⁿ, n = 0.33, $\tau_c = 33$ Pa)



Flow models

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Dimensionless depth-averaged equations (e.g. Craster & Mater, 2009)

$$\frac{\partial h}{\partial t} + \frac{\partial h\bar{u}}{\partial x} = 0,$$

$$\epsilon Re\left(\frac{\partial h\bar{u}}{\partial t} + \beta \frac{\partial h\bar{u}^2}{\partial x}\right) + \epsilon \cot\theta h \frac{\partial h}{\partial x} = h - \tau_b + \frac{\epsilon^3}{Ca}h\frac{\partial^3 h}{\partial x^3}.$$

Several simplifications developed:

- kinematic wave model (Huang & Garcia, 1994): balance between the driving and 'viscous' forces;
- diffusive wave model: balance between the driving and 'viscous' forces + pressure gradient;
- Saint-Venant model: in the limit of $Ca \to \infty$ and $\epsilon \ll 1$, with a closure equation for τ_b .



Kinematic wave model

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A simple nonlinear advection equation (hyperbolic)

 $\frac{\partial h}{\partial t} + \frac{\partial h\bar{u}}{\partial x} = 0$

Analytical solutions (using the method of characteristics) in an implicit form. For a 25° slope





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A nonlinear advection-diffusion equation (parabolic)

$$\frac{\partial h}{\partial t} + nK\frac{\partial}{\partial x}\left[\left(\tan\theta - \frac{\partial h}{\partial x}\right)^{1/n}\frac{h(1+n) + nh_c}{(n+1)(2n+1)}Y_0^{1+1/n}\right] = 0$$

No analytical solution. For a 25° slope





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Hyperbolic partial differential equations

$$\frac{\partial h}{\partial t} + \frac{\partial h\bar{u}}{\partial x} = 0,$$
$$\frac{\partial h\bar{u}}{\partial t} + \frac{\partial h\bar{u}^2}{\partial x} + gh\cos\theta\frac{\partial h}{\partial x} = gh\sin\theta - \frac{\tau_b}{\rho},$$

Coussot's closure equation:

$$\tau_b = \tau_c \left(1 + 1.93 G^{3/10} \right) \text{ with } G = \left(\frac{\mu}{\tau_c} \right)^3 \frac{\bar{u}}{h}$$



For a 25° slope



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A density-matched suspension of particles within a Newtonian carrier fluid is assumed to be quasi-Newtonian:

• effective viscosity given by empirical laws, e.g.,

$$\eta(\phi) = \frac{\mu(\phi)}{\mu_f} = \left(1 - \frac{\phi}{\phi_m}\right)^{-\beta}$$

 ϕ_m the maximum concentration and β a constant : $\beta = \frac{5}{2}\phi_m$ or $\beta = 2$ (Krieger & Dougherty 1959)

occurrence of normal stress effects (Zarraga *et al.* JOR, 2001; Boyer et al. JFM 2001; Couturier *et al.* JFM 2011)

Problem: particle migration occurs even for $\Delta \rho = 0$ (effect exacerbated when sedimentation or creaming occurs, depending on $\Delta \rho > 0$ or $\Delta \rho < 0$), so this results in a nonhomogeneous spatial distribution of the particles, thus viscosity.



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Migration: stratification of particles as a result of shear. Two approaches:

Phenomenological model developed by Leighton & Acrivos (JFM 1987)

$$oldsymbol{j} \propto -\phi^2
abla \dot{\gamma}$$
 avec $oldsymbol{j} = \phi(oldsymbol{u}^p - oldsymbol{u})$

the particle flux relative to the bulk velocity

 Microstructural approach by Nott & Brady (JFM 1994) and Morris & Boulay (JOR 1999)

$$oldsymbol{j} \propto -
abla \cdot oldsymbol{\Sigma}^p$$

The theoretical underpinning is still disputed (Lhuillier PoF 2009; Nott et al. PoF 2011).

Respective merits subject of fierce debate... with no winner



Shear-induced migration

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Phillips et al.'s model (PoF 1992) in a dimensionless form

$$\boldsymbol{j} = -\phi K_c \frac{\epsilon_a^2}{\epsilon} \nabla(\phi \dot{\gamma}) - K_\mu \dot{\gamma} \phi^2 \frac{\epsilon_a^2}{\epsilon} \frac{\mathrm{d} \ln \eta}{\mathrm{d} \phi} \nabla \phi,$$

Two processes at play:

- diffusion of particles resulting from the anisotropy in the probability of encounter between two particles,
- Fick-like diffusion

Nonlinear advection diffusion equation for ϕ

$$\frac{\partial \phi}{\partial t} + u \frac{\partial \phi}{\partial x} + v \frac{\partial \phi}{\partial y} = -\nabla \cdot \boldsymbol{j}.$$

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With $\beta=2$ and $\alpha=3/2$ (otherwise the solution is implicit), one gets the quasi-explicit solution

$$\phi = \frac{\phi_w h}{\phi_m h - (\phi_m - \phi_w)y}$$
 with $\frac{\phi_w}{\phi_m - \phi_w} \log \frac{\phi_m}{\phi_w} = \overline{\phi}$

By integrating the conservation of momentum

$$\dot{\gamma} = \frac{\bar{\eta}}{\eta(\phi)}(h-y),$$

we determine velocity profiles numerically. They take the form

$$u = \kappa \left(h^n - (h - y)^n \right)$$

n = 2 for a Newtonian fluid. Another index to characterise the deviation of the computed velocity profile from the Newtonian profile (m = 2/3)

$$m = \frac{\int_0^h u(y,t) \mathrm{d}y}{hu(h,t)} = \frac{\bar{u}}{u(h,t)},$$

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Variation of n with the particle concentration



Two parameter sets: $\beta = 2$ and $\alpha = 3/2$ (solid line) ; $\beta = 2$ and $\alpha = 1.042\phi + 0.1142$ (dashed line) using the model proposed by Tetlow *et al.* (JOR 1998)



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$$\frac{\partial \phi}{\partial t} = K_c \bar{\eta} \frac{\epsilon_a^2}{\epsilon} \frac{\partial}{\partial y} \left(\phi \frac{\partial}{\partial y} \left(\frac{\phi}{\eta(\phi)} (h - y) \right) + \alpha \frac{\phi^2}{\eta(\phi)} (h - y) \frac{\mathrm{d} \ln \eta}{\mathrm{d} \phi} \frac{\partial \phi}{\partial y} \right).$$

Steady state is reached at time

Using the assumption $\partial_x \phi = 0$

$$t_c \sim \frac{\epsilon}{\epsilon_a^2} = \frac{H_*^3}{a^2 L_*}$$

$$t_{ss} = 2t_c \left(1 - \frac{\phi}{\phi_m}\right)^{-1/3}$$



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Variation of n (and m) at short times



Trend $n = 2.2t^{1/4}$. Computations done for $\bar{\phi} = 52\%$, $K_c \bar{\eta} \epsilon_a^2 / \epsilon = 1$, $\beta = 2$, and $\alpha = 3/2$. Computed steady state time: $t_{ss} = 3.56$.



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Variation of $x_f(t)$ (dimensionless)

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For a 25 $^\circ$ slope



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For a 25° slope, but $\phi=0.45$



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For a 25° slope, but $\phi=0.57$



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Variation of n with time (dimensionless)



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For $\phi > 0.575$, behaviour is complicated, with three phases observed:

- macro-viscous regime at short times: $x_f \propto t^{1/3}$, parabolic profile of u,
- fracture regime: wavy free surface, fracture, and en-masse flow,
- *plastic regime*: intermittent motion (stick-slip).





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For 25° slope, $\phi=0.595$



Stick-slip regime

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Variation of the flow depth and the pressure 'head' $h_p = p/(\varrho g \cos \theta)$





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 $\phi = 0.56$

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The dam-break for different rheologies. Our main findings:

- For Newtonian fluids: fairly good agreement between theory/experiment.
- For viscoplastic fluids: the simplest models perform better than more elaborate models such as Saint-Venant.
- For granular suspensions: we observed *macro-viscous* behaviour for concentrations as high as $\phi = 0.57$, then for higher concentrations, complicated behaviour, including stick-slip motion (resulting from a pore-pressure diffusion mechanism that is still poorly understood).



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References

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