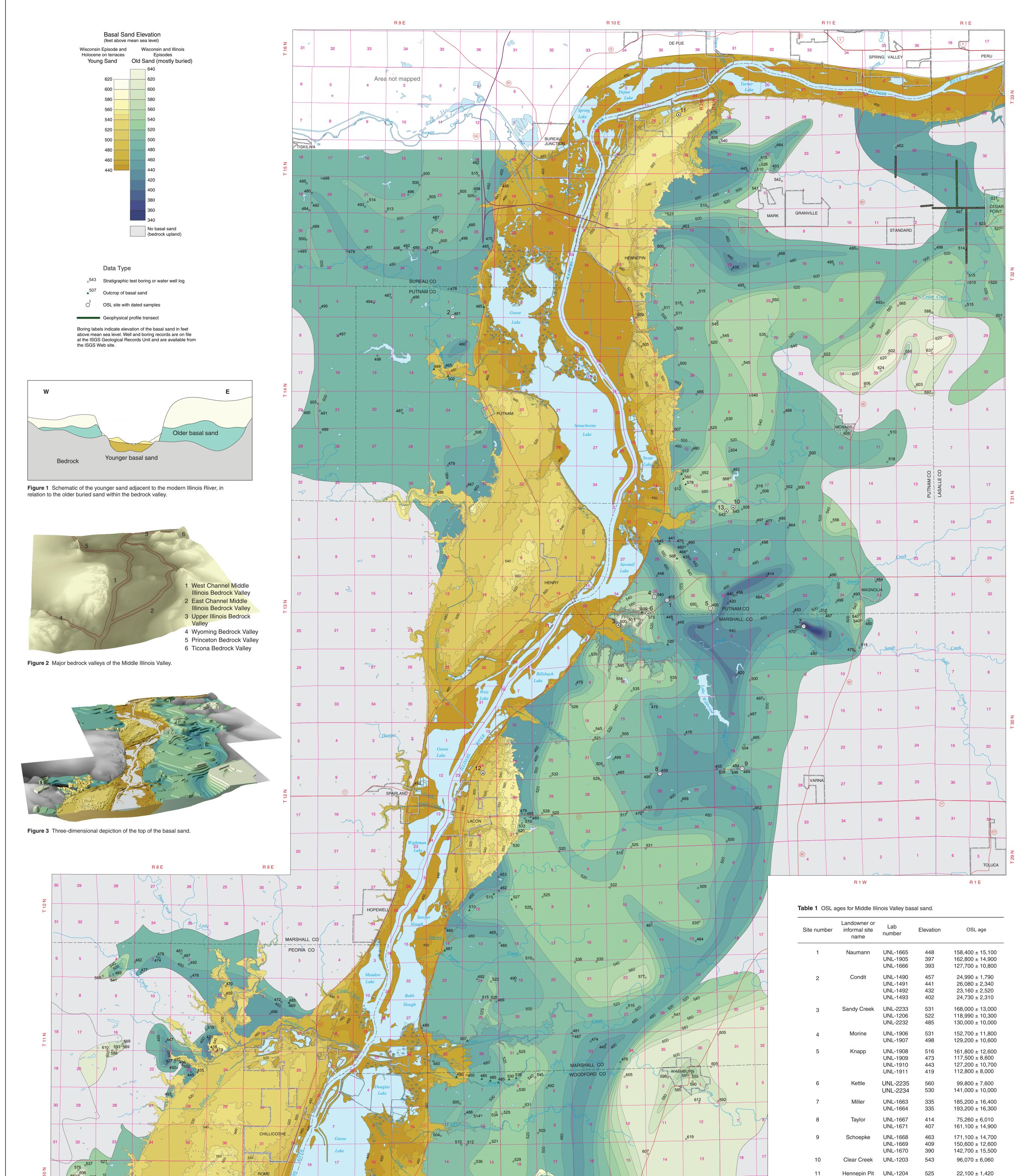
Illinois Map 17

2012

ELEVATION OF THE BASAL SAND AND GRAVEL OF THE MIDDLE ILLINOIS RIVER VALLEY

Bureau, LaSalle, Marshall, Peoria, Putnam, and Woodford Counties, Illinois

Richard C. Berg, E. Donald McKay III, and Barbara J. Stiff 2012



Base map compiled by Illinois State Geological Survey from digital and paper data provided by the United States Geological Survey.

North American Datum of 1983 (NAD 83)

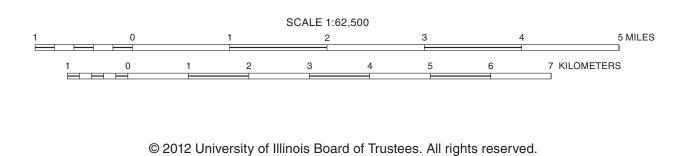
Projection: Transverse Mercator

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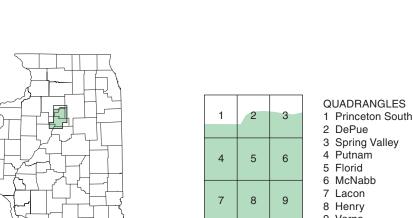
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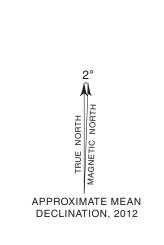
10 Rome

11 Chillicothe

12 Washburn

R 3 W









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19,450 ± 1,450

 $93,000 \pm 8,800$

Figure 4 Basal sand and gravel exposed at Site 3 on the map (looking to the north). Note southeasterly dipping beds near the bottom of the slope.

Introduction

The top elevation of the basal sand and gravel is depicted for the middle Illinois River region from east of the "Big Bend" of the river near Hennepin to south of Chillicothe. This sand and gravel resides directly above the bedrock surface, is often greater than 100 feet thick, and constitutes a major drinking water resource for the region. This map provides insight into the erosional and depositional history of the ancient Mississippi River (AMR), the ancient Illinois River, and the modern Illinois River. The topography of the sand and gravel deposit reflects numerous periods of sand and gravel deposition and erosion associated with glacial advances and retreats across the region and associated glacial-fluvial and modern fluvial events. This map is an essential precursor to an aquifer sensitivity map for the region, which will be based on the depth to and thickness of this basal deposit.

This map covers over 500 square miles and includes nine 1:24,000-scale U.S. Geological Survey 7.5-minute quadrangles—Putnam, Florid, McNabb, Lacon, Henry, Varna, Rome, Chillicothe, and Washburn—and the southern half of three quadrangles—Princeton South, DePue, and Spring Valley. Detailed mapping of the basal sand and gravel in this region was an outgrowth of geologic mapping for a proposed highway improvement project along Illinois Route 29, funded by the Illinois Department of Transportation, on the west side of the Illinois River north of Chillicothe (Berg et al. 2002, 2003). This map is part of a series that includes surficial geology (McKay et al. 2010), bedrock topography (Berg et al. 2009), and drift thickness (in preparation) maps of the same region along the middle Illinois River valley (MIV).

Methodology

The basal sand and gravel is divided into two categories (fig. 1):

1. Sand and gravel is exposed at land surface on terraces and on the floodplain along

the modern Illinois River and its major tributaries.

2. Older sand and gravel, buried by younger glacial deposits (diamictons, silts and clays, other thinner sands and gravels, and paleosols), occur east of the main Illinois River valley beneath uplands and also above the Princeton Bedrock Valley, Ticona Bedrock Valley, Wyoming Bedrock Valley, and the east channel of the Middle Illinois Bedrock Valley (fig. 2).

The elevation of the younger sand and gravel was determined from land surface contours depicted on topographic maps. The areal extent is bounded by the scarp of the modern Illinois River valley and the walls surrounding the mouths of the major tributaries. Figure 1 shows the relationship of this younger sand and gravel to the older buried sand and gravel within the bedrock valley. Figure 3 is a three-dimensional depiction of the surface of the basal sand and gravel within its bedrock valley.

For the older surface, the top elevation of the buried basal sand and gravel was determined by evaluating logs of borings and by seismic profiling. A total of 1,497 logs of water wells, engineering borings, and coal test borings as well as 23 ISGS exploratory borings and some field-described outcrops were evaluated to determine the top of this basal sand and gravel. These data are on file in the Illinois State Geological Survey (ISGS) Geological Records Unit. In the northeastern portion of the map, seismic reflection profiles, recorded along 5.15 miles of roads, were used to determine the top of the basal sand and gravel within the Ticona Bedrock Valley (Murphey 2005). Logs of nearby water wells were used as a basis for estimating seismic velocities of glacial-fluvial deposits, and travel times measured at selected points (~100 feet apart) along the transect were used to estimate the top of the basal sand and gravel.

The approximate age of the deposit, including the sand on the younger terraces, was determined for 28 samples using optically stimulated luminescence (OSL) of fluvial quartz sand (McKay and Berg 2008) obtained at the 13 sites shown on the map. Sites consisted of outcrops and cores of ISGS test borings (table 1).

To supplement borehole and seismic information, numerous cross sections were constructed to best visualize the continuity of the basal sand and gravel between logs throughout this large MIV region. The elevation of the buried surface was hand-contoured to conform with borehole data. The elevation contours of the buried sand and gravel surface were clipped along the scarp of the modern Illinois River valley and merged with the elevation contours of the younger sand and gravel deposit to produce the surface elevation contours for all MIV sand and gravel deposits.

Observations

Recent geologic mapping and OSL determinations along the MIV have provided new data to better characterize Wisconsin Episode and older sediments associated with glacial and meltwater events that affected the valley several times during the last several hundred thousand years. The AMR reoccupied the valley after each glacial retreat until the river was blocked and diverted by a glacier to its present Mississippi River course at $24,770 \pm 250$ calendar years BP (McKay et al. 2008). This date was derived from a 14 C sample at the top of lake sediment from a core on the east side of the river (site 1 on map). Burial of the AMR valley, first by the basal sand and gravel, followed by Illinois and Wisconsin Episode glacial and glacial-fluvial sediments, and then by postglacial sand and gravel, silt, and clay reflects the complexity of erosion and sedimentation associated with these events. Particularly, the buried topographic expression of the basal sand and gravel provides a first glimpse of the morphology of a patchwork of different aged surfaces.

These different aged surfaces include the following:

1. A prominent buried "terrace" resides at $480{\text -}500$ feet in the northwestern portion of the map area, west of the Illinois River and above the Princeton Bedrock Valley. Four OSL determinations of the buried basal sand and gravel from an ISGS test drilling site (site 2) reveal dates with a mean weighted average of $24{,}800 \pm 1{,}100$ years BP, which almost perfectly matches the 14 C date of AMR diversion from site 1 east of the river. This surface is the floodplain of the AMR just prior to its diversion and subsequent burial by Wisconsin Episode deposits. The surface may have some expression above the Wyoming Bedrock Valley in the southwestern portion of the map and also just east of the main Illinois River valley in the southern portion of the map. However, there are no confirmatory OSL ages for these regions.

2. A prominent buried sand and gravel upland is above an elevation of 500 feet east of the Illinois River. This upland, intersected by channels, extends north-south for about 25 miles and is roughly 1 to 8 miles wide. In the central portion of the map area, the upland is commonly more than 540 feet in elevation and rises, perhaps reflective of a "high terrace remnant," to more than 600 feet at site 3 on the north side of present-day Sandy Creek. Eleven OSL determinations from sites 3, 4, 5, and 6 suggest that this surface is a combination of late Illinois Episode (OSL determination of 141,000 \pm 10,000 years BP from a lower sample at site 6) and early Sangamon Episode ages. The latter is perhaps reflective of an interglacial AMR with basal sand and gravel ages of $118,990 \pm 10,300$ to $130,000 \pm 10,000$ years BP at site 3, $152,700 \pm 11,800$ years BP at site 4, $112,800 \pm 8,000$ to $161,800 \pm 12,600$ years BP at site 5, and $99,800 \pm 7,600$ years BP at site 6. The OSL dates at sites 10 and 13 of $96,070 \pm 6,060$ years BP and $93,000 \pm 8,800$ years BP, respectively, also suggests a Sangamon age. Some of the uppermost samples at these sites revealed older ages and are inconclusive. The Sangamon Geosol appears to be pervasive over this surface.

3. A 600-foot "high terrace remnant" is observed elsewhere on the map: (a) in the northeast portion about 5 miles south-southwest of the buried Ticona Bedrock Valley, (b) in the extreme southeastern corner, and (c) in the southwestern corner of the map as two small (<1 mi²) remnants. The relationship of these surfaces to one another is unknown as no OSL ages have been determined. Also data are sparse, and thin diamicton may underlie the basal sand and gravel in some portions of these

4. A channel cut into the buried basal sand and gravel upland surface at an elevation of less than 480 feet is observable east of the river. The channel has continuity from north to south. However, at its northern end, the channel seems to be possibly filled with younger deposits constituting the eventual 500+-foot "terrace" as described in the second item of this list. Ten OSL determinations from sites 1, 7, 8, and 9 suggest that the deposit is early Illinois Episode. At site 1, the uppermost sand has an OSL age of $158,400 \pm 15,100$ years BP. At site 7, two OSL ages are $185,200 \pm 16,400$ years BP and $193,200 \pm 16,300$ years BP. These are the oldest ages for sand that have been dated in the MIV, suggesting that the AMR first occupied this portion of the valley before it began a migration westward. At site 8, an uppermost OSL age on fine sand of $75,260 \pm 6,010$ years BP suggests a Sangamon Episode alluvial fill succession overlying older sands dated at $161,100 \pm 14,900$ years BP. The OSL ages at site 9 are $142,700 \pm 15,500$, $150,600 \pm 12,600$, and $171,100 \pm 14,700$ years BP.

Unexplainable are lower elevation portions of the channel, particularly at site 7 where an ISGS test hole revealed the top of the basal sand and gravel at an elevation of 335 feet, overlain by thick Illinois Episode diamictons. Channel scour below an elevation of 460 feet is observable throughout its northern and southern portions. The northwest to southeast channel trend in the northern portion and its alignment with the main AMR channel west of the river above the Princeton Bedrock Valley suggest that the AMR flowed on this eastern side of the valley and could be responsible for the scouring. Southeasterly, steeply dipping, and cemented beds of sand and gravel near the base of site 3 (fig. 4) north of Sandy Creek also suggest major river flow toward the southeast. The southern portion of the buried basal sand and gravel channel directly overlies a channel carved into the bedrock surface (Berg et al. 2009), suggesting that a channel has existed for a long time or perhaps re-occupied itself on the east side of the valley.

Secondary channels are also observable incised into the buried basal sand and gravel surface. The distribution of basal sand and gravel and a channel in the northern portion of the map suggest that there was flow from the Ticona River to the AMR. Two miles southeast of site 8 is a subtle channel below an elevation of 520 feet that perhaps drained the main channel westward.

5. Following the diversion of the AMR to its present Mississippi River course, the modern Illinois River was established as it drained lands from the north and east within the mapped area. Terraces up to elevations of 560 feet were constructed, and the modern floodplain is at an elevation of 450 feet. Two OSL determinations were made on terrace sands from sites 11 and 12. Sand from site 11 at the Hennepin pit had an OSL age of $22,100 \pm 1,420$ years BP, whereas that from the Lacon Pit at site 12 had an OSL age of $19,450 \pm 1,450$ years BP. Both are reflective of maximum Wisconsin Episode meltwater flowing down the Illinois River.

Acknowledgments

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