

Reconciling tracer and float observations of the export pathways of Labrador Sea Water

S. F. Gary¹, M. S. Lozier¹, A. Biastoch², C. W. Böning²

A. Biastoch, GEOMAR Helmholtz Centre for Ocean Research, Düsternbrooker Weg 20, 24105 Kiel, Germany. (abiastoch@geomar.de)

C. W. Böning, GEOMAR Helmholtz Centre for Ocean Research, Düsternbrooker Weg 20, 24105 Kiel, Germany. (cboening@geomar.de)

S. F. Gary, Division of Earth and Ocean Sciences, Nicholas School of the Environment, Duke University, Box 90227, Durham, NC 27708, USA. (stefan.gary@duke.edu)

M. S. Lozier, Division of Earth and Ocean Sciences, Nicholas School of the Environment, Duke University, Box 90227, Durham, NC 27708, USA. (mslozier@duke.edu)

¹Division of Earth and Ocean Sciences,
Nicholas School of the Environment, Duke
University, Durham, NC, USA.

²Division of Ocean Circulation and
Climate Dynamics, GEOMAR Helmholtz
Centre for Ocean Research, Kiel, Germany.

For more than fifty years, it has been generally accepted by oceanographers that the Deep Western Boundary Current (DWBC) is the principal conduit of recently-convected Labrador Sea Water (LSW) exported from the high-latitude North Atlantic to the equator. Supporting this supposition is observational evidence that the waters of the DWBC have consistently greater equatorward velocities, higher concentrations of passive tracers, and younger ages compared to ocean interior waters. However, recent observations and simulations of floats launched in the DWBC in the Labrador Sea show that most water parcels are quickly ejected from the DWBC and follow instead interior pathways to the subtropics. Here, we show that tracer observations from the last three decades are compatible with the existence of both DWBC and basin-interior export pathways. From analyses of observational data and model output, we find that equatorward transport in the basin interior is consistent with the large-scale vorticity balance at mid-depth. Furthermore, from the modeling analysis we show that despite higher, localized concentrations of tracer and particles in the DWBC, only 5% of particles released in the Labrador Sea are transported from the subpolar to subtropical gyre via a continuous DWBC pathway. Thus, the interior pathway is a significant contributor to LSW export.

1. Introduction

Atmospheric gases are mixed into the high-latitude surface waters which are brought to depth by wintertime convection and then transported equatorward, through the North Atlantic [*Doney and Jenkins, 1994*] [*Smethie, 1993*] [*Smethie et al., 2000*] [*Sabine et al., 2004*]. The deep ocean, therefore, represents a sink of anthropogenic CO₂ and chlorofluorocarbons (CFCs). Since CO₂ and CFCs are greenhouse gases and since dissolved CO₂ contributes to ocean acidification, an understanding of the mechanisms distributing these gases in the deep ocean is of interest.

Until recently, the dominant paradigm describing the circulation of the deep waters of the North Atlantic was based on a first order vorticity balance [*Stommel, 1958*]. This balance predicts that all deep flow in the basin interior is poleward and that the only possible equatorward flow at depth is along the western boundary in the Deep Western Boundary Current (DWBC). Another explanation for the DWBC is eddies interacting with steeply sloping topography [*Vallis and Maltrud, 1993*]. Regardless of its driving mechanism, the DWBC has been widely viewed as the main, continuous conduit for the equatorward export of deep waters.

Direct velocity measurements of the DWBC, conducted at several locations along the western boundary of the North Atlantic, confirm that velocities in the DWBC are equatorward and greater than the velocities in the basin interior [*Dengler et al., 2006*] [*Fischer et al., 2004*] [*Lee et al., 1996*] [*Schott et al., 2004*]. Furthermore, DWBC waters contain consistently higher concentrations of CFCs than the interior, an indication that DWBC

waters are closely tied to the high-latitude regions where deep convection acts to bring CFCs to depth [*Doney and Jenkins, 1994*] [*Smethie, 1993*] [*Smethie et al., 2000*].

In contrast to the Eulerian observations cited above, Lagrangian observations achieved with neutrally-buoyant float trajectories show that particles launched in the DWBC are likely to exit the DWBC and transit the basin interior as they move from high to low latitudes [*Bower et al., 2009*] [*Fischer and Schott, 2002*] [*Lavender et al., 2000*] [*Bower and Hunt, 2000*].

The central question of this paper stems from the apparent contradictory views of the deep circulation in the North Atlantic as revealed by tracer and float observations: Are recently ventilated waters exported to low latitudes along a dominant DWBC pathway, as suggested by high CFC concentrations in the boundary current, or does the DWBC coexist with significant interior pathways, as suggested by float observations?

2. Data Sources and Processing

2.1. Observations

CFC-11 observations are from the CARINA V1.0 [*The CARINA Group, 2009*] and GLODAP V1.0 [*Key et al., 2004*] databases. Cruises that are not in CARINA and GLODAP, but are in WOD09 [*National Oceanographic Data Center, 2010*], were also included in our database. We included data from Transient Tracers in the Ocean North Atlantic Sector (TTO-NAS) Leg 7 cruise (J. Bullister and T. Tanhua, personal communication) and Line W [*Toole et al., 2011*]. Finally, we use the updated hydrographic database of *Gary et al.* [2011] including all bottle, CTD, and Argo hydrographic data in WOD09 [*Na-*

tional Oceanographic Data Center, 2010] up to 2011. A climatology was constructed from this database with Hydrobase2 [*Lozier, 1995*] [*Curry, 2010*].

2.2. Model and Simulated Lagrangian Trajectories

The model used in this study is the ORCA025 global, $1/4^\circ$ eddy-permitting resolution, general circulation model [*Barnier et al., 2006*] [*The DRAKKAR Group, 2007*]. This particular model run is a 1958 to 2004 hindcast and is described in more detail by *Gary et al.* [2011]. The strength and variability of the ORCA025 DWBC, all along the boundary, is similar to observed values.

For the ORCA025 tracer simulation, velocity and diffusivity fields at every integration time step were used to update temperature, salinity, and CFC-11 fields. Identical operators and parameterizations were applied to all tracers. There was no CFC-11 during the first year of the simulation, 1958, to allow the model to reach an initial geostrophic balance after the shock from the initial conditions. For 1959 to 2004, the ocean surface took up CFC-11 [*Lachkar et al., 2007*] and monthly mean tracer distributions were saved.

Three-dimensional particle trajectories are computed offline with ORCA025 5-day average velocity fields [*Gary et al., 2011*] [*Blanke and Grima, 2010*]. Particles are launched every ten days from 1980 to 2004 at the Labrador Sea surface between 62°W to 48°W and 53°N to 65°N with temperatures between 3.0°C and 4.0°C , generating 609,055 trajectories. Float positions, at 30-day intervals, are calculated for 50 years by cycling through the 1980-2004 velocity fields with a temporal discontinuity at Dec 31, 2004 to Jan 1, 1980.

Particle position probability volumes are constructed by subdividing the North Atlantic into 0.25° by 0.25° by 200 m boxes, counting the number of times particles are present

in each box, including repetitions, and dividing by the number of particle positions. Different bin sizes (0.1° to 1° horizontally, 50 to 500 m vertically) generate similar results. Since we compute probabilities using full length trajectories, probabilities reflect the integrated effect of the passage of particles over the simulation duration. Therefore, probability volumes constructed from 50-year trajectories are comparable to CFC concentrations measured decades after the start of tracer accumulation.

3. Tracers in the boundary and interior

A representative section of CFC concentrations in the North Atlantic (Figure 1a) reveals high concentrations near the surface and a local maximum centered at 1000 m. This mid-depth maximum is associated with LSW that was brought to depth in the Labrador Sea during winter convection [Smethie *et al.*, 2000] [Smethie, 1993]. This CFC signature is aligned with the LSW density range $\sigma_\theta = 27.68$ to 27.80 kg m^{-3} [Dengler *et al.*, 2006]. A smaller, deeper, local maximum of CFC, centered at 3000 m, is attributed to dense overflow waters originating in the Norwegian and Greenland Seas [Smethie *et al.*, 2000] [Smethie, 1993]. The highest CFC concentration below the surface layer (green star, Figure 1a) is against the continental slope and resides in the upper portion of the DWBC. There is also a significant amount of tracer in the interior in addition to the DWBC.

The mid-depth, high CFC concentration core in the DWBC is reproduced in the CFC simulation within ORCA025. As in observations, the model exhibits elevated levels of CFC in the interior and a strong CFC signal along the boundary. The lower CFC concentrations in ORCA025 reflect the fact that CFC was added to the simulation after January 1, 1959, while the real ocean has been taking up small amounts of CFC since the 1930s.

One feature of the observed CFC distribution that is not reproduced by the model is the weaker, deeper core of CFC concentrations associated with the overflow waters. ORCA025 uses a partial step representation of bathymetry and a bottom boundary layer parameterization. This combination, while generally advantageous, does not adequately resolve overflow plumes. In case of overflow waters, additional mixing erases the tracer signal [*Griffies et al.*, 2005]. However, CFC in LSW is not significantly altered by this effect because LSW resides at intermediate depths.

We define the DWBC CFC core as the highest CFC-11 concentration below 500 m and inshore of the 4000 m isobath, the latter an estimate of the offshore edge of the DWBC since absolute velocity observations are not always available. The sorting of DWBC CFC observations by distance along the western boundary and date reveals a DWBC tracer core that appears to propagate continuously and steadily equatorward (Figure 2a). In contrast, the DWBC tracer core in the model exhibits a discontinuity near the Tail of the Grand Banks (TGB) (Figure 2b). However, when the model CFC fields are subsampled at the approximate location and date of each observation, the discontinuity at TGB is no longer resolved (Figure 2c). Therefore, we surmise that the steady progression of the observed tracer signal in the DWBC is likely an artifact of limited observations.

The discontinuity of the tracer signal at the TGB (Figure 2b), at the approximate boundary between the subpolar and subtropical gyres, is located where observations and simulations in ORCA025 and other models [*Bower et al.*, 2009] [*Gary et al.*, 2011] show that Lagrangian particles are ejected from the DWBC. While strong advective pulses of CFC in the mid to late 1990s, present in the model (Figure 2b) and suggested by

observations (Figure 2a), likely reflect intense LSW production in the early 1990s, these pulses extend to only a few 1000 km along the boundary and are intermittent. Though the loss of tracer is maximized in the vicinity of the TGB, it occurs along the whole western boundary, consistent with the well-known discrepancy between observed tracer signal propagation rates (1 cm s^{-1}) and observations of pure advection in the DWBC (10 cm s^{-1}) [Smethie et al., 2000] [Doney and Jenkins, 1994] [Bower and Hunt, 2000].

4. Lagrangian pathways of LSW

To understand differences between the distribution of LSW described by these tracer fields, modified by advection and diffusion (Figure 1a,b), and that described by Lagrangian pathways, which provide a purely advective view, we compute trajectories of 6.1×10^5 particles launched at the surface of the Labrador Sea in ORCA025, construct a particle position probability volume from the resulting trajectory ensemble, and compare the probabilities to simulated and observed tracer fields. The comparison is quite favorable: the greatest particle probabilities are in the DWBC (Figure 1c), coincident with the highest tracer concentrations (Figure 1a,b). These probabilities decrease, as do tracer concentrations, toward the interior. The favorable comparison suggests that the observed and modeled tracer signal is generally dominated by advection. As such, our interpretation of the tracer field should not markedly deviate from our interpretation of particle pathways. We do, however, note one difference: the relatively high particle probabilities at 19°N coincide with only a weak local tracer concentration maximum, a difference attributed to the accumulated impact of small scale diffusion a long distance away from the LSW source region.

Overall, we note from this comparison that although there is a strong probability and tracer signal at the boundary, many particles and much tracer reside in the basin interior.

A qualitative comparison of the spreading of tracer and particles is presented in the accompanying movie of simulated tracer and particle column inventories (Animation 1). At the beginning of the animation, particles and tracer move from the Labrador Sea into the Irminger Sea, Newfoundland Basin, and Eastern Atlantic: the three traditionally accepted pathways for LSW [*Talley and McCartney, 1982*]. After several years have passed in the animation, the export of LSW from the subpolar to subtropical gyres is visible in both the DWBC and the interior.

The relative importance of the DWBC and interior export pathways is quantified by inventories of simulated particles in the DWBC (Figure 2d) [*Bower et al., 2009*]. We define the boundary between the interior and the DWBC as the 4000 m isobath. Since our focus is the export of LSW, only trajectories of the 5.3×10^4 particles that crossed 25°N within 50 years and never went onto the continental shelf are used for this calculation. Furthermore, our inventory only includes particle first arrival positions at each 0.5° bin in the DWBC. Any recirculation is ignored. Although 75% of the particles cross 48°N in the DWBC, they are quickly ejected from the DWBC as noted above. Most particles in the DWBC have strayed into the interior and less than 5% of particles follow a direct and continuous pathway along the boundary. Even including the particles exchanged between the DWBC and the interior, only 25% reside in the DWBC at each location between 44°N and 30°N . South of 26°N , 65% of the particles have rejoined the DWBC.

Lagrangian data also allow for the quantification of particle spreading time scales. We calculate Lagrangian age by averaging the particle transit times reaching each 0.25° by 0.25° by 200 m box in the North Atlantic. The ages constructed from our simulation (Figure 3a,b) are similar to ages from tracer observations at the depth of LSW [*Smethie et al.*, 2000] [*Fine et al.*, 2002]. The DWBC is clearly the fastest LSW export pathway to the subtropics (Figure 3a,b).

5. Vorticity balance in the interior

Finally, we test if the observed equatorward movement of LSW in the basin interior [*Bower et al.*, 2009] is consistent with large-scale dynamical theory using the first order vorticity balance between water column stretching and planetary vorticity:

$$\beta v = -f \nabla_H \cdot \mathbf{u} \quad (1)$$

where f is the Coriolis parameter, $\mathbf{u} = (u, v, w)$ is velocity, $\beta = \partial f / \partial y$, y is latitude, z is the vertical coordinate, and the H subscript denotes the horizontal divergence in z -level or isopycnal coordinates [*Stommel*, 1958] [*Gary et al.*, 2011]. This balance applies to waters at depth isolated from baroclinic effects, forcing and dissipation. Equation 1 states that a local divergence of the horizontal flow, corresponding to water column compression, is dynamically consistent with equatorward velocity. In both observations and ORCA025, the layer thickness decreases as a particle follows an equatorward path (Figure 4). This equatorward flow, which is not necessarily purely meridional, is consistent with Equation 1.

Layer potential vorticity, f/h , where h is the thickness between two isopycnals, is conserved for fluid isolated from surface forcing and/or dissipation. In the absence of eddies,

we expect particles to follow contours of mean f/h . However, regions of homogenized or closed contours of mean f/h (Figure 4) are signatures of an eddy-driven mean flow across mean f/h contours, allowing equatorward flow in the interior [Lozier, 1997].

The large difference in Figure 4 between the observations and ORCA025 in the northern subpolar gyre is due to the ORCA025 $\sigma_\theta = 27.80 \text{ kg m}^{-3}$ isopycnal being much shallower than observations in that region. We note that the modeled f/h field exhibits weaker gradients compared to the observed field in this region, suggesting that the model may overestimate the spread of LSW into the interior. However, the spread of observed float pathways in the interior are indistinguishable from that of the simulated trajectories [Gary *et al.*, 2011].

6. Summary and Conclusions

We have shown that ocean tracer, density, and Lagrangian observations are consistent with the export of LSW via an interior, equatorward, pathway in addition to the traditionally accepted DWBC pathway. Our simulations, in agreement with observations, show that the DWBC is on average the fastest pathway for LSW export. In the model, 75% of simulated LSW particles are transported in the interior between 45°N and 30°N . Although the DWBC gradually regains most of the particles it lost by about 30°N , fewer than 5% of particles followed a continuous path along the DWBC. These conclusions agree with other studies which show that eddy-driven deep recirculation gyres are a plausible mechanism for the existence of an interior pathway [Gary *et al.*, 2011] [Lozier, 1997] and these large-scale, interior flow patterns shape the spreading of tracer [Lozier, 1999].

Acknowledgments. The authors are thankful for financial support from the US National Science Foundation and the German Academic Exchange Service (DAAD). S.F.G. is thankful to John Bullister and Toste Tanua for sharing CFC-11 data.

References

- Barnier, B., et al. (2006), Impact of partial steps and momentum advection schemes in a global ocean circulation model at eddy permitting resolution, *Ocean Dynamics*, *56*, 543–567.
- Blanke, B., and N. Grima (2010), Ariane v2.2.6, *Tech. rep.*, Laboratoire de Physique des Océans, <http://stockage.univ-brest.fr/~grima/Ariane/>.
- Bower, A. S., and H. D. Hunt (2000), Lagrangian observations of the Deep Western Boundary Current in the North Atlantic Ocean Part I: Large scale pathways and spreading rates, *Journal of Physical Oceanography*, *30*, 764–783.
- Bower, A. S., M. S. Lozier, S. F. Gary, and C. W. Böning (2009), Interior pathways of the North Atlantic Meridional Overturning Circulation, *Nature*, *459*, 243–247.
- Curry, R. G. (2010), Hydrobase2: A global database of hydrographic profiles and tools for climatological analysis, *Tech. rep.*, WHOI, <http://www.whoi.edu/science/PO/hydrobase/>.
- Dengler, M., J. Fischer, F. A. Schott, and R. J. Zantopp (2006), Deep Labrador Current and its variability in 1996-2005, *Geophysical Research Letters*, *33*, doi: 10.1029/2006GL026702, l21S06.
- Doney, S. C., and W. J. Jenkins (1994), Ventilation of the Deep Western Boundary Current and abyssal western North Atlantic: Estimates from tritium and ^3He distributions,

Journal of Physical Oceanography, 24, 638–659.

Fine, R. A., M. Rhein, and C. Andri  (2002), Using a CFC effective age to estimate propagation and storage of climate anomalies in the deep western North Atlantic Ocean,

Geophysical Research Letters, 29, doi:10.1029/2002GL015618.

Fischer, J., and F. A. Schott (2002), Labrador Sea Water tracked by profiling floats - from the boundary current to the open North Atlantic, *Journal of Physical Oceanography*,

32, 573–584.

Fischer, J., F. A. Schott, and M. Dengler (2004), Boundary circulation at the exit of the Labrador Sea, *Journal of Physical Oceanography*, 34, 1548–1570.

Gary, S. F., M. S. Lozier, C. W. B ning, and A. Biastoch (2011), Deciphering the pathways for the deep limb of the Meridional Overturing Circulation, *Deep Sea Research II*, 58,

1781–1797.

Griffies, S. M., et al. (2005), Formulation of an ocean model for global climate simulations, *Ocean Sciences Discussions*, 2, 165–246.

Key, R. M., et al. (2004), A global ocean carbon climatology: Results from Global Data Analysis Project (GLODAP), *Global Biogeochemical Cycles*, 18, doi:

10.1029/2004GB002247.

Lachkar, Z., J. C. Orr, J.-C. Dutay, and P. Delecluse (2007), Effects of mesoscale eddies on global ocean distributions of CFC-11, CO₂ and $\delta^{14}\text{C}$, *Ocean Science*, 3, 461–482.

Lavender, K. L., R. E. Davis, and W. B. Owens (2000), Mid-depth recirculation observed in the interior Labrador and Irminger seas by direct velocity measurements, *Nature*,

407, 66–69.

- Lee, T. N., W. E. Johns, R. J. Zantopp, and E. R. Fillenbaum (1996), Moored observations of western boundary current variability and thermohaline circulation at 26.5°N in the subtropical North Atlantic, *Journal of Physical Oceanography*, *26*, 962–983.
- Lozier, M. S., W. B. Owens, and R. G. Curry (1995), The climatology of the North Atlantic, *Progress in Oceanography*, *36*, 1–44.
- Lozier, M. S. (1997), Evidence for large-scale eddy-driven gyres in the North Atlantic, *Science*, *277*, 361–367.
- Lozier, M. S. (1999), The impact of the mid-depth recirculation on the distribution of tracers in the North Atlantic, *Geophysical Research Letters*, *26*, 219–222.
- National Oceanographic Data Center (2010), World Ocean Database Select and Search, *Tech. rep.*, National Oceanic and Atmospheric Administration, <http://www.nodc.noaa.gov/OC5/SELECT/dbsearch/dbsearch.html>.
- Sabine, C. L., et al. (2004), The oceanic sink for anthropogenic CO₂, *Science*, *305*, 367–371.
- Schott, F. A., R. Zantopp, L. Stramma, M. Dengler, J. Fischer, and M. Wibaux (2004), Circulation and deep-water export at the western exit of the subpolar North Atlantic, *Journal of Physical Oceanography*, *34*, 817–843.
- Smethie, W. M. (1993), Tracing the thermohaline circulation in the western North Atlantic using chlorofluorocarbons, *Progress in Oceanography*, *31*, 51–99.
- Smethie, W. M., R. A. Fine, A. Putzka, and E. P. Jones (2000), Tracing the flow of North Atlantic Deep Water using chlorofluorocarbons, *Journal of Geophysical Research*, *105*, 14,297–14,323.

- Smith, W. H. F., and P. Wessel (1990), Gridding with continuous curvature splines in tension, *Geophysics*, *55*, 293–305.
- Stommel, H. (1958), The abyssal circulation, *Deep Sea Research*, *5*, 80–82.
- Talley, L. D., and M. S. McCartney (1982), Distribution and circulation of Labrador Sea Water, *Journal of Physical Oceanography*, *12*, 1189–1205.
- The CARINA Group (2009), Carbon in the Atlantic Ocean Region - the CARINA project: Results and data, version 1.0, *Tech. rep.*, Carbon Dioxide Information Analysis Center, Oak Ridge National Laboratory, US Department of Energy, Oak Ridge, TN, doi: 10.3334/CDIAC/otg.CARINA.ATL.V1.0.
- The DRAKKAR Group (2007), Eddy-permitting ocean circulation hindcasts of the past decades, *CLIVAR Exchanges*, *12*(42), 8–10.
- Toole, J. M., et al. (2011), Transport of the North Atlantic Deep Western Boundary Current about 39°N, 70°W: 2004–2008, *Deep Sea Research II*, *58*, 1768–1780.
- Vallis, G. K., and M. E. Maltrud (1993), Generation of mean flows and jets on a beta plane and over topography, *Journal of Physical Oceanography*, *23*, 1346–1362.

Figure 1. Observed and modeled CFC-11 concentrations and particle position probability near 66°W. a, CFC-11 concentrations measured along R/V Knorr cruise 316 from 10/24/2003 to 11/12/2003. b, November 2003 mean CFC-11 concentration from ORCA025. Green stars are the instantaneous locations of the DWBC CFC core. c, Particle position probability computed from 609,055 simulated particles launched at the surface of the Labrador Sea in ORCA025. Gray lines are the average locations of the $\sigma_\theta = 27.68$ and 27.80 kg m^{-3} isopycnals (the characteristic LSW density range) from the hydrographic climatology (a) or from the model (b and c). Inset maps show section locations.

Figure 2. Spreading rates of the CFC-11 signal in the DWBC. a, Magnitude of the DWBC CFC core at each observed section of the DWBC. The shaded and contoured field is interpolated [*Smith and Wessel, 1990*] between observations (black dots) onto a 200 km by 6-month resolution grid. Each dot is a CFC-11 maximum in the DWBC. b, Magnitude of the DWBC CFC core in ORCA025. Black contour lines at $0.4 \text{ picomole kg}^{-1}$ intervals represent a smoothed version of the color shaded background at the same resolution as panel a. c, Magnitude of the DWBC CFC core in ORCA025 output that has been subsampled at the approximate time and place of the observations (black dots). The surrounding values are determined by interpolation as in panel a. d, Percent of particle first arrival positions from 48°N to 25°N in the DWBC, defined by being inshore of the 4000 m isobath. The dashed line represents particles that are continuously inshore starting from 48°N . The solid line is an inventory of all particles that go into the DWBC. Two standard deviation envelopes, constructed from 1000 iterations of 500 random trajectory subsamples of the ensemble, are shaded around each line. The distance along the boundary starts at 0 km at 55°N on the western boundary of the Labrador Sea and geography is indicated by FC (Flemish Cap), TGB (Tail of the Grand Banks), CC (Cape Cod), A (Abaco), PR (Puerto Rico), and B (Barbados). Gray lines represent a slope of 1 cm s^{-1} .

Figure 3. Lagrangian age of simulated LSW. a, Average particle age at 1500 m. b, Vertical section of average age at 66°W . Blank regions experienced less than 100 particle visits over 50 years. Black lines on maps are section locations. Gray lines on sections are the average location of the $\sigma_\theta = 27.68$ and 27.80 kg m^{-3} isopycnals in ORCA025 from 1980 to 2004, as in Figure 1.

Figure 4. LSW layer thickness and potential vorticity. a, Average thickness, h , between the $\sigma_\theta = 27.68$ and 27.80 kg m^{-3} isopycnals from the hydrographic climatology. b, Average thickness between the same isopycnals as in panel a but in the ORCA025 model. Average layer potential vorticity, f/h , for each is shown with solid (dashed) black lines at 1.5×10^{-8} (1×10^{-7}) $\text{m}^{-1}\text{s}^{-1}$ intervals (zero at the equator).







