



The role of localised compressional Ultra-Low Frequency waves in energetic electron precipitation

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RESEARCH ARTICLE

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Key Points:

- We detail a new mechanism for direct modulation of electron precipitation via localized compressional waves
- Electrons encountering a time-varying and spatially localized ULF wave can break the third invariant
- This localized mechanism has not previously been considered and may be important for radiation belt losses

Supporting Information:

- Supporting Information S1

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The Role of Localized Compressional Ultra-low Frequency Waves in Energetic Electron Precipitation

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Abstract Typically, ultra-low frequency (ULF) waves have historically been invoked for radial diffusive transport leading to acceleration and loss of outer radiation belt electrons. At higher frequencies, very low frequency waves are generally thought to provide a mechanism for localized acceleration and loss through precipitation into the ionosphere of radiation belt electrons. In this study we present a new mechanism for electron loss through precipitation into the ionosphere due to a direct modulation of the loss cone via localized compressional ULF waves. We present a case study of compressional wave activity in tandem with riometer and balloon-borne electron precipitation across keV-MeV energies to demonstrate that the experimental measurements can be explained by our new enhanced loss cone mechanism. Observational evidence is presented demonstrating that modulation of the equatorial loss cone can occur via localized compressional wave activity, which greatly exceeds the change in pitch angle through conservation of the first and second adiabatic invariants. The precipitation response can be a complex interplay between electron energy, the localization of the waves, the shape of the phase space density profile at low pitch angles, ionospheric decay time scales, and the time dependence of the electron source; we show that two pivotal components not usually considered are localized ULF wave fields and ionospheric decay time scales. We conclude that enhanced precipitation driven by compressional ULF wave modulation of the loss cone is a viable candidate for direct precipitation of radiation belt electrons without any additional requirement for gyroresonant wave-particle interaction. Additional mechanisms would be complementary and additive in providing means to precipitate electrons from the radiation belts during storm times.

1. Introduction

Energetic electron losses are a critical component of electron dynamics in the inner magnetosphere and outer radiation belt. Although electron dynamics in this region are only partially understood, electron losses can essentially occur either when their drift trajectories intersect with the magnetopause, termed magnetopause shadowing (e.g., Ozeke et al., 2014; Turner et al., 2012; West et al., 1972), or when their bounce trajectories lead them to be lost to the upper atmosphere (e.g., Millan et al., 2002). Traditionally, wave-particle interaction and in particular gyroresonant interaction with plasmaspheric hiss (e.g., Meredith et al., 2007) or whistler mode waves (Horne et al., 2003) are invoked as being responsible for pitch angle scattering of electrons into the loss cone and enhanced precipitation (e.g., Clilverd, Rodger, et al., 2015; Kennel & Petschek, 1966; Millan & Thorne, 2007; Rodger et al., 2012). Indeed, whistler mode chorus waves are thought to provide the source of the lower-energy diffuse aurora (e.g., Thorne et al., 2010), and observations of whistler mode chorus have been shown to be clearly linked to modulation of the diffuse aurora (e.g., Nishimura et al., 2010). However, large whistler mode chorus amplitudes are waves typically limited to the nighttime and morning sectors of the magnetosphere (Li et al., 2010; Meredith et al., 2012), whereas diffuse auroral processes occur at all magnetic local times (MLTs). More recently, other gyroresonant interactions have also been identified as being a major loss mechanism for relativistic electrons (Breneman et al., 2015), first invoked by Thorne and Kennel (1971). Electromagnetic ion cyclotron (EMIC) wave-driven electron loss is now considered to also be a significant contributor to radiation belt losses (e.g., Hendry et al., 2012, 2016, 2017;

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Rodger et al., 2015; Zhang et al., 2016). The source of auroral particle precipitation across all MLT sectors remains to be determined.

Ultra-low frequency (ULF) waves have been proposed to provide both resonant (e.g., Elkington et al., 1999; Mann et al., 2013) and diffusive (e.g., Brautigam & Albert, 2000; Schulz & Lanzerotti, 1974) acceleration and transport of electrons. ULF wave precipitation signatures have been observed since the early 1960s (e.g., Anger et al., 1963; Brown, 1964; Ziauddin, 1960) and have been observed in riometer (e.g., Beharrell et al., 2010; Heacock & Hunsucker, 1977; Olson et al., 1980; Spanswick et al., 2005), auroral (e.g., Rae, Mann, Dent, et al., 2007; Roldugin & Roldugin, 2008), and X-ray-related precipitation (e.g., Brito et al., 2012; Halford et al., 2015; Motoba et al., 2013). However, although ULF wave signatures have been observed in precipitation across a wide range of energies from keV to MeV, these waves have only been proposed to be indirectly involved in energetic electron losses. This ULF modulation of precipitation in the Pc4–Pc5 frequency range (e.g., Jacobs et al., 1964) is often discussed in terms of the ULF modulation of other wave modes, principally ULF modulation of whistler mode wave growth rates (Breneman et al., 2015; Li et al., 2011; Millan & Thorne, 2007; Nishimura et al., 2013; Rae, Mann, Dent, et al., 2007; Spanswick et al., 2005; Watt et al., 2011). In these circumstances, ULF modulation of VLF wave intensities would provide an energy-dependent response where those electrons that are able to resonate with specific VLF wave frequencies would be precipitated. However, the remaining part of the phase space density (PSD) would most likely be unchanged, at least over ULF wave periods (~minutes) where pitch angle diffusion time scales are usually very long (hours to days; e.g., Horne et al., 2005). Hence, any broadband precipitation must also be explained in full, again at all local times. Recently, Brito et al. (2012, 2015) proposed a new mechanism whereby global ULF waves could be directly implicated in radiation belt losses. The radial motion of relativistic electrons within global-scale compressional ULF waves would mean that electrons would experience larger magnetic fields and shorter field line lengths during the inward motion phase, where conservation of the first and second adiabatic invariants would lead to a gain of parallel energy, altering pitch angles of some electrons sufficiently such that they move into the local loss cone. Although not discussed explicitly by Brito et al. (2012, 2015), this mechanism may work over a large range of electron energies.

In this paper, we explore a related but new mechanism of electron precipitation directly driven by highly localized compressional ULF waves simply by modulating the equatorial loss cone appreciably from its average, or typical, value. We apply this mechanism initially at geosynchronous (GEO) orbit and find that the average geosynchronous loss cone can increase by up to 50% during large-amplitude compressional ULF waves. Depending upon the shape of the PSD close to the loss cone, this mechanism can provide a significant additional amount of precipitating flux without the requirement for any wave-particle interaction processes. We show clear experimental evidence of ULF wave-modulated precipitating electron fluxes across a wide range of energies (20–400 keV) and conclude that compressional ULF waves should be considered a direct, rather than an indirect, candidate precipitation mechanism for radiation belt electrons, or indeed all electrons close to the loss cone with bounce periods less than the wave period. This mechanism has the potential to directly drive electron precipitation across the entire outer radiation belt and over a wide range of energies and is not limited to geostationary magnetic latitudes where our observations are concentrated.

2. Instrumentation

In this paper, we primarily utilize data from the GOES fluxgate magnetometers at 0.512 s cadence (Singer et al., 1996). However, we also augment this with ground magnetometer data from the CARISMA (Canadian Array for Realtime Investigations of Magnetic Activity; Mann et al., 2008), together with Northern Solar Terrestrial Array (NORSTAR) riometer data (<http://aurora.phys.ucalgary.ca/norstar/rio/>), both at 1 s cadence. We further utilize southern hemispheric measurements of bremsstrahlung X-rays, related to the precipitation of energetic electrons from the Balloon Array for Radiation belt Relativistic Electron Loss (BARREL; Millan et al., 2013) Campaign 1, Payload 1H (1H) at both 50 ms resolution (fast X-ray spectrum channel 1 at <180 keV X-ray energy) and 32 s cadence (slow X-ray spectrum, for ~30 keV–10 MeV X-ray energies).

3. Motivation: 26 January 2013 Case Study

Figure 1 shows an overview of in situ and ground-based observations for a case study of ULF-modulated precipitation during the first NASA BARREL campaign on 26 January 2013. BARREL 1H is in the southern

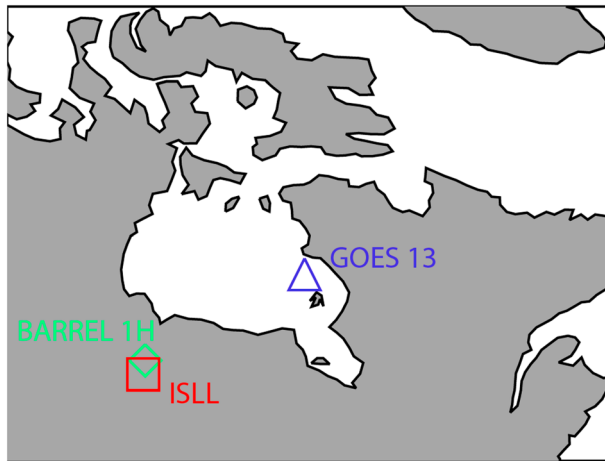


Figure 1. An overview of the relevant instrumentation for a case study on 26 January 2013 from 1930 to 2130 UT, projected into the ionosphere using the Tsyganenko 96 magnetic field model (Tsyganenko, 1995). During this event, the BARREL 1H balloon, situated in the southern hemisphere, was immediately conjugate to the ISLL NORSTAR riometer, at dipole L of ~ 5.2 , with the geostationary GOES 13 satellite around 1 h of local time to the east.

hemisphere immediately conjugate to the NORSTAR Island Lake (ISLL) riometer, determined using the T96 magnetic field model. At the period of interest, 1930–2130 UT, GOES 13 is at slightly higher latitudes and around 1 h of magnetic local time (MLT) to the east.

From top to bottom, Figure 2 shows (a) the GOES 13 and conjugate ground magnetometer magnetic field magnitudes, with (b) the modulation of the loss cone using measured and modeled ionospheric magnetic fields (to be discussed later). Figure 2c shows channel 1 from the BARREL fast spectra of <180 keV X-rays, and Figure 2d shows BARREL 1H slow spectra from 50 to 300 keV X-rays. Figure 2e shows riometer absorption from ISLL. Figure 2f shows the normalized frequency content of each of these data sets calculated within the vertical lines using Fast Fourier Transform (FFT) analysis of the respective time series from GOES 13 (black), BARREL fast (blue) and slow (green and yellow) spectra, and the ISLL riometer (red).

Figure 2a shows that large-amplitude (25 nT peak to valley on a background ~ 85 nT) compressional ULF waves are observed at geosynchronous orbit between 4 and 5 mHz (Figures 2a and 2f) in a temporally localized period between ~ 1950 and 2030 UT. Around 20–25 min later in UT (2015–2050 UT) and in the southern hemisphere, BARREL 1H

measures a clear and large-amplitude ULF-modulated electron precipitation event (Figures 2c, 2d, and 2f), as does the ISLL riometer, which is conjugate in the northern hemisphere (Figures 2e and 2f). Fourier analysis of these time series for the respective periods of modulation (denoted by the dashed vertical black lines) reveals that GOES magnetic field and BARREL 1H precipitation signatures share a dominant common frequency of 4–5 mHz, and the ISLL riometer a slightly lower dominant frequency of 3–4 mHz. Hence, there are common frequencies observed in both geosynchronous magnetometer data and modulated precipitation in the ionosphere. In section 4 we discuss the relevance of the slightly lower frequency observed in the precipitation seen through riometer absorption changes. We note that the precipitation signatures actually correlate best with the compressional ULF wave signature if shifted by 0:30 UT, which suggests that there is in fact a localized source of ULF wave activity drifting slowly westward or sunward, for example, ULF waves driven via an internal source such as unstable ion distributions drifting through that region (Yeoman & Wright, 2001). Further observations from the GOES 15 magnetometer, and McMurray (MCMU) riometer stations (not shown), both located further west from the GOES 13, ISLL, and BARREL-1H measurements also indicate that the ULF wave activity is localized in space and persists for at least 2 h of universal time.

We calculate the linear correlation coefficients for the period of ULF wave activity, noting that there is a large background perturbation to both the ISLL and BARREL data, and that a slightly lower frequency is observed at ISLL for reasons we discuss in the section 7, but which are primarily due to an ionospheric decay effect. Peak correlation coefficients between GOES and ISLL are 0.5, and between GOES and BARREL is 0.57. Correlation between both ionospheric measurements is significantly better given that both are ionospheric measurements and hence are subject to the same ionospheric decay, peaking at 0.87 between ISLL and BARREL slow spectra at 53 keV.

In summary, this case study exhibits localized compressional ULF wave observations from GOES at ~ 20 UT and ~ 1430 MLT in addition to localized ULF wave-modulated precipitation at ISLL and BARREL 1H at ~ 2015 UT at ~ 1315 MLT. Given that the ULF wave signatures are at the same frequency, our hypothesis is that a localized ULF wave field drives ULF-modulated precipitation. The changes in MLT of the localized ULF wave activity as time progresses indicate that these ULF wave signatures must be slowly moving westward, in keeping with an ion-generated compressional ULF wave.

4. What Processes could Drive Localized ULF-Modulated Precipitation?

Given that the ULF signatures are observed in the same local time region, but are temporally limited in extent, we interpret these combined measurements as clear evidence of a large-amplitude, spatially localized ULF wave field in a highly limited spatial range in the postnoon sector (14–15 MLT). We discuss the potential

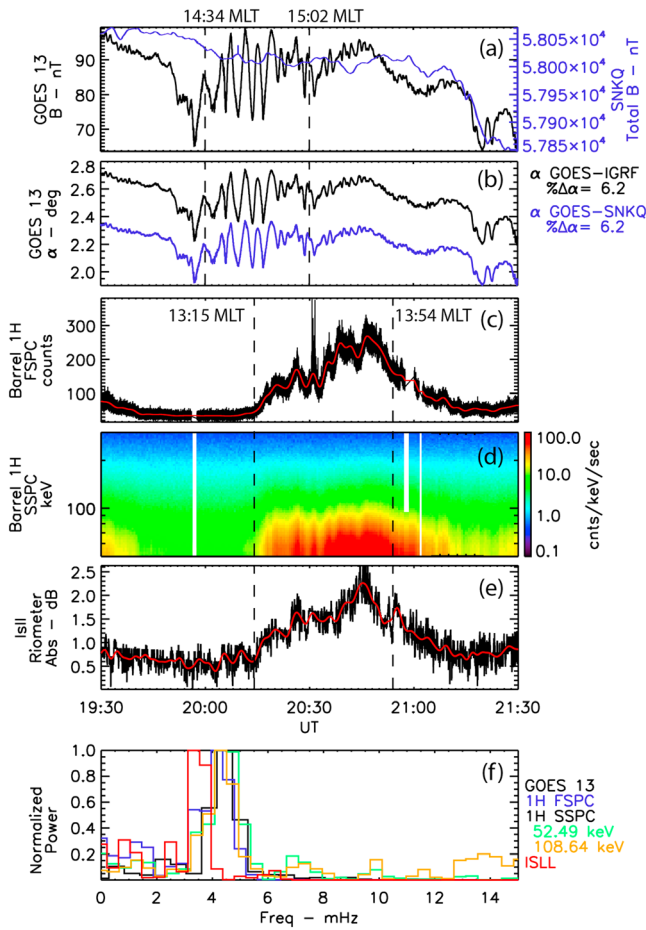


Figure 2. Measurements from the case study on 26 January 2013 from 1930 to 2130 UT. (a) GOES 13 and conjugate ground magnetometer magnetic field magnitudes, (b) modulation of the loss cone using measured and modeled ionospheric magnetic fields (to be discussed later), (c) channel 1 from the BARREL fast spectra of <180 keV X-rays, (d) BARREL 1H slow spectra from 50 to 300 keV X-rays, (e) riometer absorption from ISLL, and (f) the normalized frequency content of each of these data sets calculated within the vertical lines using FFT analysis of the respective time series from GOES 13 (black), BARREL fast (blue) and slow (green and yellow) spectra, and the ISLL riometer (red).

mirror point close to the ionosphere. Throughout this paper, we assume that the variation in magnetic field strength observed by GOES can be interpreted as the temporal variation of minimum magnetic field strength along the field line that threads its location. The values of magnetic field in this definition should be understood to be averages over time scales greater than the electron bounce time, which is short compared to a ULF wave period. To estimate the time evolution of equatorial BLC using the observed equatorial magnetic field at GOES for B_G , we must first estimate the magnetic field strength at the ionosphere B_I . Note that of the two magnetic field strengths required for equation (2), $B_G \ll B_I$. It is likely that both B_G and B_I vary as a result of the ULF wave, but the variations in B_G are a significant fraction of B_G , whereas the variations in B_I are very small compared to the magnitude of B_I . Therefore, the average B_I in the vicinity of the field line foot point mapped from the Geostationary Operational Environmental Satellite (GOES) spacecraft could be used in equation (2) with very little loss in accuracy. In this case study, we have compared two estimates for B_I : the projected International Geomagnetic Reference Field (IGRF) at 100 km of the location of the magnetic field foot point of the GOES position, as mapped using the Tsyganenko T89 (Tsyganenko, 1989) magnetic field model, and the magnetic field strength measured at the Sanikiluaq (SNKQ) ground-based magnetometer (being the magnetometer closest to the foot print of GOES West). Figure 2b shows the estimated modulation of α_G when using the IGRF field (black) and the measured field at SNKQ (blue). Regardless of the source of the

source of these waves in section 7 but conclude that whatever mechanism leads to the ULF-modulated precipitation event is highly localized in space, and not in time. The question then becomes, what drives this ULF-modulated precipitation?

Whistler mode waves are invoked to drive precipitation across a wide range of energies (e.g., Miyoshi et al., 2015). In the case of ULF-modulated precipitation, whistler mode waves are assumed to already exist, and the ULF waves modulate the growth rates of the waves due to a preexisting source of free energy (e.g., Coroniti & Kennel, 1970). Alternately, the VLF spectral distribution is modified via wave-wave interaction between ULF and VLF waves (e.g., Chen, 1974) leading to a ULF-modulated precipitation signature being observed. However, Figure 2 demonstrates that precipitation is not observed by BARREL or ISLL riometer outside of the bounds of the ULF event above the background level, implying that whatever processes cause the precipitation only exist inside the region of ULF waves, indicated in this figure by vertical dashed lines. If whistler mode waves are present outside of this spatial window, then one would certainly expect to observe unstructured, or differently structured, precipitation to be occurring when the ULF wave field is not present.

In this case study, we show that large-amplitude ULF wave fields are localized to only a fraction of the drift trajectory of an electron, meaning that an energetic electron will encounter a rapid step change in local magnetic field as it undertakes gradient-curvature drift. If the time scale of this wave is shorter than the drift period, the third adiabatic invariant is likely to be violated. We explore the effects of localized perturbations in magnetic field on the conservation or otherwise of all invariants.

The equatorial bounce loss cone (BLC) characterizes the maximum pitch angle of particles that would precipitate into the ionosphere within one bounce period and is defined as

$$\sin^2 \alpha_G = B_G / B_I \tag{1}$$

where α_G is the equatorial bounce loss cone angle and B_G is the magnetic field strength in the equatorial plane. The value of B_G is approximated by the magnetic field magnitude at GOES situated close to the equatorial plane, and B_I is the magnetic field strength at the particle

estimated ionospheric field, there is little difference to the modulation of the loss cone; it is only the average size of the loss cone that is different. Since we are interested in the modulation of the loss cone, we will for simplicity use the IGRF field at the location of the GOES foot print to determine B_i in the subsequent analysis, noting that this simplification of dipolar L shell determination of the first and second adiabatic invariants illustrates the utility of this calculation for enhanced modulation of precipitation, and which becomes increasingly appropriate for locations inside geosynchronous orbits and closer to the radiation belt region.

We now consider how the ULF wave alone could affect the pitch angle of individual particles. Since ULF wave time scales are of order minutes, we can assume that the first and second adiabatic invariants are conserved, but the third is not. Previous studies have investigated how conserving the first and second invariants affects the change in pitch angle and loss cone, under the assumption of a relatively dipolar magnetic field (e.g., Foster et al., 2015; Halford et al., 2015; Li et al., 1993; Wygant et al., 1994). For example, Halford et al. (2015) showed that the change in the equatorial pitch angle of a particle in a slowly changing and dipolar magnetic field configuration was independent of mass or energy and could be written as

$$\sin \alpha_{\text{eq},f} = \frac{-L_f^{1/2} \cos^2 \alpha_{\text{eq}0}}{2L_0^{1/2} \sin \alpha_{\text{eq}0}} + \frac{1}{2} \left(\frac{L_f \cos^4 \alpha_{\text{eq}0}}{L_0 \sin^2 \alpha_{\text{eq}0}} + 4 \right)^{1/2} \quad (2)$$

where $\alpha_{\text{eq}0}$ and $\alpha_{\text{eq},f}$ are the initial and final equatorial pitch angles and L_0 and L_f are the initial and final L values of the particle in dipolar L . This equation is valid for the action of a sufficiently low-frequency ULF wave. We use this idealized equation to make a comparison between the changes in particle pitch angle due to a slowly changing magnetic field and the changes in loss cone due to the same slowly changing magnetic field. We note here that at the location of GOES, a dipolar approximation is a simplification of the real measured magnetic field. However, as can be seen from Figures 2 and S1 in the supporting information, this is a reasonable assumption given that the measured magnetic field magnitude is ~ 100 nT during the event at geosynchronous orbit.

Figure 3 shows how the equatorial pitch angles $\alpha_{\text{eq},f}$ vary in a ULF wave-modulated magnetic field according to equation (1), where both the first and second adiabatic invariants are conserved. These changes (solid lines) are shown relative to the changes in the BLC α_G according to equation (2) (dashed lines). Figure 3a shows a range of low particle pitch angles (colored lines). The expected change in the BLC is a dashed line, and pitch angles that fall within the BLC are shaded in grey. For a slowly varying magnetic field, the change in BLC is far greater than any change in particle pitch angle conserving the first and second invariants.

Figure 3b shows selected “important” pitch angles for the case study shown in Figure 2. If only the average magnetic field is considered, the vertical location where the blue (upper) solid line crosses the dashed line would indicate the largest particle pitch angle to be lost into the BLC. The vertical location where the green (lower) solid line crosses the dashed line indicates the largest pitch angle that would be lost under the action of the ULF wave, which in this case has an amplitude of 13 nT. We will discuss in section 5 how even these small changes in equatorial loss cone can lead to large changes in precipitating flux.

Figure 3c shows the fractional change in $\alpha_{\text{eq},f}$ (solid lines) and α_G (dashed line) to demonstrate that the changes in the BLC are indeed much larger than the changes in the particle equatorial pitch angles and that for larger ULF wave fields, this effect becomes increasingly pronounced.

5. Implications for Precipitating Electron Flux

The previous section showed that compressional ULF waves can significantly modify the size of the equatorial BLC. In this section, we discuss the implications for driving or enhancing electron precipitation across all energies, likely impacting radiation belt electron dynamics.

Figure 4 illustrates the concept of localized ULF wave-driven precipitation, and how this precipitation mechanism is affected by the localization of the wave and shape of the equatorial distribution in pitch angle. Figure 4a shows a schematic demonstrating how drifting electrons might interact with localized compressional ULF waves and result in electron precipitation, with the Sun to the right of the figure. Electrons undergoing gradient curvature drift around the Earth will encounter a localized region of compressional ULF wave activity, such that electrons that were previously just outside of the bounce (and potentially drift) loss cones

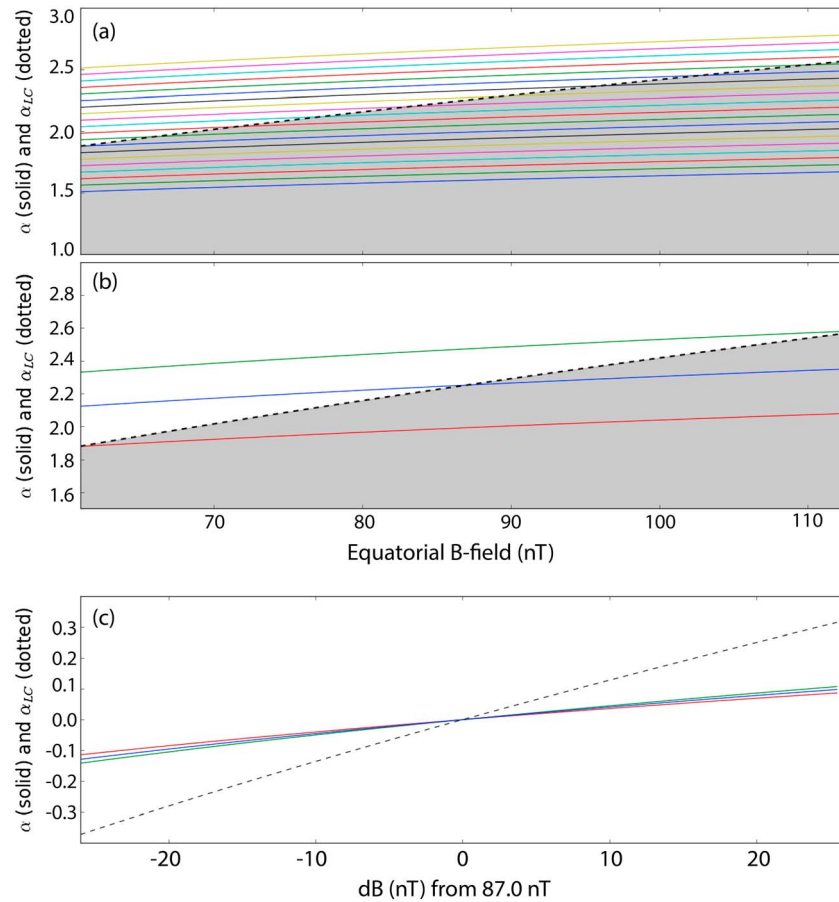


Figure 3. Demonstration of how equatorial pitch angles $\alpha_{eq,G}$ vary in a slowly varying magnetic field under the assumption of conservation of the first and second adiabatic invariants but not the third invariant (after Halford et al., 2015). (a) The range of small particle pitch angles (colored lines) and their variation according to equation (1). The expected change in the BLC due to a ULF wave with an amplitude of 13 nT is denoted as a dashed line, with pitch angles less than that and hence within the BLC shaded in grey. Colored lines denote sample pitch angles and their variation due to conservation of the first and second. (b) Important pitch angles for the case study shown in Figure 2. If only the average magnetic field is considered, the pitch angle where the blue (upper) solid line that crosses the dashed line would indicate the largest particle pitch angle to be lost into the BLC, and where the green (lower) solid line that crosses the dashed line indicates the largest pitch angle that would be lost under the action of the ULF wave. (c) The fractional change in $\alpha_{(eq,t)}$ (solid lines) and α_G (dashed line) to demonstrate that the changes in the BLC are indeed much larger than the changes in the particle equatorial pitch angles and that for larger and larger ULF wave fields, this effect becomes more and more pronounced.

and hence were trapped, then find themselves within the loss cone. We reiterate that this is a consequence of conservation of the first and second invariants and the violation of the third adiabatic invariant due to the spatially localized nature of the ULF waves.

If there are no additional electron sources to replenish those electrons that have been precipitated, and the region of ULF wave activity persists over time scales longer than a drift period, a range of resultant effects may be experienced, from a large precipitation spike into the atmosphere to a longer-lived ULF modulated precipitation signature (see Figure 4b). The precipitation signature as detected in the ionosphere depends upon the energy of the electron (i.e., how much time it spends within the ULF wave region) and the phase of its drift orbit relative to the phase and localization of the ULF oscillation. A single pulse of precipitation would indicate that a compressional ULF wave is acting over a large range of MLT such that electrons across a large fraction of the drift orbit at all energies within the enhanced loss cone would precipitate within the first wave cycle. For more localized compressional wave activity, the ionospheric electron precipitation signature may depend upon (i) the azimuthal wavenumber of the wave, (ii) the phase of the wave as the electron passes through the active region, and (iii) the azimuthal extent of the localized wave region. Hence, each drift

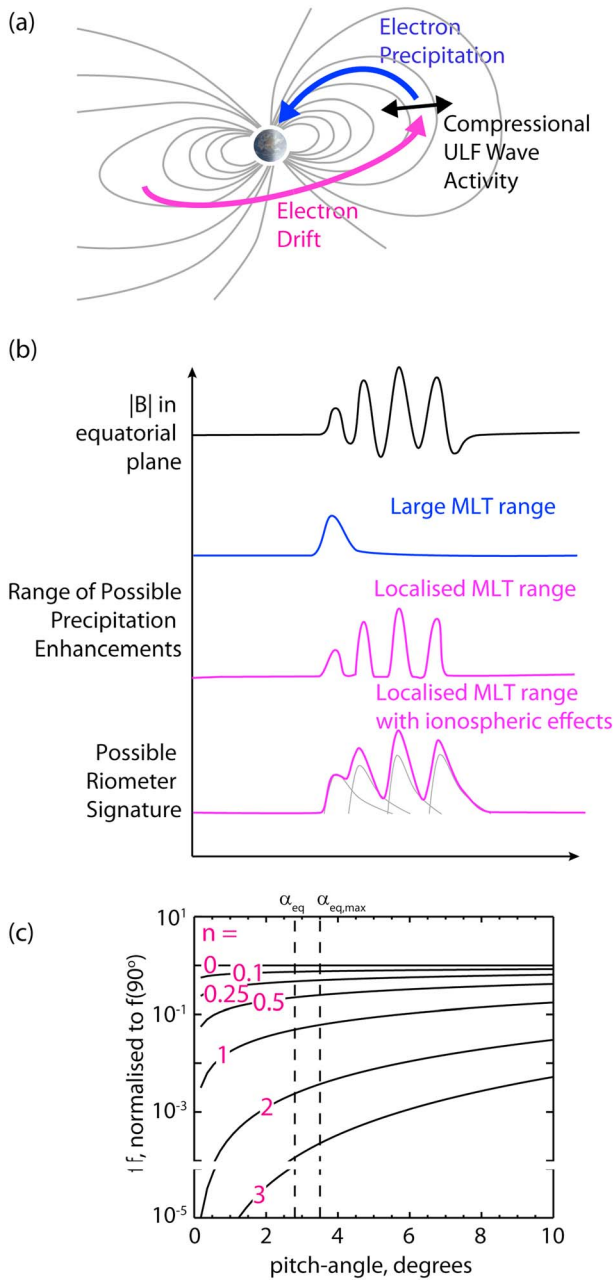


Figure 4. (a) Schematic representation of the noon-midnight meridian of Earth’s magnetosphere indicating the direction of electron drift, a possible region of compressional ULF wave activity on the dayside, the proposed generation mechanism for localized compressional waves, and the resulting electron precipitation. (b) Theoretical variations of phase space density at a particular energy as a function of equatorial pitch angle. Following Gu et al., (2011), we use $f = f_0 \sin^n \alpha_G$, where possible values of n are indicated in the figure. The median and maximum bounce loss cone angles for the event shown in Figure 1 are indicated by vertical dashed lines. (c) Predicted enhancements in precipitation as a function of time in response to a perturbation in $|B|$ in the equatorial plane due to a compressional ULF wave. Different scenarios are demonstrated—one where global ULF compressional oscillations lead to an enhanced loss cone at all local times and one where either localization of the ULF waves or an external source of drifting electrons, leads to modulation in the precipitation. The bottom trace indicates the predicted riometer signature given the likely ionospheric response to energetic particle precipitation.

shell up to $\alpha_{G, \max}$ will not necessarily be fully depleted after a single drift period. For localized ULF wave activity, the ULF modulated precipitation signature would be maintained as long as the ULF wave was maintained, and until the flux in each drift shell is fully depleted.

The expected precipitation signatures also depends upon whether electrons with pitch angles close to the edge of either the typical or enhanced BLC are replenished from elsewhere in the magnetosphere, that is, there are additional processes providing a source of electrons on particular drift shells (e.g., the source/seed populations discussed in Jaynes et al., 2015). Substorm injections (e.g., Reeves et al., 1990) and enhanced convection (e.g., Walach & Milan, 2015) can be responsible for the refilling of drift shells. Electron flux can also be replenished through local wave-particle interaction processes (e.g., Summers & Thorne, 2003).

Note that precipitation as measured in the ionosphere by a riometer or any other instrument which senses atmospheric ionization changes will not necessarily depend upon time in the same manner as the precipitating flux. The ionospheric recovery times for the conductivity changes must also be taken into account (Clilverd et al., 2007; Rodger et al., 2007). In this instance, each periodic enhancement in the precipitation flux magnitude would have an associated ionospheric decay time, such that additional pulses of precipitation would add to the previous ionospheric enhancement. ULF modulation in the riometer signal would therefore appear as only a small perturbation on a background enhancement as shown in Figure 4b. In addition, a long ionospheric decay time relative to the period of wave would result in the ULF modulation of the riometer signal having a slightly lower frequency response than the original ULF wave. We propose this simple explanation for the results shown in Figure 2: a 4–5 mHz precipitation signature is observed by BARREL, but a slightly lower frequency signature is observed in the precipitation as measured by a ground-based riometer.

Naively, for an isotropic distribution, one might expect that a given percentage increase in α_G might result in a similar percentage increase in the amount of precipitating flux. However, magnetospheric electron distributions are not generally isotropic with respect to pitch angle, particularly close to the loss cone (e.g., Gu et al., 2011). Typically, electron flux at a constant energy varies as $f = f_0 \sin^n \alpha$ where n can take a range of values, for example, $n = 0, 0.1, 0.25, 0.5, 1, 2, 3 \dots$ and f_0 indicates the value of the flux at 90° . For example, $n = 0$ would correspond to the naive isotropic assumption discussed above. However, Figure 4c shows how the pitch angle variations due to compressional ULF waves can drive increased precipitation for increasing values of n , using the compressional wave example shown in Figure 1, where $\alpha_G = 2.8^\circ$ and $\alpha_{G, \max} = 3.3^\circ$. From Figure 4c, it can be seen that varying the shape of the pitch angle distribution close to the loss cone can drive significantly more precipitation loss than that implied by the given percentage increase in α_G . For a close to isotropic distribution, that is, for n values between $n = 0$ and $n = 0.5$, an 18% increase in α_G would render a similar ~18% increase in precipitation. However, if the shape of the PSD is closer to the $n = 3$ example, an 18% increase in α_G would render significantly larger percentage increase in precipitating flux, closer to a 100% increase.

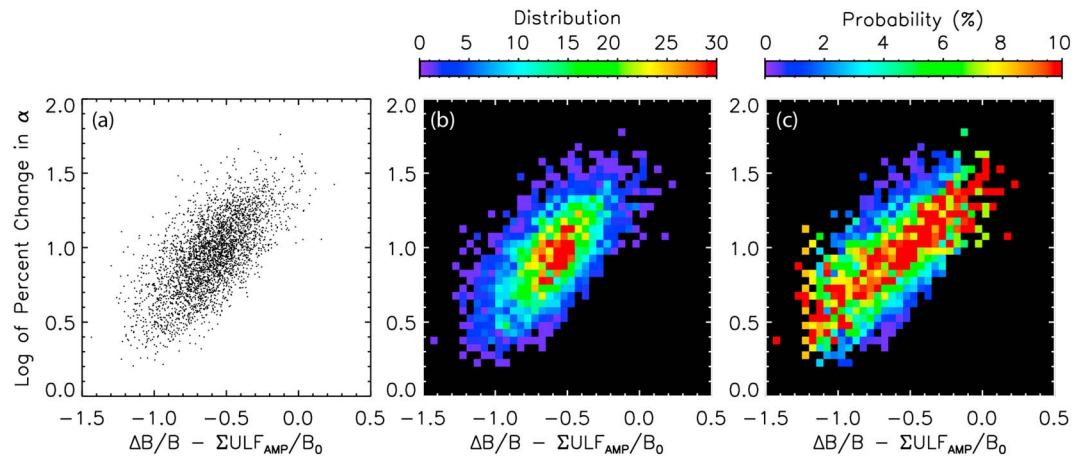


Figure 5. Statistical analysis of GOES West compressional ULF wave events from 1995–2008. (a) A scatterplot of the percent change in α_G against $\Delta B/B_0$ on a logarithmic scale. (b) A two dimensional histogram of the events in Figure 1a, where color represents the number of events in 0.1×0.1 bins, in log space. (c) The probability distribution function (PDF) of observing a given percent change for a given $\Delta B/B_0$, also in 0.1×0.1 bins on a log scale, such that each column adds up to 100%.

6. Statistical Results of GOES Bounce Loss Cone Variations

We employ 14 of geosynchronous Geostationary Operational Environmental Satellite (GOES) magnetometer measurements at 1 min cadence (Singer et al., 1996) to statistically study the variation in the BLC during compressional ULF wave events. Since the GOES satellites are in the geographic equatorial plane, we limit our statistical analysis to satellites located at the GOES West location, since these satellites are closer to the magnetic equator than their GOES East counterparts. As in the previous section, we calculate the variation in BLC using equation (2). The equatorial magnetic field strength B_G is obtained from the GOES measurements, and the ionospheric magnetic field strength B_I is estimated from the IGRF. In order to compile a large database of compressional wave events, we use a 14 year (1995–2008 years) database of GOES data (Ozeke et al., 2012). We limit our analysis to the dayside magnetosphere (06–18 MLT) to concentrate specifically on ULF wave activity and avoid the large-scale topological changes associated with magnetospheric substorms that occur on time scales in the ULF wave band. However, we note that, in principle, our analysis is also relevant to any significant and localized geomagnetic field magnitude variation (as discussed in Section 7).

We define a localized compressional ULF wave event as a quasi-periodic modulation in the magnetic field magnitude above a given amplitude threshold during a 1 h analysis window. We calculate the wave amplitude from the power spectral density at each frequency and identify discrete peaks above a 2 nT threshold using a peak finding algorithm. The 2 nT threshold minimizes the chance of the detection of any sudden impulses, or small ULF wave packets, using the same approach adopted in Watt et al. (2011). Any 1 h window with a discrete peak is flagged as an event containing a compressional ULF wave. In order to avoid overlapping windows or double counting, if the hour analyzed contains a compressional ULF wave that fits this criteria within it, this 1 h analysis window is shifted by an hour. If the hour analyzed does not contain a compressional ULF wave, the analysis window is stepped by 15 min in order to identify the highest number of unique ULF wave events possible. Finally, any hour for which the GOES-measured magnetic field contained a geosynchronous $B_{Z\text{GSM}} < 30$ nT was considered to be potentially affected by magnetopause encounters and were thus discarded (cf. Watt et al., 2011). In total, through this approach we find 3,591 compressional wave events that satisfy our criteria over this 14 year period.

For each of the 3,591 identified events, we determine the median and maximum magnetic field magnitude from the GOES 60 min observations, as well as the median and maximum BLC angle $\alpha_{G,\text{MAX}}$ during the hour. In order to determine the relationship between the equatorial loss cone variations with compressional ULF wave activity, we express the percentage change in the BLC (i.e., the maximum change in BLC as a fraction of the median BLC angle) during an hour as a function of compressional wave amplitudes normalized to the background magnetic field magnitude ($\Delta B/B_0$). Note that the maximum magnetic field strength observed at GOES is equivalent to $B_G = B_0 + \Delta B$, and so $\Delta B/B_0$ is a direct measure of the ULF wave amplitude, but not a direct measure of the change in the loss cone.

Figure 5a shows the ULF wave amplitudes as a function of background field strength and their corresponding change in α_G , on a log-log scale. There is a strong linear relationship between α_G and dB/B_0 , demonstrating that the changes in α_G are indeed linearly related to the fractional change in the magnetic field magnitudes from localized compressional wave activity. Figure 5b reinforces this relationship by displaying a two-dimensional histogram of these points. Finally, Figure 5c shows a probability distribution function (PDF) of these events as a function of dB/B_0 , where each (vertical) column sums to 100%. Figure 5c shows that there is a strong linear correlation between the size of the compressional wave activity and a most likely given change in the equatorial loss cone. From Figure 5 it can be seen that in the 14 year period studied there are certainly events whereby a narrowband ULF fluctuation occurs that is of order the background magnetic field strength, and which would correspond to around 50% increase in the size of the ambient BLC. Although fluxes are small at these small pitch angles relative to the core radiation belt population which have pitch angles closer to 90° , we discuss how a direct ULF modulation of the BLC can provide additional precipitation.

7. Discussion

Traditionally, ULF waves are not considered a direct precipitation mechanism for energetic electrons, and instead, the ULF modulation of VLF growth rates is invoked to explain precipitation modulated at ULF frequencies (Coroniti & Kennel, 1970). This is despite clear observational links between ULF magnetic field oscillations and a variety of auroral (Rae et al., 2014), riometer (Spanswick et al., 2005) and bremsstrahlung-related (Breneman et al., 2015; Halford et al., 2015) electron precipitation signatures. A primary reason for this is that, essentially, global-scale ULF wave fields vary much more slowly than electron bounce times and therefore cannot force bouncing electrons to violate their second adiabatic invariant (e.g., Olson et al., 1980). However, in the case where localized ULF wave fields exist only for a fraction of an electron's drift orbit, it is likely that drifting electrons would rapidly encounter magnetic fields that are not varying smoothly or slowly enough to satisfy conservation of the third adiabatic invariant.

Previous work has focused upon resonant global ULF wave processes such as field line resonance-driven auroral particle precipitation (e.g., Milan et al., 2001; Rae, Mann, Dent, et al., 2007; Rae et al., 2014; Rankin et al., 2005, 2007; Samson et al., 1991, 1996, 2003; Xu et al., 1993), as opposed to any direct modulation of the conditions for particle precipitation by the ULF wave itself. Under these circumstances, it is largely electrons with energies less than a few keV that are involved in the Field Line Resonance (FLR)-electron interaction. FLRs have been shown to be linked to periodic auroral arc structuring (e.g., Rae, Mann, Dent, et al., 2007; Samson, 1994; Samson et al., 1991, 1996), are capable of modulating existing auroral arcs (e.g., Lotko et al., 1998), or are directly powering auroral displays via parallel electric fields accelerating auroral energy electrons (e.g., Rankin et al., 2005, 2007). More complex auroral structuring can also be explained as a result of two harmonically related FLRs that result of field-aligned current element "braiding" (Milan et al., 2001). However, it is unlikely that electrons above approximately keV energies could be accelerated in the field-aligned direction in any of these scenarios, as toroidal mode FLRs have no compressional component, although they have recently been postulated to play a secondary role (e.g., Motoba et al., 2013).

At electron energies above approximately keV, a plethora of observations exist that link ULF waves in ground magnetometer and riometer absorption (e.g., Anger et al., 1963; Beharrell et al., 2010; Brown, 1964; Heacock & Hunsucker, 1977; Olson et al., 1980; Rae, Mann, Dent, et al., 2007; Roldugin & Roldugin, 2008; Spanswick et al., 2005; Ziauddin, 1960). Spanswick et al. (2005) used statistics of NORSTAR riometer measurements to investigate the relationship between Pc5 wave power observed in riometer data and FLRs observed in ground magnetometer data, finding that when significant ULF wave power was observed in riometer absorption, there was always generally a corresponding Pc5 wave signature in ground magnetometer data. In addition, it was found that FLR Pc5 activity was more efficient at producing the riometer modulation than non-FLR Pc5 activity. Spanswick et al. (2005) concluded that the most likely scenario was that when a suitable energetic electron population in the inner magnetosphere was present, resonant ULF waves could play a role in their precipitation but that pitch angle scattering from some other plasma wave (for example whistler mode waves) was required as well before both ground magnetometer and riometer would observe a ULF modulated signal. From a theoretical perspective (Coroniti & Kennel, 1970; Watt et al., 2011), a variation in magnetic field strength (i.e., a compressional component of the wave magnetic field) is required to

modulate VLF growth rates. Moreover, a variation in magnetic field strength that is in direct antiphase with the cold plasma number density (cf. Li et al., 2011 and Watt et al., 2011) is required to modify VLF growth rates sufficiently to account for the changes in precipitation. Since FLRs are where energy from a propagating compressional wave couples to the shear mode (Samson et al., 1992), it is not clear whether the wave properties necessary to modify VLF wave growth rates are satisfied in an FLR. We postulate in this study that the reason is that the compressional component of the FLR driver may be the direct generator of ULF-modulated riometer absorption, rather than the action of the FLR itself. Specific case studies of a simultaneous compressional ULF wave and an FLR have been presented in the literature (e.g., Rae, Mann, Dent, et al., 2007). The evidence presented here suggests that an alternative explanation for the modulation of ULF-precipitation in this and other cases is the direct modulation of the equatorial BLC by the compressional component of the ULF wave.

Direct enhancement of the local equatorial bounce loss cone enhances other mechanisms for precipitation of electrons from the magnetosphere. Brito et al. (2012, 2015) used MHD simulations to show that the radial displacement of electrons due to global-scale compressional ULF waves can itself lead to enhanced precipitation. The radial motion of the electrons encountering a compressional ULF wave causes their trajectories to move closer to the Earth into a stronger magnetic field, where the loss cone is larger. Additionally, the inward radial motion of the electrons leads them into regions with shorter field lines, where they gain perpendicular energy due to conservation of the first adiabatic invariant and parallel energy due to conservation of the second adiabatic invariant. In this paper we show additional precipitation effects if these ULF wave fields are localized; under these circumstances, the loss cone is locally and abruptly modified as a function of time through the action of the compressional ULF waves themselves. Compressional magnetospheric ULF waves at geosynchronous orbit can have sufficient amplitudes to locally enhance the size of the bounce loss cone by over 50%. Of course, this effect (shown in Figure 4) depends upon the ratio of the wave amplitude to the background magnetic field, and the background magnetic field varies as a function of radial distance r as roughly r^{-3} . Our observations are confined to geosynchronous orbit, to which the majority of riometer absorption modulation also map (Spanswick et al., 2005). However, closer to the heart of the outer radiation belts at $L = 4-5$, where the field strength increases and ULF modulated precipitation is often seen (e.g., Breneman et al., 2015; Brito et al., 2015), the fractional enhancement in the traditional loss cone will become smaller for a given ULF wave amplitude. However, again, there are competing effects to be considered, given that equatorial BLC also increases with decreasing radial distance; this means that both the equatorial loss cone and compressional ULF wave amplitudes must be computed across all radial distances in order to determine their effect across the entire outer radiation belt region.

The direct enhancement of the BLC by a localized compressional ULF wave will also greatly enhance any precipitation mechanism that is due to pitch angle scattering. Whistler mode chorus (see Millan & Thorne, 2007, for a comprehensive review) is often invoked to pitch angle scatter radiation belt electrons outside of the plasmapause, with plasmaspheric hiss acting in a similar way inside of the plasmapause (e.g., Breneman et al., 2015). Electromagnetic ion cyclotron (EMIC) waves have also been shown to play a role in enhanced relativistic electron precipitation (e.g., Carson et al., 2013; Clilverd, Duthie, et al., 2015; Rodger et al., 2008) through cyclotron resonant interactions. Pitch angle scattering rates depend upon the wavenormal angle and power spectral densities of the whistler mode chorus (e.g., Ni et al., 2011). However, pitch angle diffusion rates for a 30 keV electron at geosynchronous orbit range from 10^{-3} to 10^{-4} s^{-1} , which is comparable to Pc5 ULF wave frequencies. By contrast, inside the plasmasphere, plasmaspheric hiss can have pitch angle diffusion rates of 10^{-2} to 10^0 s^{-1} (e.g., Breneman et al., 2015). As EMIC waves are the left-hand counterpart of whistler mode waves, there would be no reason not to expect that EMIC wave growth would also be affected by large-amplitude monochromatic changes of the magnetic field magnitude and number density either, as Loto'aniu et al. (2009) discussed.

Whistler mode precipitation will be enhanced by a temporally varying loss cone for two reasons. First, pitch angle scattering increases the flux at pitch angles close to the bounce loss cone, while our mechanism increases the size of the bounce loss cone thus leading to enhanced loss. Second, by increasing the amount of precipitation, the anisotropy that drives whistler mode waves unstable may also increase during different phases of the wave leading to either enhanced wave amplitudes or longer lifetime and thus increased precipitation. This explanation provides additional insight into events discussed by Halford et al. (2015) and Breneman et al. (2015), where ULF waves were proposed to be modulating the resonance condition,

leading to both an enhanced background level of precipitation and modulation at ULF frequencies. This symbiotic relationship, comparable to that espoused by Baumjohann et al. (2000) regarding whistler mode waves inside mirror mode waves in the dawn sector magnetosphere (e.g., Rae, Mann, Watt, et al., 2007), is ripe for further exploration.

Most importantly with regard to the results in this paper, it remains to be established whether ULF waves and ULF-modulated precipitation are observed without the presence of whistler mode chorus (e.g., Nishimura et al., 2013) or plasmaspheric hiss (e.g., Breneman et al., 2015). Our results suggest that such a precipitation mechanism is possible in theory and offers a suggested mechanism for the case study shown in Figure 2, in a region typically associated with limited VLF wave activity and where no enhanced precipitation outside of the compressional ULF wave region is observed. Future work will explore the wealth of ground and space-based observations available in the Van Allen Probe era to identify whether ULF-modulated precipitation can indeed exist without any VLF pitch angle scattering mechanism.

The localization of the ULF pulsation appears to be very important for the precipitation of electrons. Localized dayside ULF wave fields are often referred to as drift-bounce resonance or “storm time Pc5 waves” and are thought to be driven by unstable ion distributions emanating from magnetotail injections (e.g., Lanzerotti et al., 1969; Southwood et al., 1969; Wright et al., 2001). They are detected mainly in the afternoon/evening sector of the magnetosphere (e.g., Anderson et al., 1990). Our case study (Figure 1) shows ULF compressional wave activity in the afternoon sector. However, in our statistical study, we show observations of compressional pulsations at geosynchronous orbit across all of the dayside magnetosphere, and so other generation mechanisms may also play a role (e.g., mirror mode waves in the dawn sector) (e.g., Constantinescu et al., 2009; Liu et al., 2016; Rae, Mann, Watt, et al., 2007; Vaivads et al., 2001; Zhu & Kivelson, 1994). We recognize that mode structure along the field is important for determining changes in BLC at any point along the geomagnetic field due to ULF wave modulation (e.g., Ozeke & Mann, 2004; Perry et al., 2005; Takahashi et al., 1987). Indeed, it is interesting to note that localized compressional waves (e.g., Liu et al., 2016; Rae et al., 2007) would act to trap particles primarily with pitch angles closer to 90° in magnetic bottles via the mirror effect. Hence, trapping of high pitch angle particles may act in concert with the enhanced precipitation of low pitch angle particles. To confirm that our assumptions are correct, future work will use electric, magnetic, and plasma density measurements to characterize mode structure and perform more accurate calculation of the change in the equatorial BLC in each case. Future work will utilize a more realistic three-dimensional magnetospheric wave model (Degeling et al., 2010) where localization of the waves in magnetic local time and realistic field-aligned structures can be reproduced. By doing this, electrons can then be traced through to see how their behavior is modified and the loss cone is modified due to the presence of the localized, compressional ULF waves.

8. Conclusion

This paper explored the potential role of localized compressional ULF waves as a candidate mechanism to directly enhance electron precipitation by simple modulation of the local bounce loss cone. Periodic magnetic compression of a localized magnetospheric region on long period time scales relative to the gyration and bounce allows conservation of the first and second adiabatic invariants but a clear opportunity to violate the third invariant. We demonstrate that the change in pitch angle of a given electron due to the conservation of the first and second invariants (Figure 3) is far smaller than the change in loss cone due to the localized ULF wave (Figures 2 and 4). In this way, we show that localized compressional ULF waves can directly contribute to electron precipitation.

Previous studies (e.g., Brito et al., 2012, 2015) have focused on the role of global compressional ULF waves in driving radial motion of radiation belt electrons to additionally precipitate. Direct modulation of the loss cone differs from any other mechanism traditionally invoked to explain, in particular, radiation belt electron losses during active times. ULF modulation of the bounce loss cone would be enhanced during active times, such as during a storm main phase where compressional ULF wave amplitudes are largest and up to ~ 2 orders of magnitude higher than statistically found (Murphy et al., 2015). We note here that this mechanism will also operate across all electron energies, but with subtly different observational characteristics, potentially explaining how low-energy auroral (e.g., Samson et al., 1991), keV (e.g., Spanswick et al., 2005), hundreds of keV (e.g., Breneman et al., 2015), and MeV (e.g., Foat et al., 1998; Millan et al., 2002; Woodger et al.,

2015) electron energies can all display ULF modulation, which no one single gyroresonant process can be invoked to explain.

Statistically, we show that large-amplitude highly localized compressional ULF waves can modulate the loss cone by $\pm 20\%$, which in turn allows a significantly greater fraction of the electron PSD to precipitate than previously thought. Importantly, this requires no other wave-particle interaction to cause precipitation of energetic electrons with pitch angles outside of the traditional loss cone, although this mechanism would be enhanced by local pitch angle scattering to refill the near-loss cone population. Hence, what fraction of this distribution is locally precipitated depends upon the strength of the perturbation, local magnetic field magnitude, shape of the pitch angle distribution close to the traditional loss cone, and the nature of any additional sources of energetic electrons into the ULF region (e.g., substorm injections) or near the loss cone (e.g., pitch angle scattering due to whistler mode waves). Since this mechanism does not require the presence or the absence of VLF wave-particle interaction, we simply point out that localized compressional waves should be considered along with other precipitation mechanisms within the current literature.

We show direct evidence of ULF wave modulated precipitation across the energy ranges measured by riometers and BARREL, which is spatially correlated with localized large-amplitude ($\sim 15\%$ of the ambient magnetic field) compressional ULF wave activity in the afternoon sector. Within this case study we show clear evidence that the ULF wave fields are spatially localized, although we note here that there is no means to investigate other precipitation sources for this case which would be expected from pitch angle scattering mechanisms such as whistler mode chorus or plasmaspheric hiss.

We believe that this mechanism warrants further study to determine whether compressional waves and ULF modulated losses are indeed causally related and whether such loss can routinely occur independently of gyroresonant interactions.

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