

Continental-scale temperature variability during the past two millennia

PAGES 2k Network*

Past global climate changes had strong regional expression. To elucidate their spatio-temporal pattern, we reconstructed past temperatures for seven continental-scale regions during the past one to two millennia. The most coherent feature in nearly all of the regional temperature reconstructions is a long-term cooling trend, which ended late in the nineteenth century. At multi-decadal to centennial scales, temperature variability shows distinctly different regional patterns, with more similarity within each hemisphere than between them. There were no globally synchronous multi-decadal warm or cold intervals that define a worldwide Medieval Warm Period or Little Ice Age, but all reconstructions show generally cold conditions between AD 1580 and 1880, punctuated in some regions by warm decades during the eighteenth century. The transition to these colder conditions occurred earlier in the Arctic, Europe and Asia than in North America or the Southern Hemisphere regions. Recent warming reversed the long-term cooling; during the past 30 years (AD 1971–2000), the area-weighted average reconstructed temperature was higher than any other time in nearly 1,400 years.

During the current interglacial period, Earth's climate has undergone significant climate variations that have yet to be quantified at the continental scale, where climate variability is arguably more relevant to ecosystems and societies than globally averaged conditions. Determining the magnitude of these changes is needed to distinguish anthropogenic impacts from the background range of natural variability¹. Reconstructing spatiotemporal patterns of past climate variability helps us to understand and quantify the influence of externally forced and intrinsic dynamics of the global climate system², and to understand natural climate variability, which needs to be considered in future climate scenarios^{3,4}. Here we present a global data set of proxy records and associated temperature reconstructions for seven continental-scale regions. We describe the most prominent features of continental-scale temperature changes at multi-decadal to millennial timescales. In contrast to other recent global-scale reconstructions^{5,6}, this study provides an inter-continental perspective of temperature evolution during the past one to two thousand years.

The PAGES 2k Network and temperature reconstructions

The '2k Network' of the IGBP Past Global Changes (PAGES) project aims to produce a global array of regional climate reconstructions for the past 2000 years (www.pages-igbp.org/workinggroups/2k-network). Nine PAGES 2k working groups represent eight continental-scale regions (Fig. 1) and the oceans. Regional representation brings critical expert knowledge of individual proxy data sets, which is essential for improving palaeoclimate reconstructions⁷. The PAGES 2k Network is coordinated with the National Oceanic and Atmospheric Administration (NOAA) World Data Center for Paleoclimatology to establish a benchmark database of proxy climate records for the past two millennia (Supplementary Note A).

Reconstruction domains for the PAGES 2k regions reported here encompass 36% of the Earth's surface (Fig. 1). Although the regions largely coincide with the continents rather than climatological criteria, the annual mean temperature averaged over these regions explains 90% of the global mean annual temperature variability in the instrumental record (Supplementary Fig. S1). Suitable proxy records from Africa (Supplementary Part II) are currently too sparse for a reliable temperature synthesis⁸, and analysis of palaeoceanographic

data by the recently formed Ocean2k group is in progress⁹. Each regional group identified the proxy climate records that they found were best suited for reconstructing annual or warm-season temperature variability within their region, using *a priori* established criteria (Supplementary Database S1). The PAGES 2k data set includes 511 time series of tree rings, pollen, corals, lake and marine sediments, glacier ice, speleothems and historical documents that record changes in biological or physical processes that can be used to reconstruct temperature variations.

Except for North America, the PAGES 2k reconstructions have annual resolution (Supplementary Fig. S2 and Database S2). Three of the temperature reconstructions span approximately 2,000 years (Arctic, Europe and Antarctica), and three cover the past 1,000–1,200 years (Asia, South America and Australasia). The North American region includes a shorter decadal resolved tree-ring-based reconstruction (back to AD 1200) and a longer 30-year-resolved pollen-based reconstruction (back to AD 360).

Each continental-scale temperature reconstruction was derived using different statistical methods, with each 2k Network group tailoring its procedures to the strengths of their regional proxy records and calibration targets (Supplementary Part II). Most groups used either a scaling approach to adjust the mean and variance of a predictor composite to an instrumental target, or a regression-based technique to extract a common signal from the predictors using principal components or distance weighting (Table 1). Some of the heterogeneities among the regional temperature reconstructions might be due to differences in reconstruction methods, which may underestimate the amplitude of low-frequency variability differently¹⁰. Furthermore, differences between regional reconstructions may reflect differences in their reconstruction targets ('land only' versus 'land and ocean'), proxy types and proxy-dating uncertainties. Temperature variability also differs between summer and annual reconstructions, although the two are correlated in meteorological records from our regions (mean r -value = 0.73 ± 0.08 ; Table 1). Screening proxy records based on their statistical similarity to an instrumental target, as was done for most regional reconstructions, can hamper interpretations that involve comparing the variability of temperature during the twentieth century to that of earlier periods¹¹. To address these issues, and to

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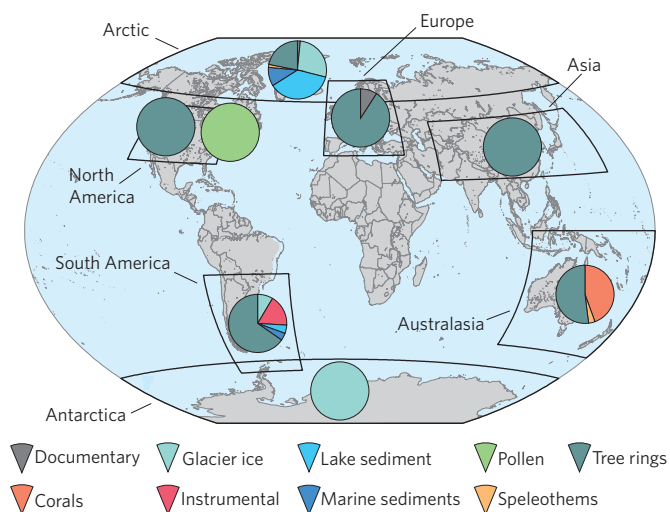


Figure 1 | The PAGES 2k Network. Boxes show the continental-scale regions used in this study. The pie charts represent the fraction of proxy data types used for each regional reconstruction. Supplementary Database S1 includes information about each study site and the proxy data for all time series used in the regional reconstructions.

assess the extent to which the choice of reconstruction method might influence our conclusions, we also applied uniform procedures to the same proxy data to generate three additional reconstructions from each region (Supplementary Note B); these are referred to as 'alternative reconstructions'.

Our analysis of multi-decadal variability focuses on 30-year-mean temperatures (three calendar decades) and their standardized values (Fig. 2; Supplementary Database S2). This temporal resolution enables the inclusion of the 30-year-resolved pollen-based reconstruction from North America. Using standardized values circumvents regional differences in the magnitude of temperature variability, which depends on geographical factors and can be influenced by seasonal biases in some proxies.

Millennial cooling

Over their respective record lengths, all regions experienced a long-term cooling trend followed by recent warming during the twentieth century, except Antarctica. Before AD 1900 (Supplementary Table S2 includes an analysis of regressions up to both AD 1900 and 1850), the cooling trend is significant ($P < 0.05$; see Methods for statistical tests) in all regions except North America, where the cooling trend is weakly significant ($P < 0.10$) in the tree-ring reconstruction, and not significant in the pollen reconstruction. The regional rate of cooling varies between about 0.1 and 0.3 °C per 1,000 years. Since AD 1000 — the interval represented in all regions — the trends are also significant, except for Europe ($P = 0.13$). In general, the overall trends in the 30-year-averaged PAGES 2k Network reconstructions agree with those in the alternative reconstructions (Table S2). They are also consistent with the cooling of global sea-surface temperatures from years AD 1 to 1800 exhibited in the PAGES Ocean2k synthesis⁹.

The individual site-level proxy records were also analysed to determine the extent to which the long-term cooling trend is a common feature, an approach that is independent of the reconstruction procedures. Generally, the longer the proxy record, the more likely it is to exhibit a significant long-term cooling trend (Fig. 3). For example, of the 24 individual records from all regions that extend back to AD 1,

Table 1 | Summary of continental-scale proxy temperature reconstructions used in this study.

Region (area: 10 ⁶ km ²)	Time period (AD)	Proxy type (number)	Reconstruction target	Annual-summer <i>r</i> value*	Reconstruction method [†]	Calibration <i>r</i> value [‡]	Major references
Arctic (34.4)	1-2000	Lake sediments (22) Ice cores (16) Tree rings (13) Marine sediments (6) Documentary (1) Speleothem (1)	Mean annual Land & ocean	0.79	PaiCo	0.56	32 18
Europe (13.0)	1-2003	Tree rings (10) Documentary (1)	Summer (JJA) Land only	0.66	CPS	0.85	33 19 34
Asia (31.1)	800-1989	Tree rings (229)	Summer (JJA) Land only	0.71	PPR	0.52	35
North America (12.5)	1204-1974 480-1950	Tree rings (146) Pollen (4 eco-regions)	Mean annual Land and ocean	0.68	PCSR PCA	0.96 [§] 0.56	36 37
South America (20.0)	857-1995	Tree rings (15) Instrumentals (4) Ice cores (2) Marine sediment (1) Lake sediment (1)	Summer (DJF) Land only	0.82	PCR/CPS	0.82	38
Australasia (37.9)	1001-2001	Coral (13) Tree rings (14) Speleothem (1)	Warm season (Sept-Feb) Land and ocean	0.80	PCR	0.80	39
Antarctica (34.4)	167-2005	Ice cores (11)	Mean annual Land only	0.62	CPS	0.60	40 41

*Correlation between annual and summer (Jun-Aug or Dec-Feb) temperatures based on HadCRUT4, 1850-2012 and using land or land and ocean as per the reconstruction targets; correlation for Australasia is for the warm season (Sept-Feb). [†]See Supplementary Information; PaiCo, pairwise comparison; PCA, principal component analysis; PCR, Principal component reconstruction; PCSR, principal components spatial regression; CPS, composite plus scaling; PPR, point-by-point regression. [‡]*r* and *P* values were calculated for the correlation between reconstructed values and target instrumental data using the seasons, calibration period, domains and gridded instrumental products specified for each region in the Supplementary Information. *P*-values for calibrations are all <0.001 as estimated by phase-randomization Monte-Carlo methods that capture the serial correlation structure of the time series (ref. 42). [§]Correlation for decadal mean reconstruction. ^{||}Square root of the explained variance (*R*²) of the fit of decadal tree-ring-based reconstruction to regional pollen-based reconstructions (0.334) multiplied by *R*² of the decadal tree-ring-based calibration itself (0.929); *P*-value not determined.

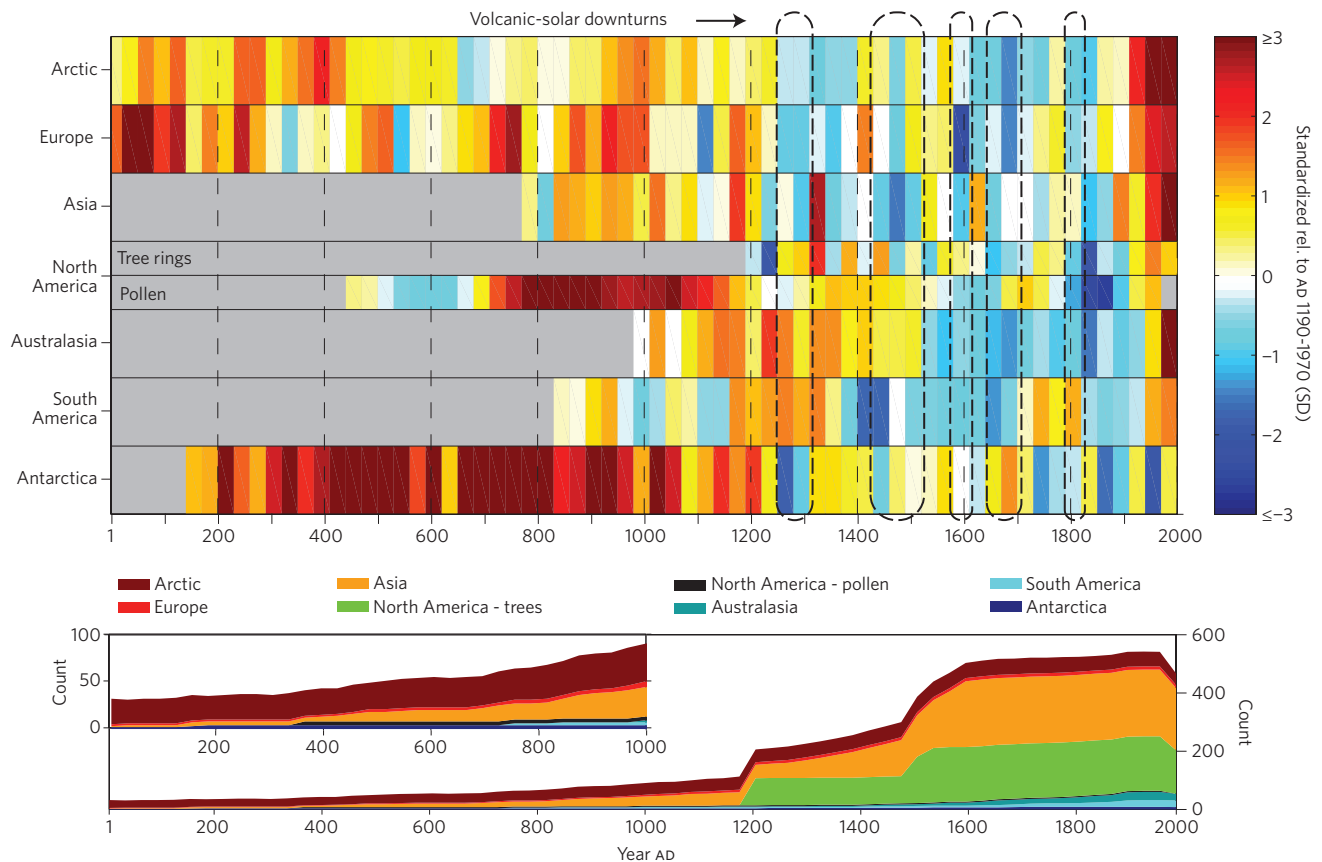


Figure 2 | Continental-scale temperature reconstructions. 30-year mean temperatures for the seven PAGES 2k Network regions, standardized to have the same mean (0) and standard deviation (1) over the period of overlap among records (AD 1190–1970). North America includes a shorter tree-ring-based and a longer pollen-based reconstruction. Dashed outlines enclose intervals of pronounced volcanic and solar negative forcing since AD 850 (see Methods). The lower panel shows the running count of number of individual proxy records by region. Data are listed in Supplementary Database S2.

20 have negative slopes between years AD 1 and 1900 (Supplementary Fig. S6), which is extremely unlikely to occur by chance (one-tailed sign test, $P < 0.001$). Of these, three-quarters of the slopes are statistically significant ($P < 0.05$). Of the 90 records that extend to AD 1000, 54 have negative slopes (sign test, $P < 0.03$), over one-third of which are statistically significant ($P < 0.05$).

The regression analysis of the cooling trend is supported by temperature differences at multi-centennial scales. The three longest reconstructions (Arctic, Europe and Antarctica) show higher temperatures during the early centuries of the first millennium compared with those of the late (pre-twentieth) centuries of the second millennium. Commensurate cooling occurred during the last millennium; for example, the first four centuries (AD 1000–1400) were warmer than the following four centuries (AD 1400–1800) in all regions, despite cooling that starts between AD 1200 and 1300 in the Arctic, Europe and Asia. The average temperature difference between the four-century intervals averaged over all the regions was about $0.1\text{ }^{\circ}\text{C}$, which is consistent with the underlying millennial-scale cooling trend (roughly $-0.2\text{ }^{\circ}\text{C}$ per 1,000 years).

Several studies have investigated the cause of cooling between the tenth and sixteenth centuries, suggesting a potential role of solar irradiance, volcanic activity, land-cover changes and orbitally driven insolation^{6,12–15}. An ensemble of simulations performed with a climate model of intermediate complexity (Supplementary Note C) shows that these four forcings are the dominant cause of the annual mean cooling trend simulated between AD 900 and 1850, with the relative contribution of each forcing varying among regions (Supplementary Table S4). The same forcings are likely to have played a similar role on a longer timescale that includes the entire first millennium. For

example, some records of volcanic¹⁶ and solar¹⁷ activities (Fig. 4f) show significant long-term negative forcing over the last 2,000 years (Mann-Kendall trend test, $P < 0.05$). Furthermore, in the Northern Hemisphere mid to high latitudes, the millennial-scale cooling can be ascribed to an orbitally driven decrease in local summer insolation, as suggested for the Arctic¹⁸ and northern Scandinavia¹⁹. In the Southern Hemisphere, a climate model of intermediate complexity simulates a multi-millennial cooling in summer as a delayed response to the decrease in local spring insolation, modulated by the thermal inertia of the Southern Ocean²⁰.

Multi-decadal to centennial variability

Temperatures did not fluctuate uniformly among all regions, highlighting the regionally specific evolution of temperature at multi-decadal to centennial time scales (Fig. 2). However, the period from around AD 830 to 1100 generally encompassed a sustained warm interval in all four Northern Hemisphere regions. In South America and Australasia, a sustained warm period occurred later, from around AD 1160 to 1370. In the Arctic and Europe, temperatures were relatively high during the first centuries AD. Most other reconstructions are too short to infer temperatures before around AD 1000.

The transition to colder regional climates between AD 1200 and 1500 is evident earlier in the Arctic, Europe and Asia than in North America or the Southern Hemisphere. Differences among regions could reflect non-forced variations involving the major modes of atmospheric variability^{21,22}. By around AD 1580, all regions except Antarctica entered a protracted cold period. Apart from intervals of relative regional warmth, especially during the eighteenth century, cold conditions prevailed until late in the nineteenth century. Again,

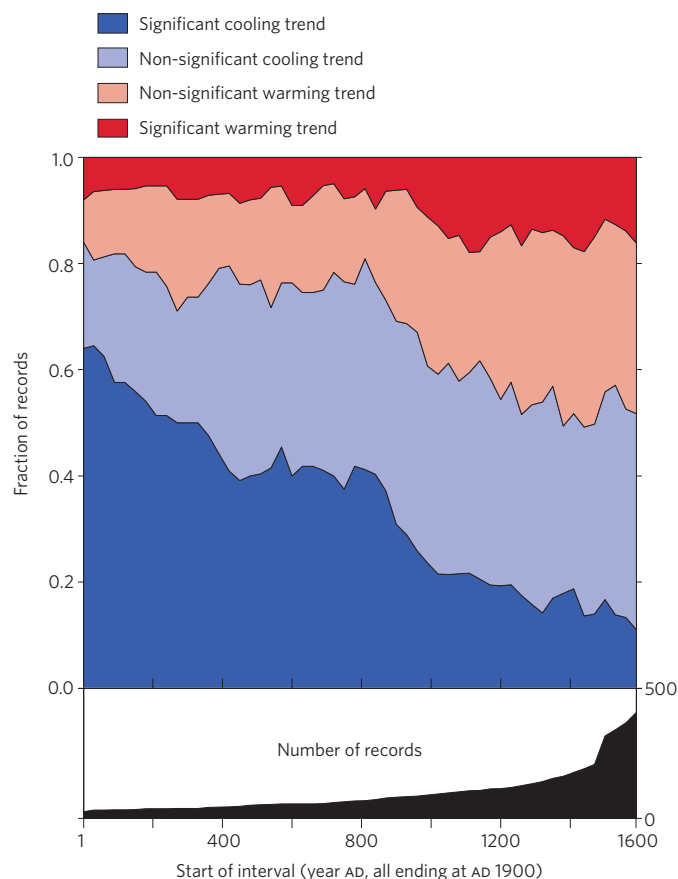


Figure 3 | Summary of long-term trends in individual site-level proxy records. Sign (positive or negative) and statistical significance of the slope of least-squares linear regression through each site-level proxy record within the PAGES 2k data set. The fraction of records that exhibit significant ($P < 0.05$) or non-significant cooling trends was evaluated for records extending back different lengths in 30-year steps. The longer the record, the more likely it is to exhibit a significant long-term cooling trend. For illustration purposes, the fraction of positive trends with magnitude smaller (light red) and larger (red) than the one-sided $P = 0.05$ level is also included.

the alternative reconstructions show patterns similar to those of the PAGES 2k regional reconstructions (Supplementary Fig. S4).

Palaeoclimate records spanning the past millennium are often characterized as including some manifestation of a warm Medieval Warm Period (MWP) followed by a cool Little Ice Age (LIA)²³. Previous reviews of these intervals have shown a tendency for centennial-scale temperature anomalies, but have also emphasized their heterogeneity through space and time^{6,24–27}. Our regional temperature reconstructions (Fig. 2) also show little evidence for globally synchronized multi-decadal shifts that would mark well-defined worldwide MWP and LIA intervals. Instead, the specific timing of peak warm and cold intervals varies regionally, with multi-decadal variability resulting in regionally specific temperature departures from an underlying global cooling trend.

Our regional temperature reconstructions can be compared with records of solar irradiance and volcanic activity to examine the extent to which temperature changes coincide with these climate forcings. The time series of area-weighted, standardized 30-year-mean temperatures averaged across all regions (Fig. 4b) displays broad similarities with records of climate forcings (Fig. 4f,g), a feature also apparent in previously published reconstructions of Northern Hemisphere average temperature (Fig. 4a). The periods of most negative volcanic-solar forcing (nine 30-year periods clustered in five distinct 30-

90-year intervals between 1251 and 1820, as determined using criteria specified in Methods) generally correspond to a drop in the averaged temperature (Fig. 4b). Monte Carlo sampling of the detrended global time series shows that the averaged temperature during these perturbed intervals was significantly lower ($P < 0.01$) than the means of nine randomly selected 30-year periods between AD 830 and 1910. At the regional scale, these volcanic- and solar-perturbed periods include the coolest 30-year period of each reconstruction, although the response in Australasia and North America seems to lag by one 30-year period during the nineteenth century (which includes the cluster of volcanic eruptions between AD 1823 and 1835¹³), and Antarctica was coldest during the middle of the twentieth century (Fig. 2). Not all regions cooled during each of the strong volcanic-solar downturns. For example, Australasia, South America and North America tree rings do not show multi-decadal cooling during the earliest interval of strong negative forcing (AD 1251–1310).

Twentieth-century reconstructed temperature

The twentieth century ranked as the warmest or nearly the warmest century in all regions except Antarctica, where the large thermal inertia of the surrounding ocean may dampen warming²⁸. Excluding Antarctica, the twentieth-century average temperature among the six regions was about 0.4 °C higher than the averaged temperatures of the preceding five centuries (Supplementary Table S3 lists centennial temperature differences based on the PAGES 2k and alternative reconstructions). Compared to the preceding five centuries, twentieth-century warming in the four Northern Hemisphere regions was, on average, about twice that of the more strongly ocean-dominated regions of Australasia and South America (about 0.5 °C compared with 0.2 °C), with the greatest differences at northern high latitudes. Twentieth-century warming in the Arctic (0.9 °C) was about three times that of the average of the other five non-polar regions.

Our best estimate of reconstructed temperature for AD 1971–2000 can be compared with all other consecutive 30-year periods within each regional reconstruction. In Asia and Australasia, reconstructed temperature was higher during 1971–2000 than any other 30-year period. The Arctic was also warmest during the twentieth century, although warmer during 1941–1970 than 1971–2000 according to our reconstruction. In South America, the AD 1971–2000 reconstructed temperature was similar to the record maximum in AD 1251–1280. In North America, the reconstructed temperature for the 1971–2000 interval does not include the warm decades since 1980, and therefore underestimates the actual temperature for that interval. In Europe, slightly higher reconstructed temperatures were registered in AD 741–770, and the interval from AD 21–80 was substantially warmer than 1971–2000. Antarctica was probably warmer than 1971–2000 for a time period as recent as AD 1671–1700, and the entire period from 141–1250 was warmer than 1971–2000. These interpretations are generally supported by the relative magnitude of recent warming in the alternative reconstructions (Supplementary Fig. S4 and Database S2).

Each individual proxy record contributing to the regional reconstructions was analysed to evaluate whether the values during 1971–2000 indicate higher temperatures than for any other 30-year period (Fig. 4d,e), independent of the procedures used for calibrating the temperature reconstructions. According to this analysis, of the 323 individual proxy records that extend to AD 1500, more sites seem warmest during 1971–2000 than during any other 30-year period, both in terms of the total number of sites and their proportion in each region. Similarly, of the 52 individual records that extend to AD 500, more sites (and a higher proportion) seem warmest during the twentieth century than during any other century. The fraction of individual records that indicates the highest temperatures during 1971–2000 decreases with increasing record length, consistent with an overall cooling trend over the past two millennia (Figs 2 and 3).

The area-weighted average of the best estimate of past temperature from all seven regions indicates that 1971–2000 was warmer than

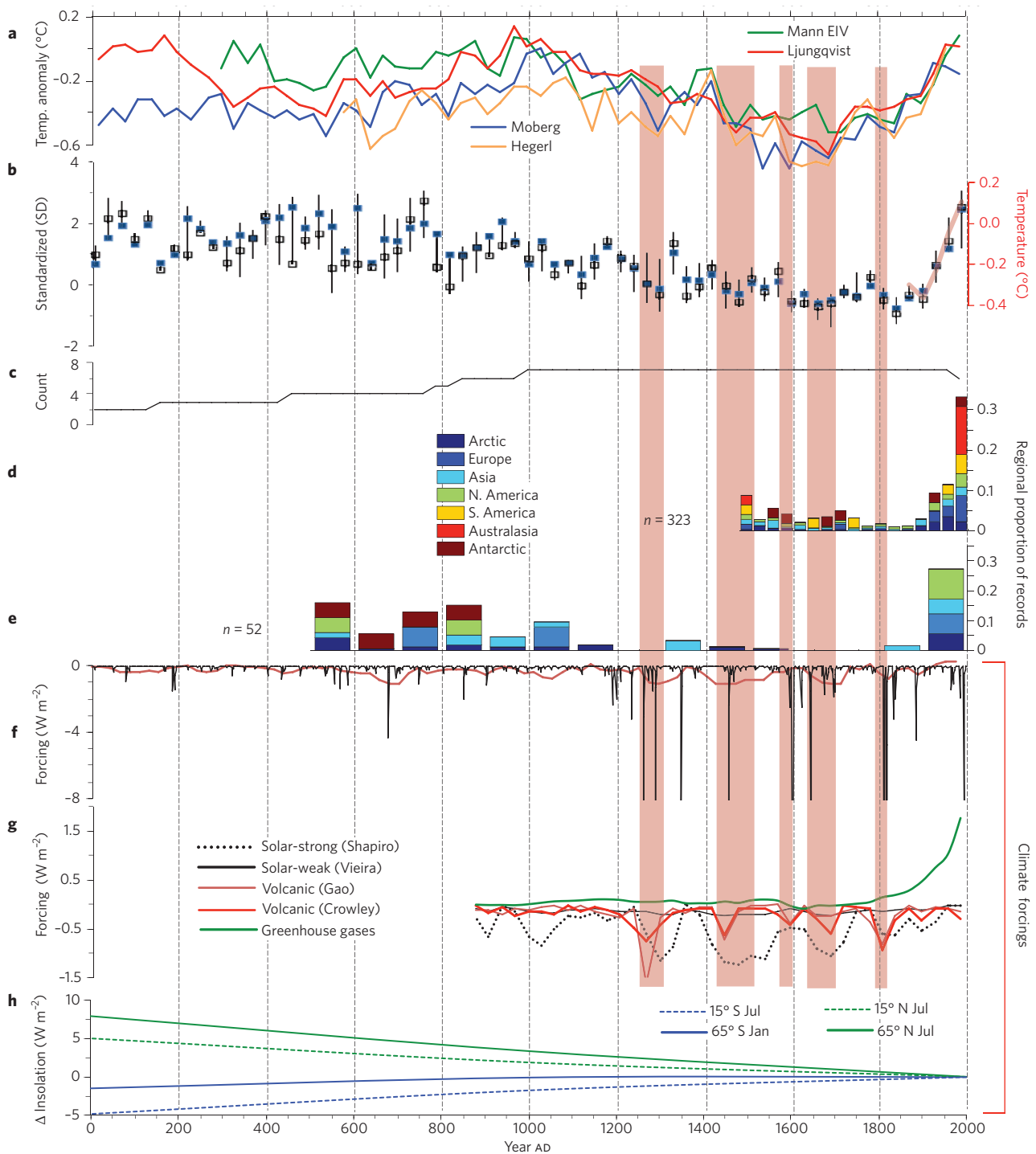


Figure 4 | Composite temperature reconstructions with climate forcings and previous hemisphere-scale reconstructions. **a**, Previously published Northern Hemisphere 30-year-mean temperature reconstructions relative to the 1961–1990 reference period^{d5,43–54}. **b**, Standardized 30-year-mean temperatures averaged across all seven continental-scale regions. Blue symbols are area-weighted averages using domain areas listed in Table 1, and bars show twenty-fifth and seventy-fifth unweighted percentiles to illustrate the variability among regions; open black boxes are unweighted medians. The red line is the 30-year-average annual global temperature from the HadCRUT4 (ref. 29) instrumental time series relative to 1961–1990, and scaled visually to match the standardized values over the instrumental period. **c**, Running count of the number of regional reconstructions. **d**, For each 30-year period since AD 1500, the proportion of individual proxy records within each region that indicate the highest temperature during that 30-year period. **e**, For each century since AD 500, the proportion of individual proxy records within each region that indicate the highest temperature during that century. **f**, Long-term volcanic forcing from ref. 16 (black curve; spikes beyond -8 W m^{-2} are truncated), and solar forcing from ref. 17 (red curve). **g**, Radiative forcings relative to AD 2000 smoothed using 30-year averages from ref. 31, including: two estimates of volcanic forcing^{46,47}; two estimates of solar forcing that span the range from strong⁴⁸ to weak⁴⁹; and well-mixed greenhouse gases relative to AD 850. **h**, Change in summer (July and January) insolation at 65°N/S and 15°N/S latitudes relative to AD 2000 from ref. 50. Vertical red bands indicate volcanic-solar downturns as defined in Methods.

any other time in nearly 1,400 years (Fig. 4b), keeping in mind that this analysis does not consider the uncertainty associated with the temperature estimates, and that the reconstructions are of different lengths. Area-weighted averages of the three alternative reconstructions generally support this result (Supplementary Fig. S5). Large uncertainties remain, especially during the first millennium, when only some regions are represented. Regardless, the global warming that has occurred since the end of the nineteenth century reversed a persistent long-term global cooling trend. The increase in average temperature between the nineteenth and twentieth centuries exceeded the temperature difference between all other consecutive centuries in each region, except Antarctica and South America.

The PAGES 2k Network reconstructions show clear regional expressions of temperature variability at the multi-decadal to centennial scales, whereas a long-term cooling trend before the twentieth century is evident globally. Centennial-scale temperature changes in Australasia and South America generally follow those of the Northern Hemisphere regions, whereas temperature changes in Antarctica do not correlate with those of other regions. The pre-twentieth-century long-term global cooling trend is evident in differences in average temperature between multi-centennial periods, and in the significance of the slopes of least-squares linear regressions for both the regional reconstructions and the individual site-level proxy records. Our reconstructions and proxy-data compilation will be useful in future studies, serving as a benchmark for comparisons with climate model simulations aimed at understanding the cause of the global cooling, and the extent to which externally forced and unforced variability can explain temperature fluctuations and trends at the continental scale.

Methods

Data sources. All proxy records used for the regional reconstructions are included in Supplementary Database S1, and are archived by the NOAA World Data Center for Paleoclimatology (www.ncdc.noaa.gov/paleo/pages2k/pages-2k-network.html). All regional temperature reconstructions are included in Supplementary Database S2.

Relation between proxy records and temperature. The sign (positive or negative) of the relation between each proxy record and temperature is listed in Supplementary Database S1. These relations were determined using different approaches for different regions. In the Arctic, the relation is adopted from the publication of the original records. In Europe, the 11 inputs to the regional reconstruction include 10 tree-ring composites and documentary evidence, all calibrated to temperature. The North America pollen-based reconstruction also relies on regional syntheses that were previously transformed into temperature. For Asia and Antarctica, only records with demonstrated positive correlations with local or regional (respectively) instrumental data were included in the reconstructions. The Asia, North America tree-ring, South America and Australasia reconstructions were based on principal component regression methods, whereby individual proxy records could contribute with either a positive or negative temperature relation with different principal components, or for ensemble members with different instrumental targets. In the North America tree-ring reconstruction, the sign of the contribution for each record could vary spatially as well. The sign of the relation with temperature for these regional reconstructions listed in Database S1 is based on the direction of the relation between the proxy records and the area-weighted average mean annual temperature from the domain areas in Fig. 1 using the HadCRUT4 (ref. 29) data series.

Significance of trends. The significance of long-term cooling trends in the reconstructions (Supplementary Table S2) and in the individual site-level proxy records (Fig. 3 and Supplementary Fig. S6) was calculated at the $P = 0.05$ confidence level using a one-sided Student's

t -test, while accounting for lag-1 autocorrelation in both the calculation of standard error and in indexing the t -values³⁰. Trends in the individual proxy records (site level) were determined by first inverting those proxy records having a negative correlation with temperature (as described above). The pre-1900 trend in each record was analysed for successively shorter segments of time, beginning with the entire time series, then truncating each series by 30-year intervals, with each segment ending in AD 1900. The trend in each truncated series was determined by the slope of its linear regression. For illustration purposes, the fraction of positive trends with magnitude larger than the one-sided $P = 0.05$ level are shown in Fig. 3.

Volcanic-solar downturns (Figs 2 and 4). We analysed two time series of solar forcing and two of volcanic forcing from ref. 31 (Fig. 4g), the volcanic series from ref. 16 (Fig. 4f), and the solar series from ref. 17 (Fig. 4e) to discern periods of strongest agreement and magnitude of forcing. Each time series was binned into 30-year intervals that coincide with the temperature reconstructions (bin centres from AD 845–1985). The eight bins (21% of each record) with most negative forcing were selected from each series; then, the subset of bins overlapping among at least one-half (≥ 3) of the six time series was identified. The nine 30-year bins defined this way comprise the five intervals that we term 'volcanic-solar downturns': AD 1251–1310, 1431–1520, 1581–1610, 1641–1700 and 1791–1820.

Temperatures during volcanic-solar downturns. To evaluate whether temperatures were significantly lower during intervals of negative climate forcing, we implemented a Monte Carlo procedure to randomly sample the area-weighted regional temperatures between AD 831 and 1910 (we excluded the warm twentieth century (AD 1911–2000) to make the test more conservative). For each of 10,000 iterations, the mean temperature of nine 30-year intervals was calculated. Because most of the intervals of negative forcing occur later in the record, the low temperatures during these intervals may be partially due to long-term cooling. To assess this possibility, we also conducted the test on the detrended area-weighted mean. For both the original and detrended tests, the mean of the nine intervals of negative solar and volcanic forcing was lower than more than 99% of the means of randomly selected intervals (that is, $P < 0.01$).

Warmest intervals within individual proxy records (Fig. 4d,e). Site-level warmest-interval analysis was conducted for 30-year intervals spanning from AD 1491 to 2000, and for 100-year intervals between AD 501 and 2000. Proxy records with negative correlation with temperature were inverted. Only records with at least one measurement in each 30- or 100-year bin were used. The three tree-ring records from northern Fennoscandia, which were used in both the Arctic and the Europe reconstructions, were included only once in the site-level analysis. The warmest intervals were identified by ranking the values of the 30- or 100-year bins for each record. To better represent regions with fewer records, the values were presented as the fraction of records in each region, rather than the number of records. The fractions were scaled such that a value of 1 corresponds to all records in all regions.

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Additional information

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