## A particle model of rolling grain ripples under waves

## Andersen, Ken Haste

Published in:
Physics of Fluids

Link to article, DOI:
10.1063/1.1332390

Publication date:
2001

Document Version
Publisher's PDF, also known as Version of record

Link back to DTU Orbit

Citation (APA):
Andersen, K. H. (2001). A particle model of rolling grain ripples under waves. Physics of Fluids, 13(1), 58-64. DOI: 10.1063/1.1332390

## DTU Library

Technical Information Center of Denmark

## General rights

Copyright and moral rights for the publications made accessible in the public portal are retained by the authors and/or other copyright owners and it is a condition of accessing publications that users recognise and abide by the legal requirements associated with these rights.

- Users may download and print one copy of any publication from the public portal for the purpose of private study or research.
- You may not further distribute the material or use it for any profit-making activity or commercial gain
- You may freely distribute the URL identifying the publication in the public portal


# A particle model of rolling grain ripples under waves 

Ken Haste Andersen ${ }^{\text {a) }}$<br>Department of Hydrodynamics and Water Resources, the Technical University of Denmark, 2800 Kgs. Lyngby, Denmark; Center for Chaos and Turbulence Studies, the Niels Bohr Institute, University of Copenhagen, 2100 Ø, Copenhagen, Denmark;<br>and Dipartimento di Fisica, Universitá degli Studi di Roma La Sapienza, Piazzale Aldo Moro 2, I-00185 Roma, Italy

(Received 20 October 1999; accepted 18 October 2000)


#### Abstract

A simple model for the formation of rolling grain ripples on a flat sand bed by the oscillatory flow generated by a surface wave is presented. An equation of motion is derived for the individual ripples, seen as 'particles," on the otherwise flat bed. The model accounts for the initial appearance of the ripples, the subsequent coarsening of the ripples, and the final equilibrium state. The model is related to the physical parameters of the problem, and an analytical approximation for the equilibrium spacing of the ripples is developed. It is found that the spacing between the ripples scales with the square-root of the nondimensional shear stress (the Shields parameter) on a flat bed. The results of the model are compared with measurements, and reasonable agreement between the model and the measurements is demonstrated. © 2001 American Institute of Physics.


[DOI: 10.1063/1.1332390]

## I. INTRODUCTION

In the coastal zone where the water is relatively shallow, a ubiquitous phenomenon is the formation of ripples in the sand. These ripples have been described in the seminal work of Bagnold, ${ }^{1}$ who called them vortex ripples. The vortex ripples have been studied recently from a pattern-forming point of view. ${ }^{2-4}$ Bagnold also described another kind of ripple that was created from an initially flat bed, as a transient phenomenon before the creation of vortex ripples. These ripples were created by the rolling back and forth of individual grains on the flat bed, and were called rolling grain ripples. It is this latter class of ripples which is the topic of this article.

The most well-known type of ripple created by oscillatory motion is the vortex ripple, so called because of the strong vortices created by the sharp crest of the ripples. These are the ripples one encounters when swimming along a sandy beach. The shape of the ripples is approximately triangular, with sides being at the angle of repose of the sand. The length of the ripples scales with the amplitude of the oscillatory motion of the water near the bed, $a$. The dynamics of the vortex ripples and the creation of a stable equilibrium pattern have been recently described. ${ }^{4}$

In contrast to the vortex ripples, the rolling grain ripples have not (to the author's knowledge) been observed in the field, and have only been studied in controlled laboratory experiments, where it is possible to have a completely flat bed as initial condition. In this article, the term 'rolling grain ripple" refers to the ridges with triangular cross section occurring on an otherwise flat bed (Fig. 1 and Fig. 2) which is in agreement with Bagnold's definition of the rolling grain

[^0]ripples (see also plate 3 in Ref. 1). The length of the triangular cross section is small (typically smaller than 1 cm ), and much smaller than the spacing between the ridges, which again is typically smaller than the spacing between the vortex ripples. As the crest of the rolling grain ripples is rather sharp, a separation zone can be induced in the lee side of the ripple (a separation zone is an area where the flow is in the opposite direction of the main flow). This has also been found from numerical simulations of the flow over triangular ridges. ${ }^{5}$

Bagnold assumed that there were no separation zones behind the rolling grain ripples, and this has led to the adoption of the term rolling grain ripples for ripples without separation. It seems as if another type of ripple can emerge from a flat bed. These ripples initially have a sinusoidal shape, and thus no flat bed between them. The ripples grow and become sharper until they are steep enough to induce separation, whereafter they grow to become vortex ripples. From the author's own experiments ${ }^{5}$ and from Ref. 6, the formation of the ripples seems to depend on the initial preparation of the bed. If the grains are small and the bed well packed, the rolling grain ripples as described by Bagnold appear ('‘Bagnold-type rolling grain ripples'). If, however, the grains are large or the flat bed is carefully prepared, so as not to be packed, i.e., by letting the grains settle through the water, a mode with sinusoidal ripples can be observed. ${ }^{6}$ These "sinusoidal rolling grain ripples" have been described by a hydrodynamic instability of the wave boundary layer. The idea was originally formulated by Sleath, ${ }^{7}$ but was developed in great detail in a series of papers by Blondeaux, Foti, and Vittori. ${ }^{8-12}$ This kind of ripple is not described by the model which will be developed here.

The model to be presented is a 'particle model'" in the sense that it sees each ripple as a "particle" that interacts


FIG. 1. An example of a rippled bed viewed from above, with rolling grain ripples (bottom) and vortex ripples invading from the top. The experimental setup is a wave tank 60 cm wide and approximately 50 cm deep, containing sand with a median diameter of 0.2 mm .
with its neighboring particles. A heuristic equation of motion is written for each particle, taking into account the movement of the particle and the presence of the neighboring particles. When two particles collide, they merge and form a new, larger particle. This continues until a steady state is achieved.

After a more detailed description of the rolling grain ripples in Sec. II, the heuristic model is developed in Sec. III. It turns out, fortunately, that the parameters entering this model can be connected to measurable quantities (Sec. III A). In Sec. III B the model is solved numerically and analytically, and comparison with experimental measurements of rolling grain ripples is made in Sec. III C. The results are discussed in Sec. IV and conclusions can be found in Sec. V.

## II. THE ROLLING GRAIN RIPPLES

An example of a flat bed with rolling grain ripples coexisting with vortex ripples is seen in Fig. 1. In the middle of the picture the ripples have not yet formed and the bed is still flat, while on the top vortex ripples are seen to invade the flat bed. The vortex ripples are typically nucleated from the boundaries or from a perturbation in the bed. In the lower part of the picture the rolling grain ripples have formed on the flat bed, and are seen as the small bands of loose grains on top of the flat bed.

The flow over the bed created by the surface wave is oscillating back and forth in a harmonic fashion. This flow creates a shear stress on the bed, $\tau(t)$, which in nondimensional form reads

$$
\begin{equation*}
\theta(t)=\frac{\tau(t)}{\rho(s-1) g d} \tag{1}
\end{equation*}
$$

where $\rho$ is the density of water, $s$ is the relative density of the sand (for quartz sand in water $s=2.65$ ), $g$ is the gravity, and $d$ is the mean diameter of the grains. $\theta$ is usually called the Shields parameter. ${ }^{13}$ When the shear stress exceeds a critical value $\theta_{c}$, the grains start to move. For a turbulent boundary layer the value of the critical Shields parameter is $\theta_{c}$


FIG. 2. A sketch of a ripple on the flat bed with related quantities. $x_{i}$ is the position of the ripple $i, A_{i}$ is the area, $h$ is the height, and $\phi$ is the angle of repose of the sand.
$\approx 0.06 .{ }^{14}$ The grains which have become loosened from the bed start to move back and forth on the flat bed, and after a while the grains come to rest in parallel bands. In the lee side of each band the bed is shielded from the full force of the flow, creating a 'shadow zone"' where the grains move more slowly than in the upstream side of the bands.

Due to this shadow zone, more grains end up in the bands than leave the bands, and they grow until they form small ridges, the rolling grain ripples. When the rolling grain ripples are fully developed, no grains will be pulled loose from the bed in the space between them, and they are stable. However, in reality the rolling grain ripples are dominated by invading vortex ripples, ${ }^{2}$ which is the reason why they have not been observed in nature.

## III. A SIMPLE MODEL

The above scenario can be formulated mathematically by writing an equation of motion for each grain/particle. In the following an equation of motion for the particles is developed. In the beginning the particles represent the grains, but as the single grains quickly merge, the particles most of the time represent the ripples. First, the velocity of each particle is found, assuming that the particle is alone on the flat bed, and then the influence of the shadow zones from neighboring particles is taken into account.

Consider $N$ particles rolling on top of a rough, solid surface. Each particle is characterized by its position $x_{i}$ and its height $h_{i}$ (see Fig. 2). As the ripples are triangular, the area of each particle $A_{i}$ and their heights are related as

$$
\begin{equation*}
h_{i}=\sqrt{A_{i} \tan \phi} \tag{2}
\end{equation*}
$$

where $\phi$ is the angle of repose of the sand (approximately $33^{\circ{ }^{14}}$ ).


FIG. 3. Geometrical illustration of the quantities involved in the derivation of the " 1 /height'" law. $q_{\text {crest }}$ is the flux of sand over the crest. $\delta x h_{i}$ is the amount of sand needed to move the ripple forward a distance $\delta x$.


FIG. 4. An illustration of the shadow zone of one ripple, in the part of the wave period where the flow is from the left to the right. Below is seen the function $f(\Delta x)$ used to describe how a particle is slowed down when it enters the shadow zone of another particle. $\alpha_{s} h$ is the total length of the shadow zone.

The ripple moves back and forth more slowly than a single grain, according to the ' 1 /height'" law. This law is well known in the study of dunes in the desert ${ }^{15}$ or subaqueous dunes, ${ }^{16}$ and can be illustrated by a simple geometrical argument. Suppose that there is a flux of sand over the crest of a ripple or a dune $q_{\text {crest }}$ (Fig. 3). To make the ripple move a distance $\delta x$, an amount of sand $h \delta x$ is needed. As the sediment flux is the amount of sand per unit time, the velocity of the ripple is $u_{\text {ripple }}=q_{\text {crest }} / h \propto 1 / h$. If the height of the initial particles (the single grains) is assumed to be equal to the grain diameter $d$, the velocity of the particles can be related to that of the single grains as

$$
\begin{equation*}
u_{i}=\frac{d}{h_{i}} U_{g} \sin (\omega t) \tag{3}
\end{equation*}
$$

where $U_{g}$ is the velocity amplitude of the motion of a single grain and $\omega$ is the angular frequency of the oscillatory motion.

In the wake of each particle/ripple there is a shadow zone (Fig. 4), which is the area behind the particle where the absolute value of the shear stress is smaller than it would be on a flat bed. The length of the shadow zone is therefore larger than the length of the separation bubble formed by the particle (note that the shadow zone would be present even in the absence of separation). If the shadow zone is much smaller than the amplitude of water motion, the flow in the lee side of the ripple can be assumed to have sufficient time to become fully developed. The fully developed flow over a triangle is similar to that past a backward facing step in steady flow, which has been extensively studied (see, e.g., Tjerry ${ }^{17}$ ). In that case the relevant quantities, i.e., the length of the separation bubble, the length of the wake, etc., scale with the height of the step. As a first assumption, the shadow zone is therefore assumed to have a length which is proportional to the height of the particle: $\alpha_{s} h_{i}$. If a particle enters the shadow zone of another particle, it is slowed down according to the distance between the particles. This means that the actual velocity of a particle is $u_{i} f(\Delta x)$, where $\Delta x$ is the distance between the grain and the nearest neighbor up-
stream, $u_{i}$ is the velocity of the particle outside the shadow zone, and $f$ is a function determining the nature of the slowing down of the particle motion. A simple linear function is used, as shown in Fig. 4. The exact form of the function $f$ is not crucial, as will become evident later; the important parameter is the extent of the shadow zone as determined by $\alpha_{s}$.

It is now possible to write the equations of motion for the particles as a system of coupled ordinary differential equations (ODEs)

$$
\begin{equation*}
\dot{x}_{i}=\frac{d}{h_{i}} u_{g}(t) f \underbrace{f\left(\frac{x_{i}-x_{i-1}}{\alpha_{s} h_{i-1}} \frac{u(t)}{|u(t)|}\right)}_{\text {positive half period }} \underbrace{\left(\frac{x_{i}-x_{i+1}}{\alpha_{s} h_{i+1}} \frac{u(t)}{|u(t)|}\right)}_{\text {negative half period }} \tag{4}
\end{equation*}
$$

for $i=1, \ldots, N$, where $u_{g}(t)=U_{g} \sin (\omega t)$. The motion of a particle is thus made up of three parts: (i) the motion of the single undisturbed particle; (ii) the effect of the shadow from the particle to the left $(i-1)$, which might affect particle $i$ in the first half period, and (iii) the effect of the shadow of the particle to the right $(i+1)$ in the second half period.

When lengths are scaled by the diameter of the grains and time by the frequency $\omega$, it is possible to identify the three relevant dimensionless parameters of the model:
(i) $\alpha_{s}$, the length of the shadow zone divided by the height;
(ii) $a_{g} / d$, the amplitude of the motion of a single grain, divided by the grain diameter $\left(a_{g}=\omega U_{g}\right)$; and
(iii) $\lambda_{i} / d$, the initial distance between the grains; $\lambda_{i}$ $=L / N$, where $L$ is the length of the domain and $N$ is the initial number of grains.

## A. Relation to physical quantities

Even though the model seems quite heuristic, the parameters entering the model, $a_{g} / d, \lambda_{i} / d$, and $\alpha_{s}$, can be related to physical parameters describing the flow and the properties of the grains. The line of arguments presented here closely follows those used to derive the flux of sand on a flat bed (the bed load), as can be found in Ref. 14, or in Ref. 18. First, the velocity of a single grain will be derived, from which $a_{g} / d$ can be inferred. Thereafter, the initial number of grains in motion is found, from which follows $\lambda_{i} / d$. Finally, the length of the shadow zone $\alpha_{s}$ is discussed.

## 1. The velocity of the grains

The velocity of the grain can be found by considering the force balance on a single grain lying on the flat bed. The grain is subject to a drag force proportional to the square of the relative flow velocity $u_{r}=u_{n b}-u_{g}$, where $u_{n b}$ is the velocity near the bed and $u_{g}$ is the velocity of the grain

$$
\begin{equation*}
F_{d}=\frac{1}{2} C_{D} \rho A\left|u_{r}\right| u_{r}, \tag{5}
\end{equation*}
$$

where $A$ is the area of the grain and $C_{D}$ is a drag coefficient. The numerical sign is used to obtain the right sign of the force. The velocity profile in the vicinity of the bed is supposed to be logarithmic. As shown in Ref. 19, this is a reasonable assumption except when the flow reverses. However,


FIG. 5. A sketch of the shear stress, $\tau$, on the bed during a steady flow over a backward facing step with height $h$.
during reversal the velocities are small anyway and the accuracy is of minor importance. The logarithmic profile over a rough bed can be written as (e.g., Ref. 14)

$$
\begin{equation*}
u(y)=\frac{u_{f}}{\kappa} \ln \left(\frac{30 y}{k_{N}}\right), \tag{6}
\end{equation*}
$$

where $\kappa=0.41$ is the von Kármán constant, $u_{f} \equiv \sqrt{\tau / \rho}$ is the friction velocity, and $k_{N}$ is the Nikuradse roughness length. It is then possible to find the near bed velocity as the velocity at $y \approx 1 / 2 d$

$$
\begin{equation*}
u_{n b}=\xi u_{f}, \tag{7}
\end{equation*}
$$

where the constant $\xi$ can be determined from Eq. (6) by assuming $k_{N}=d$. Opposing the drag on the grain is the friction of the bed

$$
\begin{equation*}
F_{f}=-\mu W, \tag{8}
\end{equation*}
$$

where $\mu$ is a friction coefficient and $W=\rho g(s-1) d^{3} \pi / 6$ is the immersed weight of the grain. By making a balance of forces, $F_{d}+F_{f}=0$, the velocity of the grain can be found

$$
\begin{equation*}
u_{g}=\xi u_{f}\left(1-\sqrt{\left|\frac{\theta_{c}}{\theta}\right|}\right), \tag{9}
\end{equation*}
$$

where

$$
\begin{equation*}
\theta_{c}=\frac{4 \mu}{3 C_{D} \xi^{2}} \tag{10}
\end{equation*}
$$

is the critical Shields parameter. Usually, $\mu=\tan \phi \approx 0.65 .{ }^{14}$

## 2. The initial spacing of grains

Now that the velocity of the grains has been calculated, it still remains to determine the number of grains per area $n$ in motion. To this end, a small volume of moving sand at the top of the flat bed is considered. The balance of the forces acting on this volume is written as

$$
\begin{equation*}
\tau_{b}=\tau_{G}+\tau_{c} \tag{11}
\end{equation*}
$$

The interpretation of the terms is as follows: The parameter $\tau_{b}$ is the shear stress on the top of the bed load layer. It is assumed that this is equal to the shear stress on a fixed flat bed. $\tau_{G}$ is the stress arising from the intergranular collisions, giving rise to " grain stresses", ${ }^{18}$ modeled as $\tau_{G}=n \mu W$. It is assumed that the intergranular stress absorbs all the stress except the critical stress $\tau_{c}$; this is the so-called "Bagnold hypothesis." ${ }^{18}$ Making Eq. (11) nondimensional by dividing with $\rho(s-1) g d$, the number of grains in motion is found as


FIG. 6. Zoom of the movement of the particles in the first two wave periods. $a_{g} / d=35, \lambda_{f}=3.23$, and $\alpha_{s}=10.0$.

$$
\begin{equation*}
n=\frac{6}{\pi d^{2} \mu}\left(\theta-\theta_{c}\right) \tag{12}
\end{equation*}
$$

If $\theta<\theta_{c}$, then there are no grains in motion and $n=0 . n$ can also be viewed as the initial density of grains, and by assuming a square packing of the grains the initial distance between the grains becomes

$$
\begin{align*}
\frac{\lambda_{i}}{d} & =\frac{1}{\sqrt{n} d},  \tag{13}\\
& =\sqrt{\frac{\pi \mu}{6\left(\theta-\theta_{c}\right)}} \tag{14}
\end{align*}
$$

## 3. The length of the shadow zone

The last parameter $\alpha_{s}$, which characterizes the length of the shadow zone, is estimated by exploring the analogy with the backward facing step, which was suggested in Sec. II. In the backward facing step there is a zone with flow separation which extends approximately 6 step heights from the step; see Fig. 5. After approximately 16 step heights there is a point where the shear stress has a small maximum. The length of the shadow zone should be longer than the separation zone, but shorter than the point of the maximum in the shear stress, i.e., $6<\alpha_{s}<16$.

## B. Numerical and analytical solutions of the model

In the following section the behavior of the model is examined. To study the detailed behavior the set of coupled ordinary differential equations (4) is integrated numerically. It will be demonstrated that the model reaches a steady state, and an analytical expression for the spacing between the ripples in the steady state is developed.

The numerical simulations in this section are based on a simple example with $a_{g} / d=35$ and $\lambda_{i} / d=3.23$ and $\alpha_{s}$ is set to 10 . As initial condition all particles have an area of 1.0 $\pm 10 \%$, to add some perturbation. The initial number of particles $N$ in this example is 800 .

In the first few periods a lot of grains are colliding and merging (Fig. 6). As the ripples are formed and grow bigger,


FIG. 7. The development of the particles until steady state is reached. The lines show the positions of the particles at the end of each period.
the evolution slows down, until a steady state is reached (Fig. 7). The spacing between the ripples in the steady state shows some scatter around the average value, which is also observed in experiments. The variation in the average spacing at the steady state, $\lambda_{e q}$, for realizations with different initial random seed turned out to be on the order of $1 / N_{e q}$ where $N_{e q}$ is the final number of ripples.

To study the behavior of the average spacing of the ripples, a number of simulations are made where the parameters are varied one at a time. Each run is started from the initially disordered state.

Changing $a_{g} / d$ only results in a minor change in the spacing of the ripples [Fig. 8(a)]. The final spacing between the ripples does depend on the length of the shadow zone $\alpha_{s}$; the longer the shadow zone, the larger the wavelength of the ripples [Fig. 8(b)]. This can be used to estimate the average equilibrium spacing between the ripples. When the distance between two ripples is longer than the shadow zone of the ripples, they are no longer able to interact. This gives

$$
\begin{equation*}
\lambda_{e q}>\alpha_{s} h_{e q}, \tag{15}
\end{equation*}
$$

where subscript $e q$ denotes an average value at equilibrium. However, if the spacing between two ripples is just barely shorter than Eq. (15), they will be able to interact and eventually they will merge. One can therefore expect to find spacings up to $\lambda_{e q}=2 \alpha_{s} h_{e q}$. Assuming that the average length is in between the two bounds, one gets

$$
\begin{equation*}
\lambda_{e q}=\gamma \alpha_{s} h_{e q}, \quad 1<\gamma<2 \tag{16}
\end{equation*}
$$

where $\gamma$ can be found by comparing the results from the full simulations with Eq. (16). The height of the ripples at equilibrium can be found by splitting the initial number of particles evenly onto the equilibrium ripples. Then the average area at equilibrium is $A_{e q}=\lambda_{e q} / \lambda_{i} d^{2}$ and from Eq. (2) it follows that the height is $h_{e q} / d=\sqrt{\tan \phi \lambda_{e q} / \lambda_{i}}$, which gives an average equilibrium spacing

$$
\begin{equation*}
\frac{\lambda_{e q}}{d}=\gamma^{2} \frac{\alpha_{s}^{2} d \tan \phi}{\lambda_{i}}=\alpha_{s}^{2} \gamma^{2} \sqrt{\frac{6 \tan \phi}{\pi}} \sqrt{\theta-\theta_{c}} . \tag{17}
\end{equation*}
$$



FIG. 8. The spacing between the rolling grain ripples at equilibrium as a function of the initial density of grains. The basic example is shown with the circles: $a_{g} / d=35$ and $\alpha_{s}=10.0$. In (a) the value of $a_{g} / d$ is varied, for pluses: $a_{g} / d=20$, crosses: $a_{g} / d=50$. In (b) $\alpha_{s}$ is varied; pluses: $\alpha_{s}=7$, crosses: $\alpha_{s}=13$.

The equilibrium spacing is therefore found to be proportional to $\sqrt{\theta-\theta_{c}}$ with the constant of proportionality being made up of $\alpha_{s}, \gamma$, and various geometrical factors. All the quantities related to the dynamical evolution of the ripples, i.e., the velocity of the ripples, the shape of the function $f(\Delta x)$, etc. do not enter into the expression.

## C. Comparison with experiments

The only parameter that has not been accurately determined is $\alpha_{s}$. The value of this parameter can be estimated by comparison with measurements.

In 1976, Sleath made a series of experiments, measuring the spacing between rolling grain ripples. ${ }^{7}$ The ripples were formed on a flat tray oscillating in still water using sand of two different grain sizes: 0.4 and 1.14 mm . To compare with the experiments the value of $\theta$ must be calculated. $\theta$ reflects the number of grains in motion, and it is assumed that the grains which are set in motion when the shear stress on the bed is at a maximum are kept in motion throughout the wave period. Therefore, the maximum value of the Shields parameter during the period, $\theta_{\text {max }}$, is used. To evaluate $\theta_{\text {max }}$, the


FIG. 9. Comparison between the measured wave lengths of rolling grain ripples (circles) (Ref. 7), with results from the model (pluses), and from Eq. (16) with $\gamma=1.40$ (the line). The value of $\alpha_{s}$ is 15.0. The critical Shields parameter is 0.04 and $\mu=0.65$.
shear stress on the bed has to be estimated. The maximum shear stress on the bed, $\tau_{\max }$, during a wave period can be found using the concept of a constant friction factor $f_{w}:{ }^{20}$

$$
\begin{equation*}
\tau_{\max }=\frac{1}{2} \rho f_{w} u_{\max }^{2} \tag{18}
\end{equation*}
$$

with $u_{\text {max }}$ being the maximum near-bed velocity. The friction factor can be estimated using the empirical relation ${ }^{14}$

$$
\begin{equation*}
f_{w}=0.04\left(\frac{a}{k_{N}}\right)^{-0.25} \tag{19}
\end{equation*}
$$

where $a=u_{\max } \omega$ and $k_{N} \simeq d$.
In this way the range of Shields parameters in the experiments was found to be from the critical Shields parameter to $\theta=0.42$. For the high Shields parameters the rolling grain ripples were reported to be very unstable and quickly to develop into vortex ripples. In these cases, the measured ripple spacing then reflects the spacing between the rolling grain ripples before they developed into vortex ripples. ${ }^{7}$

In Fig. 9 the experimental results are compared with runs of the model using $\alpha_{s}=15.0$ (the reason for this particular value will be shown shortly) and $N=10000$. By fitting all the runs to Eq. (16) it was found that $\gamma=1.40$. The results using Eq. (17) and $\gamma=1.40$ are shown with a line.

First, it is seen that Eq. (17) predicts the results from the full model Eq. (4) well. The correspondence between the model and the experiments is reasonable, but there are some systematic discrepancies, which will be discussed.

There are a few points with small ripple spacing for which the model does not fit the measurements. These measurements have a Shields parameter very near the critical (i.e., just around the onset of grain motion), which implies some additional complications. The grains used in the experiment were not of a uniform size; rather, they were part of a distribution of grain sizes, and the grain size reported is then the median of the distribution, $d_{50}$. The Shields parameter is calculated using the median of the distribution, but actually one could calculate a Shields parameter for different fractions of the distribution, thus creating a $\theta_{10}$, a $\theta_{50}$, etc. When $\theta_{50}$ is smaller than the critical Shields parameter, $\theta_{10}$ might still be higher than the critical Shields parameter. This implies that grains with a diameter smaller than $d_{50}$ will be in motion, while the larger grains will stay in the bed. As only $d_{50}$ is used in the calculation of the equilibrium ripple spacing, the distance between the grains $\lambda_{i}$ will be overestimated near the critical Shields parameter, where the effect of the polydispersity is expected to be strongest. An overestimation of $\lambda_{i}$ will lead to an underprediction of the ripple length, which is exactly what is seen in Fig. 9.

There are also three points from the experiments taken at very large Shields parameters that are not well predicted by the model. As already mentioned, these points are probably doubtful because of the very fast growth of vortex ripples. It is therefore reasonable to assume that vortex ripples invaded the rolling grain ripples before these had time to reach their full length.

To find a reasonable value of $\alpha_{s}$ Eq. (17) was fitted to the experimental points. To avoid the points which might be of doubtful quality, as discussed above, only the points in the
range $0.075<\theta<0.3$ were used. This gave the value of $\alpha_{s}$ $=15.0$, in agreement with the qualitative arguments in Sec. III A 3 .

## IV. DISCUSSION OF THE RESULTS

From the comparison of the model with measurements the model confidently reproduces the experiments.

In the model the number of grains in motion is constant (even though the number of particles changes). In an experimental situation, however, new grains might be lifted from the bed and added to the initial number of grains in motion. As the part of the flat bed between the ripples is covered by the shadow zones of the particles, these stretches will be shielded from the full force of the flow, and only very slowly will new grains be loosened here. This small addition of new grains will result in a slow growth of the rolling grain ripples, and they eventually grow into vortex ripples. This slow growth is very well illustrated by recent measurements, ${ }^{2}$ but not covered by the present model.

## V. CONCLUSION

In conclusion, a model has been created which explains the creation and the equilibrium state of rolling grain ripples of the type described by Bagnold. The final distance between the ripples is proportional to $\sqrt{\theta-\theta_{c}}$. The model has been compared with measurements with reasonable agreement.

## ACKNOWLEDGMENTS

It is a pleasure to thank Professor Tomas Bohr, Dr. Clive Ellegaard, Dr. Enrico Foti, Professor Jørgen Fredsøe, and Dr. Vachtang Putkaradze for useful discussions.
${ }^{1}$ R. A. Bagnold, 'Motion of waves in shallow water, interaction between waves and sand bottoms,'’ Proc. R. Soc. London, Ser. A 187, 1 (1946).
${ }^{2}$ A. Stegner and J. E. Wesfreid, 'Dynnamical evolution of sand ripples under water,' 'Phys. Rev. E 60, R3487 (1999).
${ }^{3}$ M. A. Scherer, F. Melo, and M. Marder, "Sand ripples in an oscillating annular sand-water cell,'" Phys. Fluids 11, 58 (1999).
${ }^{4}$ K. H. Andersen, M.-L. Chabanol, and M. v. Hecke, ''Dynamical models for sand ripples beneath surface waves," arXiv.org/abs/ cond-mat/0003475, 2000 (submitted).
${ }^{5}$ K. H. Andersen, 'Ripples beneath surface waves and topics in shell models of turbulence," Ph.D. thesis, the Niels Bohr Institute, University of Copenhagen, http://www.isva.dtu.dk/~ ken/Thesis.html (1999).
${ }^{6}$ E. Foti, personal communication.
${ }^{7}$ J. F. A. Sleath, ''On rolling-grain ripples,'" J. Hydraul. Res. 14, 69 (1976).
${ }^{8}$ P. Blondeaux, 'Sand ripples under sea-waves. I. Ripple formation," J. Fluid Mech. 218, 1 (1990).
${ }^{9}$ G. Vittori and P. Blondeaux, 'Sand ripples under sea waves. II. Finiteamplitude development,' J. Fluid Mech. 218, 19 (1990).
${ }^{10} \mathrm{G}$. Vittori and P. Blondeaux, ''Sand ripples under sea waves. III: Brickpattern ripple formation,'" J. Fluid Mech. 239, 23 (1991).
${ }^{11} \mathrm{E}$. Foti and P. Blondeaux, "Sea ripple formation: The turbulent boundary layer case,'" Coastal Eng. 25, 227 (1995).
${ }^{12} \mathrm{E}$. Foti and P. Blondeaux, ''Sea ripple formation: The heterogeneous sediment case,'" Coastal Eng. 25, 237 (1995).
${ }^{13}$ I. A. Shields, Anwendung der Aehnlichkeitsmechanik und der Turbulenzforschung auf die Geschiebebewegung (Mitt. Preuss. Versuchsanstalt, Berlin, 1936), p. 26.
${ }^{14}$ J. Fredsøe and R. Deigaard, Mechanics of Coastal Sediment Transport (World Scientific, Singapore, 1992).
${ }^{15}$ H. Nishimori, M. Yamasaki, and K. H. Andersen, "A simple model for the various pattern dynamics of dunes," Int. J. Mod. Phys. B 12, 257 (1998).
${ }^{16}$ J. Fredsøe, "'The stability of a sandy river bed,"’ in Issues and Directions in Hydraulics, edited by T. Nakato and R. Ettema (Iowa Institute of Hydraulic Research, 1996), pp. 99-113.
${ }^{17}$ S. Tjerry, 'Morphological calculation of dunes in alluvial rivers," Ph.D. thesis, ISVA, the Technical University of Denmark, 1995.
${ }^{18}$ A. Kovacs and G. Parker, "A new vectorial bedload formulation and its application to the time evolution of straight river channels,"
J. Fluid Mech. 267, 153 (1994).
${ }^{19}$ B. L. Jensen, B. M. Sumer, and J. Fredsøe, ''Turbulent oscillatory boundary layers at high Reynolds numbers," J. Fluid Mech. 206, 265 (1989).
${ }^{20}$ I. G. Jonsson and N. A. Carlsen, 'Experimental and theoretical investigations in an oscillatory turbulent boundary layer," J. Hydraul. Res. 14, 45 (1976).


[^0]:    ${ }^{\text {a) }}$ Electronic mail: ken@isva.dtu.dk

