1	Optical absorption, infrared, Raman and EPR spectral studies on natural Iowaite
2	mineral
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14	Abstract
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16	Natural iowaite, magnesium-ferric oxychloride mineral having light green colour
17	originating from Australia has been characterized by EPR, optical, IR and Raman
18	spectroscopy. The optical spectrum exhibits a number of electronic bands due to both Fe(III)
19	and Mn(II) ions in iowaite. From EPR studies the parameters derived, g values for Fe(III)
20	and g and A values for Mn(II) confirm optical absorption studies and are in distorted
21	octahedron. The bands that appear both in NIR and Raman spectra are due to the overtones
22	and combinations of water and carbonate molecules. Thus EPR, Optical and Raman
23	spectroscopy has proven most useful for the study of the chemistry of natural iowaite and
24	chemical changes in the mineral.
25	
26	Keywords: Iowaite mineral, hydrotalcite, optical absorption spectrum, EPR spectrum,
27	Fe(III), Mn(II). Raman spectroscopy, infrared spectroscopy
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- **30 1. Introduction**
- 31

32 Magnesium based hydrotalcite is also known as iowaite. Iowaite is a hydrous 33 magnesium-ferric oxychloride mineral. It occurs as bluish-green close-green crystals in 34 calcite-brucite veiniets and cavities in magnesian slam and altered dolostones. . The usual formula is $Mg_6Fe_2^{3+}(OH)_{16}Cl_2 4H_2O$. Oxidation of iowaite results in the formation of Mg_4FeO 35 36 (Cl₂CO₃)(OH)₈ 4H₂O On exposure to the atmosphere it turns whitish green with a rusty 37 overtone, owing to loss of zeolitic water. Iowaite belongs to R/3m space group. The 38 hexagonal unit cell parameters are a = 0.3116 nm, c = 2.4113 nm.[1-4]. Its structure 39 consisting of three brucite like layers and three disordered inter layers. The water molecules 40 are migrating within the inter layers [4]. The mineral iowaite does not show any marked 41 variation in composition. Iowaite from the Korshunovskoe, Siberia contain MgO = 39.80, 42 $Fe_2O_3 = 24.71$ and Cl = 5.41 wt%. Further ESR studies on it indicate that the presence of Fe(III) in the crystal structure and Mn(II) as an impurity[1]. Also Raman and IR spectra were 43 Hydrotalcites has many applications such as nano-44 reported on this compound [2]. 45 composites, in the removal of environmental hazards in acid mine drainage, anion exchangers, disposal of radio-active wastes, heavy metal removal from contaminated water 46 47 etc., [2,5-7]. Research into the use of these hydrotalcites results from their potential 48 application as catalysts, adsorbents and anion exchangers [5]. In this study, the authors made 49 an attempt to determine the valence state of iron and the site symmetry using electron 50 paramagnetic resonance (EPR), optical absorption, Raman and infrared spectroscopic 51 features of natural iowaite from Australia.

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53 **2. Experimental**

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A light green coloured iowaite originating from Australia is used in the present work. EPR spectra of the sample in powder form are recorded at room temperature (RT) and liquid nitrogen temperature (77 K) on a JEOL JES-TE100 ESR spectrometer operating at X-band frequencies (v = 9.41095 GHz), having a 100 KHz field modulation to obtain a first derivative EPR spectrum.

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61 Optical absorption spectrum of the sample was recorded at room temperature in the 62 region 200- 2500 nm on a Varian Cary 5E UV-Vis- NIR spectrophotometer. The spectra were transformed according to the Kubelka-Munk algorithm. NIR spectrum of the sample
was recorded on a Nicolet Nexus FT-IR spectrometer with a Nicolet near IR fibreport
accessory. A white light source was employed with a quartz beam splitter and TECNIR.

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The powdered mineral sample was placed and oriented on the stage of an Olympus BHSM microscope, equipped with 10x and 50x objectives and part of a Renishaw 1000 Raman microscope system. Raman spectra were excited by He-Ne laser (633nm) at a resolution of 2 cm⁻¹ in the range between 100 and 4000 cm⁻¹. Other details of the experimental technique have already been reported [8,9]

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73 **3. Results and Analysis**

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75 3.1 EPR Results

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77 EPR studies are carried out at room temperature (RT) on a JEOLJE TES 100 ESR 78 spectrometer operating at X band frequencies (v = 9. 41095 GHz) having a 100 kHz field 79 modulation and a phase sensitive detection to obtain a first derivative EPR signal. The EPR 80 spectrum of powdered iowaite is shown in Fig.1. The spectrum can be divided into two parts. 81 First part, having resonances at low field region, centered around 170 mT and another part, 82 consisting of a broad line situated around 350 mT. The low field resonances have g values of 83 4.279 and 3.911. These lines can be ascribed to Fe(III) resonances in the mineral sample [10], based on previous observations. Fe(III) is having five unpaired electrons and hence S =84 The spin states in the absence of an external magnetic field are $|\pm 1/2\rangle$, $|\pm 3/2\rangle$ 85 5/2. and $\pm 5/2$. These three Kramer's doublets are separated by 2D and 4D respectively. 86 Application of external magnetic field, splits these Kramer's doublets into six levels labelled 87 as $|-5/2\rangle$, $|-3/2\rangle$, $|-1/2\rangle$, $|1/2\rangle$, $|3/2\rangle$ and $|5/2\rangle$ respectively. EPR transitions between 88 these levels will give rise to five resonances, known as fine structure lines. However, in case 89 90 of Fe(III), the resonances are ranging in the whole magnetic field range of X-band EPR 91 spectrometer, i.e., g values range very largely. This type of observation has been explained, 92 by considering that the three Kramers doublets are populated at temperature of measurement. 93 As mentioned in the literature [10], the observed g values range from 30/7 to 2/7, 2 to 6 and 2/7 to 6 depending upon the population of a particular Kramers doublet. However, in case of 94 95 a powder spectrum, the resonances at 30/7 and 2 are generally observed, due to large

anisotropy in g values. In the present case, we have observed resonances at g = 4.279 and 96 97 3.911. In addition, the broad line at g ≈ 2 also arises from Fe(III) impurity. The literature 98 survey indicates that the observation of g=30/7 resonances confirms that the Fe(III) impurity 99 is under the influence of a strong tetragonal distortion [10]. A close look at high field resonances indicates a sextet, in addition to the broad bond, already discussed earlier. The g 100 value of this sextet is 2.007 and the hyperfine coupling constant is 9.04 mT. These spin 101 102 Hamiltonian parameters confirm that the impurity responsible for these resonances is Mn(II). 103 The low intensity of the sextet indicates a very low concentration of impurity. The other 104 resonances can be thought of arising from Mn(II) lines, due to non-zero value of D. 105 However, due to their low intensity, the value of D cannot be estimated.

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107 The molecular formula indicates the presence of Fe(III), as observed by EPR. The 108 small amount of Mn(II) impurity might have entered the lattice in place of Mg(II), 109 considering the ionic radii and charge of these two ions.

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111 3.2 Optical absorption Results

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Ferric iron, Fe(III), has the electronic configuration Ar $(3d^5)$ with a half field d-shell having one unpaired electron in each of the orbital. Hence, ground state configuration is $t_{2g}^3 e_g^2$ and has only spin forbidden d-d transitions. These occur from the ground state ${}^6A_{1g}(S)$ to the excited states ${}^4T_{1g}(G)$, ${}^4T_{2g}(G)$, ${}^4A_{1g}(G) {}^4E(G)$, ${}^2T_{1g}(D)$, ${}^4E_g(D)$ and ${}^4T_{1g}(P)$ states in regular octahedron sites. The degeneracy of E and T states being lifted with lower symmetry.

Optical absorption spectrum of iowaite recorded in the mull form at RT is shown in 119 Fig.2. The spectrum shows energies at12660, 13070, 13425, 16130 cm⁻¹ in the UV-Vis 120 region. The very sharp band observed at 13070 cm⁻¹ with split component on either side (at 121 12660 and 13425 cm⁻¹ with average of 13043 cm⁻¹) is assigned to the transition ${}^{6}A_{1s}(S) \rightarrow$ 122 ${}^{4}T_{1g}(G)$. Similar assignment is also made in carbonate mineral (dolomite) [11] whereas the 123 band at 16130 cm⁻¹ is assigned to ${}^{4}T_{2g}(G)$ transition. The band visible around 21000 cm⁻¹ due 124 to the transition of ${}^{6}A_{1g}(S) \rightarrow {}^{4}A_{1g}(G)$, ${}^{4}E_{g}(G)$ transition is not seen. This may be due to low 125 126 concentration of iron in the mineral. These bands are characteristic of Fe(III) ion in 127 octahedral symmetry in the mineral. This is also further supported by EPR studies. Using the Tree's polarization term [12] the energy matrices of the d^5 configuration are solved for 128

various B,C and Dq values. A comparison is also made between the calculated andobserved energies of the bands and these are presented in Table 1.

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Table 1

Band headed data with assignments for Fe(III) in natural iowaite.

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 $Dq = 900, B = 600 \text{ and } C = 2595 \text{ cm}^{-1} \alpha = 90 \text{ cm}^{-1}$

Wave length	Wave number (cm ⁻¹)		Transition
(nm)	Observed	Calculated	from
			${}^{6}A_{1g}$
790	12660		
765	13070	13087	${}^{4}T_{1g}(G)$
745	13425		
620	16130	16852	${}^{4}T_{2g}(G)$
		20739	${}^{4}E(G)$
			${}^{4}A_{1g}(G)$

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137 3.3. NIR and infrared spectroscopy

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139 Water has three fundamental modes. They are symmetric OH stretch (v_1) , H-O-H bending mode (v_2) and asymmetric OH stretch (v_3) . In solid these will appear at 3220 (v_3) , 140 1620(v_2) and 3400 (v_1) cm⁻¹ [13]. The shifting of v_1 and v_3 towards lower frequency side and 141 v_2 towards higher frequency is due to hydrogen bonding [14] The carbonate ion has six 142 normal modes of vibration, the symmetric stretching mode v_1 , which is infrared active, the 143 144 out of plane bending mode v_2 , the asymmetric stretching mode v_3 and in-plane bending mode v₄. v₃ and v₄ are doubly degenerate. The fundamental frequencies of CO_3^{2-} ion are v₁ = 1063, 145 $v_2 = 879$, $v_3 = 1415$ and $v_4 = 680$ cm⁻¹ [15,16] 146 147

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151 **T**

The 10000 to 6500 cm^{-1} spectral region

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Fig. 3 shows near infrared spectrum of natural iowaite mineral from 10000 - 6500 cm⁻¹ 153 region. In this region the two bands observed at 9678 and 8661 cm⁻¹ are broad and intense 154 and is assigned to the two components of the transition ${}^{5}T_{2g} \rightarrow {}^{5}E_{g}$ which is clearly due to 155 trace of Fe(II) present in the sample. The average of these bands 9170 cm⁻¹ is taken as 10Dq. 156 The group of bands at 7174 with maximum intensity are assigned to $2v_3$ the asymmetric OH 157 stretch in the mineral. Hydroxyl exists as part of the structure and the stretching mode 158 appears whenever water is present in any form. The v_{OH} overtone gives rise to a band in the 159 NIR spectrum [14]. Accordingly the bands observed at 7681 and 6986 cm⁻¹ are assigned to 160 161 the first overtone of OH.

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163 The 6000 to 4000 cm⁻¹ spectral region

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165 The spectral region from 6000 to 4000 cm⁻¹ is made up of two parts the 4000 – 4600 166 cm⁻¹ region in which OH combination bands are to be found and 4600 - 5600 cm⁻¹ in which 167 water combination bands are observed. Fig.4 shows profile of bands around 4300 cm⁻¹ with 168 large intensity (4401,4347, 4294, 4129 cm⁻¹) and are assigned to the water OH combination. 169 Where as the intense band observed at 5080 cm⁻¹ is attributed to the water OH overtone. The 170 band observed at 5554 cm⁻¹ with weak intensity is attributed to the first fundamental overtone 171 OH stretching mode [15].

172

173 The IR spectra as studied also be subdivided into three spectral regions: (a) 3700 to 174 2900 cm⁻¹ region where OH fundamental vibrations are observed (Fig.5); (b) 1750 to 1250 175 cm⁻¹ range where the water bending modes are observed (Fig. 6); (c) 1125 to 525 cm⁻¹ region 176 shows bands due carbonate stretching vibrations (Fig. 7). As Raman spectroscopy of iowaite 177 has already been reported [1], the analysis is not discussed here.

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179 The 3700 to 2900 cm⁻¹ spectral region

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The OH stretching region of iowaite is shown in Fig.5. The IR spectrum shows several bands at 3685, 3662, 3647, 3633,3572,3433,3269 and 3051 cm^{-1} . The very sharp bands at 3685, 3662 and 3633 cm⁻¹ with component at 3647 cm⁻¹ are assigned to Mg₃OH, Mg₂FeOH and Fe₃OH bands. Similar assignments were already reported in the literature [2]. This assignment is due to OH stretching vibrations of water (v_3). Whereas other two broad and intense bands at 3572 and 3433 cm⁻¹ and two weak bands at 3269 and 3051 cm⁻¹ are assigned

187 to the stretching vibrations of inter layer water.

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189 *The 1750 to 1250 cm⁻¹ spectral region*

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191 This region is significant for the observation of H₂O bending modes. The v₂ water-bending 192 modes are more pronounced in the IR spectrum. Fig.6 shows a broad and sharp band at 1639 193 cm⁻¹ is noticed in spectrum. The appearance of v₂ mode on the lower wave number side is an 194 indication of weakly hydrogen bonded water. This complexity of bands provides evidence for 195 the existence of water in several states in the structure of iowaite. Broad and intense profile 196 centered at 1468 cm⁻¹ with maximum half width is identified as due to v₃ CO_3^{2-} anti-197 symmetric stretching vibrations.

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199 The 1125 to 525 cm⁻¹ spectral region

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The infrared spectra for the fundamental vibrations of carbonate in the iowaite mineral are shown in Fig. 7. All the spectral features in this region are characteristic of carbonate groups. The band at 1016 cm⁻¹ in the spectrum is attributed to CO symmetric stretching mode. The intense band at 1065 cm⁻¹ with component at 1088 cm⁻¹ is identified as v_1 the symmetric stretching mode of carbonate.[15]. Another group of bands observed around 600 cm⁻¹ to 640 cm⁻¹ are identified as due to v_4 of carbonate ion.

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208 Raman spectroscopy

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210 The low energy Raman bands in the range 3740-3600 cm⁻¹ observed for this compound are

shown in Fig 8. The bands in this region are often attributable to water fundamentals and Mg-

212 O/Fe-O lattice vibrations. The observed Raman energies of the bands are similar to the

reported bands [2].

- 214
- 215 Conclusions
- 216

217	1. Natural iowaite crystal structure is similar to synthetic iowaite and is compared with
218	already reported date which is similar.
219	
220	2. EPR results are indicating that Fe(III) and Mn(II) are present in the mineral in distorted
221	octahedral environment. The $g = 4.279$, 3.911 and 2.0 are ascribed to Fe(III) and g and A
222	value observed in the spectrum are 2.007 and 9.04 mT are due to Mn(II) in the mineral.
223	This results are supporting the molecular formula that the sample contains Fe(III) and
224	Mn(II) impurity might have entered the lattice in place of Mg(II).
225	
226	3. The optical absorption studies are indicating that only Fe(III) ion in mineral is present in
227	distorted octahedral environment.
228	
229	4. Near infrared spectrum is due to overtones and combinations of water molecules and
230	carbonate ions. Further IR studies are due to water and carbonate ions in the mineral.
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232	5. Raman spectrum is compared with already reported data and is identical.
233	
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Fig.1. ESR spectrum of natural iowaite at room temperature (v = 9.41095 GHz)
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Fig.2 Optical absorption spectrum of natural iowaite at room temperature











Fig. 4. NIR spectrum of natural iowaite in the range 6000 to 4000 cm⁻¹







321 Fig. 7 IR Spectrum of iowaite at room temperature from 1125 to 525 cm⁻¹

