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Microstructural deformation features in artificially deformed ice and in polar ice cores (EDML, Antarctica)

Introduction

Ice of polar ice sheets is an important information source for studying the past climate. For a correct interpretation of the age of the ice knowledge on flow and deformational processes following the deposition is essential.

Recently obtained data on sub-grain boundaries (sGB) of the deep ice core from EPICA-Dronning Maudiand (Hamann et al. in prep. Figure 1 and 2) reveal that an onset of polygonizationsub-grain rotation recrystallization or an onset of migration recrystallization cannot be found, indeed this study shows that sub-grain formation is active in any depth of EDML tec core, which indicates that the classical tripartition of recrystallization regimes (1. Grain growth, 2. Polygonization/sub-grain rotation recrystallization, 3. Migration recrystallization), which is the standard conception (e.g. Duval 2000), is not easily applicable here and has to be reconsidered.

In order to understand the sGB evolution with ongoing deformation under controlled conditions, creep experiments follow standard procedures have been conducted and creep test samples have been prepared as thin sections to investigate microstructures ((dipstuh) et al. submitted).



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With increasing strain from 0.5 to 2% sub-grain boundary density increases, reaching a stable value at approximately 3% strain. As strain rates reached a steady state at ca. 2 to 3% during these experiments, an explanation might be the following. The decreasing trend of the creep rate corresponds to the evolution of the substructure of the crystal, namely the production of disocation suble and sub-grain boundaries. During primary transient creep the strain rate decreases continuously, which means that the deformation speed slows down, i.e. the material hardens. Taking find account that the deformation is to a large extent accomplished by introduction and movement of disocations, one cause for this decrease of deformation rate is the production of disocation wills and sub-grain boundaries. But findens. Taking find excerns the deformation rate is absen thre as pure ice is used) are dislocation will as dus ub-grain boundaries. As the production of such obstacles continues the deformation rate keeps decreasing until a steady arount of obstacles is reached. This probably falls together with the achievement of maximum value of the sub-grain boundary density. As dislocations cannor move freely stress must be accumulating around obstacles. Due to the fact that deformation rate increases again at some point a new process must take over the main deformation calve. The grain brotation or grain boundary migration. That grain boundary migration intensifies and starts to dominate can be observed by the grain shape parameter.

Observations of microstructures from samples deformed by creep experiments reveal information on the formation of these features. During experiments under conditions, which allow grain growth, the grain shape changes significantly with increasing strain. The grains become more inregular with bulging grain boundary migrations is at least partly caused by internal strain energy differences. This is the case in experiments conducted at high temperature (5°C) as well as under low temperatures (2°C), where migration expectialisation was not expected according to standard conceptions.

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References

Bons, P.D., Jessell, M.V. 1999. Micro-shear zones in experimentally deformed OCP. J. Struct. Geol. 21, 323-334. Daval, P. 2000. Deformation and dynamic recrystallization of ice in polar ice aheets. In: Physics of loce core records p. 103-113. ed. T. Hondoh. Hokkado University Press. Sapporo. Forl H. 2000. Deformation and dynamic recrystallization of ice in polar ice aheets. In: Physics of loce core records p. 103-113. ed. T. Hondoh. Hokkado University Press. Sapporo. Forl H. 2000. Deformation and S.H. Azuma, N. Sak-gain boundaries in EPRAC Dromming Maculated (EDML) deep ice core. In preparation. Heamann, L, Spithult, S. Fana, S.H. Azuma, N. Sak-gain boundaries in EPRAC. Dromming Maculated (EDML) deep ice core. In preparation. Kytohauf, S. Hamann, L, Lambeed, L., Arrelag, J. Fanis, S.H. Gorgover, D. Azumann. 2008. Microheature manging a new method for imaging deformation-induced microstructural features of ice on the grain scale, submitted to J. Glacici. Wang, Y. Kipfstuhl, S., Azuma, N., Thorsteinsson, Th., Miller, H. 2003. Ice-tabrics study in the upper 1500 m of Dome C (East Antarctica) deep ice core. Ann. Glacol. 37, 97-104.