

1 2 3	CO2 seawater acidification by CCS-simulated leakage: Kinetic modelling of Zn, Pb, Cd, Ni, Cr, Cu and As release from contaminated estuarine sediment using pH-static leaching tests
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12	Highlights
13	• Kinetic element release in a pH-static leaching test using CO <sub>2</sub> for acidification
14	• The model fits experimental release including the influence of Fe and other ions
15	• There is a leaching simulation over time under acidic scenarios up to pH=6
16 17	• CO <sub>2</sub> and HNO <sub>3</sub> acidifications lead to different kinetics during element release
18 10	• The generalized model proposed here is also useful for areas conteminated by
19	• The generalised model proposed here is also useful for areas containinated by
20	iron
21	
22	Keywords: Kinetic modelling; CO2 acidification; contaminant release; sediment; pH-
23	static leaching; Fe influence
24	
25	Abstract
26	A modified pH-dependent leaching test with continuous pH control that employed CO <sub>2</sub>
27	to acidify a seawater-sediment mixture is used to address Zn, Pb, Cd, Ni, Cr, Cu and As
28	release from contaminated estuarine sediments under the influence of acidification

processes. Long-term (480 h) leaching experiments at pH values of 7.0, 6.5 and 6.0 are performed. The different evolutionary patterns of the redox potential and Fe release at pH=6 with respect to the other pH values shows the need to assess the influence of the initial Fe content in seawater upon elemental release.

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Hence, assays at pH=6.0 are conducted using natural seawater with Fe concentrations between 9.02 and 153  $\mu$ g/L. A set of in-series reactions for trace elements, Fe and other ions associated with Fe is proposed to model a Fe/multi-ion-dependent mechanism for trace metal release. The maximum concentration of each contaminant that can be released from the sediment and the kinetic parameters of the proposed model are completed for the studied pH values, for good consistency between the experimental and simulated mobilisation of each studied element.

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#### 43 1. INTRODUCTION

Carbon capture and storage (CCS) in geological formations is one of the most promising strategies for curbing global climate change (IPCC, 2014). One of the primary risks in the case of ocean storage technology is the potential direct leakage of  $CO_2$  gas or  $CO_2$  dissolved in seawater, which provokes a decrease in the pH of the medium. This acidification might mobilise contaminants from marine and estuarine sediments (Rodríguez-Romero et al., 2014; Ardelan et al., 2009).

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Laboratory leaching tests allow for the assessment of contaminant releases under different scenarios and conditions. Column leaching tests (Bateman et al., 2005; Frye et al., 2012; Lawter et al., 2016; Payán et al., 2012a) and batch or semi-batch leaching

tests (Ardelan and Steinnes, 2010; Ardelan et al., 2009; de Orte et al., 2014; Kirsch et al., 2014; Little and Jackson, 2010; Lu et al., 2010; Payán et al., 2012b) have been performed to assess elemental mobilisation from different types of matrices such as sediments, sandstones or rocks when the addition of CO<sub>2</sub> decreases the pH of the medium.

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The release of experimental contaminants from sediments and rocks by using different leaching tests to mimic the effects of potential CO<sub>2</sub> leakages has been studied before. Equilibrium conditions have been modelled using geochemical software such as PHREEQC or Visual MINTEQ (de Orte et al., 2014a; Martín-Torre et al., 2015a), whereas simulations for the period before equilibrium is achieved have been performed using reactive transport models as performed with the TOUGHREACT code (Zheng et al., 2016; Zheng et al., 2009) or PHREEQC (Cahill and Jakobsen, 2015).

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68 Reactive transport models combine mineral dissolution/precipitation, aqueous 69 complexation, acid-base, redox, cation exchange and surface complexation processes. Therefore, the use of this type of models for simulating contaminant mobilisation 70 implies that the modelled has knowledge of the different characteristics of the solid 71 72 matrix under study, such as crystalline phases, that are usually not possible to know in the case of sediments because of their great complexity and low contaminant 73 concentrations. Moreover, the need to pre-treat the sediment sample complicates the 74 75 determination of element speciation in a sensitive manner.

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The pH is one of the variables that most influences the mobility and availability of inorganic contaminants from solid matrices (Coz et al., 2007), and thus, pH-dependent

leaching tests seem to be an appropriate assay for assessing its influence. Standard or
modified pH-dependent leaching tests with continuous pH control such as CEN/TS
14997 (2015) have been widely used to evaluate the release of contaminants from soil
and sediments (Cappuyns and Swennen, 2005; Cappuyns et al 2004a,b; Centioli et al.,
2008; Horckmans et al., 2007; Shtiza et al., 2009; Van Herreweghe et al., 2002).

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To study the release of contaminants that occurs when environmental conditions in 85 marine and estuarine media vary, a modified pH-dependent leaching test with 86 continuous pH control has been previously performed to assess the contaminants 87 released from contaminated sediment when a total mixture of sediment-seawater is 88 acidified by nitric acid (HNO<sub>3</sub>) (Martín-Torre et al., 2015b). Based on these 89 experimental results, a kinetic mathematical model has been proposed to obtain the 90 91 generalised kinetic constants of the studied elemental release. Owing to the 92 experimental difficulties involved in characterising the sediment and the many chemical 93 reactions that occurred during the process, global phenomena rather than specific 94 reactions have been considered.

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It is known that inorganic contaminants, primarily metals and metalloids, are largely affected by acidification processes (Basallote et al., 2014), and that their effects on marine organisms are higher when using  $CO_2$  gas than mineral acids (Ishimatsu et al., 2004; Kikkawa et al., 2004). Hence, the performance of a pH-dependent leaching test with  $CO_2$  gas for acidification and the subsequent assessment and modelling of the triggered contaminant mobilisation from sediment would be useful for a CCS technology impact assessment.

The aim of the present work is to obtain, analyse and simulate the experimental release 104 105 of Zn, Pb, Cd, Ni, Cr, Cu and As from contaminated marine sediments during the pH-106 static leaching test with continuous pH control by using seawater as a leaching liquid 107 and CO<sub>2</sub> gas to acidify the medium. The kinetic behaviour of Zn, Pb, Cd, Ni, Cr, Cu 108 and As is studied by using long-term experiments lasting 480 hours at pH values of 7.0, 6.5 and 6.0. The influence of the Fe content of the seawater on the redox potential 109 evolution and elemental release is highlighted at pH=6.0 because of the different 110 111 evolution of these parameters with respect to the experimental data obtained at pH values of 7.0 and 6.5. 112

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Starting with the mathematical model by Martín-Torre et al. (2015b) in which HNO<sub>3</sub> acidification was initially proposed, a modified kinetic model is proposed here to explain the behaviour of long-term experimental data at different pH values and different Fe concentrations in seawater. The generalised kinetic expression and its kinetic parameters are obtained to fit the experimental and modelled release of the elements.

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## 122 2. MATERIALS AND METHODS

The selected estuarine sediment samples were collected in the Suances estuary (Cantabrian region, northern Spain), which is a representative area for a possible CCS (BOE, 2008). As explained in detail in Martín-Torre et al. (2015b), surface sediment (0-5 cm) and its initial water content was collected using a plastic paddle, and it was sieved through a 2 mm plastic mesh to remove the gravel fraction. Afterwards, the sediment was homogenised and frozen in plastic bags until use. An X-ray diffraction analysis (Siemens D5000 diffractometer) using Cu Kα radiation
and operating at 30 mA and 50 kV was used to determine the crystalline phases of the
selected sediment (Payán et al. 2012b). The total sediment content was determined by
an external laboratory (Activation Laboratories, Canada). The As and Cr contents were
measured by Instrumental Neutron Activation Analysis (INAA) and the Cd, Cu, Ni, Pb
and Zn contents were analysed by Total Digestion-Inductively Coupled Plasma/ Optical

135 Emission Spectrometry (TD-ICP/OES) method.

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137 The results of the sediment characterisation are explained in more detail in Martín-Torre

et al. (2015b), but they are described briefly in Table 1.

Table 1. Principal crystalline phases, redox potential and trace element concentrations of the studied
 sediment

Principal crystalline phases	Trace element	Content (mg/kg)					
Quartz	Zn	$5220 \pm 140$					
Aluminium oxide	Pb	$564 \pm 2.22$					
Calcite	Cd	$12.6\pm0.732$					
Dolomite	Ni	$36 \pm 1.86$					
	Cu	$48\pm3.43$					
	Cr	$72\pm5.31$					
As $59 \pm 1.39$							
Redox potential (mV): -150±37							

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The entire Cantabrian region and the Suances estuary in particular was an important mining area, with a considerable number of iron ore mines and abandoned landfills containing solid mining wastes. Therefore, the presence of this metal in rivers and effluents that lead in the estuary and its accumulation in sediment is appreciable. Therefore, the concentration of iron in the area from which the seawater is taken should be influenced by tides and other atmospheric conditions. To prevent large variations in the concentration determinations, the seawater was chemically analysed by using

149	inductively coupled plasma-mass spectrometry (ICP-MS) equipment in helium collision
150	mode according to the same procedure that was used for the sample leachates, prior to
151	use. Table 2 lists the average elemental concentrations in the seawater that was used in
152	the leaching assays at pH=7.0, 6.5 and 6.0. Seawater samples with iron concentrations
153	([Fe] <sub>seawater</sub> ) of 9.02, 46.1 and 153 $\mu$ g/L were sampled at the same place that the
154	sediment was taken from but under different natural conditions of tide, meteorology and
155	season.

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Table 2. Concentration of the studied elements and redox potential in the seawater leachant

Element	Content (µg/L)
As	2.58±0.981
Cd	$0.204 \pm 0.0797$
Cr	0.920±0.631
Cu	$2.46{\pm}1.49$
Ni	1.21±1.37
Pb	$0.929 \pm 0.727$
Zn	11.6±8.21
Fe	6.45±2.91
Redox potential (mV)	$160 \pm 16.0$

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The experimental equipment consisted of a glass-made 2-L jacketed vessel, a 160 temperature controller (Polyscience) and a system for maintaining a constant pH (Fig. 161 162 1). In this study, a pH controller (AT Control systems) with an electrode that was 163 suitable for samples containing suspended solids was used; it was accurate to 0.01 pH units and was calibrated against standard solutions. This device was used to monitor the 164 pH and inject pure CO<sub>2</sub> bubbles into the suspension as needed, with a permitted 165 hysteresis of 0.1 pH units. When the suspension reached the required pH, the gas supply 166 167 was automatically stopped.



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Fig. 1. Experimental equipment used in the pH-static leaching test with continuous pH control.

171 Seawater and sediment were placed in a reactor with an L/S ratio of 10, considering the moisture of the sediment (51.26%). To make a homogenised mixture at the beginning of 172 173 the assay, the reactor was shaken for 15 minutes at the natural pH before starting the CO<sub>2</sub> addition (CO<sub>2</sub> supplied by AL Air Liquide España, S.A., Zamudio, Vizcaya, País 174 Vasco, Spain). In this work, the studied pH values are 7.0, 6.5 and 6.0. It was not 175 possible to achieve lower pH values because of the high buffering capacity of the 176 177 sediment-seawater system (Martín-Torre et al., 2015a). Additionally, an assay without pH control was conducted to observe the pH evolution over time. 178

The pH-static leaching tests conducted in this work lasted almost 480 hours to achieve near-equilibrium conditions, which were not obtained over shorter assay durations (Martín-Torre et al., 2015b). Nevertheless, the assay at pH=7.0 lasts 264 hours, when the consumption of the buffer capacity of the mixture makes the pH uncontrollable. The experiments were performed in duplicate. Samples at 0, 0.5 h, 1 h, 3 h, 6 h, 12 h, 24 h,

48 h, 72 h, 96 h and afterwards for every 48 hours were taken using a syringe and
without stopping the mixing. The redox potential (Eh) of the leachate was measured
using a Basic 20 pH metre (Crison) with a special electrode for samples containing
suspended solids; each sample was filtered through a 0.45-µm pore size nitrocellulose
filtration membrane and acidified to determine the concentrations of Fe, Zn, Pb, Cd, Ni,
Cr, Cu and As. The corresponding amount of solids was put back in the medium to
maintain a constant L/S ratio during the whole assay.

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The elemental concentrations were determined by an external laboratory (Mass 193 Spectrometry Unit, University of Oviedo, Spain) using Agilent 7500CE inductively 194 coupled plasma-mass spectrometry (ICP-MS) equipment in helium-collision mode 195 (Agilent Technologies, California, EEUU). The samples were diluted (1:20) with HNO<sub>3</sub> 196 197 (VWR International, Fontenay-sous-Bois, France) 1% prior to analysis, and Rh was 198 added as an internal standard to correct for the eventual signal drift during analysis. The 199 metal concentrations were calculated using external calibration with internal standard 200 correction. The certified reference material NASS-5 (Seawater reference material for 201 trace metals, NRCC (Ontario, Canada)) was spiked with the elements at two different concentration levels of 1 and 10 µg/L. The certified reference material and the two 202 203 spiked concentrations were measured every 6 samples as a quality control. The 204 detection limits for the elements under study (Fe, Zn, Pb, Cd, Ni, Cr, Cu and As) were 0.79; 0.75; 0.02; 0.06; 0.23; 0.03; 0.21 and 0.16 µg/L, respectively. Prior to the 205 206 experiments, all the sampling and laboratory material was precleaned, acid-washed (10% HNO<sub>3</sub>), and rinsed with Milli-Q water (Direct-Q 5 UV, Merck Millipore, USA). 207

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The modelling of this study and the estimation of the corresponding parameters are completed using Aspen Custom Modeler software (Bedford, Massachusetts, USA) which solves rigorous models and simultaneously estimates parameters. The adjustment of the model parameters was performed using an NL2SOL algorithm for the leastsquare minimization of the deviation between the experimental and theoretical data.

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#### 215 **3. RESULTS**

## 216 3.1. Evolution of redox potential and Fe release at pH=7.0, 6.5 and 6.0

The evolution of the redox potential ( $E_h$ ) over time is followed in all the assays. Similar to the findings of Cappuyns and Swennen (2005), this evolution is as useful as the exact values when considering that the measured value is a mixed redox potential. It cannot be a true equilibrium potential because of the different redox couples in the sediment and the slow kinetics of redox reactions (Sigg, 2000).

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To compare the evolution of the redox potential at the different pH values under study better, Fig. 2 shows the evolution of the redox variation ( $E_h$ -  $E_{h,0}$ ) over time. During the first 30-60 minutes of the assay there is a rapid increase in the  $E_h$  value, possibly because of the dissolution of major ions from the sediment. After this initial increase, a common trend is observed for pH=7.0 and 6.5. In these cases, the  $E_h$  value increases from t=1 h until t=50-72 h. Afterwards, the value of this parameter remains almost constant.

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At pH=6.0 and the Fe concentration in the selected seawater ([Fe]<sub>seawater</sub>) of 9.02  $\mu$ g/L, the E<sub>h</sub> value shows an irregular and different time behaviour it decreases from 1 to 12 hours. Afterwards, it remains almost constant until 48 h, when a rapid increase begins.

After a maximum  $E_h$  of 96 h, a decrease occurs. There is a subsequent increase until 192 h, when it remains almost constant until the end of the assay.



Fig. 2. Evolution of redox variation (mV) over time in pH-static leaching tests.  $E_{h,0}$  and  $E_h$  are the redox potentials at t=0 h and t=t, respectively. • pH=7; • pH=6.5; • pH=6.0, [Fe]<sub>seawater</sub>=9.02 µg/L; • pH=6.0, [Fe]<sub>seawater</sub>=46.1 µg/L; and  $\Box$  pH=6.0, [Fe]<sub>seawater</sub>=153 µg/L. Connecting lines were added for clarity.

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241 As shown in Fig. 3 a and b, at pH values of 7.0 and 6.5, the dissolved Fe concentration 242 increases during the first three hours because of the mixing of the wet sediment with the seawater and possibly because of the oxidation and later release of different iron 243 compounds. Afterwards, there is a decrease in the dissolved Fe concentration as a 244 consequence of the Fe(II) to Fe(III) oxidation and the subsequent precipitation of 245 Fe(III), which is much less soluble than Fe(II). When Fe(III) precipitates, it might form 246 various compounds such as Fe oxyhydroxides (Appelo et al., 1999; Wang et al., 2016). 247 A minimum value of approximately 10 µg/L, which remains constant until the end of 248 249 the assay, is achieved at t=24 h when the pH is 7.0 and at 96 hours at pH=6.5.

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The release of Fe at pH=6.0 and  $[Fe]_{seawater}$ = 9.02 µg/L (Fig. 3 c) is higher than it is at more neutral pH values because of its higher solubility at lower pH values (Johnston et al., 2016). Moreover, the experimental results at pH=6.0 indicate that high concentrations of Fe remain dissolved until 48 hours of the assay has passed, most

likely because of the slower oxidation of Fe(II) to Fe(III) at pH=6.0. This finding is 255 completely consistent with previous studies showing that the rate constant of Fe(II) 256 oxidation increases at higher pH values as well as in the presence of HCO<sub>3</sub><sup>-</sup> ions 257 (Millero and Izaguirre, 1989; Millero et al., 1987). The concentration of HCO<sub>3</sub><sup>-</sup> at pH 258 values of 7.0 and 6.5 (0.90 and 0.82 molar fraction, respectively) is higher than it is at 259 pH=6.0 (0.6 molar fraction) (Payán et al., 2013). In coinciding with the high increase of 260 the E<sub>h</sub> value from 48 to 96 h, the dissolved Fe concentration decreases because of its 261 262 oxidation to Fe(III) and its subsequent precipitation. A low Fe release, from 40 to 60  $\mu$ g/L, is maintained from t=192 h to the end of the assay. This final dissolved 263 concentration is higher at pH=6.0 than at more neutral pH values, likely because of the 264 265 increase of in the Fe solubility at lower pH values.

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The different behaviour of both parameters ( $E_h$  and Fe release) at pH=6.0 leads us to assess the influence of the initial Fe concentration in the seawater on the elemental release and to consider it as an important variable in leaching tests at pH=6.0.



Fig. 3. Fe release (µg/L) over time in pH-static leaching tests: ● pH=7; ■ pH=6.5; ◇ pH=6.0,
[Fe]<sub>seawater</sub>=9.02 µg/L; ○ pH=6.0, [Fe]<sub>seawater</sub>=46.1 µg/L; and □ pH=6.0, [Fe]<sub>seawater</sub>=153 µg/L. Connecting
lines were added for clarity.

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3.2. Evolution of redox potential and Fe release at pH=6.0 using seawater with
different Fe content
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In accounting for the fact that the Fe concentration of the selected seawater from the studied estuarine area fluctuates because of local tides and atmospheric conditions, assays at pH=6.0 are conducted using seawater with higher Fe concentrations, namely 46.1 and 153  $\mu$ g/L, and lower E<sub>h</sub> values of 61 and 116 mV, respectively. The concentrations of the other studied elements do not present as much variation, and they fall within the concentration range shown in Table 2.

In Fig. 2, the value of  $E_h$  increases in the assays conducted at pH=6.0 when using 285 286 seawater with higher concentrations of Fe are also shown. During the first 24 hours of the assay, the evolution of the Eh of the three assays that were performed at pH=6.0 287 288 follows the same pattern: the increase during the first half hour is followed by a slight decrease and a plateau that lasts until 24 h when the [Fe]<sub>seawater</sub>=46.1 µg/L or 48 hours in 289 the other two cases. This difference might be more strongly influenced by the 290 291 concentration of the major ions that are present in the seawater and the heterogeneity of 292 the sediment rather than by the Fe content of the seawater. When the plateau ends, a rapid increase occurs until 72-96 h, depending on the experiment. At t=72 h and 293  $[Fe]_{seawater} = 46.1 \,\mu g/L$ , the E<sub>h</sub> begins to decrease and a plateau is observed until t=168 294 h. At the lowest and highest [Fe]seawater, the decrease begins at 96 h and it lasts from 96 295 to 144 h in the case of  $[Fe]_{seawater}=153 \mu g/L$ . Afterwards, there is an increase, which is 296 297 less pronounced at higher Fe concentrations in the seawater, which ends in a constant 298 value that is maintained until the end of the assay. The time at which the redox potential 299 begins to remain constant seems to depend on the Fe content, because at the higher 300 seawater Fe concentration, a constant E<sub>h</sub> value is achieved later (192 h, 216 h and 432 301 h).

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In all the assays conducted at pH=6, high Fe concentrations are in dissolution during the first 24 or 48 hours (Fig. 3 c, d and e). Afterwards, the precipitation of Fe(III) decreases the dissolved Fe concentration, and a minimum release is observed at 72 h ([Fe]<sub>seawater</sub> = 46.1 and 153  $\mu$ g/L) or 96 h when [Fe]<sub>seawater</sub>=9.02  $\mu$ g/L. In assays with [Fe]<sub>seawater</sub> values of 46.1 and 153  $\mu$ g/L, an increase in the Fe mobilisation is observed at 168 and 288 h, respectively; it is more pronounced at the highest Fe concentration in the initial seawater. Once the dissolved Fe concentration decreases after the second peak, likely because of the Fe(II) oxidation and later precipitation of Fe(III), similar Fe release values are obtained at the end of the three assays that were performed at pH=6.0  $(36.5\pm8.5 \,\mu\text{g/L})$ .

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The fact that there is a second dissolved Fe concentration peak suggests a mobilisation that could be induced by different phenomena such as the destabilisation of Fe(III) oxyhydroxides or ionic competition and displacement reactions.

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The destabilisation of Fe(III) oxyhydroxides could be caused by their reductive 318 dissolution (Root et al., 2007). In this case, the abiotic reactions promoted by organic 319 compounds present in the medium and H<sub>2</sub>S might provoke this reductive dissolution 320 (Hering and Stumm, 1990; Schwertmann, 1991; Thamdrup, 2000). Luther III et al. 321 322 (1992) identify acidification as one of the different impacts that cause iron oxyhydroxides to become unstable. H<sub>2</sub>S could be formed when Fe-sulphides are 323 dissolved, and S<sup>2-</sup> reacts with protons from the acidification (Cappuyns and Swennen, 324 325 2005). In the presence of sulphides, destabilised iron oxyhydroxides are converted into iron sulphides (Salomons, 1995; Luther III et al., 1992) that might be oxidised and 326 released into the medium. 327

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Iron monosulphides are partially soluble in water, with higher solubility at lower pH values, whereas other metal monosulphides are less soluble than Fe monosulphides. Hence, the displacement of Fe from monosulphides occurred as well as the inclusion of other divalent metals in iron monosulphides, as shown in Eqs. 1 and 2 respectively, could be the cause of an increase in the concentration of dissolved Fe<sup>2+</sup> (Di Toro et al., 1990; Morse and Arakaki, 1993; Wong et al., 2013).

$$Fe_2S + Me^{2+} \rightarrow Fe(Me)S + Fe^{2+}$$
(1)

$$FeS + Me^{2+} \rightarrow MeS + Fe^{2+}$$
(2)

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337 Ionic competition is a phenomenon that might be occurring constantly because of the huge number of ions present in the medium, because seawater was used as well as the 338 339 acidification caused by CO<sub>2</sub> gas. In contrast to the assays at pH=6.0 with HNO<sub>3</sub> by Martín-Torre et al. (2015b), in this study the displacement reactions might be an 340 341 important phenomenon because CO<sub>2</sub> gas was used to acidify the mixture. The use of this gas instead of HNO<sub>3</sub> to acidify- the mixture affects the different equilibria present in 342 343 the suspension, and the newly formed ions could highly influence the release of the 344 studied contaminants (Tokoro et al., 2010).

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#### 346 **3.3. Trace element release**

347 The average value of the experimental results obtained for the release of Zn, Pb, Cd, Ni, 348 As, Cu and Cr from the pH-static leaching test and the error bars between both 349 replicates are shown in Fig. 4. The relative error between both replicates under the same leaching conditions is lower than 20% for any of the studied contaminants. Moreover, 350 351 the relative error of more than 85% of the experimental data for Zn, Pb, Cd, Ni and As 352 is lower than 10%. Most of the experimental data from the Cr and Cu release show errors between 10% and 20%, likely as a consequence of their low mobilisation from 353 354 the sediment in any of the studied conditions. The release of Cr is lower than  $2 \mu g/L$  and 355 does not present a clear trend over time. Something similar occurs in the case of Cu in that its release does not indicate a clear pattern and its dissolved concentration is lower 356

357 than 10  $\mu$ g/L in all of the studied scenarios with respect to the pH and Fe 358 concentrations.

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The release of Ni does not present any initial delay in any of the assays. Moreover, after an initial rapid release, a near-equilibrium condition is reached after 300 h of the assay at pH values of 7.0 and 6.5. It is necessary to extend the experiment from 432-480 h to reach this condition in assays that were conducted pH=6.0. As shown here, in comparing the three assays at this most acidic pH value, lower Fe concentrations in seawater cause a faster mobilisation of Ni as well as a higher dissolved Ni concentration under the near-equilibrium condition.

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Fig. 4. Elemental release over time in all assays. • pH=7.0; • pH=6.5; • pH=6.0, [Fe]<sub>seawater</sub>=9.02 µg/L; • 372 pH=6.0, [Fe]<sub>seawater</sub>=46.1 µg/L; and **•** pH=6.0, [Fe]<sub>seawater</sub>=153 µg/L. Connecting lines were added for 373 clarity. Error bars are also shown.

374 375	Regarding Cd, Pb and Zn mobilisation, an initial delay is observed that might be the
376	consequence of the association of these metals with sulphur and the slow oxidation
377	kinetics of these metal sulphides during the assay (Cappuyns and Swennen, 2008; Ho et
378	al., 2012). This finding is supported by the evolution of the pH over time in the assay
379	without pH control as shown in Fig. 5. Without $CO_2$ addition, the acidification of the
380	medium that was observed during the first 65 h of assay, might be caused by the $H^+$
381	released from the oxidation reactions of reduced compounds such as sulphides
382	(Cappuyns and Swennen, 2005; Eggleton and Thomas, 2004; Hwang et al., 2011). At
383	pH=6.0, higher concentrations of Fe in the seawater cause longer delays in the release
384	of these metals as a consequence of the higher solubility of Fe monosulphides than the
385	cation monosulphides, which favours the precipitation of dissolved metal ions (Di Toro
386	et al., 1990; Morse and Arakaki, 1993) at the same time that Fe <sup>2+</sup> is released (Eq. 1 and
387	2). Moreover, the delay caused by the different solubility products might also be
388	influenced by ionic competition to form compounds with other ions in the medium,
389	such as $CO_3^{2-}$ , $SO_4^{2-}$ or $Cl^-$ (Millero, 2009; Millero et al., 1995; Wong et al., 2013).







Fig. 5. pH evolution over time in the leaching test without pH control.

After an initial rapid release of As from the sediment, which became more pronounced 396 397 at more neutral pH values, there is a decrease in the As concentration. The removal of 398 this element from the solution has been widely believed to be caused by the 399 coprecipitation and adsorption provoked by the production of iron oxyhydroxides (Cappuyns et al., 2005; Omoregie et al., 2013; Wallmann et al., 1996; Zhang et al., 400 401 2007); therefore, the oxidation rate of Fe(II) highly influences the time over which the 402 oxyanion As remains in solution. Hence, at the most acidic pH value in the study, the 403 presence of dissolved As is longer than that of the other two pH values under study. Moreover, As precipitation as trace metal arsenates represents a potential mechanism 404 405 that contributes to this decrease (Vaca-Escobar et al., 2015). In the case of pH=6.0, after a minimum dissolved concentration of approximately 48-72 h, the mobilisation of As 406 407 increases. This mobilisation might be a consequence of the destabilisation of Fe(III) 408 oxyhydroxides, and the As might be adsorbed. The lower adsorbing capacity of iron 409 sulphides, which are formed with the destabilisation of Fe(III) oxyhydroxides, causes an 410 increase in the dissolved As concentration (Salomons, 1995). Moreover, sorbed As could be released after being displaced by the action of other oxyacids such as 411

phosphate, sulphate, carbonate, bicarbonate and silicate (Appelo et al., 2002; Arai et al.,
2014; Jain and Loppert, 2000; Meng et al., 2002, 2000).

414

# 415 4. MODELLING AND DISCUSSION

The pH-static leaching test used here allows for the study of elemental release from contaminated sediment as a function of pH and time. Modelling the obtained experimental results is useful for assessing the evolution of contaminant over time at the pH values of interest.

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Owing to the impossibility of performing a rigorous characterisation of all the species present in the sediment, the difficulties involved in analysing the major seawater ions mobilised from the sediment and the high number of chemical reactions that occurred during elemental release, simplified kinetic models that consider general reaction schemes to interpret contaminant release from sediments are usually proposed (Martín-Torre et al., 2015b).

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Therefore, generalised mathematical models are useful for studying contaminant release and for determining the principal processes that influence their mobilisation, without specifying all the phenomena that occurred within the sediment. The proposed mathematical model is not applied to the experimental results of Cu and Cr because their release over time does not present a clear trend, and it is very low, at lower than 10  $\mu$ g/L in all cases.

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# 435 4.1. Generalised kinetic model as applied to Zn, Pb, Cd, Ni, Cr, Cu and As

The kinetic model proposed in Martín-Torre et al. (2015b) for HNO<sub>3</sub> acidification, is used here. This model considers that the contaminant (M) is associated with an oxidised fraction (MOx) and with a reduced fraction (MRed) of the sediment that must be oxidised before the release of the element. Moreover, it includes the adsorption or precipitation of the released element through a third reaction in series. The reaction scheme and mass balances, when considering first-order reactions, are shown in Eqs. 3-7.

$$MRed(s) \xrightarrow{k_{1,i}} MOx(s) \xrightarrow{k_{2,i}} M(aq) \xrightarrow{k_{3,i}} M_{Ad}(s)$$
(3)

$$\frac{d[MRed]_i}{dt} = -k_{1,i} [MRed]_i$$
(4)

$$\frac{d[MOx]_{i}}{dt} = k_{1,i} [MRed]_{i} - k_{2,i} [MOx]_{i}$$
(5)

$$\frac{d[M]_{i}}{dt} = k_{2,i} [MOx]_{i} - k_{3,i} [M]_{i}$$
(6)

$$\frac{d[MAd]_i}{dt} = k_{3,i} [M]_i$$
(7)

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where  $[MRed]_i$ ,  $[MOx]_i$ ,  $[M]_i$  and  $[MAd]_i$  are the concentrations of element i in the reduced sediment fraction, the oxidised sediment fraction, the aqueous phase and in the adsorbed or precipitated phase, respectively. The k<sub>j,i</sub> are the rate coefficients of element i in reaction j (oxidation, release or adsorption/precipitation), and t is the reaction time.

- 448
- 449 The integral equation of the set from Eqs. 4-7 is shown in Eq. 8.

$$\frac{\text{LS}}{1000} [M]_{i} = \left(\frac{k_{1,i} k_{2,i} [M\text{Red}]_{i,0}}{(k_{2,i} - k_{1,i}) (k_{3,i} - k_{1,i})}\right) \exp(-k_{1,i}t) \\ + \left(\frac{k_{1,i} k_{2,i} [M\text{Red}]_{i,0}}{(k_{1,i} - k_{2,i}) (k_{3,i} - k_{2,i})} \\ - \frac{k_{2,i} [M\text{Ox}]_{i,0}}{(k_{2,i} - k_{3,i})}\right) \exp(-k_{2,i}t)$$

$$+ \left(\frac{\text{LS} [M]_{i,0}}{1000} + \frac{k_{2,i} [M\text{Ox}]_{i,0}}{(k_{2,i} - k_{3,i})} \\ + \frac{k_{2,i} k_{1,i} [M\text{Red}]_{i,0}}{(k_{1,i} - k_{3,i}) (k_{2,i} - k_{3,i})}\right) \exp(-k_{3,i}t)$$
(8)

where LS corresponds to the Liquid/Solid ratio of the experiment,  $[M]_{i,0}$ , which is expressed in units of  $\mu$ g/L, is the concentration of element i in the liquid at t = 0 and [MRed]<sub>i,0</sub> and [MOx]<sub>i,0</sub> are the maximum concentration as expressed in mg/kg, of element i that can be released from the reduced and oxidised fractions of the sediment respectively.

456

The resolution of the model implies the estimation of the rate coefficients and the initial concentrations of each contaminant i in the oxidised and reduced fractions ( $[MOx]_{i,0}$ and  $[MRed]_{i,0}$ , respectively) based on the experimental results. The experimental and simulated results obtained at pH values of 7.0, 6.5 and 6 at  $[Fe]_{seawater}=9.02 \mu g/L$  are shown in Fig. 6 and 7, and the initial concentrations are listed in Table 3. To better represent the elemental release, a dimensionless quotient (Eqs. 9-10) is represented over time.

$$x_{i} = \frac{[M]_{i} - [M]_{i,0}}{\frac{[MS]_{i,0}}{LS}}$$
(9)

$$[MS]_{i,0} = [MOx]_{i,0} + [MRed]_{i,0}$$
(10)

464 where  $[M]_i$  is the dissolved concentration of the element i at any time (in units of  $\mu g/L$ ), 465  $[M]_{i,0}$  is the dissolved concentration of the element i at t = 0 (in units of  $\mu g/L$ ), 466  $[MRed]_{i,0}$  and  $[MOx]_{i,0}$  are the maximum concentrations of the element i that can be 467 leached from the reduced and oxidised fractions of the sediment, respectively (in units 468 of mg/kg),  $[MS]_{i,0}$  is the maximum concentration of element i that can be released from 469 the sediment (in units of mg/kg) and LS is the Liquid/Solid ratio of the experiment (in 470 L/kg).

472 Higher concentrations are released from the reduced fraction than from the oxidised
473 fraction, except in the case of Ni, because of the initial reduced state of the sediment. In
474 the case of Cd, its concentration in the oxidised fraction is zero.

475

Table 4 lists the kinetic rate coefficients of the reactions  $(k_{j,i})$ . According to the experimental results obtained here, the third adsorption or precipitation reaction is only included in the case of As. For the pH values under study, a good fit between experimental and simulated release concentrations is indicated by the percentage variation-explained values (R<sup>2</sup>) of 99.4 (pH =7.0), 99.6 (pH = 6.5) and 98.3 (pH = 6.0 and [Fe]<sub>seawater</sub>=9.02 µg/L).

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Using the model published in Martín-Torre et al. (2015b), the simulated results for Zn, Pb, Cd, Ni and As fit more poorly with the experimental data at pH=6.0 than at the higher pH values, as shown in Fig. 7 with dashed lines for the three assays at this pH value. Among the different assays at pH=6.0, the better fit is obtained at the lowest concentration of seawater Fe (9.02  $\mu$ g/L), and it seems that the initial concentration of

488	Fe in the seawater highly influences the characteristics of the medium, modifying the
489	release behaviour of the contaminants. The experimental release of Zn, Pb and Cd at
490	pH=6.0 present a longer delay than that simulated with this model. This delay in the
491	release could be a consequence of the displacement reactions because of the higher
492	solubility of the Fe(II) compounds at the most acidic pH value under study. The studied
493	trace elements, except As, present a rapid release once Fe is removed from the aqueous
494	phase. Hence, a modified mathematical model that considers the initial Fe
495	concentration, its release and its precipitation is proposed.



Fig. 6. Trace element release  $x_i = \{[M]_i - [M]_{i,0}\}/\{[MS]_{i,0}/LS]\}$  over time at pH values of 7.0 and 6.5, where i is the trace element,  $[M]_i$  is the dissolved concentration of the element i,  $[M]_{i,0}$  is the dissolved concentration of the element i at t = 0 and  $[MS]_{i,0}$  i is the maximum concentration of the element that can be released from the sediment. The experimental release (● pH=7; ■ pH=6.5), error bars and simulated curves using the model by Martín-Torre et al. (2015b) (\_\_\_) are represented. 



510 511

Fig. 7. Trace element release  $(x_i = \{[M]_i - [M]_{i,0}\}/\{[MS]_{i,0}/LS]\})$  over time at pH=6.0, where i is the trace element,  $[M]_i$  is the dissolved concentration of the element i,  $[M]_{i,0}$  is the dissolved concentration of the element i at t = 0 and  $[MS]_{i,0}$  i is the maximum concentration of the element that can be released from the sediment. Experimental release ( $\diamond$ pH=6.0, [Fe]<sub>seawater</sub>=9.02 µg/L;  $\circ$  pH=6.0, [Fe]<sub>seawater</sub>=46.1 µg/L; and  $\Box$  pH=6.0, [Fe]<sub>seawater</sub>=153 µg/L), error bars and simulated curves are represented: \_\_\_\_\_ Martín-Torre et al. (2015b) model; and \_\_\_\_ modified model

# 518Table 3. Estimated values of the maximum concentrations that can be released from the reduced and oxidised fractions of the sediment, $([MRed]_{i,0} \text{ and } [MOx]_{i,0} \text{ respectively})$ ,519as expressed in units of mg/kg.

_	2	n
Э	Z	υ

	Assay					Contar	ninant					
	[Fe] <sub>seawater</sub> (µg/L)		Z	n	Р	b	C	d	N	li	Α	S
рн		[ZnRed] <sub>0</sub>	[ZnOx] <sub>0</sub>	[PbRed] <sub>0</sub>	[PbOx] <sub>0</sub>	[CdRed] <sub>0</sub>	[CdOx] <sub>0</sub>	[NiRed] <sub>0</sub>	[NiOx] <sub>0</sub>	[AsRed] <sub>0</sub>	[AsOx] <sub>0</sub>	
Martín-To	Martín-Torre et al. (2015b) model											
7.0	6.80	17.6	2.06	0.280	0.0359	0.0676	0	0.0908	0.280	0.971	0.136	
6.5	3.50	373	10.8	0.390	0.0733	0.178	0	0.635	0.390	0.298	0.0557	
6.0*	9.02	828	0	2.96	0	8.06	0	1.14	2.96	5.01	2.43	
Modified model for pH=6												
6.0	46.1	1150	33.9	3.07	0.012	8.53	0	1.65	3.07	4.37	7.70	
6.0	153	1110	70.6	5.11	0.424	8.71	0	1.082	5.11	4.40	2.37	

522 \*Estimated concentrations at pH=6 and [Fe]<sub>seawater</sub>=9.02 μg/L take the same value when using the model by Martín-Torre et al. (2015b) and the modified model at pH=6.

Table 4. Estimated kinetic rate coefficients for each contaminant in all the assays.

	Assav		Contaminant							
рН	[Fe] <sub>seawater</sub> (µg/L)	$k_{j,i} \ (h^{\text{-}1})$	Zn	Pb	Cd	Ni	As			
Martín-T	orre et al. (2015b)									
model		I	2 10 10-3	<b>5 00 10</b> -3	77710-3	0.05.10-3	2 5 4 10-3			
7.0	< 00	к <sub>1,i</sub>	5.10 10 <sup>-3</sup>	5.02 10 5	1.00.10-3	9.95 10 <sup>-3</sup>	$3.54 \ 10^{-3}$			
7.0	6.80	К <sub>2,i</sub>	3.30 10-5	6.34 IU <sup>-5</sup>	1.09 10-5	3.06 10-5	3.62 10 <sup>-3</sup>			
		К <sub>3,i</sub>	0	0	0	0	2.92 10-2			
		k <sub>1,i</sub>	1.96 10-3	4.64 10-3	6.41 10-3	2.54 10-3	2.08 10-3			
6.5	3.50	k <sub>2,i</sub>	1.70 10 <sup>-3</sup>	8.60 10-3	1.04 10-3	2.91 10 <sup>-1</sup>	2.86 10 <sup>-3</sup>			
		k <sub>3,i</sub>	0	0	0	0	1.64 10-2			
		k <sub>1,i</sub>	2.30 10-3	1.60 10-3	3.01 10-4	7.69 10 <sup>-3</sup>	2.18 10-3			
6.0	9.02	$\mathbf{k}_{2,i}$	3.68 10-3	1.09 10-3	3.46 10-4	5.88 10-3	1.74 10-2			
		$\mathbf{k}_{3,i}$	0	0	0	0	1.15 10-1			
Modified	model for pH=6									
		k <sub>1,i</sub>	2.30 10-3	1.60 10-3	3.01 10-4	7.69 10 <sup>-3</sup>	2.18 10-3			
< 0	0.02	k <sub>2,i</sub>	5.28 10-4	4.36 10-3	2.57 10-5	0	1.14 10-2			
6.0	9.02	$k_{2,i} = k_{2,i}^{1}$	1.74 10-5	4.08 10-5	1.65 10-6	4.05 10-5	2.34 10-3			
		$\mathbf{k}_{2i}^2$		0		2.03 10-3	0			
		k <sub>3,i</sub>	0	0	0	0	1.15 10-1			
		k <sub>1.i</sub>	2.30 10-3	1.60 10-3	3.01 10-4	7.69 10-3	2.18 10-3			
		k <sub>2.i</sub>	4.87 10-4	2.46 10-3	4.89 10 <sup>-5</sup>	1.59 10 <sup>-3</sup>	1.25 10-3			
6.0	46.1	$k_{2,i} = k_{2,i}^1$	4.28 10-6	8.91 10-4	6.62 10-7	1.57 10-6	5.29 10-4			
		k <sub>2,i</sub>		3.91 10-4		2.47 10-3	5.57 10 <sup>-3</sup>			
		k <sub>3.i</sub>	0	0	0	0	1.15 10-1			
6.0	153	k <sub>1,i</sub>	2.30 10-3	1.60 10-3	3.01 10-4	7.69 10 <sup>-3</sup>	2.18 10-3			
		k <sub>2 i</sub>	1.19 10-4	8.38 10-6	0	0	1.14 10-2			
		$k_{2,i} = k_{2,i}^{2,i}$	1.01 10-6	1.58 10-6	1.78 10-7	3.50 10-6	4.35 10-3			
		$k_{2i}^{2i}$		3.91 10-4		5.41 10-4	0			
		k <sub>3.i</sub>	0	0	0	0	1.15 10-1			
		0,2								

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#### 540 **4.2. Kinetic model of Fe release**

541 When using the kinetic model by Martín-Torre et al. (2015b), the good fit between the 542 experimental and simulated Fe release at pH values of 7.0 and 6.5 is shown in Fig. 8. 543 Moreover, the Fe concentrations in the reduced and oxidised fractions, or  $[FeRed]_0$  and 544  $[FeOx]_0$ , respectively, and the rate coefficients for the simulated Fe release at these pH 545 values are listed in Table 5.

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However, this model cannot explain the behaviour of Fe at the most acidic pH value
under study. At pH=6.0, the simulated curves that employ this model explain the release
of Fe within short periods but do not explain its decrease, nor is the second peak
observed in Fe mobilisation (dashed curves in Fig. 8).

551

More complex kinetic schemes, which always consider first-order reactions with respect 552 to Fe, have been tested to find a mathematical model that better predicts the behaviour 553 554 of this element at pH=6.0. The possibility that Fe is released directly from the reduced fraction was considered; iron sulphides such as pyrite could be released into the medium 555 without being oxidised. Another hypothesis was that most of the dissolved Fe 556 precipitates took the form of Fe(III), whereas the rest of it precipitates as Fe(II) before 557 558 being released again to the medium; afterwards, this re-dissolved Fe could precipitate as 559 Fe(III). These kinetically Fe-dependent schemes do not improve the previous fitting, and so we propose that there is a dependency between the Fe release and the presence of 560 other ions in association. Therefore, a model that includes the kinetics of ions associated 561 with Fe is introduced. The global model of Fe consists of two series of reactions in 562 parallel (Eqs. 11 and 12) and includes the influence of other ions on the release of Fe. 563

564

$$\operatorname{FeRed}(s) \xrightarrow{k_{1,Fe}} \operatorname{FeOx}(s) \xrightarrow{k_{2,Fe}} \operatorname{Fe}^{2+}(aq) \xrightarrow{k_{3,Fe}} \operatorname{Fe}^{3+}(s) \tag{11}$$

$$AC1(aq) \xrightarrow{k_{1,a}} AS1(s) \xrightarrow{k_{2,a}} AC2(aq) \xrightarrow{k_{3,a}} AS2(s)$$
(12)

Eq. 11 considers that reduced Fe should be oxidised before being released from the 566 sediment into the medium, and that dissolved Fe precipitates as Fe(III) whereas Eq. 12 567 568 represents an additional scheme for ions that are associated with to Fe. Ions that are 569 associated with Fe in solution (AC1) could precipitate with Fe (AS1) and be released together (AC2) and precipitated (AS2) because of the different solubility of products 570 from the compounds present in the medium. This scheme is feasible because of the high 571 concentration of ions dissolved in the seawater-sediment system and the ionic 572 573 interactions associated with this situation.

574

The association between Fe and these ions is taken into account by assuming that the release and precipitation rate coefficients, namely  $k_{2,Fe}$  and  $k_{3,Fe}$ , depend on the dissolved concentration of these ions through Eqs. 13 and 14.

$$k_{2,Fe} = k_{2,Fe}^{0} + k_{2,Fe}^{1} x_{A}$$
(13)

$$k_{3,Fe} = k^{0}_{3,Fe} + k^{1}_{3,Fe} (1 - x_{A})$$
(14)

578

where  $k_{2,Fe}^{0}$ ,  $k_{2,Fe}^{1}$ ,  $k_{3,Fe}^{0}$  and  $k_{3,Fe}^{1}$  are kinetic rate coefficients and  $x_{A}$  is the fraction of ions released from the sediment, as calculated as  $x_{A} = \frac{AC2-AC1_{0}}{AC_{max}}$ . AC2 refers to the concentration of ions released with Fe, AC1<sub>0</sub> is the concentration of seawater ions associated with Fe and AC<sub>max</sub> is the maximum concentration of ions associated with Fe in the aqueous phase.

In Fig. 8, the simulated curves at pH=6.0 are shown using a continuous line. The proposed model fits well with the experimental data, including the second peak observed at the higher Fe concentrations in the selected seawater. The good fit is corroborated by the percentage variation-explained values ( $R^2$ ) (99.3, 95.8 and 96.1 when the Fe concentrations of the seawater are 9.02, 46.1 and 153 µg/L, respectively). Additionally, Table 5 lists the maximum concentration of Fe that can be released from each fraction of the sediment and kinetic rate coefficients using the latter model.

592

This more complex model has been applied to the Fe release at pH values of 7.0 and 6.5 without obtaining an improvement in the simulated results. The global percentage variation-explained value of 97.8 and a relative standard deviation of 0.325 could be concluded for the Fe release model at the pH values and Fe seawater concentrations under study.



Fig. 8. Fe release  $(x_{Fe} = [Fe]/[Fe]_0)$  over time at pH values under study, where [Fe] is the dissolved concentration of Fe and  $[FeS]_0$  is the maximum Fe concentration that can be released from the sediment. Experimental release ( $\bullet$  pH=7;  $\blacksquare$  pH=6.5;  $\diamond$  pH=6.0, [Fe]<sub>seawater</sub>=9.02 µg/L;  $\circ$  pH=6.0, [Fe]<sub>seawater</sub>=46.1

 $\begin{array}{ll} \textbf{602} & \mu g/L; \text{ and } \square \text{ pH=6.0, [Fe]}_{seawater} = 153 \ \mu g/L) \text{ and simulated curves are represented: } \underline{\qquad} \text{most suitable model} \\ \textbf{603} & \text{for each assay; } \underline{\qquad} \underline{\qquad} \text{Martín-Torre et al. (2015b) model applied to assays at pH=6.0.} \end{array}$ 

Table 5. Estimated values for the maximum concentrations of Fe that can be released from the reduced ( [FeRed]<sub>0</sub>) and oxidised ( [FeOx]<sub>0</sub>) sediment fractions and kinetic rate coefficients in all assays.
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рН	[Fe]seawater	[FeRed] <sub>0</sub>	[FeOx] <sub>0</sub>	k <sub>1,Fe</sub>	$k_{2,Fe}^0$	$k_{2,Fe}^1$	k <sup>0</sup> <sub>3,Fe</sub>	k <sup>1</sup> <sub>3,Fe</sub>	k <sub>1,a</sub>	k <sub>2,a</sub>	k <sub>3,a</sub>
	(µg/L)	(mg	g/kg) ——				(h	1 <sup>-1</sup> ) —			
7.0	6.80	276	39.0	1.08 10-4	2.74 10-1		2.95 10-1				
6.5	3.50	86.0	623	3.50 10-4	2.01 10-1		2.35 10-1				
6.0	9.02	42.8	8040	1.88 10-3	1.42 10-2	1.11 10-1	9.73 10-2	0	1.16 10-2	1.15 10-2	6.95 10-4
6.0	46.1	26.1	6620	1.88 10-3	1.09 10-2	1.37 10-1	1.32 10-1	7.61 10-2	1.71 10-2	1.72 10-2	1.62 10-3
6.0	153	58.6	18400	1.88 10-3	9.90 10-3	4.12 10-2	1.69 10-1	6.80 10-2	1.89 10-2	1.89 10-2	1.52 10-3

## 612 **4.3. Trace element modelling at pH=6.0**

The contaminant release at pH=6.0 is first simulated according to the kinetic scheme used at higher pH values (Eq. 3). However, the release reaction is influenced by the dissolved Fe concentration so the rate coefficient  $k_{2,i}$  is modified according to Eq. 15 to account for this effect.

$$k_{2,i} = k_{2,i}^{0} + \frac{k_{2,i}^{1}}{x_{Fe}}$$
(15)

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618 where i represents the element (Zn, Pb, Ni, Cd, or As),  $k_{2,i}^{0}$  and  $k_{2,i}^{1}$  are kinetic rate 619 coefficients, whereas the fraction of Fe released from the sediment is defined as  $x_{Fe} =$ 620  $\frac{[Fe]}{[FeS]_{0}}$  with [Fe] being the dissolved concentration of Fe and [FeS]\_0 being the maximum 621 concentration of Fe that can be released from the sediment. The former coefficient 622  $(k_{2,i}^{0})$  considers the characteristics of the aqueous medium that might modify the 623 solubility of the different species present in the sediment whereas  $k_{2,i}^{1}$  represents the 624 direct influence of Fe on the release rate.

625

Including the influence of Fe, the simulated results of Zn and Cd fit well with their experimental release. However, the simulated release of Pb, Ni and As (not shown) is much lower than the experimental one for short periods (t<96 h) so the ionic competition with other cations should be considered in a similar way as the Fe release.

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To avoid a high increase in the number of parameters, only the influence of Zn, which is the major trace element in the studied sediment, is also included when simulating the release of Pb, Ni and As. Hence, a further term is added and the rate coefficient of the release reaction of Pb, Ni and As includes the combined influence of Fe and Zn asshown in Eq. 16.

$$k_{2,i} = k_{2,i}^{0} + \frac{k_{2,i}^{1}}{x_{Fe}} + \frac{k_{2,i}^{2}}{x_{Zn}}$$
(16)

636

637 where i represents the element (Pb, Ni and As),  $k_{2,i}^2$  is the rate coefficient which 638 represents the direct influence of Zn on the release rate and the fraction of Zn released 639 from the sediment is defined as  $x_{Zn} = \frac{[Zn]}{[ZnS]_0}$  with [Zn] being the concentration of 640 dissolved Zn and [ZnS]<sub>0</sub> being the maximum concentration of Zn that can be released 641 from the sediment. The variables  $k_{2,i}^0$ ,  $k_{2,i}^1$  and  $x_{Fe}$  were previously explained.

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The fit between the simulated and experimental release for the different assays at pH=6.0 when using this modified model is shown in Fig. 7 with continuous lines. The percentage variation-explained values values ( $R^2$ ) for the different assays at pH=6.0 are, from the lowest to the highest concentrations of Fe in the seawater, 99.5, 96.1 and 96.3. These values are influenced by the simulated release of Fe, leading to higher  $R^2$  values of elemental release for the better Fe fittings.

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The concentrations of each element in the reduced and oxidised sediment fractions are estimated by using this modified model and are listed in Table 3. At pH=6 and [Fe]<sub>seawater</sub>=9.02  $\mu$ g/L, the estimated concentrations take the same value as they do when using the model by Martín-Torre et al. (2015b). Similar to what occurred at most neutral pH values, higher concentrations are released from the reduced fraction than from the oxidised fraction of the sediment, except the Ni and As in the assay in which the seawater contains 46.1  $\mu$ g/L of Fe.

Regarding the pH=6 assay with the lowest seawater Fe concentration (9.02  $\mu$ g/L), a better fitting is obtained when using the modified model than the previous one. Percentage variation-explained values (R<sup>2</sup>) of 99.5, 98.6, 99.7, 98.4 and 81.2 instead of 99.3, 98.4, 98.3, 97.0 and 74.1 are obtained for Zn, Pb, Cd, Ni and As, respectively. However, a trade-off between the estimated work and the accuracy should be decided by the user because the use of this modified model implies the estimation of more parameters.

666 Given that ionic competition, based on the different solubility products, is a process with a great influence on the contaminant release, the interactions among all of the ions 667 could be added through additional elemental fractions  $(x_{Pb}, x_{Cd}, x_{Ni} \text{ and } x_{As})$ . This 668 approach has been attempted, and despite the increasing the number of estimated 669 670 parameters, the results do not present enough sensitivity to better fit the experimental release. Hence, the dependence of Fe is considered for all the studied trace elements at 671 pH=6.0, but the influence of Zn is only contemplated for Pb, Ni and As release at this 672 673 pH value.

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Something similar occurs when this model, which was proposed for the simulation of contaminant release at pH=6.0, is used to fit experimental results from assays at pH values of 7.0 and 6.5 a better fit is obtained when the parameters that are not included in the initial proposed model employ a value of zero.

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As expected, higher concentrations of the studied trace elements are released from the sediment  $([MS]_{i,0} = [MOx]_{i,0} + [MRed]_{i,0})$  at lower pH values because of the greater

solubility of the studied contaminants at acidic pH values, except for As, which presents
the highest mobilisation at pH=6.5 (Fig. 9 a). In comparing the mobilisation at pH=6.0,
Fig. 9b indicates that the higher Fe concentrations in the selected seawater causes a
slightly higher mobilisation of all the studied contaminants.

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687 688

Fig. 9. a) Maximum concentration of contaminant that can be released from the sediment as a function of pH; and b) maximum concentration of contaminant that can be released from the sediment at pH=6.0 as a function of the Fe concentration in the seawater. ● Zn; ■ Pb; ○ Ni; □ Cd; and ▲ As

As previously stated, the rate coefficients of the different assays are presented in Table 4, in which As is the only element with an adsorption or precipitation reaction  $(k_3 \neq 0)$ . Moreover, at pH=6.0 the oxidation and adsorption/precipitation processes of this oxyanion do not depend on the initial Fe concentration in the seawater, and thus the rate coefficients have the same value in the three assays conducted at this pH value.

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Although the model considers the influence of Zn on the release of Pb, Ni and As, it takes a nonzero value ( $k_{2,i}^2 \neq 0$ ) in all the cases for Ni release, for Pb release at the two highest concentrations of Fe in seawater (153 and 46.1 µg/L) and for As mobilisation at [Fe]<sub>seawater</sub>= 46.1 µg/L; this last case likely occurs because the released concentration of this oxyanion from the oxidised fraction is higher than it was from the reduced fraction at this seawater Fe content. However, the obtained rate kinetic parameters  $(k_{2,i}^0, k_{2,i}^1)$ and  $k_{2,i}^2$ ) do not present a clear trend as a function of [Fe]<sub>seawater</sub>.

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707 The type of acidification (CO<sub>2</sub> or HNO<sub>3</sub>) used in the pH-static assay influences the rate 708 coefficients obtained from the Martín-Torre et al. model (2015b). Contrary to what 709 occurs when the mineral acid is used (Martín-Torre et al., 2015b), rate coefficients obtained using CO<sub>2</sub> do not present a clear trend with respect to the pH and cannot be 710 711 fitted to any polynomial equation. HNO<sub>3</sub> is a strong oxidising acid that generates soluble salts in the medium, and it might be completely dissociated.  $CO_2$  is a weak 712 oxidising acid that forms partial insoluble salts, the solubility products of which modify 713 714 the characteristics of the medium through ionic competition and therefore influence the 715 element release rates.

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The maximum concentrations of the elements that can be released from the sediment might depend on the type of acidification in addition to the oxidation state of the sediment. Hence, higher values are generally obtained in assays that employ  $CO_2$ instead of HNO<sub>3</sub> (Martín-Torre et al., 2015b). Exceptions to this trend are Pb concentration at pH values of 7.0 and 6.5 (similar values are independent of the selected acid) and As at the most neutral pH values, where higher concentrations are released from the sediment in assay samples acidified by HNO<sub>3</sub>.

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The parity plots with the percentage variation-explained values  $(R^2)$  and relative standard deviation (RSD) included a consideration of the global modified model of this work for all the assays performed, as obtained for each contaminant (Fig. 10). Parity plots are useful for the validation of the model in terms of the released element 729 concentration at any time and pH value. In general, at release higher than 25% of the 730 maximum dissolved concentration, differences between experimental and simulated mobilisation are lower than 20% although a higher dispersion is observed in the case of 731 As. Better fittings are obtained at pH values of 7.0 and 6.5 because of a lack of the 732 733 influence from the Fe concentration and ionic competition on the elemental release. The percentage variation-explained values  $(R^2)$  for each contaminant are higher than 97.8 for 734 all the elements under study, except for As, in which  $R^2$ =89.4. The high values of this 735 736 statistical parameter show the good fit of the proposed model.





Fig. 10. Parity plots of the element concentrations from the experiment (Cexp) and simulated results
(Csim) of the elements under study. The area inside the dashed square includes concentrations lower than
25% of the maximum dissolved concentration. ● pH=7.0; ■ pH=6.5; ◇pH=6.0, [Fe]<sub>seawater</sub>=9.02 µg/L; ○
pH=6.0, [Fe]<sub>seawater</sub>=46.1 µg/L; and □ pH=6.0, [Fe]<sub>seawater</sub>=153 µg/L. Additionally, the data number (N),
percentage variation-explained value (R<sup>2</sup>) and relative standard deviation (RSD) are shown.

## 747 5. CONCLUSIONS

This work presents the experimental results of Zn, Pb, Cd, Ni, Cr, Cu and As mobilisation over time when sediment and seawater are totally mixed and acidified by CO<sub>2</sub> gas in a pH-static leaching test. Long-term 480-h leaching tests are performed at
pH values of 7.0, 6.5 and 6.0 to assess different levels of acidification.

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The evolution of the redox potential and Fe release over time at pH=6.0 is different than the evolution at the more neutral pH values under study. Therefore the influence of the seawater Fe concentration is assessed by conducting assays at pH=6.0 using natural seawater with different Fe contents: 9.02; 46.1 and 153  $\mu$ g/L.

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A set of three in-series reactions for trace elements, for Fe and for other ions associated with Fe is proposed to model a Fe/multi-ion-dependent mechanism for trace metal release. The model uses global  $R^2$  values of 98.9-89.4 and RSD < 0.325 to explain the release behaviour over time for Zn, Pb, Cd, Ni and As at pH values of 7.0, 6.5 and 6.0, which mimic the potential CCS acidification, and three seawater iron concentrations, which represent estuarine seawater in the iron mine areas.

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The model proposed in this work extends the scope of previous models to the acidification of seawater with  $CO_2$ , at different iron concentrations; however, a tradeoff between estimation and accuracy is made because there are higher numbers of parameters in the proposed model. Using the Aspen Custom Modeler, the maximum concentrations of each element that can be released from the sediment, and the kinetic rate coefficients are estimated for all the cases.

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Additionally, the kinetic rate coefficients obtained in this work are compared with those that were obtained previously when  $HNO_3$  was used to acidify the medium. Different trends are observed because of the impact that  $CO_2$  has on the ionic competition and

contaminant release, highlighting that the displacement reactions should be considered when acidification from  $CO_2$  leakages originating from CCS sites is assessed.

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778 The parity plots of the pH-static leaching test when using CO<sub>2</sub> to acidify the medium show a good fit between the experimental and simulated element release, confirming 779 that the proposed model can be applied to simulate contaminant mobilisation from 780 contaminated sediment to seawater under total mixed acidic conditions at pH values of 781 782 7.0, 6.5 and 6.0. Hence, a useful, broader generalised kinetic model that explains the contaminant release over time in sediment-seawater mixtures that were acidified by 783 784 HNO<sub>3</sub> or CO<sub>2</sub> was made to obtain acidic pH values and different concentrations of iron. The proposed model is flexible enough to work with sediments that have different 785 contaminant contents through model-fitting parameters such as the oxidised and 786 787 reduced fractions and the kinetic rate coefficients.

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