

1 **Characterising the Land Surface Phenology of Africa using 500 m MODIS EVI**

2 **Abstract**

3 Vegetation phenological studies at different spatial and temporal scales offer better understanding of
4 the relationship between the global climate and the global distribution of biogeographical zones.
5 These studies in the last few decades have focussed on characterising and understanding vegetation
6 phenology and its drivers especially using satellite sensor data. Nevertheless, despite being home to
7 17% of the global forest cover, approximately 12% of the world's tropical mangroves, and a diverse
8 range of vegetation types, Africa is one of the most poorly studied regions in the world. There has
9 been no study characterising land surface phenology (LSP) of the major land cover types in the
10 different geographical sub-regions in Africa, and only coarse spatial resolution datasets have been
11 used for continental studies. Therefore, we aim to provide seasonal phenological pattern of Africa's
12 vegetation and characterise the LSP of major land cover types in different geographical sub-regions in
13 Africa at a medium spatial resolution of 500 m using MODIS EVI time-series data over a long
14 temporal range of 15 years (2001 – 2015). The Discrete Fourier Transformation (DFT) technique was
15 employed to smooth the time-series data and an inflection point-based method was used to extract
16 phenological parameters such as start of season (SOS) and end of season (EOS). Homogeneous
17 pixels from 12 years (2001 – 2012) MODIS land cover data (MODIS MCD12Q1) was used to
18 describe, for the first time, the LSP of the major vegetation types in Africa. The results from this
19 research characterise spatially and temporally the highly irregular and multi-annual variability of the
20 vegetation phenology of Africa, and the maps and charts provide an improved representation of the
21 LSP of Africa, which can serve as a pivot to filling other research gaps in the African continent.

22

23

24 **Keywords**

25 Climate; Land cover; Ground-based; Remote sensing; Vegetation.

26

27 **1. Introduction**

28 The study of vegetation phenology, which deals with the timing of plant growth stages and their inter-
29 annual variation, can increase our understanding of global climate-vegetation relationships, and in
30 particular can be used to characterise the impact of climate change on terrestrial ecosystem
31 (Chmielewski & Rötzer, 2001; Cleland *et al.*, 2007; Richardson *et al.*, 2013; Broich *et al.*, 2014;
32 Clinton *et al.*, 2014). Consequently, the study of vegetation phenology has received increased
33 attention in recent years, providing detailed characterisation of spatio-temporal changes in terrestrial
34 biogeochemical cycles.

35

36 Ground-based observations of vegetation phenology, offer detailed and fine temporal resolution data
37 for different vegetation types (Rodriguez-Galiano *et al.*, 2015b). However, these observations are
38 limited in spatial coverage (Studer *et al.*, 2007). On the other hand, satellite-based remote sensing
39 techniques, which measure *land surface phenology* (LSP) (defined “*as the seasonal pattern of*
40 *variation in vegetated land surfaces observed from remote sensing*” (Friedl *et al.*, 2006)), offer wide
41 spatial coverage, and can monitor the inter-annual variability of vegetation dynamics in areas without
42 ground data (Julien & Sobrino, 2009; Guan *et al.*, 2013; Zhang *et al.*, 2014; Rodriguez-Galiano *et al.*,
43 2015a). These techniques also offer the capability of quantifying vegetation response to climate
44 variability (Ma *et al.*, 2008; Zhu *et al.*, 2012; Broich *et al.*, 2014; Guan *et al.*, 2014b). Other
45 advantages can be seen in studies covering ecosystem processes and diversity, for example, in studies
46 of the phenology of bird communities from space (Cole *et al.*, 2015), and understanding transhumance
47 patterns (Butt *et al.*, 2011; Brotem *et al.*, 2014).

48

49 In the northern high latitude regions such as Europe and North America, numerous studies have
50 detailed the characteristics of vegetation phenology at both fine and coarse temporal and spatial
51 resolutions, either through ground-based measurements or by remote sensing techniques
52 (Chmielewski & Rötzer, 2001; Zhang *et al.*, 2004; Menzel *et al.*, 2006; Ganguly *et al.*, 2010; Wu *et*
53 *al.*, 2012; Jeganathan *et al.*, 2014; Walker *et al.*, 2014; Rodriguez-Galiano *et al.*, 2015a). There are
54 also robust ground-based observation networks in these regions. Examples of such networks are: the

55 US National Phenology Network, the Woodland Trust, UK, International Phenological Gardens (IPG)
56 in Europe and the German phenological network (Chmielewski *et al.*, 2004; Graham *et al.*, 2010;
57 Boyd *et al.*, 2011; Zhang *et al.*, 2012; Menzel, 2013; Wolkovich *et al.*, 2014).

58

59 In Africa, there have also been several phenological studies, both ground-based and satellite-based
60 (Adole *et al.*, 2016). However, despite being home to 17% of the world's forest cover (Food and
61 Agriculture Organization of the United Nations, 2010), approximately 12% of the world's tropical
62 mangroves (Giri *et al.*, 2010; Donato *et al.*, 2011), and with a diverse range of vegetation types
63 (Figure 1), compared to other continents, the number of phenological studies in Africa is very limited
64 (Adole *et al.*, 2016). Similarly, unlike other regions, there are no phenological networks in Africa
65 (Adole *et al.*, 2016).

66

67 A recent systematic review by Adole *et al.* (2016) revealed that of 9,566 articles on vegetation
68 phenology globally, only 130 focused on Africa. Moreover, despite the advances in LSP, particularly
69 with the availability of fine spatial resolution data, and knowing that at coarser spatial resolutions
70 phenological information may be misread (Fisher & Mustard, 2007), only 15 studies evaluated LSP at
71 a continental scale using coarse spatial resolution (ranging from 1 to 8 km) data (Adole *et al.*, 2016).
72 Adole *et al.* (2016, Table 1) found that studies over longer periods used coarse spatial resolution
73 datasets while those with a shorter duration of five years or less commonly used a spatial resolution of
74 1 km. Additionally, the temporal resolutions of most of these studies were relatively coarse (10 – 16
75 day), thereby increasing the potential for errors in vegetation phenology estimation (Zhang *et al.*,
76 2009). Although the MODIS Land Cover Dynamics product (MCD12Q2) provides global LSP
77 information at a spatial resolution of 500 m there are large uncertainties, and sometimes unrealistic
78 LSP parameter values, associated with this product (Ganguly *et al.*, 2010; Vintrou *et al.*, 2012) and,
79 thus, may not be reliable for detail characterisation of LSP. Also, this product which was last released
80 in 2012 is not as recent as other MODIS data and does not benefit from the recent reprocessing of
81 MODIS data products. Based on these findings, we have summarized the identified research gaps
82 which are relevant to this below:

83

84 (1) There has been no study characterising LSP of the major land cover types in the different
85 geographical sub-regions in Africa.

86 (2) At a continental scale, only coarse spatial resolution datasets ranging from 1 to 8 km have
87 been used for LSP studies in Africa, and

88 (3) 10 – 16 day temporal resolution datasets were used with the exception of only two studies
89 which used daily datasets, albeit at coarse spatial resolutions of 3 and 5 km (see Table 1).

90

91 In addition to the above highlighted gaps, Africa is known to have complex vegetation dynamics

92 (Favier *et al.*, 2012) and its vegetation types are very distinct in their responses to climatic factors,

93 resulting in great variability in phenological patterns. Although there are generally two major

94 maximum rainfall seasons in Africa (the June-to-August season in the northern latitudes and the

95 December-to-February season in the southern latitudes) (Griffiths, 1971), the distribution of these

96 seasons varies considerably across the continent. This can be seen in the rainfall seasons in the

97 extreme north falling into the December-to-February season and southwestern Africa falling into the

98 June-to-August season (Griffiths, 1971). Also, the Horn of Africa, which is greatly affected by the

99 Inter-Tropical Convergence Zone (ITCZ) (Thompson, 1965), and the Guinea coast in West Africa

100 exhibit a unique double peak or two seasonal rainfall patterns (Herrmann & Mohr, 2011; Liebmann *et*

101 *al.*, 2012). This variation in the climate of the different geographical sub-regions in Africa (see Figure

102 1) plays a significant role in the vegetation dynamics in these regions, hence the requirement to

103 characterise LSP regionally.

104

105 In view of the above, it is apparent that there is a need to provide more detailed LSP information for

106 the African continent. This detailed LSP information is likely to be very important in climate-

107 vegetation modelling and can potentially help in increasing our understanding of carbon, energy and

108 water cycles, characterisation of soil-vegetation-atmospheric feedbacks, and predictive phenology

109 modelling. This would also aid in-depth monitoring of agricultural production and livestock

110 management practices which would be unique to the different geographical regions in African

111 farmlands and rangelands. Therefore, the aim was to characterise the spatial distribution of LSP in
112 Africa using medium spatial and temporal resolution (500 m, 8-day) MODIS EVI time-series data
113 with a long temporal range of 15 years (2001 – 2015). The specific objectives were to:

114

- 115 (1) establish a baseline of LSP over Africa at a fine spatial resolution of 500 m
- 116 (2) determine the latitudinal variation and inter-annual variability of LSP in Africa at a fine
117 spatial resolution of 500 m compared to previous work.
- 118 (3) Using these data, characterise the LSP of the major land cover types in different
119 geographical sub-regions in Africa, and
- 120 (4) demonstrate the advantages of the medium spatial resolution of 500 m.

121

122 Comprehensive ground-based validation of the LSP maps from this research is not possible presently
123 due to the absence of a broad-scale ground-based observation network across the African continent.
124 Therefore, comparisons were made between the estimated LSP and previous vegetation phenology
125 studies, and the ground-based vegetation phenology data for the few areas for which data were
126 available.

127 **Table 1: Number of LSP studies in Africa undertaken at a continental scale with the Advanced Very High**
 128 **Resolution Radiometer (AVHRR), the Moderate-resolution Imaging Spectroradiometer (MODIS) and**
 129 **the Spinning Enhanced Visible and InfraRed Imager (SEVIRI) sensors.**

Authors	Period	Temporal frequency	Sensor	Spatial Resolution (km)	Index	Research findings
Brown <i>et al.</i> (2010)	1981 - 2008	15-day	AVHRR	8	NDVI	LSP is significantly affected by climate oscillations
Camberlin <i>et al.</i> (2007)	1981 - 2000	15-day	AVHRR	8	NDVI	Significant correlation between annual NDVI values and rainfall variations
Guan <i>et al.</i> (2013)	2000 - 2012	16-day	MODIS	5	EVI	Strong seasonality coupling between vegetation function and structure which is controlled by precipitation in tropical forest
Guan <i>et al.</i> (2014a)	2007 - 2011	Daily	SEVIRI	3	LAI	New algorithm that can be used to derive LSP across other carbon related datasets
Guan <i>et al.</i> (2014b)	2000 - 2011	Daily	MODIS	5	NDVI	Distinct responses of African savannas and deciduous woodlands LSP to rainy season
Jönsson & Eklundh (2002)	1982 - 2000	10-day	AVHRR	8	NDVI	New algorithm for estimating LSP
Jönsson & Eklundh (2004)	1998 - 2000	10-day	AVHRR	8	NDVI	TIMESAT programme for processing time-series of satellite data
Justice <i>et al.</i> (1989)	1981	15-day	AVHRR	8	NDVI	Microwave polarization difference temperature (MPDT) relationship with NDVI seasonal variations
Linderman <i>et al.</i> (2005)	2000 - 2004	16-day	MODIS	1	EVI	Interannual changes in vegetation activity not linked to shifts in phenology
McCloy & Tind (2011)	1982 - 2008	15-day	AVHRR	8	NDVI	Changes in vegetation phenology overtime
Stroppiana <i>et al.</i> (2009)	1990 - 2002	10-day	AVHRR	8	NDVI	A new anomaly indicator (AI) for abstract environmental status assessment and monitoring using phenological data
Vrieling <i>et al.</i> (2008)	1981 - 2006	15-day	AVHRR	8	NDVI	Temporal trend analysis of crop phenology showing both positive and negative yield across Africa
Vrieling <i>et al.</i> (2011)	1982 - 2006	15-day	AVHRR	8	NDVI	Understanding variability and trends in seasonal cumulated NDVI (cumNDVI) is important in characterising farming systems
Vrieling <i>et al.</i> (2013)	1981 - 2011	15-day	AVHRR	8	NDVI	The variability and trend of length of growing period (LGP) in Africa
Zhang <i>et al.</i> (2005)	2000 - 2003	16-day	MODIS	1	EVI	Vegetation green-up strongly dependent on rainfall seasonality in Africa

131 **2. Methodology**

132 **2.1. Data acquisition and pre-processing**

133 **2.1.1. MODIS land surface reflectance data**

134 MODIS data, which are significantly improved in terms of spatial and spectral resolution, atmospheric
135 corrections, cloud screening and sensor calibration (Soudani *et al.*, 2008) compared to AVHRR, were
136 acquired for this study. 16 years (18 Feb 2000 – 24 June 2016) of 44 MODIS/Terra Surface
137 Reflectance 8-Day L3 Global 500 m SIN Grid V005 data (MOD09A1) tiles were downloaded from
138 NASA’s LP DAAC (<https://lpdaac.usgs.gov/>). These data provide a long temporal record of a
139 medium spatial resolution product. Apart from the seven spectral bands [bands 1 (620-670 nm), 2
140 (841-876 nm), 3 (459-479nm), 4 (545-565 nm), 5 (1230-1250 nm), 6 (1628-1652 nm), and 7 (2105-
141 2155 nm)], this product has an additional 32-bit Quality Assurance (QA) layer which was used for
142 quality assessment. To filter out residual atmospheric and sensor effects, only pixels with the highest
143 quality of band 1 – 7 which had adjacency and atmospheric correction performed, and all possible
144 corrections of MODIS land Quality Assessment (MODLAND QA), were retained. (see
145 https://lpdaac.usgs.gov/sites/default/files/public/modis/docs/MODIS_LP_QA_Tutorial-3.pdf for
146 details on the QA assessment procedures).

147

148 The Enhanced Vegetation Index (EVI), which overcomes the saturation problems of the Normalized
149 Difference Vegetation Index (NDVI), especially in areas with large amounts of vegetative biomass
150 (Huete *et al.*, 2002), was selected as the vegetation index for use in this study. It was developed with
151 the inclusion of the blue reflectance band (B) to correct for atmospheric and soil background
152 influences (Huete *et al.*, 2011; Rowhani *et al.*, 2011), and is derived according to the following
153 equation:

154

155
$$EVI = G * \frac{(NIR - Red)}{(L + NIR + C1 * Red - C2 * Blue)}$$

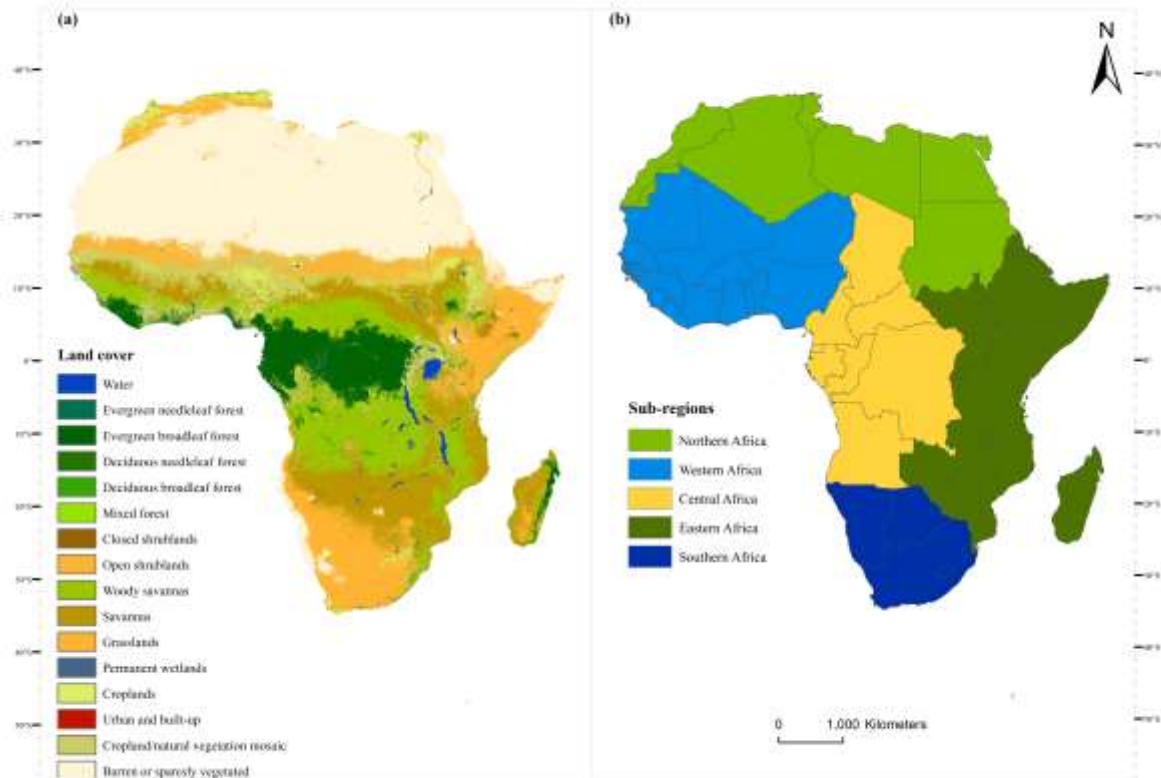
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157 The coefficients of the EVI equation are $L=1$ (canopy background adjustment factor); $C1= 6$ and $C2 =$
158 7.5 (aerosol correction factors); and $G = 2.5$ (gain factor) (Huete *et al.*, 2002, 2011; Reed *et al.*, 2009;
159 Rowhani *et al.*, 2011).

160

161 **2.1.2. MODIS Land Cover Type data**

162 To represent the land cover of Africa, 12 years (2001 – 2012) of 44 tiles MODIS/Terra Land Cover
163 Type Yearly L3 Global 500 m SIN Grid V005 data (MCD12Q1) (h16v05 to h22v11) were
164 downloaded from NASA’s LP DAAC (<https://lpdaac.usgs.gov/>). This product has five different land
165 cover classification schemes. The 17-class International Geosphere Biosphere Programme (IGBP)
166 global vegetation classification scheme, shown to be the best among the five schemes, was selected
167 for this analysis (Scepan & Estes, 2001; Friedl *et al.*, 2010) (see figure 1).



168

169 **Figure 1:** (a) Land cover map of Africa derived from the 500 m MODIS land cover type product
 170 (MCD12Q1) data for 2012, downloaded from NASA’s LP DAAC (<https://lpdaac.usgs.gov/>). (b) Map
 171 of Africa, showing the five different geographical sub-regions (Griffiths, 1971; United Nations,
 172 2014).

173

174 **2.2. Data analysis**

175 **2.2.1. LSP estimation**

176 To begin LSP estimation, EVI data were stacked into 86 layers (Figure 2) (a layer being one
 177 composite EVI image), which defined a “cycle” to include two years (i.e., July of year 1 to June of
 178 year 3). This is to account for the non-uniform growing seasons across Africa, where start of season
 179 is much earlier in the northern latitudes compared to southern latitudes, ensuring that seasonal
 180 phenological parameters are estimated yearly.

181 Four steps were carried out to estimate LSP from the EVI time-series data (Figure 3).

182 (1) Removal of drop outs in the EVI time-series with a temporal moving average window

- 183 (2) Linear interpolation for gap filling (Dash *et al.*, 2010)
- 184 (3) Data smoothing to further reduce residual noise in data using the inverse Discrete Fourier
- 185 Transform (DFT)
- 186 (4) A search process to find the phenological parameters (e.g., minima in the smoothed time-
- 187 series).

188 The Discrete Fourier Transform (DFT), a frequency-based smoothing technique was applied to the

189 EVI time-series. This method undertakes a frequency decomposition of the temporal profile of a time-

190 series using Fourier analysis and then reconstructs back to the temporal domain via an inverse Fourier

191 transform, in the present case based on only the smoother components (Moody & Johnson, 2001;

192 Atkinson *et al.*, 2012). One major advantage of this technique is the minimal user input, as users need

193 to specify only the number of harmonics required to reconstruct the time-series (Dash *et al.*, 2010). It

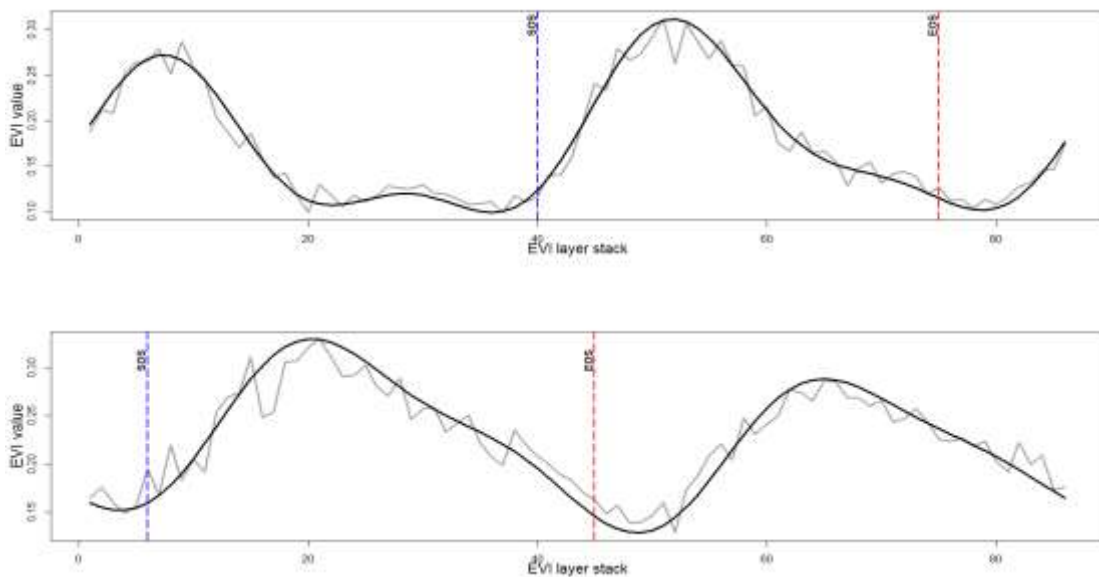
194 has been established that the first two harmonics can adequately represent annual or semi-annual

195 cycles (Jakubauskas *et al.*, 2001). Considering the bimodal seasonality and double cropping

196 agricultural systems found in some parts of Africa, the first six harmonics, as used in Dash *et al.*

197 (2010), were used to generate the smoothed time-series (Figure 2).

198

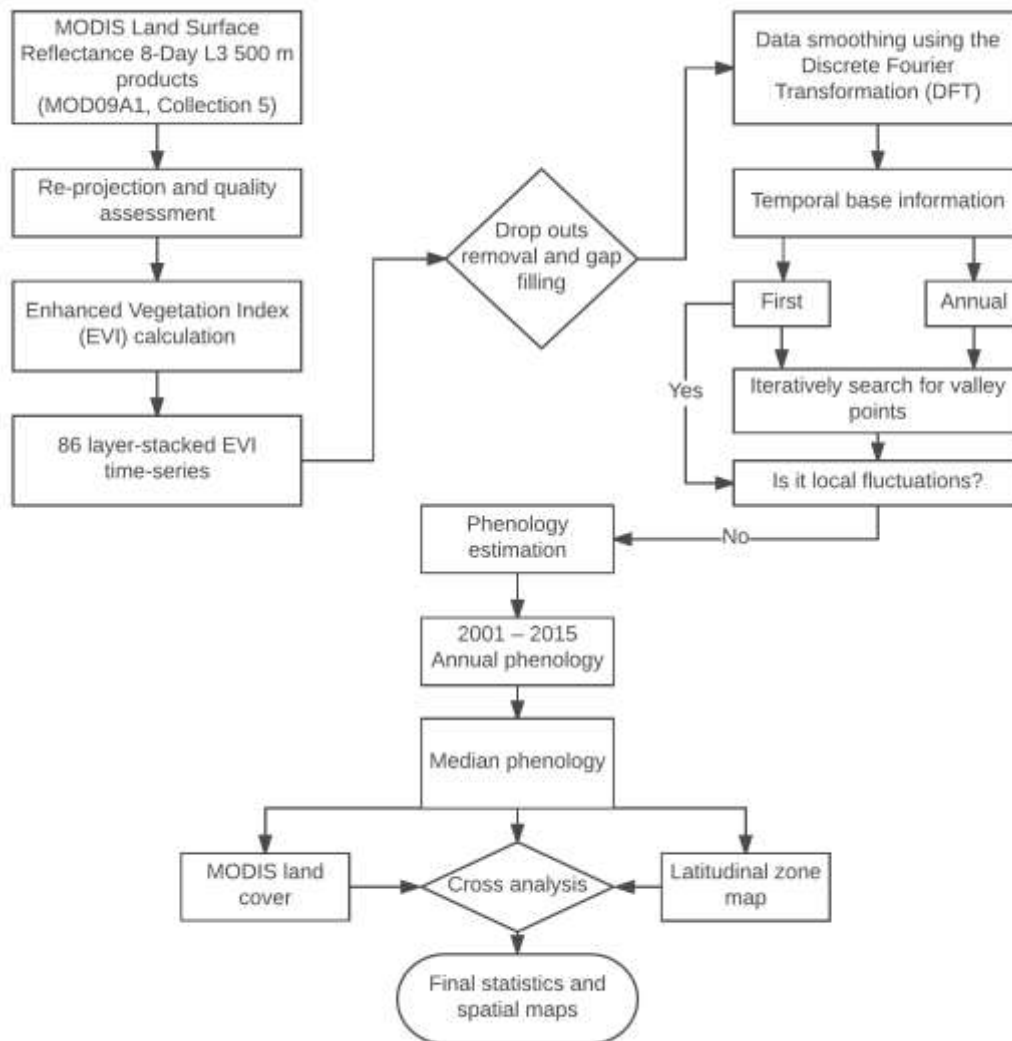


199

200 **Figure 2:** Example of pixels showing the smoothed temporal profile of an 86 layer-stacked EVI time-
201 series in black superimposed on the raw EVI data in grey. Blue dotted lines are the SOS and red
202 dotted lines are EOS estimated for each time-series.

203

204 Finally, LSP parameters were estimated using the inflection point method based on points of maximal
205 curvature in the time-series (Figure 2) (Reed *et al.*, 1994; Moulin *et al.*, 1997; Zhang *et al.*, 2001;
206 Dash *et al.*, 2010). We used an algorithm which departs at the maximum peak, and iteratively
207 searches for valley points (change in derivative value) at the beginning of the growing cycle (Start of
208 Season (SOS), i.e. a change in derivative value from positive to negative) and at the decaying end of
209 the phenology cycle (End of Season (EOS), i.e. a change in derivative value from negative to positive)
210 (Dash *et al.*, 2010; Pastor-Guzman *et al.*, 2018). The length of season (LOS) was determined as the
211 difference between the estimated SOS and the EOS, converted to number of days. The median values
212 for these parameters for the period of 2001 to 2015 were estimated and then converted to their
213 corresponding Julian days (i.e. day of year (DOY)).



214

215 **Figure 3:** Schematic diagram illustrating the research methodology adopted in this study.

216

217 2.2.2. MODIS land cover masking

218 A further reclassification was carried out on the MODIS land cover 17-class International Geosphere

219 Biosphere Programme (IGBP) global vegetation classification scheme, by merging classes with very

220 similar phenological behaviour into broad vegetation classes. For example, evergreen needleleaf

221 forest and evergreen broadleaf forest were merged together to give one class of “evergreen forest”.

222 Pixels belonging to other land cover types that are not vegetation were masked out. Additionally,

223 pixels which remained as the same class over the time-series of 12 years were extracted and used to

224 mask the phenology estimates based on the geographical sub-regions in Africa.

225

226 **2.3. Analysis of LSP**

227 To analyse the variation in phenology with latitude, the majority (i.e. modal values) of LSP
228 parameters were estimated per degree increase in latitude. Thereafter, a simple linear regression
229 model was used to estimate the expected change in phenological parameter per degree increase in
230 latitude (LSP parameters as the dependent variable and latitude as the independent variable) and the
231 significance of the models assessed.

232

233 To determine the inter-annual variability of LSP parameters over the entire time-series, the temporal
234 standard deviation (STD) values for each LSP parameter in each pixel were estimated. A large
235 magnitude of STD can reveal areas that have unstable seasons in Africa. Additionally, to quantify the
236 spatial distribution of LSP parameters across Africa the percentage of pixels of LSP parameters
237 belonging to each land cover type in the different geographical sub-regions was determined. Finally,
238 to demonstrate the effect of spatial resolution, the STD of the SOS values were estimated with spatial
239 resolutions of 1 km, 3 km, 5 km and 8 km obtained by image degradation (linear averaging).

240

241 **3. Results**

242 **3.1. Spatiotemporal variation in vegetation phenological parameters**

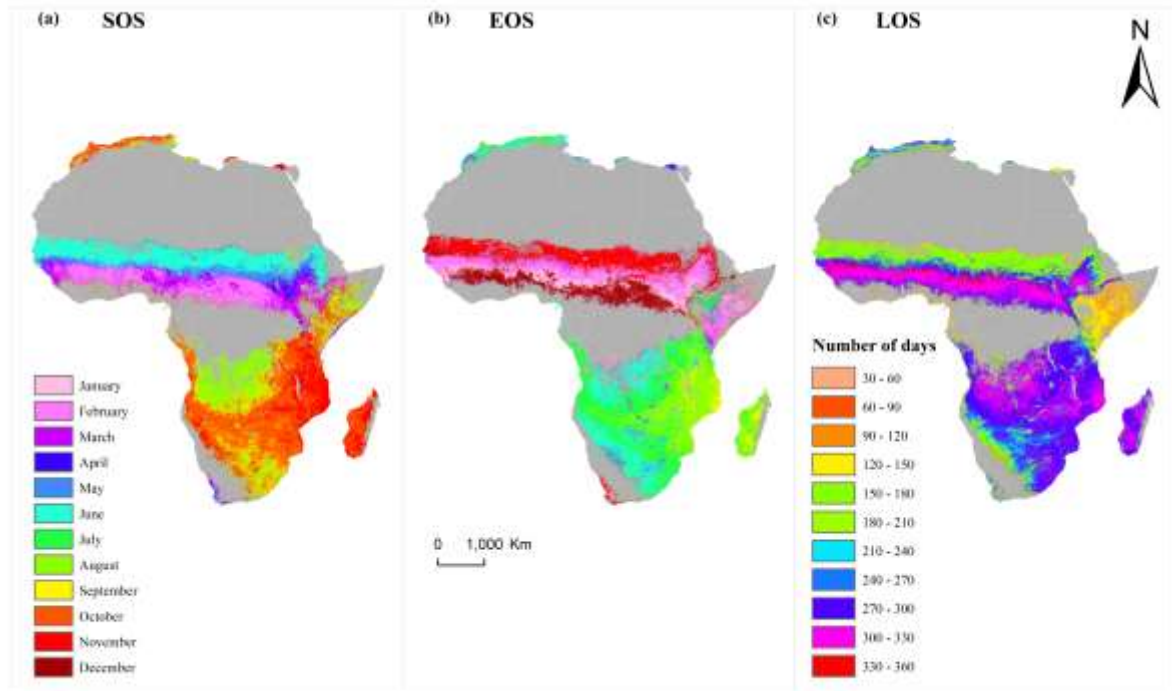
243 Maps produced indicating the median start and end dates, as well as the length of the growing season
244 for the study period 2001-2015 across Africa showed high variability throughout the continent (Figure
245 4). Between the latitudes of 0° and 20°N which covers the Sahel, Sudan and Guinean regions of
246 Africa, the beginning of the growing season (SOS) has a wide range between late February and early
247 August with most SOS estimates occurring in late February and June. The end of the growing season
248 in these regions falls between late November and the following February, with a long growing season
249 of 150 – 310 days. These very long growing seasons have also been observed by Yan *et al.* (2016).
250 However, some parts of Eastern Africa have SOS dates that are between August and October and
251 EOS between late June and August of the following year. Further north, above 27°N , most SOS dates
252 occurred between September and November. The corresponding EOS dates are between May and

253 August. This can be attributed to the different seasonal rainfall patterns observed in the extreme north
254 which begins around September with peaks in December and February (Griffiths, 1971; Liebmann *et*
255 *al.*, 2012). No clear seasonality was detected in most parts of Central Africa, due to the presence of
256 very dense canopies of evergreen forest, and persistent cloud prohibited sufficient cloud free data
257 collection.

258 In contrast to most areas in the north, for the south of Africa, between latitudes 0⁰ and 34⁰S, the
259 majority of SOS dates fell between August and November and corresponding EOS dates between
260 May-June and August of the following year. In the southwestern region, different SOS and EOS dates
261 were observed; February to April for SOS and November to the following year February for EOS.
262 This can be explained by the distinct rainfall pattern observed in this region (rainfall peaks in June to
263 August) (Griffiths, 1971; Liebmann *et al.*, 2012).

264 Bimodality was also observed in the Horn of Africa and some parts of Western Africa particularly in
265 the coast of Guinea (Figure 5). This could be as a result of dual seasonal rainfall patterns, with peaks
266 in April-May and October-November observed in these regions (Herrmann & Mohr, 2011; Liebmann
267 *et al.*, 2012) or artificial bimodality due to residual noise in the EVI data, especially where the
268 bimodality lacks consistency in space and time. Vegetation growth for this second season starts
269 between late August and November and ends between December and February. A shorter LOS of 112
270 – 144 days was also observed in the Horn of Africa for both the first and second seasons (see figure 4
271 and 5).

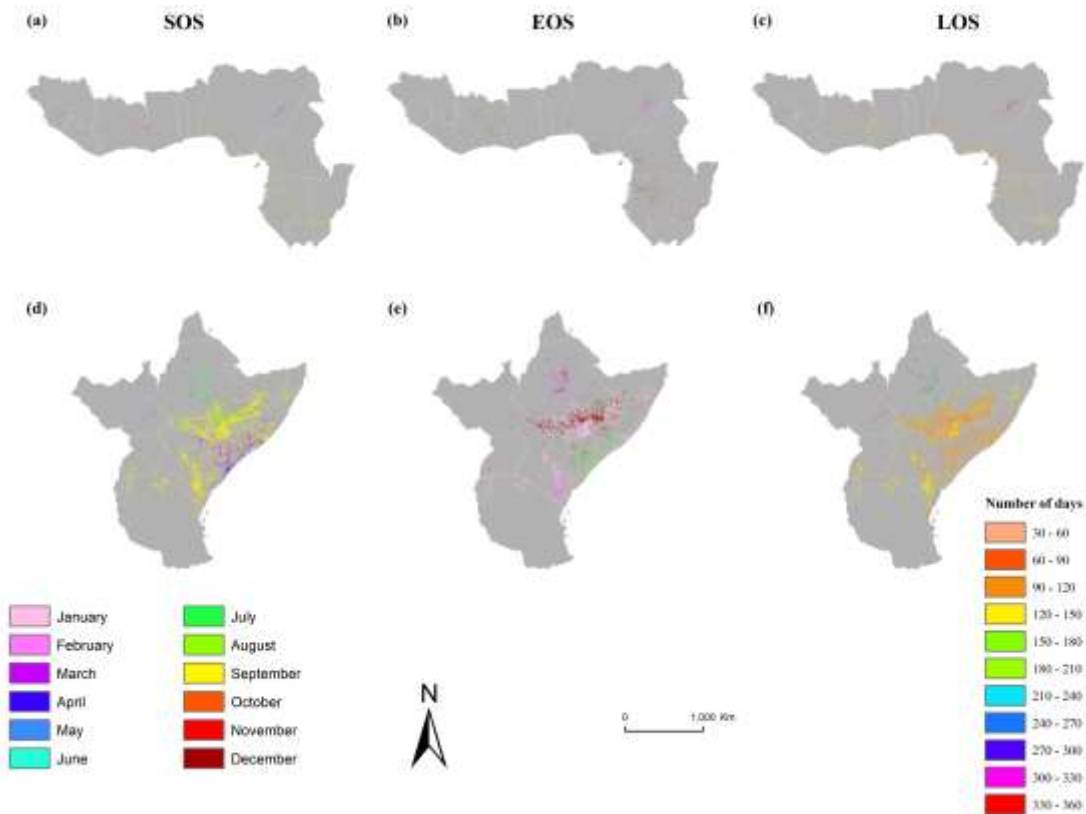
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274 **Figure 4:** The median values of phenological patterns derived from MODIS EVI data. (a) Start of
 275 Season (SOS) and (b) median End of Season (EOS) shown in months; (c) median Length of Season
 276 (LOS) shown in number of days.

277



278

279 **Figure 5:** The median values of phenological patterns derived from MODIS EVI data. (a,b,c) median
 280 (a) start, (b) end and (c) length of season for areas with second seasonal cycle in Western Africa and
 281 (d,e,f) median (d) start, (e) end and (f) length of season for areas with second seasonal cycle in
 282 Eastern Africa.

283

284 3.2. Latitudinal gradient

285 The variability of the majority values of LSP parameters was observed across the African latitudinal
 286 gradient (Figure 6). Latitude had more influence on SOS and EOS in the northern part of Africa than
 287 in the south. Approximately 49% of SOS dates and 59% of EOS dates north of the equator can be
 288 explained by latitude ($p < 0.0001$). A one degree increase in latitude will result in an approximately 5
 289 days delay in SOS and 5 days advance in EOS dates ($0.05 \text{ days km}^{-1}$) (see Table 2). However, the
 290 correlation between LOS and latitude was not significant ($p = 0.870$).

291

292

293

294 **Table 2: y-intercept, slope and coefficient of determination for linear regression between LSP parameters and latitude.**

Latitude	y-intercept			Slope			R ²			p (Sig.)		
	SOS	EOS	LOS	SOS	EOS	LOS	SOS	EOS	LOS	SOS	EOS	LOS
North	104.019	320.978	216.349	5.118	5.511	0.185	0.485	0.590	0.001	<0.0001	<0.0001	0.870
South	296.467	556.786	225.771	1.434	1.155	1.035	0.212	0.029	0.044	0.005	0.325	0.225

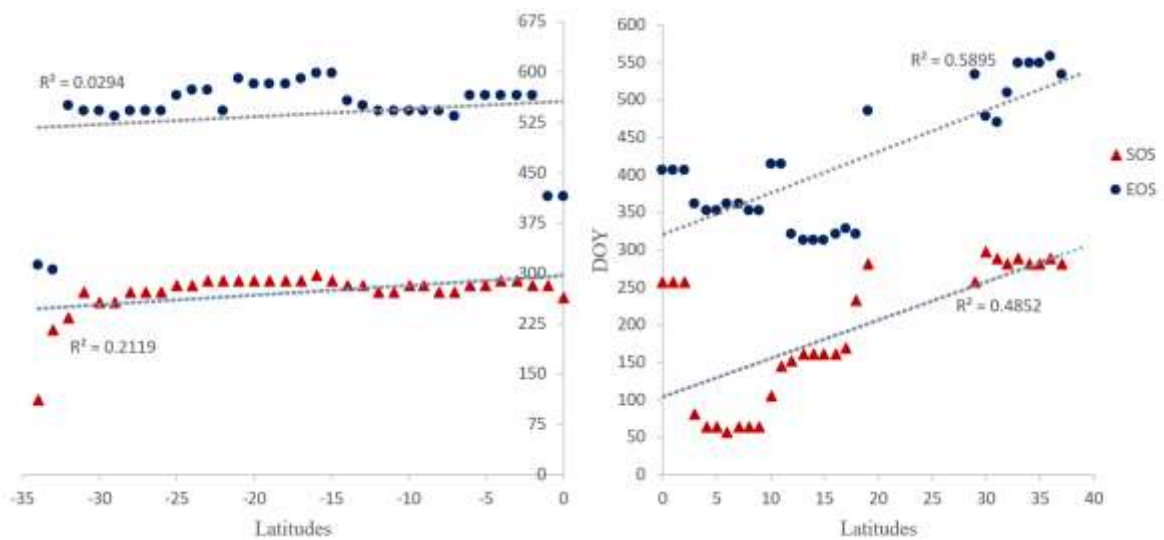
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296 However, no significant relationship was observed between EOS and latitude south of the equator
 297 ($R^2= 0.029$, $p=0.325$), while a relatively small correlation was observed between SOS and latitude
 298 ($R^2= 0.212$, $p=0.005$).

299 For a specific land cover type, the latitudinal variation in the phenology also follows the same pattern
 300 as explained before (i.e. a very small phenology-latitude correlation in the Southern hemisphere, and a
 301 large dependence on latitude in the Northern hemisphere of Africa). However, this trend was
 302 interrupted at latitude 30°N northwards and latitude 31°S southwards. This could be because of the
 303 different climatic conditions operating in these regions (see section 3.1).

304

305



306

307 **Figure 6:** Latitudinal variation in the LSP parameters, SOS and EOS, the left plot showing variation
 308 in the southern hemisphere and the right plot showing variation in the northern hemisphere.

309

310

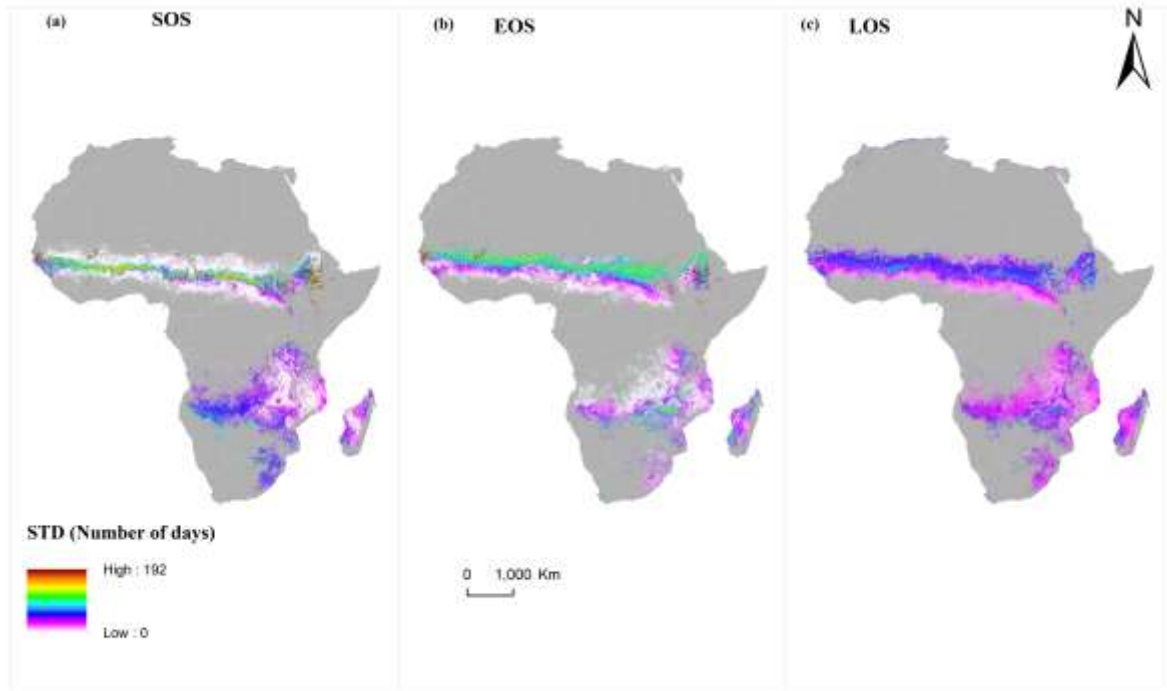
311

312 **3.3. Variability in LSP parameters**

313 Across the whole of Africa the STD of all LSP parameters for 15 years ranges from 0 – 80 days.
314 However, greater variability was observed in SOS compared to EOS and LOS, and this occurred mostly
315 in the Sahelian region, and croplands (see Figure 6). Although representing less than 1% of the total
316 number of pixels, some areas in Western Africa and the Horn of Africa, mainly croplands, produced
317 very large standard deviations for SOS of up to 128 days. The same large standard deviation was
318 observed for both EOS and LOS.

319 No significant inter-annual variability was observed for the evergreen and deciduous forest across
320 Africa as standard deviation values were very small, of less than 10 days. The same observation was
321 recorded for STD of SOS for shrublands and grasslands, with the exception of a few locations in Eastern
322 and some parts of Western Africa that had SOS STD values of up to 128 days. Nevertheless, EOS and
323 LOS for both land cover types had STD values ranging from 0 to 48 days and these were mainly in the
324 Sahelian and eastern sub-regions. On the other hand, the STD of SOS for savannas (woody
325 savannas/savannas) ranged from 0 to 40 days, and the number of days increased in EOS (0 to 48 days)
326 and LOS (0 to 56 days).

327 Contrasting with the first season, no significant variability was observed in LSP for the entire second
328 season, as STDs were very small, with values of less than a day.



329

330 **Figure 7:** Standard deviation of LSP parameters in number of days for the period of 2001 to 2015 for
 331 (a) SOS, (b) EOS, and (c) LOS.

332

333 **3.4. Characterisation of the LSP of the major land cover types in different geographical sub-**
 334 **regions**

335 The spatial and temporal variability of the vegetation phenological pattern in Africa is greatly
 336 influenced by different climatic factors (rainfall, temperature and insolation) in the geographical sub-
 337 regions, and vegetation type. Different patterns were observed in the LSP parameters across the six
 338 types of land cover based on the five geographical sub-regions in Africa (Figure 8).

339 Croplands/natural vegetation in Western Africa and some parts of Eastern Africa had over 70% of the
 340 SOS dates (homogeneous pixels) from late February to June (with over 36% occurring in June), and
 341 EOS between November and February. In geographical sub-regions south of the equator, there was
 342 an observed shift in SOS dates, occurring later between August to November, with their
 343 corresponding EOS dates between June and August. However, some locations in Northern Africa also
 344 exhibited similarly advanced SOS dates (see section 3.1 for explanation). When LOS is compared to

345 other vegetation land cover types, croplands/natural vegetation had the longest growing season of
346 approximately 12 months and these were mostly located in Western Africa.

347 One unique feature of croplands/natural vegetation is the bimodality observed in Eastern Africa.
348 Although this was seen in very few pixels (see Figure 5), this nevertheless indicates double cropping
349 activities made possible by bimodal rainfall regimes.

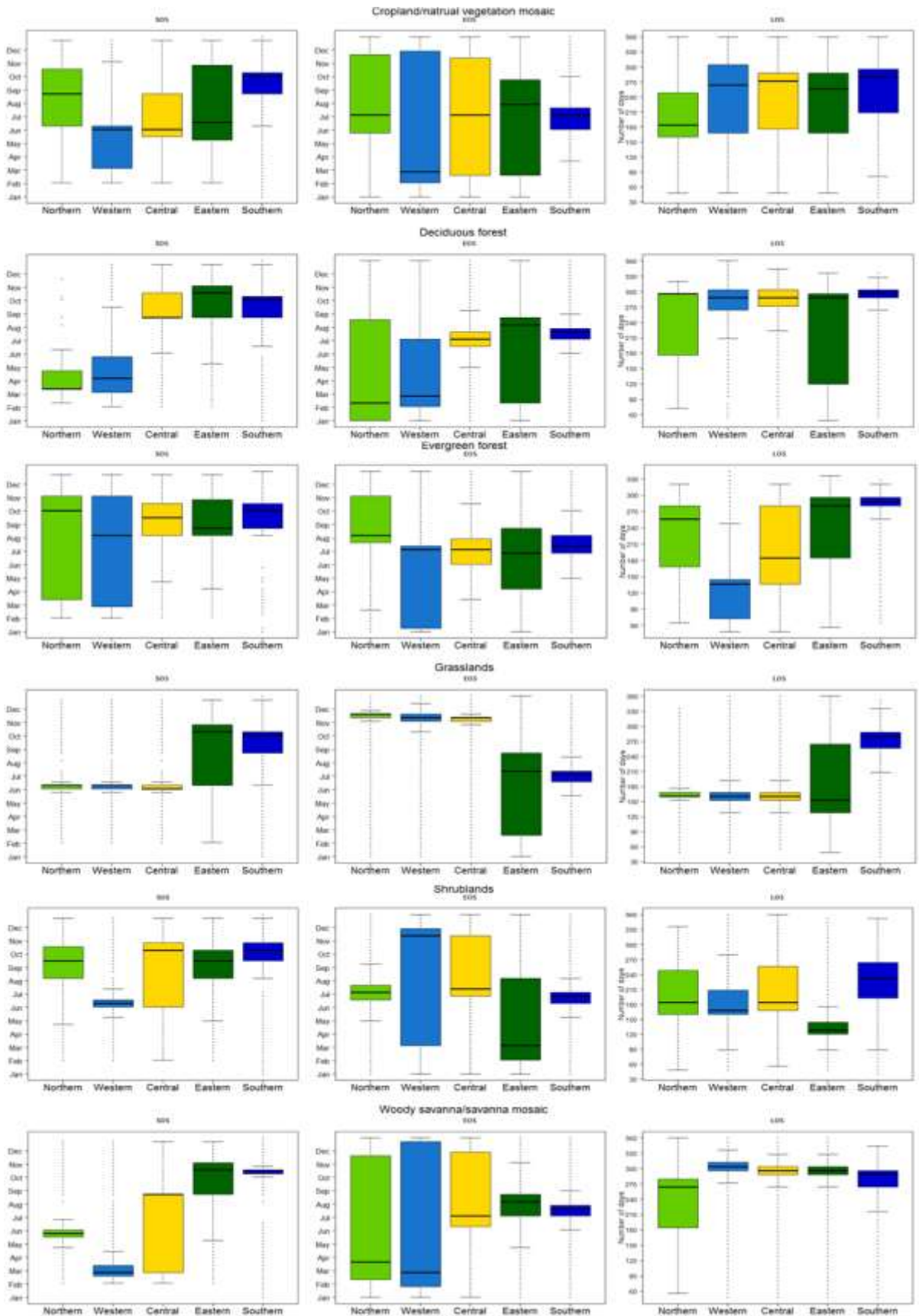
350 The phenologies of deciduous forest and evergreen forest are somewhat similar, especially in the
351 southern regions of Africa, with both having growing seasons starting mostly between August and
352 November, and ending mostly in January, June, July and August. The average LOS of both land cover
353 types is 10 months. As expected with most land cover types, the spatial location influences the
354 phenology of both forest types, as SOS dates are much earlier in Northern, Western Africa and parts
355 of Eastern Africa.

356 Grasslands, unlike most other land cover types, exhibit very distinct SOS and EOS dates, occurring
357 mainly in the month of June and November, respectively, for all geographical sub-regions in the north
358 of Africa, while southern and some eastern grasslands have a diverse range of SOS and EOS dates.

359 Shrublands also have very diverse SOS and EOS dates across the geographical sub-regions, especially
360 Southern Africa, resulting in a wide range of LOS from 3 to 11 months. However, shrublands in
361 Western Africa have a distinct LSP, with the growing season beginning in early June and mostly
362 ending towards late November.

363 Woody savanna and savanna are very different from most land cover types. In Western Africa, their
364 SOS dates were mainly in February and March, unlike grasslands, which have most SOS dates in
365 June. Over 85% of the homogeneous pixels of the woody savanna/savanna land cover type have a
366 growing season length between 9 to 10 months.

367



368

369 **Figure 8:** Box plots showing the distribution of pixels of LSP parameters in the six major land cover

370 types based on five geographical sub-regions.

371

372 **3.5. Heterogeneity of LSP parameters at coarse spatial resolutions**

373 The effect of spatial resolution on LSP parameters is demonstrated in Table 3. The table shows the
 374 range of STD of SOS in grids of 8 km, 5 km, 3 km and 1 km, and the percentage of pixels having
 375 those values. 22% of the pixels with an 8 km resolution have STD values ranging from 37 – 180
 376 (DOY). As expected, this number reduces as spatial resolution increases: 19% for 5 km, 16% for 3 km
 377 and 6% for 1 km. The reverse was observed for percentage of pixels with smaller STD values (i.e., the
 378 finer the spatial resolution the greater the number of pixels with smaller STD deviation values) (see
 379 Table 3).

380

381 **Table 3: Percentage of pixels falling into different STD ranges shown for four different spatial resolutions**
 382 **of 8 km, 5 km, 3 km and 1 km.**

STD (SOS)	8000 m	5000 m	3000 m	1000 m
0 - 18	62.69	67.36	74.29	89.86
19 - 36	15.77	13.99	10.20	4.58
37 - 54	5.27	4.44	3.64	0.98
55 - 72	3.43	2.95	2.43	0.35
73 - 90	3.48	3.00	2.46	0.60
91 - 108	3.98	3.41	2.81	0.93
109 - 126	4.17	3.59	2.88	1.60
127 - 144	1.06	1.10	1.12	0.88
145 - 162	0.11	0.12	0.14	0.18
163 - 180	0.02	0.03	0.03	0.04

383

384

385 **4. Discussion**

386 The phenological pattern of vegetation across different land covers and across different African sub-
 387 regions is important in understanding the vegetation dynamics of different biomes especially in
 388 relation to climate changes. This research provides a detailed characterisation of the LSP of the major
 389 land cover types in Africa at a continental scale based on the different geographical sub-regions at the
 390 finest spatio-temporal resolution to-date.

391

392

393 4.1. Latitudinal variation in LSP

394 Latitude was found to have some controlling effect on phenological patterns which is consistent with
395 results from previous studies (Zhang *et al.*, 2005; Butt *et al.*, 2011; Brottem *et al.*, 2014; Guan *et al.*,
396 2014b). The latitudinal variation in phenology across Africa revealed greater spatial variability at
397 lower latitudes, that is, the Southern Hemisphere of the African continent. However, advances in SOS
398 dates were observed as latitude increases, especially in the northern hemisphere (see Figure 3 & 5).
399 Similar results were found in Seghieri *et al.*(2009); with the same coverage of shrubs in West Africa,
400 leafing dates were earlier at lower latitudes when compared to leafing dates at higher latitudes. Also
401 Guan *et al.* (2014b) and Zhang *et al.* (2005) showed that in Northern Africa LSP parameters were
402 more correlated with latitude than in the Southern Hemisphere. This research, which used a much
403 finer spatial resolution, not only confirms this phenology-latitude relationship but also provides the
404 average rate of increase per one degree increase in latitude. This average rate of a 0.05 days km⁻¹ for
405 both SOS and EOS is supported by previous studies: (0.12 days km⁻¹ and 0.05 days km⁻¹ for the
406 period 2000 to 2003 in Zhang *et al.* (2005), 0.05 days km⁻¹ and 0.03 days km⁻¹ for the period 2000 to
407 2008 in Bobée *et al.* (2012) and 0.09 days km⁻¹ and 0.05 days km⁻¹ for the period 2000 to 2010 in Butt
408 *et al.* (2011), respectively).

409 One major reason for this North-South discrepancy in response to latitudinal gradient is the climatic
410 factors operational in these regions. The North is mostly controlled by the northwards movement of
411 the Intertropical Convergence Zone (ITCZ) which migrates latitudinally defining the seasonality of
412 rainfall in the northern region (Giannini *et al.*, 2008). However, the south has multiple climatic factors
413 at play: the east-west oriented component of the African ITCZ, the North Atlantic Oscillation index
414 (NAO), the Pacific Decadal Oscillation (PDO) (Nicholson, 2001, 2003; Brown *et al.*, 2010) and the
415 Agulhas and Benguela current systems (Walker, 1990), each exerting their influence along the east-
416 west to the south-west coasts.

417 The present results show that in some places in the African continent LSP does not vary linearly with
418 latitude, and more importantly quantify the degree of variation.

419

420

421 **4.2. Inter-annual variability**

422 The broad spatial distribution of inter-annual variability of LSP demonstrated in this research is
423 consistent with the outcomes from previous studies. Interesting is the different pattern of inter-annual
424 variability shown by the different geographical sub-regions and the different land cover types. Inter-
425 annual variability was greater in Eastern and some parts of Western Africa, which corresponds with
426 some areas identified as hotspots of change by Linderman *et al.* (2005) (Figure 7). Most land cover
427 types in these regions had a large STD representing inter-annual variability except for the evergreen
428 and deciduous forest types. These vegetation types across Africa were found not to have significant
429 changes in LSP parameters; a similar outcome was reported for vegetation activity by Linderman *et*
430 *al.* (2005) for the period 2000 to 2004. In contrast, croplands had a large STD, with SOS having the
431 largest values. This confirms results from previous studies of crop failures in the Sahelian region and
432 Eastern Africa (Vrieling *et al.*, 2013; Landmann & Dubovyk, 2014; Meroni *et al.*, 2014). Similarly,
433 shrublands and grasslands across Africa had moderately large STDs for EOS and LOS, but large STD
434 for SOS in the Eastern and Western sub-regions. This implies that between 2001 and 2014, some
435 factors may have affected the onset of growing season in these regions. Factors that could be
436 responsible, and have been identified by previous studies are: human-induced land transformations
437 (Landmann & Dubovyk, 2014), climatic factors like droughts and rainfall anomalies (Anyamba &
438 Tucker, 2005; Meroni *et al.*, 2014), and vegetation-type transitions occasioned by both climatic and
439 human factors (Linderman *et al.*, 2005; Mitchard *et al.*, 2009).

440 Contrary to Vrieling *et al.* (2013), no heteroscedasticity was observed in LOS. Our results showed no
441 relationship between the duration of LOS and STD values of LOS. Additionally, no significant
442 relationship was detected between inter-annual variability and latitudinal gradient.

443

444 **4.3. Comparison with ground-based studies**

445 Owing to the absence of a comprehensive ground-based observation network in Africa and the very
446 limited number of ground-based studies (Rutherford & Panagos, 1982; Childes, 1989; Seghieri *et al.*,
447 2009; February & Higgins, 2016; Whitecross *et al.*, 2017a,b), direct or indirect validation of the
448 results of this study was not possible. Hence, a comparison was made with the limited existing

449 literature on ground-based studies. In Western Africa, several species of shrubs and woodland
450 savannah, and mosaic of crops and natural vegetation have been found to start leafing in February just
451 before the rainy season, and in June during the rainy season (Seghieri *et al.*, 2009). This agrees well
452 with our findings as results from our study in the same geographical locations showed SOS to begin in
453 DOY 57 – 65 and DOY 161 - 169. This early onset of growing season before the rains has also been
454 reported to occur in numerous evergreen and mostly woody plants in the African Sahel by Seghieri
455 and Do (2012); Guan *et al.* (2014b) and Brandt *et al.* (2016). More recent studies have reported the
456 ubiquitous nature of this pre-rain onset in southern Africa (Ryan *et al.*, 2017; Whitecross *et al.*,
457 2017a,b). Similarly, in Southern Africa, some species of savanna trees were found to begin their
458 growing season and attain tree canopy fullness between October and November. These savanna trees
459 were also found to have no leaves at the end of the dry season in October (February & Higgins, 2016;
460 Whitecross *et al.*, 2017a,b). Again, our results for the same geographical location are in agreement
461 with these findings.

462 Increased air temperature and atmospheric vapour pressure/relative humidity, with scleromorphic
463 features and access to deeper groundwater or stored water in plants have been proposed to be
464 responsible for this early onset of greening (De Bie *et al.*, 1998; Do *et al.*, 2005; Seghieri *et al.*, 2012).
465 Comparison was not possible with all the existing literature on ground-based studies due to the type of
466 vegetation phenological parameters measured. For example, plant phenophases such as budding,
467 shoot growth, flowering and fruiting measured by some studies (Chapman *et al.*, 2005; Do *et al.*,
468 2005; O'Farrell *et al.*, 2007; Sekhwela & Yates, 2007; Yamagiwa *et al.*, 2008; Wang'ondy *et al.*,
469 2010, 2013; Seghieri *et al.*, 2012; Polansky & Boesch, 2013) cannot be compared directly to onset of
470 greenness or leaf emergence/leafing in remote sensing studies. Regardless of this limitation, the
471 phenological patterns of major vegetation types from these ground-based studies are very similar to
472 the results presented here. This limitation further drives home the need for more ground-based
473 observations and a phenological network for the African continent.

474

475

476

477 **4.4. Comparison with other remote sensing studies**

478 The present results differ from most earlier remote sensing studies of LSP over Africa (Brown *et al.*,
479 2010, 2012; Jacquin *et al.*, 2010; Vrieling *et al.*, 2013) which used a threshold method in estimating
480 LSP parameters. In comparison, the present analysis detected SOS approximately 30 to 60 days
481 earlier across Africa. Similarly, EOS was detected approximately 30 – 60 days later. Consequently,
482 the present study produced longer LOS values of about 30 – 90 days. This supports the findings of
483 Vrieling *et al.* (2008) and de Beurs & Henebry (2010), that threshold methods estimate SOS later and
484 EOS earlier because the point of maximum curvature may be below the user-defined threshold.
485 On the other hand, the present results are in agreement with remote sensing studies (Zhang *et al.*,
486 2005; Archibald & Scholes, 2007; Butt *et al.*, 2011; Bobée *et al.*, 2012; Brottem *et al.*, 2014; Guan *et*
487 *al.*, 2014a,b; Ryan *et al.*, 2014) that applied the inflection point or the function model fitting methods
488 in estimating LSP. This consistency was very evident in the early green-up observed before the rainy
489 seasons, especially in evergreen forest and woodlands (Archibald & Scholes, 2007; Guan *et al.*,
490 2014b), and the distinct phenological pattern observed in the extreme northern and southern tips of
491 Africa (Guan *et al.*, 2014a).

492 While there exists strong agreement with previous studies, minor discrepancies of an estimated 5 – 20
493 days were observed. This could be the result of the different spatial resolution used in the studies. At
494 coarser spatial resolutions, phenological parameters are usually averaged across an area that may have
495 different vegetation types with distinct phenological patterns. This can be seen in the STDs of SOS
496 with spatial resolutions of 1 km, 3 km, 5 km and 8 km. As the spatial resolution becomes finer, the
497 STD in number of days reduces (see Table 3). This suggests that with a finer spatial resolution there is
498 less conditional bias (under-estimating highs and over-estimating lows) from spatial averaging and
499 aggregation.

500 Aside from the type of estimation technique and the spatial resolution of data, the smoothing
501 techniques (Atkinson *et al.*, 2012), sensor type (Atzberger *et al.*, 2013) and the temporal resolution of
502 data (Zhang *et al.*, 2009) could also be responsible for such discrepancies between outcomes.

503

504

505 **Conclusion**

506 The LSP of the major vegetation types in Africa was described for the first time using homogeneous
507 pixels from 12 years (2001 – 2012) MODIS land cover data (MODIS MCD12Q1) and EVI derived
508 from the MODIS MOD09A1 product at a medium spatial resolution of 500 m and a high temporal
509 frequency of 8-days. Indeed, the maps of LSP parameters (SOS, EOS, LOS) produced here represent
510 the finest spatial resolution and most detailed maps of the phenology of Africa to-date. Additionally,
511 the inter-annual variability of all LSP parameters for all of Africa was reported for the first time.
512 The well-known phenology-latitude relationship in Africa was quantified at an unprecedented fine
513 resolution, with a greater correlation found in northern latitudes. Moreover, the dependence of the
514 LSP parameters (SOS, EOS and LOS) on land cover type and geographical sub-region was analysed
515 in detail (Figure 8), revealing a complex interaction between the three dimensions of vegetation
516 timing, geographical location and land cover type.

517 The results reported here support previous studies while providing a more refined quantification with
518 some significant variations to existing maps. The spatial detail (500 m) with which the LSP
519 parameters are mapped here provides a platform to support further applied environmental research in
520 the African continent. In particular, it is anticipated that the mapped outputs from this research will be
521 important for ecosystem management and climate-related research and can be of value for further
522 studies on climate change impacts and phenology-climate modelling.

523 While it was not possible to conduct an extensive empirical validation of the maps of LSP produced
524 (due to the lack of a comprehensive African ground observation network measuring vegetation
525 phenology), comparison of the results with the available ground-based studies published in the
526 literature found close agreement. Moreover, the methods applied in this research to estimate LSP
527 parameters have been applied widely and tested extensively in other studies, including through
528 comparison with empirical ground data in those studies. Further studies should be undertaken to
529 provide a comprehensive, continental scale validation of the LSP predictions across Africa when
530 suitable ground data become available.

531

532 **Acknowledgments**

533 The authors would like to thank the Commonwealth Scholarship Commission in the UK for funding
534 and support provided to Tracy Adole.
535

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