



Thermal regime of the Costa Rican convergent margin: 1. Along-strike variations in heat flow from probe measurements and estimated from bottom-simulating reflectors

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[1] The thermal structure of convergent margins provides information related to the tectonics, geodynamics, metamorphism, and fluid flow of active plate boundaries. We report 176 heat flow measurements made with a violin bow style probe across the Costa Rican margin at the Middle America Trench. The probe measurements are collocated with seismic reflection lines. These seismic reflection lines show widespread distribution of bottom-simulating reflectors (BSRs). To extend the spatial coverage of heat flow measurements we estimate heat flow from the depth of BSRs. Comparisons between probe measurements and BSR-derived estimates of heat flow are generally within 10% and improve with distance landward of the deformation front. Together, these determinations provide new information on the thermal regime of this margin. Consistent with previous studies, the margin associated with the northern Nicoya Peninsula is remarkably cool. We define better the southern boundary of the cool region. The northern extent

of the cool region remains poorly determined. A regional trend of decreasing heat flow landward of the deformation front is apparent, consistent with the downward advection of heat by the subducting Cocos Plate. High wave number variability at a scale of 5–10 km is significantly greater than the measurement uncertainty and is greater south of the northern Nicoya Peninsula. These heat flow anomalies vary between approximately 20 and 60 mW m⁻² and are most likely due to localized fluid flow through mounds and faults on the margin. Simple one-dimensional models show that these anomalies are consistent with flow rates of 7–15 mm yr⁻¹. Across the margin toe variability is significant and likely due to fluid flow through deformation structures associated with the frontal sedimentary prism.

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1. Introduction

[2] Determinations of heat flow yield valuable information for understanding the thermal regime of shallow subduction zones and provide context for interpreting many hydrogeologic and geodynamic processes. Hydrogeologic processes of interest include sediment compaction, pore fluid expulsion, and other processes leading to fluid migration, as well as dehydration reactions within the subducting sediment and crust. Geodynamic processes of interest include modes of deformation along the plate interface, seismicity, and volcanism. These processes are not mutually exclusive since fluid production and migration can modify stresses and influence deformation.

[3] The Costa Rica subduction zone is one of the best studied erosive margins. It has been the focus of many studies including those of the Deep Sea Drilling Project, Ocean Drilling Program, and Integrated Drilling Program [von Huene *et al.*, 1985; Kimura *et al.*, 1997; Morris *et al.*, 2003], the location of a German Science Foundation funded Priority Program project, and a focus site of the U.S. Margins program. Offshore data include 2-D and 3-D seismic data [Shibley *et al.*, 1992; Hinz *et al.*, 1996], wide-angle seismic refraction [Ye *et al.*, 1996; Christeson *et al.*, 1999; Walther, 2003], swath bathymetry [von Huene *et al.*, 2000; Ranero *et al.*, 2008], magnetic

surveys [Barckhausen *et al.*, 2001] and seafloor observations [e.g., McAdoo *et al.*, 1996; Sahling *et al.*, 2008].

[4] These studies show that the margin is characterized by long-term large-scale slope subsidence and associated tectonic extension characteristic of erosive margins [Ranero and von Huene, 2000; Ranero *et al.*, 2000; Vannucchi *et al.*, 2003]. The morphology of the margin (Figure 1) is segmented along strike with a more stable portion where relatively smooth seafloor is subducted offshore the northern Nicoya Peninsula and a more dynamic portion where rough seafloor is subducting to the south [e.g., von Huene *et al.*, 2000]. Both the erosive nature and segmentation of the margin affects its thermal and hydrologic state. Characterizing the thermal regime of the margin provides insights into tectonic processes and the style, location, and magnitude of subsurface fluid flow.

[5] In this paper, the first of two, we report heat flow values based on probe measurements and the location of bottom-simulating reflectors (BSRs) along multiple transects spanning the Costa Rican margin. The quality and geographic extent of these thermal data provide an unprecedented opportunity to investigate along-margin variations in heat flow. The data are used to construct a heat flow map of the margin. The focus of this paper is on high-

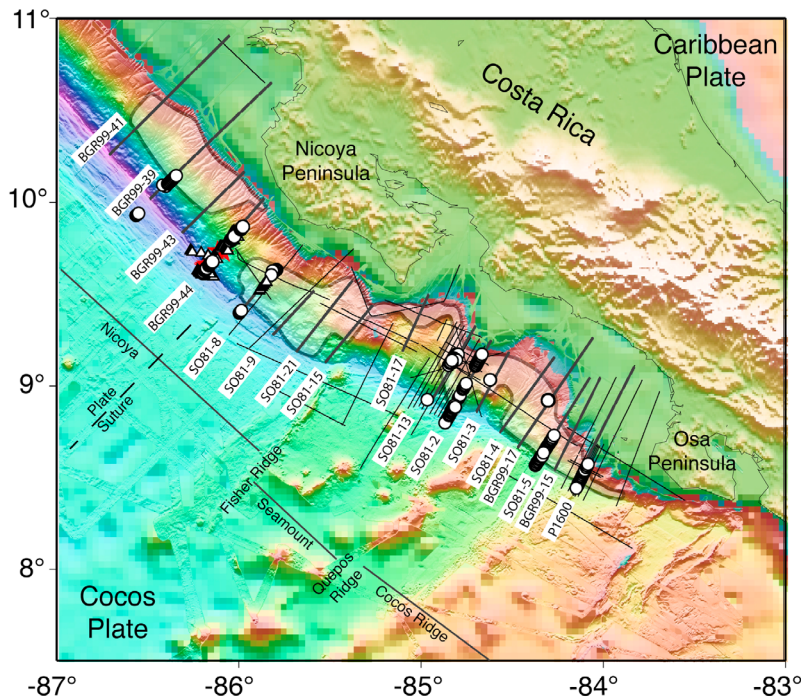


Figure 1. Map showing location of seismic lines (bold seismic lines are those analyzed in this study and by *Harris et al.* [2010]), heat flow data of *Langseth and Silver* [1996] (triangles), and new heat flow data (circles). DSDP Site 565 and ODP Sites 1039, 1040, 1041, and 1043 are shown as red stars along seismic line BGR99-44. Morphologic segments are labeled. Areas of observed bottom-simulating reflectors are outlined and shaded white. The approximate location of the plate suture is marked with a dashed line.

wave number heat flow variations that provide insight into surficial processes. *Harris et al.* [2010] generate finite element models to interpret low-wave number variations in heat flow and estimate temperatures along the subduction plate boundary.

2. Regional Setting

[6] Offshore Costa Rica the oceanic Cocos Plate subducts under the Caribbean plate forming the southern end of the Middle America Trench (Figure 1). Subduction rates vary from 70 mm/yr offshore Guatemala to 90 mm/yr offshore southern Costa Rica [*DeMets*, 2001]. The age of the Cocos Plate at the Middle America Trench decreases from 24 Ma offshore the Nicoya Peninsula to 15 Ma offshore the Osa Peninsula [*Barckhausen et al.*, 2001].

[7] Seismic data and drilling offshore the Nicoya peninsula show that the margin is composed of a thick slope sediment apron, a few hundred meters to ~2 km thick, unconformably overlying a margin of crystalline basement rock that is likely the offshore extension of igneous rocks cropping out on the Nicoya Peninsula. These low-permeability crystalline rocks are cut by numerous landward

dipping normal faults. A small frontal sedimentary prism, typically 10 km wide or less lies seaward of the margin and is compositionally similar to the sedimentary apron [*Kimura et al.*, 1997; *Morris et al.*, 2003]. The incoming oceanic plate is overlain by a maximum sediment thickness of 300–500 m [*Kimura et al.*, 1997; *Morris et al.*, 2003].

[8] The along-strike morphology of the margin is variable and largely reflects the relief of the incoming plate [*von Huene et al.*, 2000]. We briefly review this morphology because it provides a framework for understanding the distribution of BSRs and has been linked to variations in the hydrology of the margin [*Sahling et al.*, 2008; *Ranero et al.*, 2008]. Three morphologically distinct segments (Figure 1) have been recognized and from north to south are termed the Nicoya, Seamount, and Cocos Ridge Segments [*von Huene et al.*, 2004].

[9] The incoming oceanic crust within the Nicoya segment offshore southern Nicaragua and the Nicoya Peninsula has a notable tectonic history [*Hey*, 1977; *von Huene et al.*, 2000; *Barckhausen et al.*, 2001]. Oceanic crust forming the northern portion of this segment was created at the fast spreading East Pacific Rise (EPR) and oceanic

crust to the south was formed at the slow spreading Cocos-Nazca spreading center (CNS). The boundary separating EPR from CNS crust is a combination of a triple junction trace and a fracture zone, collectively comprising a “plate suture” (Figure 1). The rough-smooth boundary from *Hey* [1977] is further southeast, as explained by *Barckhausen et al.* [2001]. Recent swath mapping north of the plate suture, however, reveals the presence of more seamounts disrupting the “smooth” sedimentary cover than previously recognized [*Fisher et al.*, 2003]. The Nicoya segment margin is characterized by an upper slope with deep canyons, a relatively gentle middle slope with shallower canyons and recent sedimentation closer to the trench. Together these features have been interpreted to indicate relative stability of this portion of the margin compared to other segments [*von Huene et al.*, 2000].

[10] The seamount segment lies between Fisher Ridge and Quepos Ridge south of the Nicoya segment [*von Huene et al.*, 2000]. Here approximately 40% of the incoming seafloor is covered with seamounts having heights and diameters above the sedimentary cover of 1 to 2.5 km and 10 to 20 km, respectively. The seamount segment margin is characterized by embayments and mass wasting deposits. Seismic reflection profiles image seamounts within the margin and reveal that their subduction leads to embayments, slumping, and mass wasting deposits [*von Huene et al.*, 2000, 2004]. Rounded uplifts and domes at the surface often indicate the current position of subducted seamounts.

[11] The Cocos Ridge segment lies along the flank of the Cocos Ridge between the Quepos Plateau and the Osa Peninsula. The Cocos ridge is a bathymetric feature more than 1000 km long and 250 to 500 km wide with a bathymetric expression 2 km shallower than surrounding seafloor representing the trace of the Galapagos hot spot [*Hey*, 1977]. This region of the margin is narrow and steep and likely reflects the continuous subduction of the Cocos Ridge [*von Huene et al.*, 2000].

3. Seismic Data

[12] Seismic data mainly come from cruises *R/V Sonne* SO-81 [*Hinz et al.*, 1996; *von Huene et al.*, 2000; *Ranero and von Huene*, 2000; *Ranero et al.*, 2007], and BGR99 [*Ranero et al.*, 2008]. Seismic line P1600 is provided by Shell [*von Huene et al.*, 2000]. All seismic lines were collected with large tuned air gun arrays and multichannel streamers,

as described in the original papers. Seismic data were processed for signal enhancement including deconvolution and multiple attenuation, and were poststack time migrated. Selected sections are prestack depth migrated. All lines provide good imaging of the structure of the overriding plate including the sediment cover strata, BSRs and plate boundary reflections.

4. Distribution of BSRs

[13] BSRs were recognized in early seismic studies of the margin [*Shipley et al.*, 1979; *Yamano et al.*, 1982; *Pecher et al.*, 1998; *Müller et al.*, 2007]. *Pecher et al.* [1998] investigated the nature and distribution of BSRs in the study area. In general they found that BSRs are widespread in areas of margin stability, but are absent in areas characterized by slumping, except where the material failed as a coherent block. Because the BSRs are used to quantify the thermal regime across the margin, we briefly review representative seismic sections across each of the morphological segments and their implications for interpretations of heat flow.

[14] BSRs associated with the northern Nicoya segment (Figure 1) are relatively continuous along seismic lines BGR99-41, BGR99-39 (Figure 2a), and BGR99-43. With the exception of seismic line BGR99-44, BSRs first appear about 20 km landward of the trench and extend approximately 45 km landward of the trench. Along seismic line BGR99-44 BSRs only appear near the landward most extent of the line.

[15] Along the southern Nicoya segment, BSRs appear near the trench axis and extend approximately 30 to 50 km landward as imaged on seismic lines SO81-8, SO81-9, and SO81-21. Seismic Line SO81-21 (Figure 2b) shows a pronounced slump deposit with an overlying BSR indicating that either the mass moved as a coherent block or that the slump is old enough for the BSR to reform [*von Huene et al.*, 2004]. In contrast, just to the south along seismic line SO81-15 (Figure 2c) the BSR is truncated by a slump headwall coincident with a subducting seamount and no BSR is observed within the slump deposit [*von Huene et al.*, 2004].

[16] Large margin embayments characterize the Seamount segment where the subduction of seamounts causes large-scale slumping along the margin [*von Huene et al.*, 2000]. In this segment BSRs are not observed seaward of slumps but are present landward of slump headwalls and appear

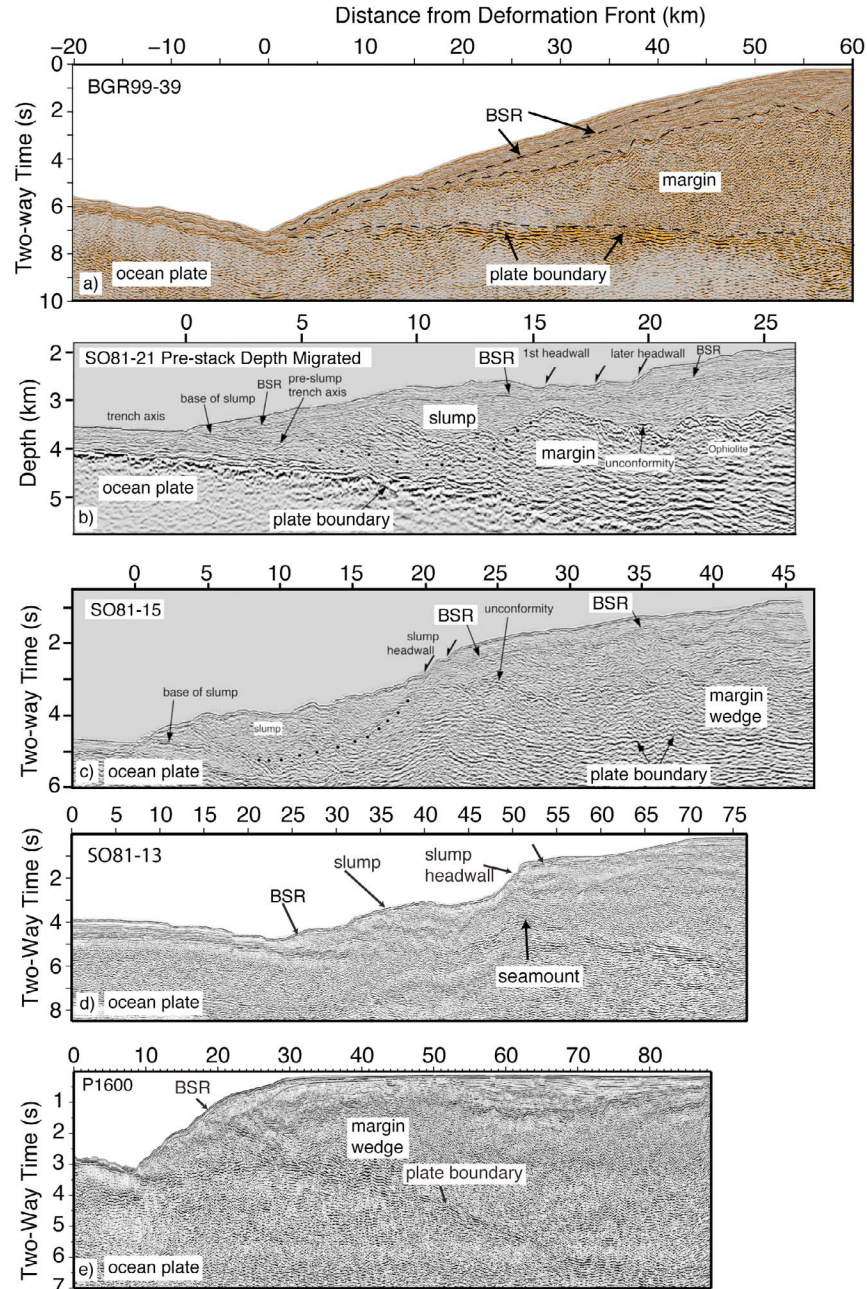


Figure 2. Poststack time migrated seismic reflection profiles across the margin, offshore Costa Rica. Seismic line locations are shown in Figure 1. Line SO81-21 is prestack depth migrated. The positions of BSRs are indicated by arrows.

truncated by them. Examples of these features can be observed on seismic lines, SO81-15 (Figure 2c) SO81-13 (Figure 2d). A rapid shallowing of BSRs toward the seafloor marks their landward extent.

[17] The Cocos Ridge segment is characterized by a steep and narrow margin slope as observed in seismic line P-1600 (Figure 2e). BSRs begin within 10 km of the deformation front and shallow toward

the seafloor. Their landward extent decreases to the south (Figure 1).

5. Heat Flow Transects Across the Margin

[18] Heat flow determinations were made at 176 points (Table 1) and all come from METEOR



Table 1. Heat Flow Data

Profile	Penetration	Longitude (deg)	Latitude (deg)	Water Depth (m)	Distance ^a (km)	Thermal Gradient (°C km ⁻¹)	Corrected Thermal Gradient ^b (°C km ⁻¹)	Thermal Conductivity (W m ⁻¹ K ⁻¹)	Heat Flow ^c (mW m ⁻²)	Instrument ^d
BGR-39	H0209P01	-86.3163	10.3072	1684	35.7	45.7	45.1	0.76 ^e	34.3	HP1
BGR-39	H0209P02	-86.3140	10.3042	1676	35.7	48.0	46.2	0.76 ^f	35.1	HP1
BGR-39	H0209P03	-86.3110	10.3013	1668	35.7	51.5	48.0	0.76 ^e	36.5	HP1
BGR-39	H0209P04	-86.3092	10.2995	1586	35.7	54.5	71.8	0.80 ^f	57.4	HP1
BGR-39	H0209P05	-86.3080	10.2985	1535	35.7	43.8	72.7	0.82 ^e	59.6	HP1
BGR-39	H0209P08	-86.3037	10.2965	1559	35.9	39.9	45.8	0.85 ^f	38.9	HP1
BGR-39	H0209P09	-86.3022	10.2963	1559	36.0	60.4	69.3	0.82 ^e	56.8	HP1
BGR-39	H0209P10	-86.3005	10.2933	1614	35.9	51.9	50.8	0.79 ^f	40.1	HP1
BGR-39	H0209P11	-86.2973	10.2902	1601	35.9	52.7	51.4	0.79 ^e	40.6	HP1
BGR-39	H0209P12	-86.2973	10.2875	1610	35.7	47.2	47.2	0.78 ^f	36.8	HP1
BGR-39	H0209P13	-86.2947	10.2937	1586	36.4	48.1	46.0	0.75 ^e	34.5	HP1
BGR-39	H0209P14	-86.3153	10.3133	1660	36.3	52.9	52.0	0.80 ^f	41.6	HP1
BGR-39	H0210A01	-86.3080	10.2985	1535	35.7	47.3	78.5	1.00 ^e	78.5	MTL
BGR-39	H0211A01	-86.3048	10.2998	1516	36.1	46.4	72.2	1.00 ^e	72.2	MTL
BGR-39	H0212A01	-86.3042	10.2968	1535	35.9	31.0	50.4	1.00 ^e	50.4	MTL
BGR-39	H0213A01	-86.3053	10.2945	1609	35.6	27.6	25.9	1.00 ^e	25.9	MTL
BGR-39	H0214P01	-86.4032	10.0873	3485	11.9	40.6	41.6	0.81 ^f	33.7	HP1
BGR-39	H0214P02	-86.4180	10.0942	3545	11.3	51.0	52.1	0.78 ^e	40.6	HP1
BGR-39	H0214P03	-86.3903	10.1005	3280	14.0	40.1	43.4	0.79 ^f	34.3	HP1
BGR-39	H0214P04	-86.3838	10.1070	3258	15.0	45.3	43.3	0.78 ^e	33.8	HP1
BGR-39	H0214P05	-86.3777	10.1133	3196	15.9	45.4	43.1	0.81 ^f	34.9	HP1
BGR-39	H0214P06	-86.3710	10.1197	3083	17.0	46.3	43.8	0.78 ^e	34.2	HP1
BGR-39	H0214P07	-86.3645	10.1262	2937	18.0	48.1	51.2	0.78 ^f	39.9	HP1
BGR-39	H0214P08	-86.3582	10.1325	2864	19.0	55.0	47.8	0.78 ^e	37.3	HP1
BGR-39	H0214P09	-86.3517	10.1390	2694	20.0	46.9	49.9	0.78 ^e	38.9	HP1
BGR-39	H0214P10	-86.3452	10.1455	2621	21.0	44.2	46.0	0.78 ^f	35.9	HP1
BGR-39	H0215A01	-86.3062	10.2977	1549	-	-	-	1.00 ^e	-	MTL
BGR-39	H0216P01	-86.3067	10.2982	1535	35.8	38.6	64.0	0.90 ^f	57.6	HP1
BGR-39	H0216P02	-86.3062	10.2978	1549	35.8	39.0	59.7	0.90 ^e	53.7	HP1
BGR-39	H0216P03	-86.3062	10.2975	1549	-	-	-	0.90 ^e	-	HP1
BGR-39	H0218P01	-86.5647	9.9288	4180	-13.0	172.7	169.5	0.74 ^f	125.4	HP1
BGR-39	H0218P02	-86.5580	9.9352	4178	-12.0	169.6	165.3	0.75 ^e	124.0	HP1
BGR-39	H0218P03	-86.5517	9.9412	4137	-11.1	118.5	129.3	0.75 ^e	97.0	HP1
BGR-44	H0219P01	-86.0567	9.8030	2512	24.7	28.0	30.1	0.75 ^e	22.6	HP1
BGR-44	H0219P02	-86.0567	9.8030	2512	24.7	28.0	30.1	0.75 ^e	22.6	HP1
BGR-44	H0219P02	-86.0507	9.8097	2463	25.7	32.5	33.1	0.75 ^e	24.8	HP1
BGR-44	H0219P03	-86.0447	9.8163	2407	26.7	32.0	31.8	0.51 ^f	16.2	HP1
BGR-44	H0219P04	-86.0388	9.8232	2371	27.6	34.9	32.5	0.75 ^e	24.4	HP1
BGR-44	H0220P01	-86.0027	9.8640	1922	33.7	36.3	32.5	0.65 ^f	21.1	HP1
BGR-44	H0220P02	-85.9968	9.8707	1806	34.7	24.3	22.3	0.75 ^e	16.7	HP1
BGR-44	H0220P03	-85.9908	9.8777	1712	35.7	26.0	22.9	0.84 ^f	19.2	HP1
BGR-44	H0221A01	-85.6802	9.8302	104	-	-	-	1.00 ^e	-	MTL
BGR-44	H0221A02	-85.6802	9.8303	104	-	-	-	1.00 ^e	-	MTL
BGR-44	H0222P01	-86.2122	9.6265	4154	-1.3	12.7	13.2	0.78 ^e	10.3	HP1
BGR-44	H0222P02	-86.2063	9.6330	4233	-0.4	13.1	14.7	0.78 ^e	11.5	HP1
BGR-44	H0222P03	-86.2005	9.6400	4345	0.7	14.0	13.4	0.82 ^f	11.0	HP1
BGR-44	H0222P04	-86.1942	9.6467	4364	1.7	18.2	17.8	0.78 ^e	13.9	HP1
BGR-44	H0222P05	-86.1885	9.6528	4379	2.6	19.6	15.6	0.77 ^f	12.0	HP1
BGR-44	H0222P07	-86.1763	9.6673	4097	4.7	45.0	49.1	0.78 ^e	38.3	HP1
BGR-44	H0222P08	-86.1703	9.6740	3981	5.7	11.9	12.5	1.44 ^f	18.0	HP1
BGR-44	H0222P09	-86.1647	9.6808	3888	6.7	19.0	19.1	0.78 ^e	14.9	HP1
BGR-44	H0222P10	-86.1585	9.6873	3785	7.6	26.4	28.0	0.87 ^f	24.4	HP1
SO81-8	H0223P01	-85.9978	9.3980	3797	-3.9	184.0	180.8	0.77 ^f	139.2	HP1
SO81-8	H0223P02	-85.9927	9.4043	3824	-3.0	167.4	167.5	0.77 ^e	129.0	HP1
SO81-8	H0223P03	-85.9865	9.4117	3859	-1.9	177.5	176.4	0.77 ^e	135.8	HP1
SO81-8	H0243P01	-85.7928	9.6340	1660	30.7	43.0	42.9	0.77 ^f	33.0	HP1
SO81-8	H0243P02	-85.7987	9.6323	1687	30.1	43.3	44.5	0.78 ^e	34.7	HP1
SO81-8	H0243P03	-85.8045	9.6253	1778	29.1	40.6	39.5	0.79 ^f	31.2	HP1
SO81-8	H0243P04	-85.8107	9.6185	1837	28.1	37.8	36.8	0.80 ^e	29.4	HP1



Table 1. (continued)

Profile	Penetration	Longitude (deg)	Latitude (deg)	Water Depth (m)	Distance ^a (km)	Thermal Gradient (°C km ⁻¹)	Corrected Thermal Gradient ^b (°C km ⁻¹)	Thermal Conductivity (W m ⁻¹ K ⁻¹)	Heat Flow ^c (mW m ⁻²)	Instrument ^d
SO81-8	H0243P05	-85.8165	9.6280	1819	28.5	41.1	42.0	0.80 ^f	33.6	HP1
SO81-8	H0243P06	-85.8222	9.6047	1928	26.1	28.8	31.1	0.80 ^e	24.9	HP1
SO81-13	H0226A01	-84.9668	8.9268	3557	-0.6	216.0	196.8	1.00 ^e	196.8	MTL
SO81-13	H0228P01	-84.8107	9.1733	785	31.0	43.3	44.3	0.89 ^f	39.4	HP1
SO81-13	H0228P02	-84.8078	9.1735	773	31.2	51.6	52.5	0.89 ^e	46.7	HP1
SO81-13	H0228P03	-84.8050	9.1737	761	31.5	65.3	68.5	0.89 ^f	61.0	HP1
SO81-13	H0228P04	-84.8043	9.1735	761	31.5	48.3	50.7	0.89 ^e	45.1	HP1
SO81-13	H0228P05	-84.8035	9.1738	753	31.6	20.7	21.7	0.90 ^f	19.5	HP1
SO81-13	H0228P06	-84.8020	9.1738	743	31.7	52.5	57.1	0.89 ^e	50.8	HP1
SO81-13	H0228P07	-84.7987	9.1742	740	32.0	46.7	50.8	0.90 ^f	45.7	HP1
SO81-13	H0229P01	-84.8242	9.1535	840	28.4	49.7	53.7	0.90 ^f	48.3	HP1
SO81-13	H0229P02	-84.8227	9.1525	827	28.4	49.7	55.6	0.89 ^e	49.5	HP1
SO81-13	H0229P04	-84.8197	9.1505	848	28.5	54.0	53.5	0.86 ^f	46.0	HP1
SO81-13	H0229P05	-84.8183	9.1497	888	28.5	45.6	38.9	0.89 ^e	34.6	HP1
SO81-13	H0229P06	-84.8170	9.1485	860	28.5	49.5	49.7	0.91 ^f	45.2	HP1
SO81-13	H0229P07	-84.8153	9.1475	890	28.6	40.6	36.3	0.89 ^e	32.3	HP1
SO81-13	H0229P08	-84.8123	9.1458	889	28.7	50.1	49.8	0.89 ^e	44.3	HP1
SO81-13	H0229P09	-84.8097	9.1438	872	28.7	34.1	34.1	0.91 ^f	31.0	HP1
SO81-13	H0229P10	-84.8067	9.1422	863	28.8	34.7	34.9	0.89 ^e	31.1	HP1
SO81-13	H0241P01	-84.8480	9.1140	2005	23.4	223.9	172.0	0.89 ^e	153.1	HP1
SO81-13	H0241P02	-84.8417	9.1237	1712	24.7	201.2	156.7	0.89 ^e	139.5	HP1
SO81-13	H0241P03	-84.8313	9.1272	1334	25.7	54.1	68.7	0.90 ^f	61.8	HP1
SO81-13	H0241P04	-84.8368	9.1315	1390	25.7	174.9	143.0	0.89 ^e	127.3	HP1
SO81-13	H0241P05	-84.8313	9.1403	948	26.8	40.8	49.6	0.92 ^f	45.6	HP1
SO81-2	H0227A01	-84.8535	8.8283	3454	-5.7	530.9	416.3	1.00 ^e	416.3	MTL
SO81-2	H0224P01	-84.8508	8.8305	3461	-5.3	516.7	444.7	0.81 ^f	360.2	HP1
SO81-2	H0224P02	-84.8557	8.8235	3296	-6.2	339.9	398.8	0.80 ^e	319.0	HP1
SO81-2	H0224P03	-84.8600	8.8155	3272	-7.2	395.1	414.3	0.80 ^f	331.4	HP1
SO81-2	H0224P04	-84.8640	8.8077	3261	-8.2	393.1	409.3	0.80 ^e	327.4	HP1
SO81-2	H0224P05	-84.8685	8.7997	3263	-9.2	316.8	291.4	0.78 ^f	227.3	HP1
SO81-2	H0225P01	-84.7957	8.9342	2882	7.7	77.0	66.1	0.82 ^f	54.2	HP1
SO81-2	H0225P02	-84.7915	8.9422	2784	8.7	46.7	47.8	0.82 ^e	39.2	HP1
SO81-2	H0225P03	-84.7872	8.9500	2758	9.7	56.3	52.0	0.80 ^f	41.6	HP1
SO81-2	H0225P05	-84.7703	8.9823	2313	13.8	51.3	48.0	0.82 ^f	39.4	HP1
SO81-2	H0225P06	-84.7662	8.9902	2151	14.7	62.6	66.2	0.82 ^e	54.3	HP1
SO81-2	H0225P07	-84.7618	8.9982	2055	15.7	63.8	65.7	0.79 ^f	51.9	HP1
SO81-2	H0225P08	-84.7573	9.0062	1931	16.8	70.6	69.8	0.82 ^e	57.2	HP1
SO81-2	H0225P09	-84.7530	9.0142	1822	17.8	49.5	52.7	0.82 ^f	43.2	HP1
SO81-2	H0230P01	-84.7015	9.1088	1232	29.7	35.8	37.0	0.89 ^e	32.9	HP2
SO81-2	H0230P02	-84.6972	9.1167	1188	30.7	36.5	36.6	0.89 ^e	32.6	HP2
SO81-2	H0230P03	-84.6928	9.1245	1149	31.7	34.3	34.5	0.89 ^e	30.7	HP2
SO81-2	H0230P04	-84.6885	9.1325	1108	32.7	34.6	34.6	0.89 ^e	30.8	HP2
SO81-2	H0230P05	-84.6840	9.1405	1082	33.7	33.0	31.6	0.89 ^e	28.1	HP2
SO81-2	H0230P06	-84.6800	9.1483	1024	34.7	27.4	27.6	0.89 ^e	24.6	HP2
SO81-2	H0230P07	-84.6753	9.1570	984	35.8	29.2	28.4	0.89 ^e	25.3	HP2
SO81-2	H0230P08	-84.6712	9.1645	897	36.8	32.0	35.2	0.89 ^e	31.3	HP2
SO81-2	H0230P09	-84.6673	9.1727	857	37.8	23.8	21.2	0.89 ^e	18.9	HP2
SO81-2	H0231P01	-84.8470	8.8392	3426	-4.3	392.2	388.9	0.82 ^e	318.9	HP1
SO81-2	H0231P02	-84.8427	8.8472	3445	-3.2	403.0	421.0	0.83 ^f	349.4	HP1
SO81-2	H0231P03	-84.8382	8.8550	3496	-2.2	442.8	448.0	0.82 ^f	367.4	HP1
SO81-2	H0231P04	-84.8345	8.8630	3521	-1.3	331.5	353.0	0.82 ^e	289.5	HP1
SO81-2	H0231P05	-84.8303	8.8708	3584	-0.3	418.8	342.9	0.82 ^e	281.2	HP1
SO81-2	H0231P06	-84.8260	8.8790	3433	0.7	229.6	228.8	1.03 ^f	235.7	HP1
SO81-2	H0231P07	-84.8215	8.8870	3398	1.8	115.9	108.5	0.88 ^f	95.5	HP1
SO81-2	H0231P08	-84.8172	8.8950	3309	2.8	124.4	132.9	0.82 ^e	109.0	HP1
SO81-2	H0231P09	-84.8132	8.8862	3406	2.1	130.7	115.5	0.82 ^e	94.7	HP1
SO81-3	H0233A01	-84.6208	9.0335	1421	21.0	67.1	73.7	0.84 ^e	61.9	MTL
SO81-3	H0234A01	-84.6215	9.0340	1421	21.0	61.2	67.1	0.84 ^e	56.4	MTL
BGR17	H0235P01	-84.3072	8.9188	1074	33.5	61.4	58.3	0.90 ^f	52.5	HP1



Table 1. (continued)

Profile	Penetration	Longitude (deg)	Latitude (deg)	Water Depth (m)	Distance ^a (km)	Thermal Gradient (°C km ⁻¹)	Corrected Thermal Gradient ^b (°C km ⁻¹)	Thermal Conductivity (W m ⁻¹ K ⁻¹)	Heat Flow ^c (mW m ⁻²)	Instrument ^d
BGR17	H0235P02	-84.3075	8.9207	1064	33.6	69.1	65.2	0.88 ^e	57.4	HP1
BGR17	H0235P03	-84.3075	8.9223	1047	33.8	47.7	47.3	0.88 ^e	41.6	HP1
BGR17	H0235P04	-84.3075	8.9240	1047	33.9	55.1	54.7	0.88 ^e	48.1	HP1
BGR17	H0235P05	-84.3075	8.9257	1033	34.1	55.7	57.5	0.87 ^f	50.0	HP1
BGR17	H0235P06	-84.3057	8.9192	1064	33.6	63.2	60.0	0.88 ^e	52.8	HP1
BGR17	H0235P07	-84.3057	8.9207	1055	33.8	68.0	64.5	0.88 ^e	56.8	HP1
BGR17	H0235P08	-84.3057	8.9223	1035	33.9	71.1	75.5	0.88 ^e	66.4	HP1
BGR17	H0235P09	-84.3057	8.9240	1035	34.1	67.2	71.2	0.89 ^f	63.4	HP1
BGR17	H0235P10	-84.3057	8.9257	1033	34.2	61.0	60.3	0.88 ^e	53.1	HP1
BGR17	H0235P11	-84.3042	8.9190	1064	33.7	29.2	27.7	0.88 ^e	24.4	HP1
BGR17	H0235P12	-84.3042	8.9207	1055	33.8	73.6	69.9	0.88 ^e	61.5	HP1
BGR17	H0235P13	-84.3040	8.9223	1024	34.0	71.0	81.3	0.88 ^f	71.5	HP1
BGR17	H0235P14	-84.3040	8.9242	1027	34.2	49.0	51.5	0.88 ^e	45.3	HP1
BGR17	H0235P15	-84.3042	8.9258	1033	34.3	65.9	65.2	0.88 ^e	57.4	HP1
BGR17	H0235P16	-84.3025	8.9257	1027	34.4	63.0	66.3	0.88 ^e	58.3	HP1
BGR17	H0235P17	-84.3023	8.9242	1027	34.3	64.4	67.7	0.87 ^f	58.9	HP1
BGR17	H0235P18	-84.3023	8.9225	1024	34.1	70.2	80.3	0.88 ^e	70.7	HP1
BGR17	H0235P19	-84.3023	8.9208	1039	34.0	71.0	73.8	0.82 ^f	60.5	HP1
BGR17	H0235P20	-84.3023	8.9192	1047	33.8	71.9	73.9	0.88 ^e	65.0	HP1
BGR17	H0235P21	-84.3010	8.9192	1042	33.9	51.1	51.6	0.91 ^f	47.0	HP1
BGR17	H0235P22	-84.3008	8.9207	1041	34.1	56.5	55.3	0.88 ^e	48.7	HP1
BGR17	H0235P23	-84.3008	8.9223	1039	34.2	52.5	50.0	0.89 ^f	44.5	HP1
BGR17	H0235P24	-84.3007	8.9240	1039	34.4	50.6	48.2	0.88 ^e	42.4	HP1
BGR17	H0235P25	-84.3007	8.9257	1035	34.5	51.4	49.3	0.87 ^f	42.9	HP1
BGR17	H0240A01	-84.3043	8.9218	1055	33.9	8.8	8.3	0.90 ^e	7.5	MTL
BGR17	H0240A02	-84.3043	8.9218	1074	-	-	-	0.80 ^e	-	MTL
SO81-5	H0236P01	-84.3203	8.6513	2235	7.3	154.4	153.7	0.82 ^f	126.0	HP1
SO81-5	H0236P02	-84.3157	8.6588	2128	8.2	138.4	130.4	0.85 ^e	110.8	HP1
SO81-5	H0236P03	-84.3105	8.6663	1969	9.3	106.7	109.1	0.85 ^f	92.7	HP1
SO81-5	H0236P04	-84.3057	8.6740	1845	10.3	90.0	87.7	0.87 ^e	76.3	HP1
SO81-5	H0236P05	-84.3008	8.6818	1749	11.3	77.4	71.1	0.85 ^f	60.4	HP1
SO81-5	H0236P06	-84.2962	8.6895	1609	12.3	70.5	73.6	0.86 ^e	63.3	HP1
SO81-5	H0236P07	-84.2913	8.6972	1545	13.3	83.0	83.3	0.86 ^f	71.6	HP1
SO81-5	H0236P08	-84.2850	8.7047	1474	14.4	58.3	62.3	0.86 ^e	53.6	HP1
SO81-5	H0236P09	-84.2818	8.7123	1409	15.3	59.8	58.1	0.86 ^f	50.0	HP1
SO81-5	H0236P10	-84.2767	8.7203	1307	16.3	53.6	52.5	0.84 ^e	44.1	HP1
SO81-5	H0236P11	-84.2720	8.7313	1137	17.6	58.8	59.8	0.84 ^f	50.2	HP1
SO81-5	H0237P01	-84.3730	8.5673	2813	-3.7	393.8	377.5	0.79 ^f	298.2	HP1
SO81-5	H0237P02	-84.3685	8.5750	2827	-2.7	401.1	407.3	0.79 ^e	321.8	HP1
SO81-5	H0237P03	-84.3635	8.5827	2849	-1.7	282.8	283.7	0.78 ^f	221.3	HP1
SO81-5	H0237P04	-84.3588	8.5903	2859	-0.7	221.8	220.0	0.79 ^e	173.8	HP1
SO81-5	H0237P05	-84.3540	8.5980	2893	0.3	200.8	183.5	0.80 ^f	146.8	HP1
SO81-5	H0237P06	-84.3493	8.6055	2789	1.3	111.9	120.6	0.84 ^e	101.3	HP1
SO81-5	H0237P07	-84.3442	8.6132	2725	2.3	122.7	118.8	0.86 ^f	102.2	HP1
SO81-5	H0237P08	-84.3393	8.6210	2654	3.3	142.9	129.5	0.85 ^e	110.1	HP1
SO81-5	H0237P09	-84.3347	8.6187	2640	3.3	138.9	129.7	0.88 ^f	114.1	HP1
SO81-5	H0237P10	-84.3300	8.6362	2428	5.3	194.9	190.1	0.87 ^f	165.4	HP1
P1600	H0238P01	-84.1208	8.5005	1713	4.5	82.6	86.3	0.89 ^f	76.8	HP1
P1600	H0238P02	-84.1247	8.4925	1866	3.5	84.4	82.2	0.87 ^e	71.5	HP1
P1600	H0238P03	-84.1287	8.4845	2013	2.5	89.1	92.2	0.87 ^f	80.2	HP1
P1600	H0238P04	-84.1327	8.4762	2190	1.5	186.4	169.3	0.83 ^e	140.5	HP1
P1600	H0238P05	-84.1367	8.4680	2257	0.5	311.1	307.5	0.80 ^f	246.0	HP1
P1600	H0238P06	-84.1407	8.4598	2229	-0.5	249.6	263.3	0.80 ^e	210.6	HP1
P1600	H0238P07	-84.1448	8.4515	2291	-1.6	162.0	136.2	0.81 ^f	110.3	HP1
P1600	H0238P08	-84.1492	8.4433	2199	-2.6	1072.0	962.8	0.85 ^f	818.4	HP1
P1600	H0239P01	-84.1210	8.5005	-	-	-	-	-	-	HP2
P1600	H0239P02	-84.1170	8.5085	1600	5.5	95.6	101.3	0.89 ^e	90.2	HP2
P1600	H0239P03	-84.1130	8.5167	1499	6.5	82.4	84.5	0.89 ^e	75.2	HP2
P1600	H0239P04	-84.1090	8.5248	1421	7.5	116.0	108.8	0.89 ^e	96.8	HP2

Table 1. (continued)

Profile	Penetration	Longitude (deg)	Latitude (deg)	Water Depth (m)	Distance ^a (km)	Thermal Gradient (°C km ⁻¹)	Corrected Thermal Gradient ^b (°C km ⁻¹)	Thermal Conductivity (W m ⁻¹ K ⁻¹)	Heat Flow ^c (mW m ⁻²)	Instrument ^d
P1600	H0239P05	-84.1052	8.5327	1289	8.5	76.5	80.4	0.89 ^e	71.6	HP2
P1600	H0239P06	-84.1012	8.5408	1170	9.5	83.8	84.6	0.89 ^e	75.3	HP2
P1600	H0239P08	-84.0928	8.5573	944	11.5	69.9	70.0	0.89 ^e	62.3	HP2
P1600	H0239P09	-84.0890	8.5653	844	12.5	66.5	60.9	0.89 ^e	54.2	HP2
P1600	H0239P10	-84.0850	8.5733	713	13.5	73.8	73.7	0.89 ^e	65.6	HP2

^aDistance is distance landward from the deformation front.

^bThermal gradient corrected for the effects of bathymetry.

^cHeat flow computed as the product of the thermal gradient and thermal conductivity.

^dInstrument types HP1 and HP1 refer to X and X m violin probes, respectively. MTL refers to measurements made with temperature loggers attached to core barrels [Pfender and Villinger, 2002].

^eEstimated values of thermal conductivity based on nearby measurements.

^fMeasured values of thermal conductivity.

cruise M54/2 conducted in 2002. Heat flow measurements were collocated with seismic reflection profiles and swath bathymetry to aid interpretations (Figure 1). In general measurement spacing along transects is about 1 km, but decreases to approximately 50 m around targets of interest. Heat flow measurements were made with a violin bow instrument design [Hyndman *et al.*, 1979; Lister, 1979; Hartmann and Villinger, 2002] that allows measurements of both thermal gradient and thermal conductivity at multiple sites with a single transit to the seafloor. Two different violin bow instruments were used, a 3 m probe with 11 thermistors and a 6 m probe with 22 thermistors. The thermistors are equally spaced along the thermistor string. Additionally, a few heat determinations were made with autonomous temperature loggers attached to the side of gravity core barrels [Pfender and Villinger, 2002]. Heat flow determinations are processed based on the formulation described by Hartmann and Villinger [2002]. Each heat flow determination consists of inserting the probe into sediments under the force of gravity. Probe insertion generates a frictional heat pulse whose decay is well described by a cylindrical source. Equilibrium temperatures are computed by fitting the cylindrical decay curve to a 7 min period of data and then extrapolating to infinite time. In situ thermal conductivity was generally measured at every other site. These measurements are made following the initial 7 min period by generating a 20 s calibrated heat pulse [Lister, 1979] along a heater wire that extends over the length of the thermistor string. The temperature decay of the calibrated heat pulse provides a measure of sediment thermal conductivity at each thermistor. Data from each thermistor is monitored in real time through a coaxial cable

between the instrument probe and the ship. Measured temperature gradients are linear indicating conductive heat transport at the resolution of the probe.

[19] Rapid sedimentation can transiently depress thermal gradients. Sedimentation rates estimated at DSDP Site 565 [von Huene *et al.*, 1985] and ODP Site 1041 [Kimura *et al.*, 1997] are similar. At Site 1041 [Kimura *et al.*, 1997] sedimentation rates are ~96 m/Myr in the late Miocene, 62 m/Myr in the early Pliocene, and 55 m/Myr in the late Pliocene. A 1-D model of sedimentation and compaction [Wang and Davis, 1992] indicates that this depositional history depresses thermal gradients by less than a few percent; we do not correct for these small sedimentation effects in our heat flow determinations.

[20] Bathymetry has the potential to distort the thermal field by compressing isotherms under bathymetric lows and stretching them under bathymetric highs. Using a ~225 m bathymetric grid [Ranero *et al.*, 2008] and temperature-depth profiles through the water column from CTD casts we estimated a seafloor temperature distribution. This temperature distribution is diffused into the subsurface using Fourier series to calculate the thermal gradient due to the bathymetry following a procedure given by Blackwell *et al.* [1980]. With the exception of thermal gradients in the immediate vicinity of Culebra Mound [Grevenmeyer *et al.*, 2004], these three-dimensional perturbations generally produce only a small change in the thermal gradient (Table 1).

[21] In situ thermal conductivity values have a mean and standard deviation of 0.84 W m⁻¹ K⁻¹

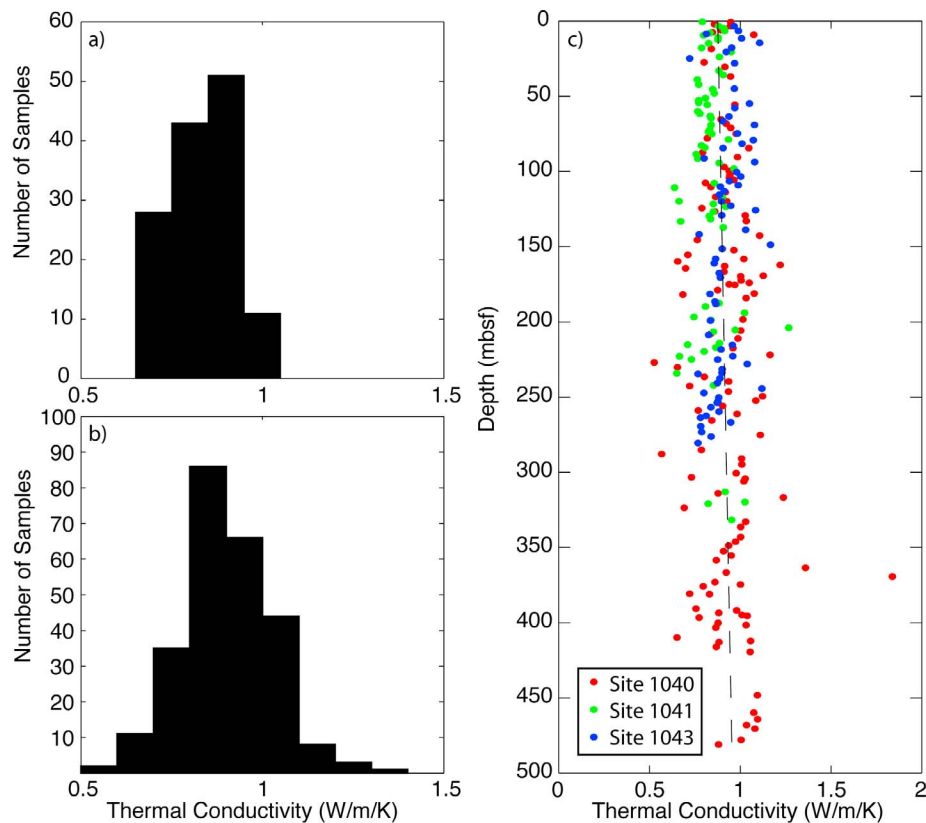


Figure 3. Thermal conductivity measurements. Histograms of thermal conductivity values from (a) shallow probes and (b) needle probe measurements ODP Leg 170 cores [Kimura *et al.*, 1997] and (c) depth distribution of needle probe measurements.

and $0.08 \text{ W m}^{-1} \text{ K}^{-1}$, respectively (Figure 3a). No significant spatial variations in thermal conductivity are observed along strike. In situ values are only slightly lower than thermal conductivity values determined on ODP cores at sites 1040, 1041, and 1043 on the margin [Kimura *et al.*, 1997]. These core values have a mean and standard deviation of 0.90 and $0.13 \text{ W m}^{-1} \text{ K}^{-1}$, respectively (Figure 3b). There is only a slight increase in thermal conductivity with depth (Figure 3c).

[22] Because observed thermal gradients are linear and thermal conductivity is approximately constant as a function of depth for each location, heat flow is computed as the product of the bathymetrically corrected thermal gradient and thermal conductivity (Table 1).

6. BSR-Derived Heat Flow

[23] The widespread presence of BSRs along the Costa Rica margin provides the opportunity to estimate geothermal gradients and when combined

with estimates of thermal conductivity can extend the coverage of heat flow data. Gas hydrates are stable within the top few hundred meters of ocean floor sediments adjacent to continental margins, and the phase boundary between hydrate and free gas is marked by a bottom-simulating reflector (BSR) [e.g., Tucholke *et al.*, 1977; Shipley *et al.*, 1979; Grevemeyer *et al.*, 2000; Pecher *et al.*, 2001]. Within the zone of hydrate stability, pore space is occupied by solid hydrate and water, while below the zone of hydrate stability, pore space is occupied by free gas and water. The transition between hydrate and water to free gas and water produces an acoustic impedance contrast that often generates a well-developed high-amplitude, reversed polarity reflection that is relatively straightforward to identify. Because the base of the hydrate stability field is primarily sensitive to temperature, with only moderate sensitivity to pressure and composition, BSRs correspond to an isotherm and provide a way to map regional estimates of average geothermal gradients on continental margins [e.g., Hyndman *et al.*, 1992; Brown and Bangs, 1995; Kaul *et al.*, 2000;

Grevenmeyer and Villinger, 2001; Minshull et al., 2005]. BSR depths are converted to estimates of heat flow based on the approach developed by *Yamano et al. [1982]* and reviewed in detail by *Grevenmeyer and Villinger [2001]*. Heat flow is calculated as,

$$q = k \frac{(T_{bsr} - T_{sea})}{(z_{bsr} - z_{sea})}, \quad (1)$$

where k is the thermal conductivity, T_{bsr} is the temperature at the BSR, T_{sea} is the temperature of the seafloor, z_{bsr} is the depth of the BSR and z_{sea} is the depth of the seafloor. This simple equation, however, is only valid for the case of a constant thermal conductivity between seafloor and BSR as indicated here (Figure 3).

[24] The geothermal gradient is based on the depth interval between the seafloor and the BSR and difference between the estimated bottom water temperature and the temperature at the BSR. Travel times between the seafloor and BSR were estimated by identifying the positive peak at seafloor and the negative peak at the BSR. The picking error is typically <20 ms, or about 5%–10%. A simple linear velocity function was used to convert this two way travel time to depth based on depth migrated seismic lines (SO81-4, SO81-5, SO81-17a, and BGR99-7),

$$v = 1550 + 0.309 z, \quad (2)$$

where v is the P wave velocity in m/s and z is the depth below seafloor in m. Although this simplified model neglects second-order velocity variations such as the potential velocity increase above the BSR in regions of strong hydrate concentration, the calculated heat flow is not very sensitive to seismic velocity. An increase in BSR depth would also increase the temperature at the BSR such that the overall change in the thermal gradient is small [*Grevenmeyer and Villinger, 2001*].

[25] Seafloor temperatures, T_{sea} , are estimated based on CTD casts from the NOAA World Ocean Data Atlas and the depth of the seafloor above each BSR pick. Seafloor temperature is assumed to be a function of depth only. Over the region of gas hydrates seafloor temperatures vary between 2°C and 15°C as a function of depth. Temporal variations in bottom water temperature are not observed. Heat probe determinations at water depths of ~1000 m show little evidence of nonlinear gradients that might reflect temporal variations in bottom water temperatures. ODP sites 1253 and 1255 with water depths of 4312 and 4376,

respectively, show annual changes less than 0.1°C [*Davis and Villinger, 2006*].

[26] The temperature at the BSR dept is calculated assuming pressure-temperature stability conditions for methane hydrate and that the primary composition is methane in seawater. We use the *Dickens and Quinby-Hunt [1994]* relationship,

$$\frac{1}{T} = 3.79 \cdot 10^{-3} - 2.83 \cdot 10^{-4} \log(P) \quad (3)$$

where T is the temperature in K and P is the pressure in MPa. Pressure at the BSR is computed assuming hydrostatic conditions and a seawater density of 1030 kg/m³.

[27] *Grevenmeyer and Villinger [2001]* show that sediment thermal conductivity is not sensitive to the presence of gas hydrates. Based on numerical models for ODP Site 1040, *Hensen and Wallmann [2005]* estimate the concentration of hydrate at 1.65 volume percent of pore space. *Grevenmeyer and Villinger [2001]* reviewed thermal conductivity, porosity and logging data from ODP sites drilled into and through gas hydrates. They find that although porosity decreases with depth, increasing the seismic velocity, thermal conductivity profiles between the seafloor and BSR can be fit using a mean value about as well as one that increases slightly with depth. Based on the lack of evidence for spatial variation in thermal conductivity along the margin and the relatively constant values of thermal conductivity determined from ODP Leg 170, we assumed a constant thermal conductivity of 0.9 W m⁻¹ K⁻¹ (Figure 3). We estimate the total error (root mean square of all errors) may reach 10%–20% and a relative error of approximately 5%–10%.

7. Comparisons Between in Situ Heat Flow and BSR-Derived Estimates of Heat Flow

[28] Heat flow observed with shallow probes and derived from BSRs are sensitive to different length scales. The 3 and 6 m probes used in this study have greater sensitivity to shallow thermal disturbances such erosion and/or sedimentation than BSR-derived values of heat flow. In contrast, BSR-derived values of heat flow are more sensitive to deeper thermal processes within the margin and their use in calculating heat flow assumes that the BSR is in thermal equilibrium. Advective processes generating curvature in the thermal gradient, such

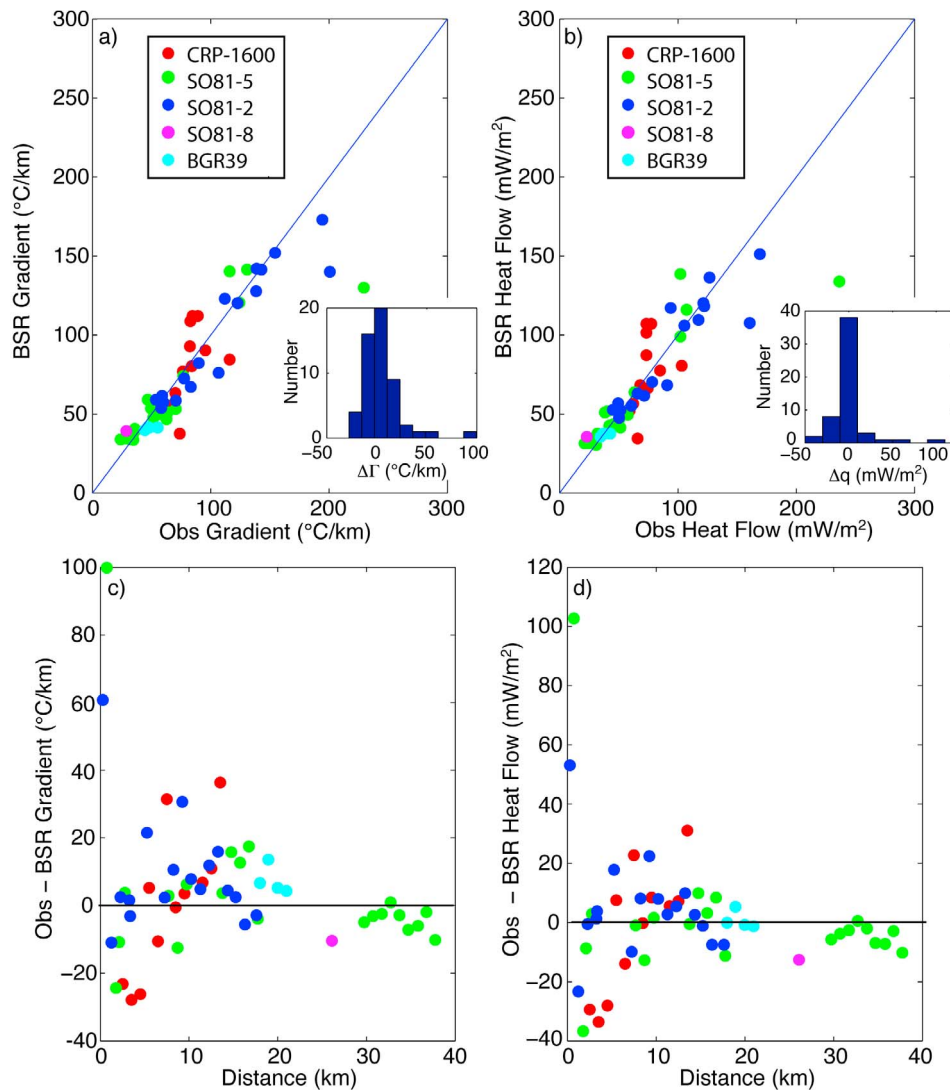


Figure 4. Correlations between probe and BSR-derived values of (a) thermal gradient and (b) heat flow. Correlations are color coded by seismic line. Insets show histograms of differences in thermal gradient and heat flow. (c and d) Thermal gradient and heat flow differences as a function of distance from the deformation front. The differences generally decrease with landward distance from the deformation front.

as block sliding or rotation, would lead to incorrect estimates of BSR-derived heat flow.

[29] Figure 4 shows comparisons between probe and BSR-derived values of thermal gradient and heat flow. These correlations are based on probe measurements within 2.5 km of the closest BSR determination. Correlations between the thermal gradient and heat flow both have r^2 values of approximately 0.80 (Figures 4a and 4b). A histogram of differences between probe and BSR heat flow values (Figure 4b, inset) show that differences are slightly skewed toward higher probe values of heat flow. This comparison contrasts with other

studies [e.g., *Ruppel, 1997; Kaul et al., 2000*] where BSR-derived values of heat flow are larger than probe values. In general values of heat flow agree to about 10%.

[30] Differences between probe and BSR values of heat flow are correlated with proximity to the deformation front (Figures 4c and 4d). Within about 5 km of the deformation front probe values of heat flow are slightly less than BSR values and between 5 and 15 km probe values of heat flow are generally somewhat higher than BSR-derived values. When points within 5 km of the deformation front are excluded, the standard deviation of

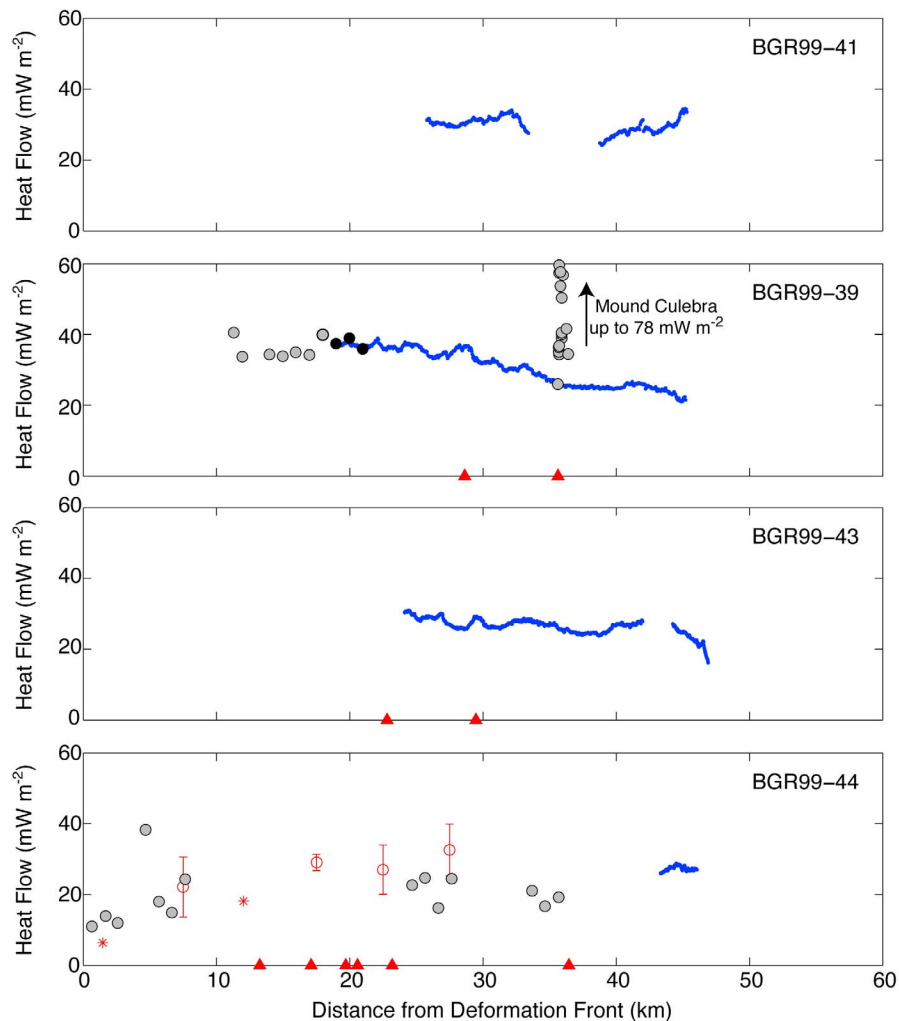


Figure 5. Heat flow profiles offshore the northern Nicoya Peninsula showing observed (black circles) and BSR-derived (small blue circles) values of heat flow. Solid black circles denote heat flow values within 2.5 km of a BSR value. Heat flow values from *Langseth and Silver* [1996] averaged over 5 km windows are shown as circles with error bars representing the standard deviation. ODP values of heat flow are shown as stars [Kimura *et al.*, 1997]. The positions of seeps along each profile are denoted as solid triangles at the base of each plot [Sahling *et al.*, 2008].

the differences decreases from 28 to 10 mW/m². The major source of these differences is likely thermal disequilibrium caused by deformation.

8. Discussion

[31] The generally good correspondence between the probe and BSR-derived values of heat flow suggests that BSR-derived values of heat flow are well determined and can be used with confidence to extend thermal mapping of the margin. Most comparisons between BSR- and probe-derived values have been made at accretionary prisms such as Cascadia [Davis *et al.*, 1990], Makran [Kaul *et al.*, 2000], and Nankai [Yamano *et al.*, 1982] and are

generally consistent to about 15%–20% [Hyndman *et al.*, 1992; Grevemeyer and Villinger, 2001].

[32] The good agreement between probe and BSR-derived values of heat flow suggest that this margin may be thermally more stable than these accretionary margins. The thermal evolution of the erosional Central America margin differs from that of Nankai, Makran, and Cascadia. Here, the crystalline rocks now beneath the middle slope of the margin subsided 3–5 km in 1–2 Ma because of erosional thinning [Vannucchi *et al.*, 2003]. In this manner the upper margin is transported into the lower margin temperature field. In contrast, at accretionary prisms, ocean basin sediment that has been drained of its pore water in the frontal prism

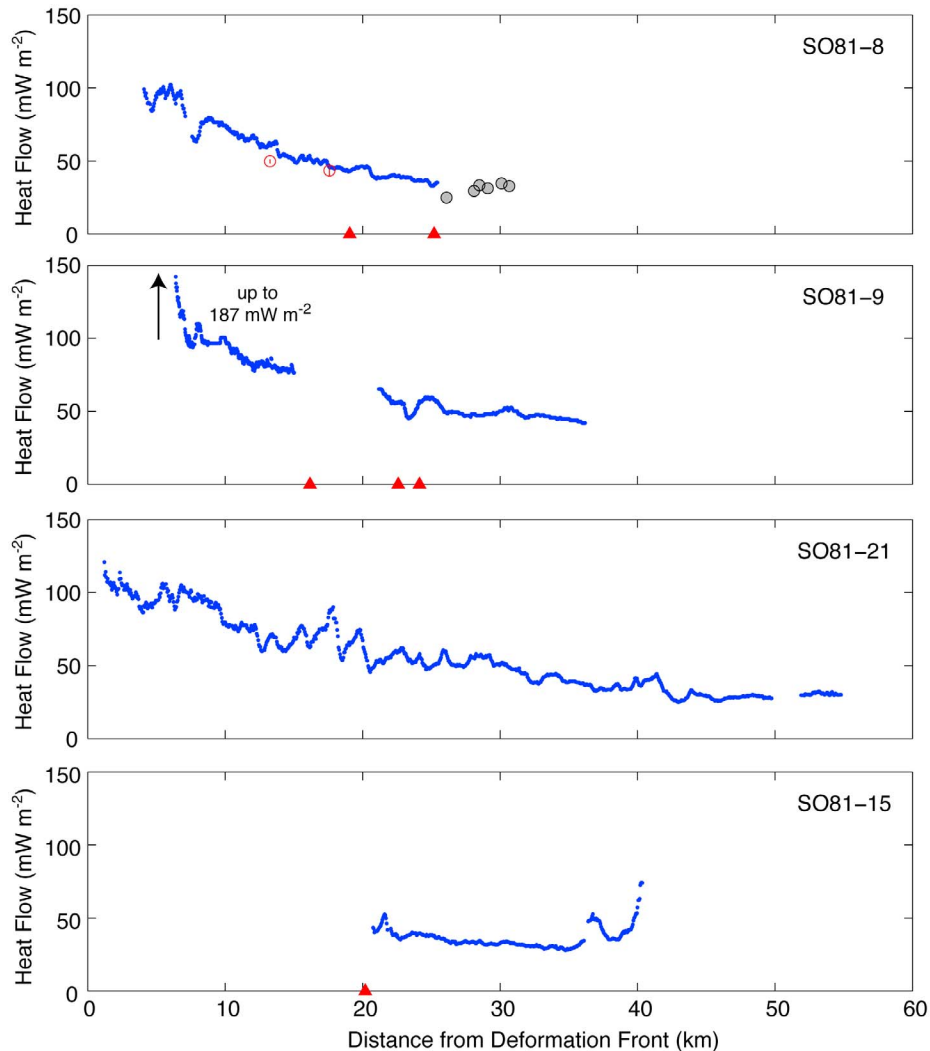


Figure 6. Heat flow profiles offshore the southern Nicoya Peninsula showing observed (solid circles) and BSR-derived (small circles) values of heat flow. Heat flow values from *Langseth and Silver* [1996] averaged over 5 km windows are shown as circles with bars showing standard deviation. The positions of seeps along each profile are denoted as solid triangles at the base of each plot [Sahling *et al.*, 2008].

becomes the middle and upper slope and is transported into the upper slope temperature field.

[33] Figures 5–8 show transects of both the in situ and BSR-derived values of heat flow along seismic lines. In general heat flow decreases with distance from the deformation front consistent with the subducting plate cooling the margin as it advects heat toward the Earth's interior.

[34] Heat flow transects along the northern portion of the Nicoya segment (Figure 5) are relatively low with maximum values less than $\sim 40 \text{ mW m}^{-2}$. These values of heat flow are remarkably low and well below the $95\text{--}105 \text{ mW m}^{-2}$ expected for 24 Ma crust. A notable exception, however, is the locally high heat flow associated with the seepage

site of Culebra Mound [Grevemeyer *et al.*, 2004] that lies a few kilometers north of seismic line BGR99-39. The absence of BSRs within approximately 20 km of the deformation front may indicate low heat flow that places the BSR depth within the oceanic crust, as suggested by deep temperature measurements made on ODP Leg 170 [Ruppel and Kinoshita, 2000], or within the ophiolitic rocks making up the margin. For example, the seaward extent of the BSR along seismic line BGR99-39 appears to be truncated where it intersects the basement rocks of the margin (Figure 2a). The magnitude of heat flow values and their pattern are consistent with exceptionally cold oceanic crust cooling the margin [Fisher *et al.*, 2003; Hutnak *et al.*, 2007].

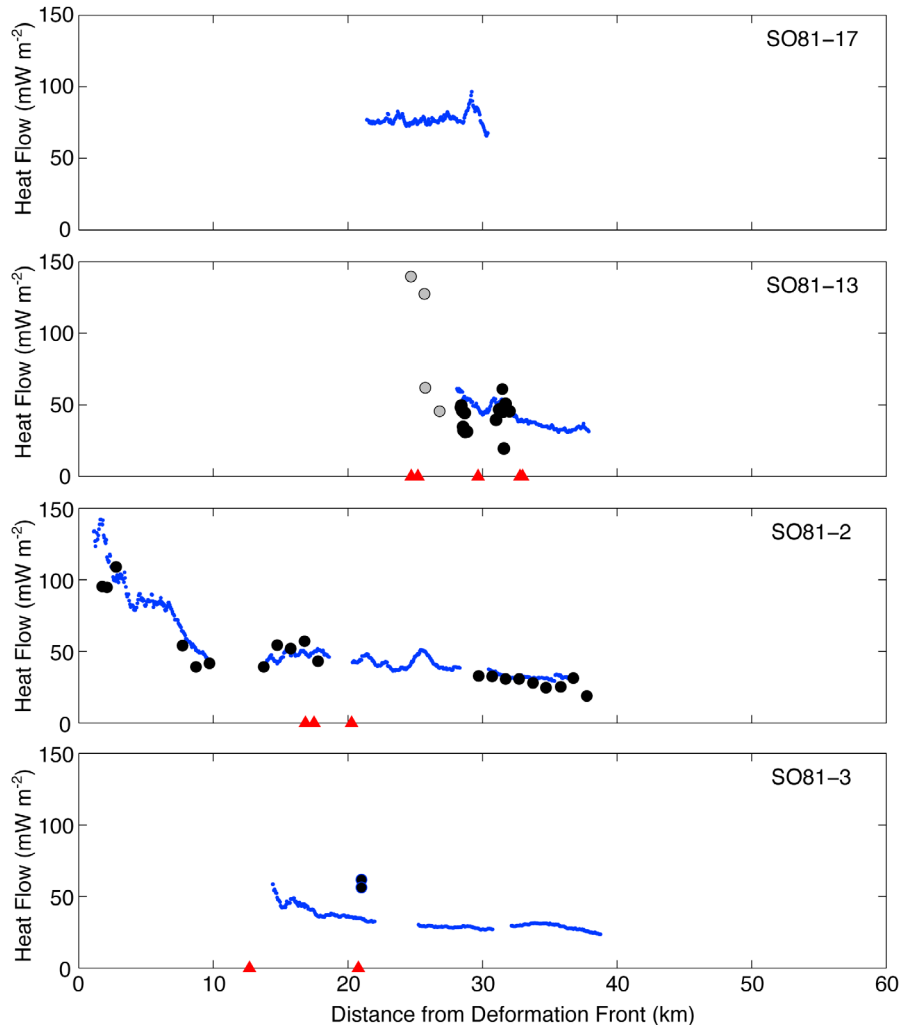


Figure 7. Heat flow profiles within the seamount segment showing observed (black circles) and BSR-derived (blue circles) values of heat flow. Solid black circles denote heat flow values within 2.5 km of a BSR value. The positions of seeps along each profile are denoted as solid triangles at the base of each plot [Sahling *et al.*, 2008].

[35] Heat flow transects along the southern Nicoya Peninsula (Figure 6) are predominantly based on BSRs. Comparisons of BSR-derived estimates of heat flow and in situ values along seismic line SO81-8 are consistent with each other. Near the deformation front heat flow values are approximately 100 mW m^{-2} and decrease landward. BSRs along seismic line SO81-15 are truncated by a slump headwall (Figure 2c).

[36] Heat flow transects within the Seamount segment show large variability (Figure 7). Heat flow along profiles SO81-17 and SO81-13 are only available for a short distance because of the disrupted margin. In situ heat flow values along profile SO81-13 are highly variable over a short distance likely due to the advective influence of seeps that was a target of this heat flow station.

Profile SO81-2 has the largest data span across the margin. Near the deformation front heat flow values are in excess of 100 mW m^{-2} and then decrease sharply to about 50 mW m^{-2} 10 km landward of the deformation front. In situ values of heat flow agree well with BSR estimates of heat flow.

[37] Heat flow profiles corresponding to the Cocos Ridge segment of the margin (Figure 8) display some of the highest observed values of heat flow along the margin. Within 5 km of the deformation front, heat flow is quite high and locally exceeds values of 150 mW m^{-2} . Along profile BGR99-17 heat flow values targeted Mound 11 [Schmidt *et al.*, 2005; Sahling *et al.*, 2008] and are quite variable. A pronounced heat flow high centered about 7 km landward of the deformation front is observed

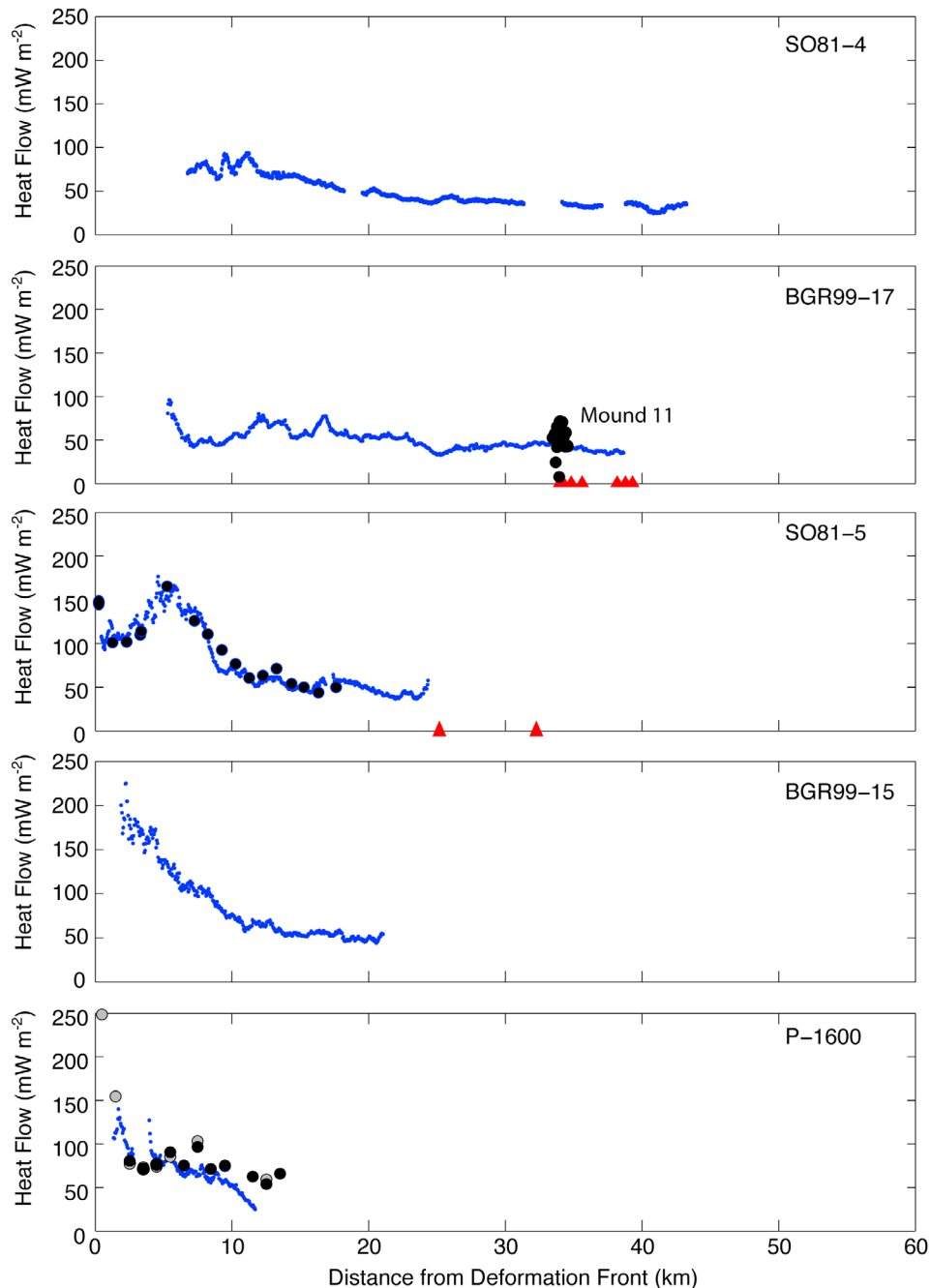


Figure 8. Heat flow profiles within the Cocos Ridge segment showing observed (black circles) and BSR-derived (blue circles) values of heat flow. Solid black circles denote heat flow values within 2.5 km of a BSR value. The positions of seeps along each profile are denoted as solid triangles at the base of each plot [Sahling *et al.*, 2008].

along profile SO81-5. This feature correlates well with a seismically imaged horst in the downgoing oceanic plate and may reflect hydrothermal circulation within this feature.

[38] Hand drawn contours of heat flow data (Figure 9) show two thermally distinct regions. The margin offshore northern Nicoya Peninsula is rel-

atively cool, especially in the area of seismic line BGR99-44 where heat flow increases landward of the deformation front before decreasing [Langseth and Silver, 1996]. This area is likely a continuation of the cool section of the Cocos plate mapped by Fisher *et al.* [2003]. To the south of the plate suture (Figure 1), heat flow increases sharply. This new data extends the transition between the warmer

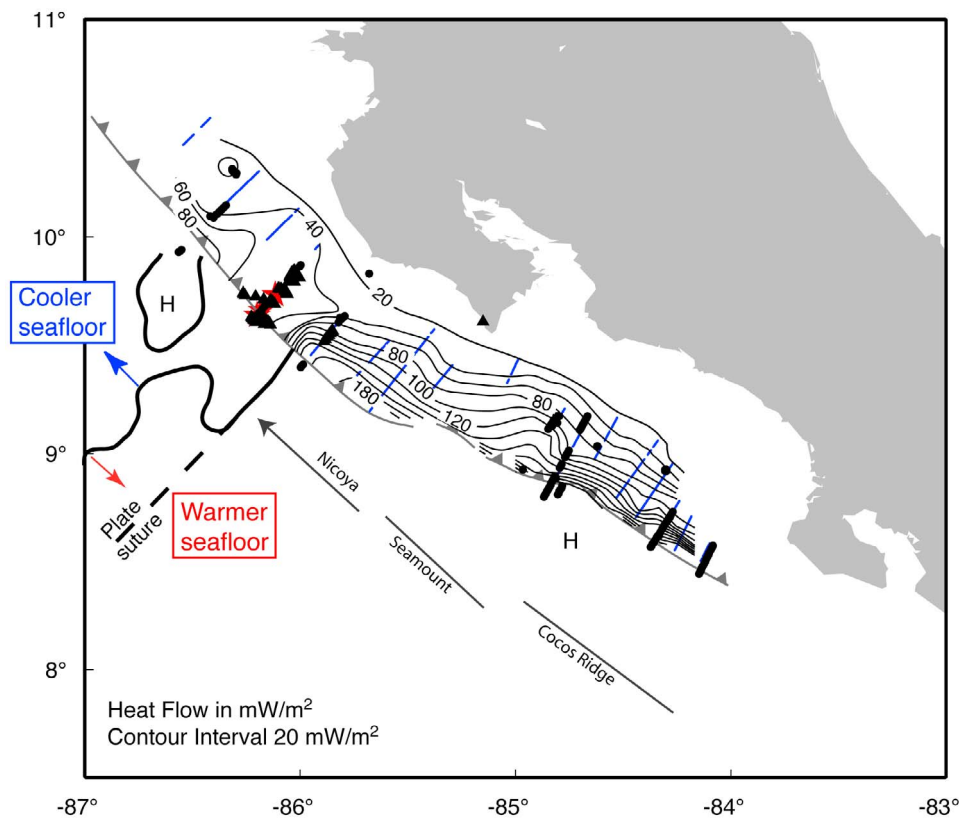


Figure 9. Heat flow map of the Costa Rica margin. Contours are in intervals of 20 mW m^{-2} . Blue lines show location of BSR-derived heat flow, and solid circles show probe values of heat flow. Triangles show data of *Langseth and Silver* [1996], and stars show ODP values of heat flow [*Ruppel and Kinoshita*, 2000]. The boundary between normal and cold EPR crust is from *Hutnak et al.* [2008].

and cooler oceanic crust, mapped by *Fisher et al.* [2003] and *Hutnak et al.* [2007] seaward of the deformation front. Between seismic lines SO81-15 and SO81-13 the margin is disrupted by seamount subduction and there is a relatively large gap in thermal data. Here a large seamount has begun to tunnel into the margin and is marked by considerable slope failure and gravity sliding. An additional buried seamount is uplifting the margin forming Cabo Blanco with rapidly uplifted shorelines [*Gardner et al.*, 2001]. Contours are interpolated through this area and are less precisely located than to the north and south. In the southern seamount segment and Cocos Ridge segment heat flow is anomalously high and decreases rapidly landward.

[39] Another feature of the data are high-wave number variations approximately 5 to 10 km wide (Figures 5–8). These variations are most pronounced where rough seafloor is being subducted south of the Nicoya segment (Figure 6). Although, steady state erosion and sedimentation are often used to explain heat flow fluctuations, the scale

of the variability and its association with subduction basement relief leads us toward alternative explanations. Candidate processes that may cause the observed variations are heat refraction caused by variations in basement relief, advective fluid flow, and the effects of deformation.

[40] In general we do not observe a positive correlation between heat flow and basement relief that suggests to us that heat refraction is a dominant process leading to the observed variability. For example seismic line SO81-21 shows a prominent basement high at a distance of approximately 16 km from the deformation front (Figure 2b). Heat flow over this feature oscillates at wave numbers significantly less than its observed width of about 4 km (Figure 6). To assess the impact of heat refraction, we parameterize basement as a sinusoid with a wave number and amplitude of 4 and 1 km, respectively, and include sediment cover with a minimum thickness of 0.5 km. The basement thermal conductivity may be as high as $3 \text{ W m}^{-1} \text{ K}^{-1}$ and the overlying sediments may be about

$0.9 \text{ W m}^{-1} \text{ K}^{-1}$, so that the ratio of thermal conductivity is about 3:1. The maximum heat flow variability generated by this model is about 15%. This numerical model indicates that while heat refraction may contribute to the observed heat flow variability, it is in general not a large component.

[41] Fluid flow through the margin is a likely source of high wave number variability. Two major pathways of fluid flow associated with the Costa Rican margin are thought to exist. The first is along the decollement [Saffer *et al.*, 2000; Silver *et al.*, 2000] and the second is through normal faults that penetrate the margin [Hensen *et al.*, 2004; Sahling *et al.*, 2008; Ranero *et al.*, 2008]. Near the deformation front, fluid flow is driven by compaction loading of the overriding margin. Saffer *et al.* [2000] estimated a fluid volume flux of $\sim 8 \text{ m}^3/\text{yr}$ per meter of trench length based on correlations of borehole density logs between reference ODP Sites 1039 and 1043, 1.6 km landward of the deformation front. Flow is thought to occur upward from the underthrust sediments to the base of the décollement zone and then through a thin zone of sediments just below the décollement [Screaton and Saffer, 2005]. The fast convergence rate coupled with the subduction of all of the incoming $\sim 400 \text{ m}$ thick sediment section suggests that fluid flow may be more sustained than at other subduction zones [Screaton and Saffer, 2005].

[42] In addition to pore fluid expulsion, fluids are generated through the dehydration of hydrous minerals and phase transformations at depth. Geochemical measurements of pore waters show low values of chlorinity and salinity consistent with fluid production through dehydration reactions at depth [Chan and Kastner, 2000; Hensen and Wallmann, 2005]. These fluids may preferentially flow through normal faults that penetrate the margin. Observations of seafloor seepage mapped at 124 sites are summarized by Sahling *et al.* [2008]. They find that mapped seeps are concentrated within a band 17 to 40 km from the deformation front. Along the Nicoya segment and to the north, the majority of seeps are related to extensional faults, and produce mud mounds ~ 0.4 to 2 km wide and ~ 20 –200 m high. Along the Seamount and Cocos Ridge segments, fluid seeps are associated with normal faults and fractures associated with upper plate uplift and slide scars caused by subducting seafloor relief. Sahling *et al.* [2008] speculate that perhaps two to three times more seeps exist than have been mapped. Examples of focused seeps with clear thermal signatures include Mound Culebra [Grevemeyer *et al.*, 2004] just north of seismic line BGR99-39 (Figures 2a and 5), Mound

11 associated with seismic line BGR99-17 (Figure 8), seeps associated with the slump deposit and head-wall along line SO81-13 (Figure 7), and the seep just off seismic line SO81-3 (Figures 2 and 7). Heat flow anomalies associated with these seeps can be quite large relative to background heat flow. Temperature-depth profiles associated with these seeps are linear suggesting flow rates of less than 1 cm yr^{-1} [Grevemeyer *et al.*, 2004; Ranero *et al.*, 2008]. Modeling of pore water chemistry collected from gravity cores at seeps indicates flow rates across most of the structure are 0.1 to 0.3 cm/yr [Hensen *et al.*, 2004; Ranero *et al.*, 2008]. Other manifestations of advective fluid flow through the margin may be gaps in BSRs associated with the presence of seeps (SO81-9, Figure 6; SO81-2 and SO81-3, Figure 7) and the high-wave number heat flow anomalies 12–22 km landward of the deformation front on observed on seismic line SO81-21 (Figure 6).

[43] We provide an estimate of the flow rate needed to generate these anomalies by solving the one-dimensional steady state advection diffusion equation,

$$\frac{\partial^2 T}{\partial z^2} - \frac{\theta \rho_w c_w}{k_e} V_z \frac{\partial T}{\partial z} = 0 \quad (4)$$

where z is depth, T is temperature, θ is porosity, k_e is effective thermal conductivity, $\rho_w c_w$ is the heat capacity of water and V_z is the mean vertical velocity. At the seafloor we use a constant temperature boundary, T_o , and a constant gradient, Γ_L , at depth, L . The solution for this model is,

$$T(z) = T_o + \frac{\Gamma_L L}{\beta} \frac{\exp\left(\frac{\beta z}{L}\right) - 1}{\exp(\beta)}, \quad (5)$$

where β is the Peclet number,

$$\beta = \frac{\theta \rho_w c_w V_z L}{k_e}. \quad (6)$$

At $z = 0$, heat flow can be expressed as,

$$q = \frac{k_e \Gamma_L}{\exp(\beta)}. \quad (7)$$

In this calculation we assume that the depth of the decollement is 2 km as interpreted from the seismic image (Figure 2), the porosity is 20% and that the background heat flux is 50 mW m^{-2} . With the exception of line SO81-2 (Figure 7) probe values of heat flow are generally 20 to 60 mW m^{-2} larger than BSR-derived estimates at slumps. These heat flow anomalies are consistent with flow rates of 7 to 15 mm yr^{-1} .

[44] Landslides and slumps can also produce variability in heat flow transects and it is interesting to consider the potential effect of a landslide or slump on heat flow. The clearest example of an erosive event is along seismic line SO81-13 (Figures 2d and 7) where a slump has removed approximately 750 m of material. This appears to be a relatively recent event because a BSR has not reformed within the downslope slump material. If we assume that the slump occurred instantaneously, and if the background thermal gradient is about 50°C/km then the top of the unroofed material would have a temperature of about 38°C just prior to unroofing. Instantaneous unroofing increases the near sub-bottom heat flow. The change in heat flow, Δq , due to a step change in temperature, ΔT , can be expressed as,

$$\Delta q = k_e \frac{\Delta T}{\sqrt{\pi \alpha t}} \exp\left(\frac{-z^2}{4\alpha t}\right), \quad (8)$$

where α is thermal diffusivity, and t is time. At $z = 0$, the change in heat flow scales with the inverse square root of time. To produce the observed heat flow anomaly of approximately 100 mW m⁻², equation (8) requires that the slump occurred within about 1000 years before present, assuming a thermal diffusivity of 32 km² Ma⁻¹ characteristic of rocks.

[45] Heat flow data within about 10 km of the deformation front shows relatively large variability (BGR99-44, Figure 5; SO81-8 and SO81-21, Figure 6; SO81-2, Figure 7; P-1600, Figure 8). This data is associated with the sedimentary frontal prism and the high wave number variability may be associated with lateral changes in the intensity of faulting during kneading of sediment, or fracturing due to deformation caused by relative low-relief subducting topography like horst and graben structures [e.g., Zwart *et al.*, 1996; Ganguly *et al.*, 2000].

[46] Two other heat flow features are important to note. The first is the heat flow high along seismic line SO81-5 that is remarkable in its amplitude, width, and good agreement between probe and BSR values of heat flow (Figures 2e and 8). A series of distinct nearly horizontal reflections depict a subducted horst associated located under this heat flow anomaly and may indicate rapid fluid flow contained within the subducting oceanic crust such that heat transfer between the BSR and the seafloor is dominantly conductive. Fluid flow through the subducting oceanic crust is modeled and discussed by Harris *et al.* [2010]. The second feature is the extremely high frontal heat flow

values associated with seismic lines BGR99-15 and P-1600 (Figures 8d and 8e). Both seismic lines are in the southern region of the Cocos Ridge segment where the buoyancy and seafloor relief of the Cocos Ridge has modified the margin so that it is short and steep. Almost no seeps have been identified in this area and it appears that fluid generated from both compaction and higher near-toe temperatures causing shallow dehydration [Ranero *et al.*, 2008] may be driving intense fluid flow along the decollement and frontal prism faults in the region. This fluid flow may help explain the anomalously high heat flow and rapid decrease in heat flow associated with the transition from the deformation front to the rock body of the margin.

9. Conclusions

[47] The combination of in situ measurements and BSR estimates of heat flow along the Costa Rican margin provides insight into the thermal regime of the margin. Our analysis leads us to the following conclusions:

[48] 1. There is an excellent correlation between in situ values of heat flow and those derived from the position of BSRs compared to accretionary prisms. This allows the conclusion that heat transfer between the BSR and seafloor is dominantly conductive.

[49] 2. Heat flow variations along strike are significant and likely reflect the thermal state of the subducting Cocos Plate. In general the margin is relatively cool offshore the northern Nicoya Peninsula where EPR generated oceanic crust is subducting. South of the plate suture the margin is relatively warm and heat flow near the deformation front increases to the south.

[50] 3. High-wave number (5–10 km) heat flow variation along transects is significant. This variation is greater south of the plate suture and is spatially associated with the subduction of basement relief and is likely due to advective fluid flow through fractures related to seamount subduction and normal faults related to upper plate extension. In some cases high heat flow may be associated with the rapid removal of seafloor material during slumping.

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References

- Barckhausen, U., C. R. Ranero, R. von Huene, S. C. Cande, and H. A. Roeser (2001), Revised tectonic boundaries in the Cocos Plate off Costa Rica: Implications for the segmentation of the convergent margin and for plate tectonic models, *J. Geophys. Res.*, *106*, 19,207–19,220, doi:10.1029/2001JB000238.
- Blackwell, D. D., J. L. Steele, and C. A. Brott (1980), The terrain effect on terrestrial heat flow, *J. Geophys. Res.*, *85*, 4757–4772, doi:10.1029/JB085iB09p04757.
- Brown, K. M., and N. L. Bangs (1995), Thermal regime of the Chile triple junction: Constraints provided by downhole temperature measurements and distribution of gas hydrate, *Proc. Ocean Drill. Program Initial Rep.*, *14*, 259–275.
- Chan, L.-H., and M. Kastner (2000), Lithium isotopic compositions of pore fluids and sediments in the Costa Rica subduction zone: Implications for fluid processes and sediment contribution to the arc volcanoes, *Earth Planet. Sci. Lett.*, *183*, 275–290, doi:10.1016/S0012-821X(00)00275-2.
- Christeson, G. L., K. D. McIntosh, T. H. Shipley, E. R. Flueh, and H. Goedde (1999), Structure of the Costa Rica convergent margin, offshore Nicoya Peninsula, *J. Geophys. Res.*, *104*, 25,443–25,468, doi:10.1029/1999JB900251.
- Davis, E. E., and H. Villinger (2006), Transient formation fluid pressures and temperatures in the Costa Rica forearc prism and subducting oceanic basement: CORK monitoring at ODP Sites 1253 and 1255, *Earth Planet. Sci. Lett.*, *245*, 232–244, doi:10.1016/j.epsl.2006.02.042.
- Davis, E. E., R. D. Hyndman, and H. Villinger (1990), Rates of fluid expulsion across the northern Cascadia accretionary prism: Constraints from new heat flow and multichannel seismic reflection data, *J. Geophys. Res.*, *95*, 8869–8889, doi:10.1029/JB095iB06p08869.
- DeMets, C. (2001), A new estimate for present-day Cocos-Caribbean plate motion: Implications for slip along the Central American volcanic arc, *Geophys. Res. Lett.*, *28*, 4043–4046, doi:10.1029/2001GL013518.
- Dickens, G. R., and M. S. Quinby-Hunt (1994), Methane hydrate stability in seawater, *Geophys. Res. Lett.*, *21*, 2115–2118, doi:10.1029/94GL01858.
- Fisher, A. T., C. A. Stein, R. N. Harris, K. Wang, E. A. Silver, M. Pfender, M. Hutnak, A. Cherkaoui, R. Bodzin, and H. Villinger (2003), Abrupt thermal transition reveals hydrothermal boundary and role of seamounts within the Cocos Plate, *Geophys. Res. Lett.*, *30*(11), 1550, doi:10.1029/2002GL016766.
- Ganguly, N., G. D. Spence, N. R. Chapman, and R. D. Hyndman (2000), Heat flow variations from bottom simulating reflectors on the Cascadia margin, *Mar. Geol.*, *164*, 53–68, doi:10.1016/S0025-3227(99)00126-7.
- Gardner, T. W., et al. (2001), Holocene fore arc block rotation in response to seamount subduction, southeastern Peninsula de Nicoya, Costa Rica, *Geology*, *29*, 151–154, doi:10.1130/0091-7613(2001)029<0151:HFBRIR>2.0.CO;2.
- Grevemeyer, I., and H. Villinger (2001), Gas hydrate stability and the assessment of heat flow through continental margins, *Geophys. J. Int.*, *145*, 647–660, doi:10.1046/j.0956-540x.2001.01404.x.
- Grevemeyer, I., A. Rosenberger, and H. Villinger (2000), Natural gas hydrates on the continental slope off Pakistan: Constraints from seismic techniques, *Geophys. J. Int.*, *140*, 295–310, doi:10.1046/j.1365-246x.2000.00009.x.
- Grevemeyer, I., et al. (2004), Fluid flow through active mud dome Mound Culebra offshore Nicoya Peninsula, Costa Rica: Evidence from heat flow surveying, *Mar. Geol.*, *207*, 145–157, doi:10.1016/j.margeo.2004.04.002.
- Harris, R. N., G. Spinelli, C. R. Ranero, I. Grevemeyer, H. Villinger, and U. Barckhausen (2010), Thermal regime of the Costa Rican convergent margin: 2. Thermal models of the shallow Middle America subduction zone offshore Costa Rica, *Geochem. Geophys. Geosyst.*, doi:10.1029/2010GC003273, in press.
- Hartmann, A., and H. Villinger (2002), Inversion of marine heat flow measurements by expansion of the temperature decay function, *Geophys. J. Int.*, *148*, 628–636, doi:10.1046/j.1365-246X.2002.01600.x.
- Hensen, C., and K. Wallmann (2005), Methane formation at Costa Rica continental margin: Constraints for gas hydrate inventory and cross-décollement fluid flow, *Earth Planet. Sci. Lett.*, *236*, 41–60, doi:10.1016/j.epsl.2005.06.007.
- Hensen, C., K. Wallmann, M. Schmidt, C. R. Ranero, and E. Suess (2004), Fluid expulsion related to mud extrusion off Costa Rica—A window to the subducting slab, *Geology*, *32*, 201–204, doi:10.1130/G20119.1.
- Hey, R. N. (1977), Tectonic evolution of the Cocos-Nazca spreading center, *Geol. Soc. Am. Bull.*, *88*, 1404–1420, doi:10.1130/0016-7606(1977)88<1404:TEOTCS>2.0.CO;2.
- Hinz, K., R. von Huene, and C. R. Ranero (1996), Tectonic structure of the convergent Pacific margin offshore Costa Rica from multichannel seismic reflection data, *Tectonics*, *15*, 54–66, doi:10.1029/95TC02355.
- Hutnak, M., et al. (2007), The thermal state of 18–24 Ma upper lithosphere subducting below the Nicoya Peninsula, northern Costa Rica margin, in *The Seismogenic Zone of Subduction Thrust Faults*, edited by T. Dixon and C. Moore, pp. 42–85, Columbia Univ. Press, New York.
- Hutnak, M., A. T. Fisher, R. Harris, C. Stein, K. Wang, G. Spinelli, M. Schindler, H. Villinger, and E. Silver (2008), Surprisingly large heat and fluid fluxes driven through mid-plate outcrops on ocean crust, *Nat. Geosci.*, *1*, 611–614, doi:10.1038/ngeo264.
- Hyndman, R. D., E. E. Davis, and J. A. Wright (1979), The measurement of marine geothermal heat flow by a multipenetration probe with digital acoustic telemetry and in situ thermal conductivity, *Mar. Geophys. Res.*, *4*, 181–205, doi:10.1007/BF00286404.
- Hyndman, R. D., J. P. Foucher, M. Yamano, A. T. Fisher, and the ODP Leg 131 Scientific Party (1992), Deep sea bottom simulating reflectors: Calibration of the base of the hydrate stability field as used for heat flow estimates, *Earth Planet. Sci. Lett.*, *109*, 289–301, doi:10.1016/0012-821X(92)90093-B.
- Kaul, N., A. Rosenberger, and H. Villinger (2000), Comparison of measured and BSR-derived heat flow values, Makran accretionary prism, Pakistan, *Mar. Geol.*, *164*, 37–51, doi:10.1016/S0025-3227(99)00125-5.
- Kimura, G., et al. (1997), *Proceedings of the Ocean Drilling Program, Initial Reports*, vol. 170, doi:10.2973/odp.proc.ir.170.1997, Ocean Drill. Program, College Station, Tex.
- Langseth, M. G., and E. A. Silver (1996), The Nicoya Convergent Margin—A region of exceptionally low heat flow, *Geophys. Res. Lett.*, *23*(8), 891–894, doi:10.1029/96GL00733.



- Lister, C. R. B. (1979), The pulse-probe method of conductivity measurement, *Geophys. J. R. Astron. Soc.*, *57*, 451–461, doi:10.1111/j.1365-246X.1979.tb04788.x.
- McAdoo, B. G., et al. (1996), Seafloor structural observations, Costa Rica accretionary prism, *Geophys. Res. Lett.*, *23*, 883–886, doi:10.1029/96GL00731.
- Minshull, T. A., R. Barolome, S. Byrne, and J. Danobeitia (2005), Low heat flow from young oceanic lithosphere at the Middle America Trench off Mexico, *Earth Planet. Sci. Lett.*, *239*, 33–41, doi:10.1016/j.epsl.2005.05.045.
- Morris, J. D., et al. (2003), *Proceedings of the Ocean Drilling Program, Initial Reports*, vol. 205, doi:10.2973/odp.proc.ir.205.2003, Ocean Drill. Program, College Station, Tex.
- Müller, C., C. Bönnemann, and S. Neben (2007), AVO study of a gas-hydrate deposit, offshore Costa Rica, *Geophys. Prospect.*, *55*, 719–735, doi:10.1111/j.1365-2478.2007.00632.x.
- Pecher, I. A., C. R. Ranero, R. von Huene, T. A. Minshull, and S. C. Singh (1998), The nature and distribution of bottom simulating reflectors at the Costa Rican convergent margin, *Geophys. J. Int.*, *133*, 219–229, doi:10.1046/j.1365-246X.1998.00472.x.
- Pecher, I. A., N. Kukowski, C. R. Ranero, and R. von Huene (2001), Gas hydrates along the Peru and Middle America trench systems, in *Natural Gas Hydrates: Occurrence, Distribution, and Dynamics*, *Geophys. Monogr. Ser.*, vol. 124, edited by C. K. Paull and W. P. Dillon, pp. 257–271, AGU, Washington, D. C.
- Pfender, M., and H. Villinger (2002), Miniaturized data loggers for deep sea sediment temperature gradient measurements, *Mar. Geol.*, *186*, 557–570, doi:10.1016/S0025-3227(02)00213-X.
- Ranero, C. R., and R. von Huene (2000), Subduction erosion along the Middle America convergent margin, *Nature*, *404*, 748–752, doi:10.1038/35008046.
- Ranero, C. R., R. von Huene, E. Flueh, M. Duarte, D. Baca, and K. McIntosh (2000), A cross section of the convergent Pacific margin of Nicaragua, *Tectonics*, *19*, 335–357, doi:10.1029/1999TC900045.
- Ranero, C. R., R. von Huene, W. Weinrebe, and U. Barckhausen (2007), Convergent margin tectonics of Middle America: A marine perspective, in *Central America: Geology, Resources and Hazards*, edited by J. Bunschuh and G. E. Alvarado, pp. 239–265, Taylor and Francis, London.
- Ranero, C. R., I. Grevemeyer, H. Sahling, U. Barckhausen, C. Hensen, K. Wallmann, W. Weinrebe, P. Vannucchi, R. von Huene, and K. McIntosh (2008), Hydrogeological system of erosional convergent margins and its influence on tectonics and interplate seismogenesis, *Geochem. Geophys. Geosyst.*, *9*, Q03S04, doi:10.1029/2007GC001679.
- Ruppel, C. (1997), Anomalously cold temperatures observed at the base of the gas hydrate stability zone, U.S. Atlantic passive margin, *Geology*, *25*, 699–702, doi:10.1130/0091-7613(1997)025<699:ACTOAT>2.3.CO;2.
- Ruppel, C., and M. Kinoshita (2000), Fluid, methane, and energy flux in an active margin gas hydrate province, offshore Costa Rica, *Earth Planet. Sci. Lett.*, *179*, 153–165, doi:10.1016/S0012-821X(00)00096-0.
- Saffer, D. M., E. A. Silver, A. T. Fisher, H. Tobin, and K. Moran (2000), Inferred pore pressures at the Costa Rica subduction zone: Implications for dewatering processes, *Earth Planet. Sci. Lett.*, *177*, 193–207, doi:10.1016/S0012-821X(00)00048-0.
- Sahling, H., D. G. Masson, C. R. Ranero, V. Hühnerbach, W. Weinrebe, I. Klauke, D. Bürk, W. Brückmann, and E. Suess (2008), Fluid seepage at the continental margin offshore Costa Rica and southern Nicaragua, *Geochem. Geophys. Geosyst.*, *9*, Q05S05, doi:10.1029/2008GC001978.
- Schmidt, M. C., M. Schmidt, C. Hensen, T. Morz, C. Müller, I. Grevemeyer, K. Wallmann, S. Mau, and N. Kaul (2005), Methane hydrate accumulation in “Mound 11” mud volcano, Costa Rica forearc, *Mar. Geol.*, *216*, 83–100, doi:10.1016/j.margeo.2005.01.001.
- Screaton, E. J., and D. M. Saffer (2005), Fluid expulsion and overpressure development during initial subduction at the Costa Rica convergent margin, *Earth Planet. Sci. Lett.*, *233*, 361–374, doi:10.1016/j.epsl.2005.02.017.
- Shibley, T. H., M. H. Houston, R. T. Buffler, F. J. Shaub, K. J. McMillen, J. W. Ladd, and J. L. Worzel (1979), Seismic evidence for the widespread occurrence of possible gas-hydrate horizons on continental slopes and rises, *Am. Assoc. Pet. Geol. Bull.*, *63*, 2204–2213.
- Shibley, T. H., K. D. McIntosh, E. A. Silver, and P. L. Stoffa (1992), Three-dimensional seismic imaging of the Costa Rica accretionary prism: Structural diversity in a small volume of the lower slope, *J. Geophys. Res.*, *97*, 4439–4459, doi:10.1029/91JB02999.
- Silver, E., M. Kastner, A. Fisher, J. Morris, K. McIntosh, and D. Saffer (2000), Fluid flow paths in the Middle America Trench and Costa Rica margin, *Geology*, *28*, 679–682, doi:10.1130/0091-7613(2000)28<679:FFPITM>2.0.CO;2.
- Tucholke, B. E., G. M. Bryan, and J. I. Ewing (1977), Gas-hydrate horizons detected in seismic-profiler data from the western North Atlantic, *Am. Assoc. Pet. Geol. Bull.*, *61*, 698–707.
- Vannucchi, P., C. R. Ranero, S. Galeotti, S. M. Straub, D. W. Scholl, and K. McDougall-Ried (2003), Fast rates of subduction erosion along the Costa Rica Pacific margin: Implications for nonsteady rates of crustal recycling at subduction zones, *J. Geophys. Res.*, *108*(B11), 2511, doi:10.1029/2002JB002207.
- von Huene, R., et al. (1985), Site 565, *Initial Rep. Deep Sea Drill. Program*, *84*, 21–78, doi:10.2973/dsdp.proc.84.1985
- von Huene, R., C. R. Ranero, W. Weinrebe, and K. Hinz (2000), Quaternary convergent margin tectonics of Costa Rica, segmentation of the Cocos Plate, and Central American volcanism, *Tectonics*, *19*, 314–334, doi:10.1029/1999TC001143.
- von Huene, R., C. R. Ranero, and P. Watts (2004), Tsunamiogenic slope failure along the Middle America Trench in two tectonic settings, *Mar. Geol.*, *203*, 303–317, doi:10.1016/S0025-3227(03)00312-8.
- Walther, C. H. E. (2003), The crustal structure of the Cocos ridge off Costa Rica, *J. Geophys. Res.*, *108*(B3), 2136, doi:10.1029/2001JB000888.
- Wang, K., and E. E. Davis (1992), Thermal effects of marine sedimentation in hydrothermally active areas, *Geophys. J. Int.*, *110*, 70–78, doi:10.1111/j.1365-246X.1992.tb00714.x.
- Yamano, M., S. Uyeda, Y. Aoki, and T. H. Shipley (1982), Estimates of heat flow derived from gas hydrates, *Geology*, *10*, 339–343, doi:10.1130/0091-7613(1982)10<339:EOHDFD>2.0.CO;2.
- Ye, S., J. Bialas, E. R. Flueh, A. Stavenhagen, and R. von Huene (1996), Crustal structure of the Middle American Trench off Costa Rica from wide-angle seismic data, *Tectonics*, *15*, 1006–1021, doi:10.1016/0012-821X(95)00244-7
- Zwart, G., J. C. Moore, and G. R. Cochrane (1996), Variations in temperature gradients identify active faults in the Oregon accretionary prism, *Earth Planet. Sci. Lett.*, *139*, 485–495, doi:10.1016/0012-821X(95)00244-7.