

1 Salinity from space unlocks satellite-based
2 assessment of ocean acidification

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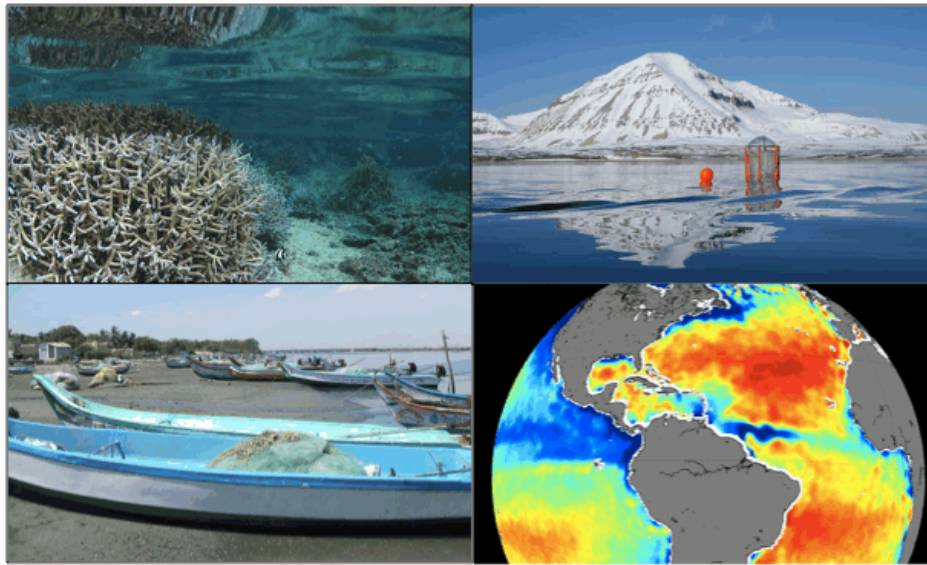
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19



Abstract artwork

Note to editor (top left to bottom right): Tropical coral; Svalbard in the Barents Sea; Beach in India on the coast of the Bay of Bengal; Salinity from space (SMOS) showing the Amazon plume. All images taken by PML staff and used with permission.

20

21 Approximately a quarter of the carbon dioxide (CO_2) that we emit into the atmosphere
22 is absorbed by the ocean. This oceanic uptake of CO_2 leads to a change in marine
23 carbonate chemistry resulting in a decrease of seawater pH and carbonate ion
24 concentration, a process commonly called ‘Ocean Acidification’. Salinity data are key
25 for assessing the marine carbonate system, and new space-based salinity
26 measurements will enable the development of novel space-based ocean acidification
27 assessment. Recent studies have highlighted the need to develop new *in situ*
28 technology for monitoring ocean acidification, but the potential capabilities of space-
29 based measurements remain largely untapped. Routine measurements from space can

30 provide quasi-synoptic, reproducible data for investigating processes on global scales;
31 they may also be the most efficient way to monitor the ocean surface. As the carbon
32 cycle is dominantly controlled by the balance between the biological and solubility
33 carbon pumps, innovative methods to exploit existing satellite sea surface temperature
34 and ocean color, and new satellite sea surface salinity measurements, are needed and
35 will enable frequent assessment of ocean acidification parameters over large spatial
36 scales.

37 **1. Introduction**

38

39 Each year global emissions of carbon dioxide (CO₂) into our atmosphere continue to
40 rise. These increasing atmospheric concentrations cause a net influx of CO₂ into the
41 oceans. Of the roughly 36 billion metric tons of CO₂ that is emitted into our
42 atmosphere each year, approximately a quarter transfers into the oceans ¹. This CO₂
43 addition has caused a shift in the seawater carbonate system, termed Ocean
44 Acidification (OA), resulting in a 26% increase in acidity and a 16% decrease in
45 carbonate ion concentration since the industrial revolution ². Recently there has been
46 recognition that this acidification is not occurring uniformly across the global oceans,
47 with some regions acidifying faster than others ^{3,4}. However, the overall cause of OA
48 remains consistent: the addition of CO₂ into the oceans, and as such, it remains a
49 global issue. Continual emissions of CO₂ into the atmosphere over the next century
50 will decrease average surface ocean pH to levels which will be deleterious to many
51 marine ecosystems and the services they provide ⁵.

52

53 While the seawater carbonate system is relatively complex, two parameters have been
54 suggested as pertinent to the monitoring and assessment of OA through time and

55 space. These are pH (the measure of acidity) and calcium carbonate (CaCO_3) mineral
56 saturation state, with aragonite generally considered to be an important CaCO_3
57 mineral to be monitored because of its relevance to marine organisms (e.g. corals) and
58 its relative solubility. Thermodynamically, CaCO_3 is stable when the saturation state
59 (an index of the concentrations of calcium and carbonate ions) is greater than one and
60 becomes unstable when seawater becomes undersaturated with these ions (saturation
61 < 1). While there is significant variability between types of organism, there is ample
62 experimental evidence that many calcifying organisms are sensitive to OA ⁶, and that
63 thresholds exist below which some organisms become stressed and their well-being
64 and existence becomes threatened ⁷. Increasingly evidence suggests that the
65 physiology and behaviour of calcifying and non-calcifying organisms can be impacted
66 by increasing OA ⁸, with cascading effects on the food chain and protein supply for
67 humans ³, and alterations to the functioning of ecosystems and feedbacks to our
68 climate ⁹.

69

70 In 2012 the Global Ocean Acidification Observing Network (GOA-ON, [www.goa-](http://www.goa-on.org)
71 [on.org](http://www.goa-on.org)) was formed in an attempt to bring together expertise, datasets and resources to
72 improve OA monitoring. At present, OA monitoring efforts are dominated by *in situ*
73 observations from moorings, ships and associated platforms. Whilst key to any
74 monitoring campaign, *in situ* data tend to be spatially sparse, especially in
75 inhospitable regions, and so on their own are unlikely to provide a comprehensive,
76 robust and cost effective solution to global OA monitoring. The need to monitor and
77 study large areas of the Earth has driven the development of satellite-based sensors.

78

79 Increasingly, as *in situ* data accumulate, attempts are being made to use *in situ*
80 hydrographic data ¹⁰⁻¹³ and/or remotely-sensed data ^{14, 15} to provide proxies and
81 indicators for the condition of the carbonate system, enabling data gaps to be filled in
82 both space and time. The increased availability of *in situ* data creates a substantial
83 dataset to develop and test the capabilities of satellite-derived products, and we
84 suggest that the recent availability of satellite-based salinity measurements provides
85 new key insights for studying and assessing OA from space.

86 **2. The complexities of the carbonate system**

87

88 The oceanic carbonate system can be understood and probed through four key
89 parameters: total alkalinity (TA), dissolved inorganic carbon (DIC), pH and fugacity
90 of CO₂ (f_{CO_2}). The latter may be replaced with the related partial pressure of CO₂,
91 p_{CO_2} , from which f_{CO_2} can be calculated, and the two are often used interchangeably.
92 In principle, knowledge of any two of these four is sufficient to solve the carbonate
93 system equations. However, over-determination, the process of measuring at least
94 three parameters, is advantageous.

95

96 The relationships between the different carbonate system parameters are
97 fundamentally driven by thermodynamics, hence influenced by temperature and
98 pressure, and knowing these is fundamental for calculating the carbonate system as a
99 whole ¹⁶. Water temperature is the major controller of the solubility of CO₂ ¹⁷, so
100 seasonal changes in sea temperature can, depending on the region, be significant for
101 driving changes in f_{CO_2} (and consequently DIC and pH). Salinity affects the
102 coefficients of the carbonate system equations. Hence to solve the equations, it is

103 necessary to estimate temperature, salinity and pressure along with carbonate
104 parameters.

105

106 The ratio between ions (the constituents of salinity) will tend to remain constant
107 anywhere in the global oceans, resulting in a strong relationship between TA and
108 salinity ¹⁸. Unfortunately, a universal relationship between TA and salinity does not
109 apply in certain regions, for instance in areas influenced by freshwater outflows from
110 rivers ⁷, or areas where calcification and/or CaCO₃ dissolution occurs, such as where
111 calcifying plankton are prevalent ¹⁹. In these regions, it is therefore critical to gain
112 additional local knowledge. For example, different rivers will have different ionic
113 concentrations (and therefore different TA concentrations) depending on the
114 surrounding geology and hydrology.

115

116 For DIC, f_{CO2} (or p_{CO2}) and pH, the other important process is biological activity ¹⁹.
117 Removal or addition of CO₂ by plankton photosynthesis or respiration can be a
118 significant component of the seasonal signal ²⁰. Biological activity, in turn, is driven
119 by factors such as nutrient dynamics and light conditions, which again are regionally
120 specific. Measurements of chlorophyll (a proxy for biomass) and/or oxygen
121 concentration can be useful for interpreting the biological component of the carbon
122 signal.

123

124 The combination of these processes means that it is extremely challenging to produce
125 a global relationship between any component of the carbonate system and its drivers.
126 To enable us to understand these dynamics, extrapolation from collected data points
127 to the global ocean is needed, and along with model predictions, empirical

128 relationships and datasets are important and need to be studied and developed. OA
 129 needs to be assessed using these relationships on a global scale, but regional
 130 complexities, particularly where riverine and coastal processes dominate ^{21, 22}, cause
 131 significant challenges for global empirical relationships.

132

133 **3. Current *in situ* approaches and challenges**

134

135 Laboratory measurements are the gold standard for assessing the carbonate system in
 136 seawater, with accuracy far in excess of that achievable from satellites.²³⁻²⁵ However,
 137 research vessel time is expensive and limited in coverage, so autonomous *in situ*
 138 instruments are also deployed, e.g. on buoys, with less accuracy ²⁶. A notable example
 139 is the Argo network of over 3000 drifters, which measure temperature and salinity
 140 throughout the deep global ocean. Interpolation of Argo data is much less challenging
 141 than for most *in situ* measurements. Argo is the closest *in situ* data have come to the
 142 global, synoptic measurements possible with satellites, but shallow or enclosed seas
 143 are not represented (there are as yet no Argo instruments in the open Arctic Ocean).
 144 Table 1 lists more examples. Of the four key parameters, only f_{CO_2} (or p_{CO_2}) and pH
 145 are routinely monitored *in situ*. As yet there are limited capabilities to measure DIC
 146 and TA autonomously, hence these parameters must be measured either in a ship-
 147 based laboratory or on land.

148

Dataset name and reference	Temporal period	Geographic location	Variables	No. of data points
SOCAT v2.0 ²⁷	1968-2011	Global*	f_{CO_2} , SSS, SST	6,000,000+
LDEO v2012 ²⁸	1980-present	Global*	p_{CO_2} , SSS, SST	6,000,000+
GLODAP ²⁹	1970-2000	Global	TA, DIC, SSS, SST, Nitrate	10,000+

CARINA AMS v1.2 ³⁰	1980-2006	Arctic	TA, DIC, SSS, SST	1500+
CARINA ATL v1.0 ³¹		Atlantic		
CARINA SO v1.1 ³²		Southern Ocean		
AMT ³³	1995-present	Atlantic	pCO _{2w} , SSS, SST, Chl, pH	1000+
NIVA Ferrybox ³⁴	2008-present	Arctic	pCO _{2w} , TA, DIC, SSS, SST	1000+
OWS Mike ³⁵	1948-2009	Arctic	TA, DIC, SSS, SST, Chl	1000+
RAMA Moored buoy array ³⁶	2007-present	Bay of Bengal	SSS, SST	1000+
ARGO buoys ³⁷	2003-present	Global	SSS, SST	1,000,000+
OOI ³⁸	2014 onwards	Global (6 sites)	pCO ₂ , SSS, SST, nitrate	New program
SOCOM ³⁹	2014 onwards	Southern Ocean	SSS, SST, pH, nitrate	New program

149

150 Table 1. *In situ* datasets and programs that can be used for the development and
151 validation of OA remote sensing algorithms.

152 4. Potential of space based observations

153

154 4.1 Advantages and disadvantages

155

156 While it has proven difficult to use remote sensing to directly monitor and detect
157 changes in seawater pH and their impact on marine organisms²², satellites can
158 measure sea surface temperature and salinity (SST and SSS) and surface chlorophyll-
159 a, from which carbonate system parameters can be estimated using empirical
160 relationships derived from *in situ* data. Although surface measurements may not be
161 representative of important biological processes, e.g. fish or shellfish, observations at
162 the surface are particularly important for OA because the change in carbonate
163 chemistry due to atmospheric CO₂ occurs in the surface first. Thus satellites have
164 great potential as a tool for assessing changes in carbonate chemistry.

165

166 SST has been measured from space with infrared radiometry since the 1960s, but the
167 data are only globally of sufficient quality for climate studies since 1991⁴⁰. Satellite
168 measurements of chlorophyll-*a* in the visible are more recent, starting in 1986 and
169 delivering high quality global data since 1997⁴¹. Both measurements are made
170 globally at high spatial and temporal resolution, but with data gaps due to effects such
171 as cloud, which can greatly affect data availability in cloudy regions. SST is measured
172 in the top few microns, and chlorophyll-*a* is generally measured to depths around 1-
173 100m, depending on water clarity. Data quality can be affected by many issues, e.g.
174 adjacent land or ice may affect both SST and chlorophyll-*a* retrievals, and suspended
175 sediment may affect chlorophyll-*a* retrievals.

176

177 Only since 2009 has a satellite-based capability for measuring SSS existed. Increasing
178 salinity decreases the emissivity of seawater and so changes the microwave radiation
179 emitted at the water surface. ESA Soil Moisture and Ocean Salinity (SMOS) and
180 NASA-CONAE Aquarius (launched in 2009 and 2011 respectively, both currently in
181 operation), are L-band microwave sensors designed to detect variations in microwave
182 radiation and thus estimate ocean salinity in the top centimeter. The instruments are
183 novel and the measurement is very challenging, and research is ongoing to improve
184 data quality⁴². The instruments can measure every few days at a spatial resolution of
185 35-100km, but single measurements are very noisy, so the instantaneous swath data
186 are generally spatially and temporally averaged over 10 days or a month, with an
187 intended accuracy around 0.1 - 0.2 g/kg for monthly 200 km data. A particular issue
188 close to urban areas is radio frequency interference from illegal broadcasts, which are
189 gradually being eliminated but still result in large data gaps, particularly for SMOS.

190 The signal can be affected by nearby land or sea ice, and the sensitivity to SSS
191 decreases for cold water, by about 50% from 20°C to 0°C⁴³.

192

193 With these challenges, a central question is whether satellite SSS can bring new
194 complementary information to *in situ* SSS measurements such as Argo for assessing
195 OA. Direct comparisons^{44, 45} indicate differences of 0.15-0.5 g/kg in a 1°x1° region
196 over 10-30 days. The two are difficult to compare directly however, as Argo measures
197 5m or more from the surface, so some differences are expected even in the absence of
198 errors, especially where the water column is stratified. A better strategy might be to
199 compare their effectiveness in estimating OA. How the uncertainties propagate
200 through the carbonate system calculations is the subject of ongoing research.

201

202 Despite biases and uncertainties, satellite measurements of SSS in the top centimeter
203 contain geophysical information not detected by Argo^{46,47}. In addition, Argo coverage
204 can be much poorer than satellite SSS in several regions such as the major western
205 boundary or equatorial currents and across strong oceanic fronts. The use of
206 interpolated Argo products presents an additional source of uncertainty due to the
207 interpolation scheme.⁴⁸ Satellite SSS can also resolve mesoscale spatial structures not
208 resolved by Argo measurements⁴⁹, and unlike Argo, satellites provide a synoptic
209 ‘snapshot’ of a region at a given time.

210

211 Regular mapping of the SSS field with unprecedented temporal and spatial resolution
212 at global scale is now possible from satellites. The impact of using satellite SSS for
213 carbonate system algorithms can now be tested, where previously there was a reliance
214 on climatology, *in situ* or model data. For example, this provides the means to study

215 the impact that freshwater influences (sea ice melt, riverine inputs and rain) can have
216 on the marine carbonate system. The use of satellite SSS data will also allow
217 evaluation of the impact on the carbonate system of the inter- and intra-annual
218 variations in SSS.

219

220 Recent advances in radar altimetry (e.g. Cryosat-2 and Sentinel 1 satellites and
221 sensors) are already enabling significant improvements in satellite sea-ice thickness
222 measurements⁵⁰. Thin sea ice thickness can now also be determined from SMOS,
223 complementing altimeter estimates mostly valid for thick sea ice⁵¹. Sea ice thickness
224 is important for OA research as it indicates whether ice is seasonal or multi-year,
225 supporting the interpretation of carbonate parameters. Altimetry is also used to
226 measure wind speeds and increases the coverage of scatterometer estimates in polar
227 regions. It provides higher-resolution (along track) estimates of surface wind stress,
228 which can potentially be used to indicate regions of upwelling. Wind-driven
229 upwelling causes dense cooler water (with higher concentrations of CO₂ and thus
230 more acidic) to be drawn up from depth to the ocean surface. This upwelling can have
231 significant impacts on local OA and ecosystems^{4, 52}, especially at eastern oceanic
232 boundaries^{53, 54}.

233

234 It is important to emphasise that the use of Earth observation data to derive carbonate
235 parameters should not be seen as a replacement for *in situ* measurement campaigns,
236 especially due to the current reliance on empirical and regional algorithms. Earth
237 observation algorithms need calibration and validation with *in situ* data such as those
238 taken by GOA-ON, and if the carbonate system response changes over time, empirical
239 and regional algorithms tuned to previous conditions may become less reliable.

240

241 4.2 Algorithms for estimating carbonate parameters

242

243 The four key OA parameters (pCO₂, DIC, TA, pH) are largely driven by temperature,
244 salinity and biological activity, allowing empirical relationships to be developed using
245 *in situ* measurements of OA parameters. Table 2 shows a range of published
246 algorithms based on such relationships, while Figure 1 shows their geographical
247 coverage. Both illustrate that most of the literature has focused on the northern basins
248 of the Pacific and Atlantic and the Arctic, especially the Barents Sea, with all other
249 regions only attracting algorithms for a single parameter or none at all.⁵⁵

250

Parameter	Dependencies	Region and references
pCO ₂	SST	Global ⁵⁶ , Barents Sea ⁵⁷
	SST, SSS	Barents Sea ⁵⁸ , Caribbean ¹⁴
	SST, Chl	N Pacific ⁵⁹
	SSS, Chl	North Sea ⁶⁰
	SST, SSS, Chl	N Pacific ⁶¹
	SST, Chl, MLD	Barents Sea ⁶²
TA	SSS	Barents Sea ⁵⁷
	SST, SSS	Global ^{18,63} , Arctic ¹⁵
	SSS, nitrate	Global ⁵⁵
DIC	SST, SSS	Equatorial pacific ⁶⁴
	SST, SSS, Chl	Arctic ¹⁵
pH	SST, Chl	N Pacific ¹⁰

251

252 Table 2. Example regional algorithms for each carbonate parameter illustrating the
253 variable dependencies. Chl is chlorophyll-*a* and MLD is mixed layer depth.

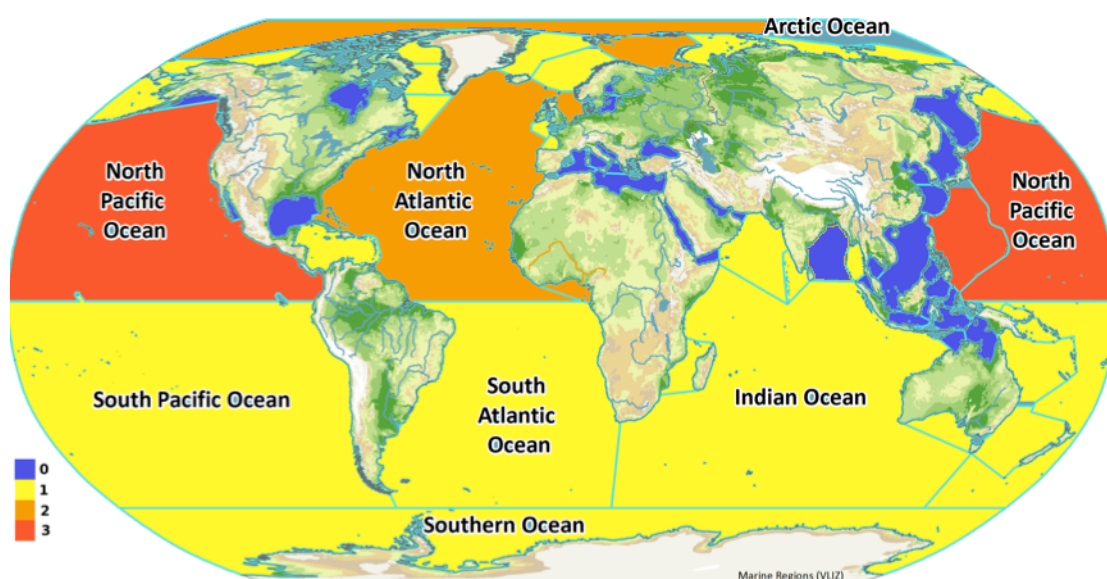


Figure 1. The number of key carbonate parameters (f_{CO_2} or p_{CO_2} , TA, DIC, pH) for which regional algorithms exist in the literature that can be implemented using just satellite Earth observation data. Regions are indicative of open ocean areas, as implementation of algorithms in coastal areas may be problematic.

255

256 NOAA's experimental Ocean Acidification Product Suite (OAPS) is a regional
 257 example of using empirical algorithms with a combination of climatological SSS and
 258 satellite SST to provide synoptic estimates of sea surface carbonate chemistry in the
 259 Greater Caribbean Region ¹⁴. p_{CO_2} and TA were derived from climatological SSS and
 260 satellite SST, then used to calculate monthly estimates of the remaining carbonate
 261 parameters, including aragonite saturation state and carbonate ion concentration. In
 262 general the derived data were in good agreement with *in situ* measured data (e.g.
 263 mean derived TA = $2375 \pm 36 \mu\text{mol kg}^{-1}$ compared to a mean ship-measured TA =
 264 $2366 \pm 77 \mu\text{mol kg}^{-1}$). OAPS works well in areas where chlorophyll-*a* is low,
 265 however in regions of high chlorophyll-*a*, where net productivity is likely to perturb

266 the carbonate system, and in areas where there are river inputs, the approach tends to
267 underestimate aragonite saturation state, for example ²¹.

268

269 A quite different approach is the assimilation of satellite data into ocean circulation
270 models ⁶⁵. The model output carbonate parameters can then be used directly. This
271 allows satellite-observed effects to be extended below the water surface, albeit with
272 the uncertainties inherent in model data. Here we seek to assess the direct use of
273 satellite data through empirical algorithms to improve OA estimates.

274

275 **4.3 Regions of interest for Earth observation**

276

277 **Arctic Seas**

278 It is increasingly recognised that the Polar Oceans (Arctic and Antarctic) are
279 particularly sensitive to OA ⁶⁶. Lower alkalinity (and thus buffer capacity), enhanced
280 warming, reduced sea-ice cover resulting in changes in the freshwater budget ⁶⁷, and
281 nutrient limitation make it more vulnerable to future OA ⁶⁸. Retreating ice also
282 provides increased open water for air-sea gas exchange and primary production ⁶⁹.

283

284 The remote nature of the Arctic Ocean provides difficulties for collecting *in situ*
285 datasets, with limited ship, autonomous vehicle and buoy access, and *in situ* data
286 collection during winter months is often impossible. Therefore the use of remote
287 sensing techniques is very attractive, if sufficient *in situ* data can be found to calibrate
288 satellite algorithms, and if the challenges of Arctic remote sensing can be overcome.

289 These waters are very challenging regions for satellite remote sensing. For instance,
290 low water temperatures reduce the sensitivity range of SSS sensors ⁴³, and sea ice can

291 complicate retrievals of SSS and chlorophyll-*a* ^{70, 71}. Improvement in the accuracy of
292 high latitude satellite SSS is expected soon by combining observations from SMOS,
293 Aquarius and the upcoming SMAP sensor, all polar-orbiting L-band radiometers.

294

295 **The Bay of Bengal**

296 This region is clearly a focus of current OA research with unique characteristics due
297 to the large freshwater influence. The flow of fresh water from the Ganges Delta into
298 Bay of Bengal (42,000 m³/sec) represents the second greatest discharge source in the
299 world. Additionally, rainfall along with freshwater inputs exceeds evaporation,
300 resulting in net water gain annually in the Bay of Bengal. Collectively these provide
301 an annual positive water balance that reduces surface salinity by 3-7 g/kg compared to
302 the adjacent Arabian Sea ^{72, 73}, resulting in distinctly different biogeochemical regimes
303 ⁷⁴. Biogeochemically, the Indian Ocean is one of the least studied and most poorly
304 understood ocean basins in the world ⁷⁴. This is particularly true for the Bay of Bengal
305 where a relatively small number of hydrographic sections and underway surface
306 observations have been undertaken, despite the notable influence of freshwater on
307 particle dynamics, air-sea carbon flux and surface carbonate chemistry ⁷⁵⁻⁷⁹. North of
308 15° S, TA increases relative to salinity ⁸⁰, indicating the presence of an important land
309 source that can broadly affect acidification dynamics.

310

311 To date there is little work on acidification dynamics and air sea exchange of CO₂ in
312 the Bay of Bengal ⁸¹⁻⁸³. In 2013, the Bay of Bengal Ocean Acidification (BOBOA)
313 Mooring was deployed for the first time in Bay of Bengal (15°N, 90°E) by PMEL
314 (NOAA) and the Bay of Bengal Large Marine Ecosystem Program (BOBLME). Data

315 from the buoy will improve our understanding of biogeochemical variations in the
316 open ocean environment of the Bay of Bengal.

317

318 It is an open question whether SSS can be used to estimate TA in the Bay of Bengal.

319 An important step towards answering this question would be to investigate the spatial

320 variability of the TA to salinity relationship in the region. Use of satellite SSS in the

321 region is also challenged by heavy radio frequency interference.

322

323 **The Greater Caribbean and the Amazon plume**

324 The reefs in the Greater Caribbean Region are economically important to the US and

325 Caribbean nations with an estimated annual net value of US\$3.1-4.6 billion in 2000 ⁸⁴.

326 At least two thirds of these reefs are threatened from human impacts including OA.

327 The skeleton of a coral is made of aragonite and the growth of their skeletons is

328 reduced by OA ⁶, and numerous studies have shown a net decline in coral

329 calcification (growth) rates in accordance with declining CaCO₃ saturation state ⁸⁵.

330 The waters of the Greater Caribbean region are predominantly oligotrophic and

331 similar to the subtropical gyre from which it receives most of its water ¹⁴. Whilst the

332 often shallow water environments of coral reefs and the plethora of small islands can

333 make it challenging for Earth observation instruments to collect reliable data, the

334 oligotrophic nature and the similarities in water type across the whole region make it

335 ideal for the development of novel products. This region therefore provides an ideal

336 case study to develop and evaluate algorithms representative of a shallow,

337 oligotrophic environment.

338

339 The Amazon plume, south of the Greater Caribbean, is the largest freshwater
340 discharge source in the world (209,000 m³/sec). It can cause SSS decreases of several
341 units many hundreds of kilometers from land, and has an area that seasonally can
342 reach 10⁶ km². These characteristics make it an ideal case study for testing and
343 evaluating remote sensing algorithms, particularly to study the space-time resolution
344 tradeoffs using SSS sensors.

345 **5. Future opportunities and focus**

346

347 The Copernicus program is a European flagship initiative, worth more than €7
348 billion, which aims to provide an operational satellite monitoring capability and
349 related services for the environment and security ⁸⁶. The launch of the Sentinel-1A
350 satellite in 2014 signaled its start. Of the five Sentinel satellite types, Sentinels 2 and 3
351 are most appropriate for assessment of the marine carbonate system ⁸⁷⁻⁸⁹. These
352 satellites will provide chlorophyll-*a* and SST with unprecedented spatial and temporal
353 coverage. The development of higher spatial resolution geostationary sensors that
354 continually monitor chlorophyll-*a* and SST over the same area of the Earth also holds
355 much potential for the future of OA assessment and research ⁹⁰. These satellites and
356 sensors are able to provide 10 or more observations per day, allowing the study of the
357 effect of tidal and diurnal cycles on OA. The societal importance of measuring and
358 observing the global carbon cycle was further highlighted with the launch of the
359 NASA Orbiting Carbon Observatory (OCO-2) in 2014. This satellite and its sensors
360 are designed to observe atmospheric CO₂ concentrations, but its potential for marine
361 carbon cycle and OA is likely to be a focus of future research.

362

363 SMOS and Aquarius have recently passed their nominal lifetimes, with SMOS now
364 extended until 2017. Based on the lifetimes of previous satellite Earth observation
365 sensors, they may well operate until the early 2020s. NASA's SMAP satellite, to be
366 launched in January 2015, should provide short-term continuity. The development of
367 the technology and the clear importance of monitoring ocean salinity are likely to
368 support the development of future satellite sensors. Also, historical time series data
369 from alternative microwave sensors hold the potential for a 10+ year time series of
370 satellite based SSS observations ⁹¹, and this sort of measurement record is likely to
371 extend into the future as it forms the basis of a global SSS monitoring effort.

372

373 In summary, satellite products developed up to now in the OA context have been
374 regional, empirical or derived with a limited variety of satellite datasets, rendering an
375 effort to systematically exploit remote sensing assets (capitalizing on the recent
376 advent of satellite salinity measurements) absolutely timely. To-date there is only
377 regional application of satellite SST to address the issue of assessing OA ⁶², along
378 with two non-peer-reviewed attempts to calculate carbonate system products using
379 satellite SSS data ^{92,93}. Supported by good *in situ* measurement campaigns, especially
380 in places with currently poor *in situ* coverage such as the Arctic, satellite
381 measurements are likely to become a key element in understanding and assessing OA.

382

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386 **Author Contributions**

387 The manuscript was written through contributions of all authors. All authors have
388 given approval to the final version of the manuscript.

389

390 **Funding Sources**

391 This work was funded by the European Space Agency Support to Science Element
392 Pathfinders Ocean Acidification project (contract No. 4000110778/14/I-BG).

393

394 **ACKNOWLEDGMENT**

395 This work was enabled by European Space Agency (ESA) Support to Science
396 Element (STSE) Pathfinders Ocean Acidification project (contract No.
397 4000110778/14/I-BG). The authors gratefully acknowledge the assistance of Diego
398 Fernandez (STSE programme manager).

399

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417 Roberto Sabia is a specialist in remote sensing of salinity working for ESA.

418

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